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Analyzing impacts of seasonality and landscape gradient on event-scale nitrate-discharge dynamics based on nested high-frequency monitoring

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Highlights

- Spatial heterogeneity of nitrate storage can affect patterns of C-Q relationships
- Hydrological connectivity on land determined seasonally varying nitrate exports
- Nested monitoring is needed to explore C-Q mechanisms in heterogeneous catchments
- Analysis of C-Q hysteresis pattern allow targeting agricultural management actions

1 Abstract

2 Increasingly available high-frequency data during storm events, when hydrological dynamics most 3 likely activate nitrate storage-flux exchanges, reveal insights into catchment nitrate dynamics. In this 4 study, we explored impacts of seasonality and landscape gradients on nitrate concentration-discharge 5 (C-Q) hysteresis patterns in the Selke catchment, central Germany, which has heterogeneous 6 combinations of meteorological, hydrogeological and land use conditions. Three nested gauging 7 stations established along the main Selke River captured flow and nitrate export dynamics from the 8 uppermost subcatchment (mixed forest and arable land), middle subcatchment (pure steep forest) and 9 lowermost subcatchment (arable and urban land). We collected continuous high-frequency (15-min) 10 discharge and nitrate concentration data from 2012-2017 and analyzed the 223 events detected at all 11 three stations. A dominant hysteresis pattern in the uppermost and middle subcatchments was counter-12 clockwise and combined with an accretion effect, indicating many proximal and mobilized distal 13 nitrate sources. However, 66% of all events at the catchment outlet experienced a dilution effect, 14 possibly due to mechanisms that vary seasonally. During wetting/wet periods (October-March), it was 15 combined mainly with a counter-clockwise pattern due to the dominance of event runoff volume from 16 the uppermost and middle subcatchments. During drying/dry periods (April-September), however, it 17 was combined mainly with a clockwise pattern due to occasional quick surface flows from lowland 18 near-stream urban areas. In addition, the clockwise hysteresis occurred mainly from May-October 19 during mostly drying/dry periods at all three sites, indicating little distal nitrate transport in response 20 to the low terrestrial hydrological connectivity, especially in the lowermost dry and flat subcatchment. 21 This comprehensive analysis (i.e., clockwise vs. counter-clockwise, accretion vs. dilution) enables in-22 depth analysis of nitrate export mechanisms during certain periods under different landscape 23 conditions. Specific combination of C-Q relationships could identify target locations for agricultural 24 management actions that decrease nitrate output. Therefore, we strongly encourage long-term 25 multisite and high-frequency monitoring strategies in heterogeneous nested catchment(s), which can 26 help understand process mechanisms, generate data for physical-based water-quality modeling and 27 provide guidance for water and agricultural management.

- 28 Key words: nitrate export dynamic, C-Q relationship, hysteresis pattern, high-frequency data,
- 29 landscape effect, seasonality effect

30 1. Introduction

31 Human activities (e.g., intensive agriculture, urbanization, deforestation) have altered the natural 32 landscape extensively and hence influenced nitrogen (N) cycling greatly (Boyer et al., 2002; Howarth et al., 2012). The large external supplies of N clearly exceed terrestrial N demands for plant/crop 33 34 growth and microbial transformation (Davidson et al., 2011). Driven by hydrological dynamics, the excess terrestrial N has been exported to surface waters and redistributed spatially and temporally 35 throughout fresh/coastal water systems (Reusch et al., 2018; Shields et al., 2008). Mitigation measures 36 have been established according to guidelines of multiple government conventions (e.g., the European 37 Union Water Framework Directive). Although diffuse nitrate pollution has been ameliorated, it 38 remains a main cause of freshwater quality degradation (EEA, 2018). Pursuing more cost-effective 39 40 measures requires better mechanistic understanding of catchment nitrate dynamics, especially in the 41 context of contrasting landscape conditions (both natural and human) and strong seasonal variability. 42 Flow and nitrate dynamics during storm events are more active due to changes in storage-flux 43 interactions and transport pathways, compared to those during hydrologically stable conditions (e.g., 44 low flow, dry periods). Therefore, the event-scale relationship between nitrate concentration and 45 discharge (C-Q relationship) has been investigated intensively to determine spatial and temporal 46 variability in catchment nitrate functioning (Baker and Showers, 2019; Dupas et al., 2016; Zimmer et al., 2019). Hysteresis is the most commonly observed pattern of the C-Q relationship (Burns et al., 47 48 2019). Hysteresis patterns vary spatially and temporally due to variable combinations of nitrate 49 sources (Bowes et al., 2015) and hydrological drivers (Vaughan et al., 2017). Celerity is well known 50 to be faster than particle transport velocity in catchment hydrology (Cheraghi et al., 2016; McDonnell and Beven, 2014; Williams et al., 2018). Therefore, proximal nitrate storages generally respond faster 51 than distal storages along the formation of hydrograph at the catchment scale, resulting in different 52 53 hysteresis loops of the C-Q relationship (i.e., clockwise vs counter-clockwise). Meanwhile, nitrate storage varies vertically (along the soil profile) and horizontally under different landscape 54 55 characteristics and anthropogenic conditions (Dupas et al., 2016; Musolff et al., 2016; Miller et al., 2017). Driven by flow generations, the mobilized terrestrial nitrate may further result in negative or 56

57 positive hysteresis slopes for stream water, representing dilution or accretion effects, respectively. In 58 turn, hysteresis analysis based on comprehensive monitoring data permits detailed explorations of the varying flow and nitrate dynamics. Continuous high-frequency data under various hydro-climatic 59 60 conditions offers the opportunity to evaluate the changes of runoff partitioning and biogeochemical 61 processes, as well as their impacts on nitrate mobilizations at multiple spatial scales (e.g., for the 62 catchment-wide scale and the local near stream scale) (Carey et al., 2014; Vaughan et al., 2017). 63 Intensive monitoring across contrasting landscape characteristics further enables detailed analysis of the interplay between heterogeneous landscape features and varying flow pathways (Fovet et al., 64 65 2018; Musolff et al., 2015; Williams et al., 2018).

66 Landscape characteristics reflect regional climate patterns, general pedological and geological 67 properties and human impacts. Therefore, the spatial heterogeneity of landscape characteristics 68 determines the spatial distribution of nitrate source areas and variable catchment mechanisms of 69 hydrology and nutrient transport (Dupas et al., 2017; Poor and McDonnell, 2007). The long history of 70 commercial fertilizer application has increased agricultural production but has also accumulated 71 excess N in terrestrial soils (Outram et al., 2016). N sources from agricultural lands have become one 72 of the main pollution sources in most rivers and caused high risks to aquatic ecosystems (EEA, 2019). 73 In forest areas, nitrate leaching likely depends on the amount of nitrate in throughfall and the C:N 74 ratio of the organic soil horizon (Borken and Matzner, 2004; MacDonald et al., 2002). Therefore, a 75 synchronous dynamic pattern between discharge and nitrate concentration is commonly observed in 76 forest catchments. Artificial N is added mainly via atmospheric wet and dry deposition (e.g., ranging from 1-60 kg N ha⁻¹ yr⁻¹), which has also been increased greatly in the past few decades (MacDonald 77 78 et al., 2002). In urban areas, extensive paved areas and artificial drainage networks can strongly alter 79 natural processes of water movement and nitrate transport (Miller et al., 2014). For example, 80 artificially drained flow easily bypasses the nitrate-rich soil and responds quickly, even under small precipitation events. Therefore, the process of N export can be misunderstood if the heterogeneity of 81 82 catchment landscape characteristics is not considered, especially in nonuniform and nested 83 catchments.

84 Driven by seasonal variations in meteorological and hydrological conditions, terrestrial export of nitrate always accompanies the changes of runoff components, and therefore, the surface nitrate C-Q 85 relationship shows strong seasonality (Sickman et al., 2003). Different runoff components (i.e., 86 surface flow, interflow and baseflow) usually have different nitrate concentrations due to their 87 88 differing degrees of interactions with soil N sources (Miller et al., 2017). Therefore, the seasonally varying characteristic of runoff partitioning can alter the C-Q relationship of specific events 89 90 considerably. Recent researches about C-Q relationship most focus on humid areas (Jacobs et al., 91 2018; Vaughan et al., 2017; Zimmer et al., 2019), where interflow plays an important role in transporting nitrate sources during events. Studies of process-based understanding of temporal nitrate 92 dynamics in dry area are still rare (Dupas et al., 2016). For example, interflow and baseflow are 93 considered the dominant runoff components during wet and dry periods, respectively, in the well-94 95 monitored Selke catchment in central Germany (Yang et al., 2018). However, quick surface flow from paved area and artificial drainage (both with relatively low nitrate concentrations) can also occur 96 97 intermittently in the lowland arable/urban area during small events and cause different C-Q 98 relationships at the outlet. The interplay among different runoff components and their effects on 99 nitrate dynamics in dry area are hence in need of improved understanding. Moreover, seasonal 100 biogeochemical processes can also influence the nitrate legacy at the catchment scale. In Western 101 Europe, winter-spring high-flow periods experience high soil moisture, which activates hydrological 102 connections between terrestrial and aquatic systems (Molenat et al., 2008; Strohmenger et al., 2020), 103 and relatively low temperatures, which do not stimulate much biogeochemical turnover (Allen et al., 104 2002). In contrast, summer-autumn high-temperature growing seasons cause high soil evaporation and 105 plant/crop transpiration, which result in low soil moisture that restricts hydrological connections 106 (Bracken and Croke, 2007) and stimulates biogeochemical transformations of N in terrestrial and instream phases (Racchetti et al., 2011; Rode et al., 2016a). Hence, these seasonal hydrological and 107 108 biogeochemical processes can characterize the variation in the C-Q relationship. 109 Overall, the mechanistic interactions between flow and nitrate dynamics at the event scale vary

spatially and temporally. Comprehensive monitoring datasets for highly heterogeneous catchments are

111 rare, but they are essential to reveal effects of landscape heterogeneity and seasonality on C-Q 112 relationships. In this study, we focused on the well-monitored Selke catchment (a subcatchment of the Terrestrial Environmental Observatories (TERENO) – Harz/Central German Lowland Observatory) 113 (Wollschläger et al., 2016; Zacharias et al., 2011). Three nested gauging stations along the main Selke 114 115 River capture the variety of catchment responses of flow and nitrate processes (Jiang et al., 2019; Yang et al., 2018). High-frequency multi-parameter sensors have been continuously deployed at each 116 117 station, ranging from upper forest area to lowland agricultural area (Rode et al., 2016a). Here we 118 collected discharge and nitrate-N concentration data that were continuously monitored during 2012-119 2017 at a 15-min interval. The objectives of this study were to (1) quantify event-scale C-O relationships among heterogeneous conditions in the nested Selke catchment, (2) analyze impacts of 120 deviating hydrological and landscape characteristics on hysteresis patterns based on subcatchments 121 122 discrepancies, and (3) investigate seasonal variability of hysteresis patterns given the contrasting wetdry conditions. With this study we show how nitrate fluxes are generated in heterogeneous 123 124 subcatchments and how the interplay of these subcatchments can modulate C-Q relationships at 125 varying seasonal conditions and event magnitudes at the whole catchment scale.

126 2. Data and methods

127 2.1. Study area and data collection

The Selke catchment (456 km²) is located in the transition area between the northern German plain 128 129 and central German uplands. The elevation ranges from ca. 590 m in the upper Harz mountain region to ca. 100 m in the lowland region (Figure 1a), with mean annual precipitation decreasing from 790 to 130 450 mm, respectively (Yang et al., 2019). Three gauging stations set up from upstream to downstream 131 (i.e., Silberhuette (SILB), Meisdorf (MEIS) and Hausneindorf (HAUS)) (Figure 1a) capture responses 132 of the heterogeneous catchments (Rode et al., 2016a). The drainage areas of the three stations are 99, 133 184, and 456 km², respectively. The uppermost and middle subcatchments lie in the Harz 134 mountainous region, which is dominated by shallow and relatively impervious schist and claystone 135 136 overlain mainly by cambisols. In contrast, the lowermost subcatchment lies in the unique central German loess-chernozem region, which has deep tertiary sedimentary rocks. Therefore, the catchment 137

has high gradients of landscape characteristics, including meteorology, hydrology, biogeochemistry
and anthropogenic impacts (Yang et al., 2019). The uppermost subcatchment is covered by wellmixed forest (60%) and agricultural (25%) areas, while most (85%) of the middle subcatchment is
covered by pure steep forest (Figure 1b). Due to the high fertility of chernozems, the lowermost
subcatchment is extensively and intensively cultivated as arable land (ca. 80%) and contains
considerable urban areas.

144 Both discharge and nitrate-N $(NO_3^- - N)$ concentration are continuously measured at the three gauging stations. We collected high-frequency (15-min interval) data from 2012-2017 for the event-145 146 scale analysis. Discharge data were provided by the State Agency for Flood Protection and Water Management of Saxony-Anhalt (LHW). $NO_3^- - N$ data were provided by the Helmholtz Center for 147 Environmental Research-UFZ, using a TRIOS ProPS-UV sensor with an optical path length of 10 148 149 mm. The sensor data were validated by biweekly parallel grab samples. For more information about 150 the high-frequency monitoring and maintenance, please refer to Rode et al. (2016b). In addition, longterm daily discharge and biweekly grab-sampled $NO_3^- - N$ data (1994 - 2011) from LHW were also 151 collected for a long-term overview of flow and $NO_3^- - N$ concentration dynamics. 152



Figure 1. (a) Elevation and (b) land use in the Selke catchment and locations of the three gaugingstations.

156 2.2. Variability in event-scale nitrate dynamics

157 2.2.1. Detecting storm events

Storm events were detected first based on automatic identification of all local maxima and minima of 158 the discharge time series. Each local maximum was considered to be the discharge peak of each event, 159 160 and the closest minima before and after the peak were selected as the preliminary start- and end-time of the event, respectively. Next, the final start- and end-times of each event were manually adjusted to 161 ensure that they had a similar discharge. Successive events without a complete recession between 162 them were merged into a single event with multiple peaks. Events with missing data (> 10%) were 163 164 excluded from further analysis. Events were detected using scripts in R software (R Core Team, 2020). 165

166 2.2.2. Calculation of hysteresis patterns

168

167 For each event, discharge and $NO_3^- - N$ were normalized following Lloyd et al. (2016a):

$$Q_{t,norm} = \frac{Q_t - Q_{min}}{Q_{max} - Q_{min}} \tag{1}$$

169
$$N_{t,norm} = \frac{N_t - N_{min}}{N_{max} - N_{min}}$$
(2)

170 where Q_t and N_t are the discharge (m³ s⁻¹) and $NO_3^- - N$ (mg l⁻¹) concentration measured at time *t*, 171 $Q_{t,norm}$ and $N_{t,norm}$ are the normalized discharge and $NO_3^- - N$ concentration, and the subscripts 172 'min' and 'max' are the minimum and maximum values of each event.

173 To quantify the hysteresis pattern of the C-Q relationship, two indices were calculated from the

174 normalized data. First, the non-dimensional hysteresis index (*HI*) was calculated as:

$$HI = \int N_{t,norm} \cdot dQ_{t,norm} \tag{3}$$

where *HI* equals the sum of hysteresis effects of the C-Q relationship during each event period (Zhang et al., 2017). *HI* ranges from -1 to 1 (*HI* > 0 indicates clockwise hysteresis, while *HI* < 0 indicates counter-clockwise hysteresis). Second, the concentration-changed index (*CI*), following Butturini et al. (2008), was calculated as:

$$CI = N_{tp,norm} - N_{ts,norm} \tag{4}$$

181 where $N_{tp,norm}$ and $N_{ts,norm}$ are the normalized $NO_3^- - N$ concentration at the discharge peak and 182 start-time of each event, respectively. *CI* ranges from -1 to 1 (*CI* > 0 indicates an accretion effect of 183 $NO_3^- - N$ concentration following flow dynamics, while *CI* < 0 indicates a dilution effect). Note that 184 if the peak discharge lasted for more than one measured time point, the first time point $NO_3^- - N$ 185 concentration was chosen as $N_{tp,norm}$.

186 2.2.3. Shared event analysis and statistic methods

Several "shared events" were specifically analyzed based on the start-, end- and discharge peak time points of each event at the SILB, MEIS, and HAUS stations. These events propagated from upstream to downstream and were detected simultaneously at all three stations. Then, the start- and end-times of each shared event were adjusted slightly to encompass the entire event duration at all three stations (i.e., using the latest start-time and the earliest end-time). Nitrate-N load (N_L , unit: kg) and runoff volume (R_V , unit: m³) of each shared event at each station were calculated using the following two equations, respectively:

194
$$N_L = \int (Q_t \cdot N_t) \cdot dt \tag{5}$$

$$R_V = \int Q_t \cdot dt \tag{6}$$

where Q_t and N_t are the measured discharge (m³ s⁻¹) and $NO_3^- - N$ concentration (mg l⁻¹), 196 respectively, during the shared period. Using values of N_L and R_V from the three nested stations, 197 nitrate-N load and runoff volume were calculated for the uppermost subcatchment (subscript U_{P}) as 198 199 the values measured at the SILB station, for the middle subcatchment (subscript 'MID') as the values measured at the MEIS station minus those at the SILB station and for the lowermost subcatchment 200 201 (subscript 'LOW') as the values measured at the HAUS station minus those at the MEIS station. In addition, nitrate-N load and runoff volume of the entire catchment (subscript 'ALL') were considered as 202 203 the values measured at the HAUS station. 204 The nonparametric Wilcoxon signed rank test and Kruskal-Wallis test were used to detect significant

205 differences in population medians of paired or multiple categories, respectively (Kruskal and Wallis,

206 1952; Wilcoxon, 1945). The distributions of samples were considered significantly different when the

207 *p*-value was below the level of significance of 0.05.

208 3. Results

209 3.1. Overview of long-term dynamics

In the long-term data (1994-2017), runoff volume increased disproportionately from upstream to 210 downstream stations; for example, although the uppermost subcatchment covered only 22% of the 211 catchment, it contributed 64% of its total mean annual runoff volume (i.e., 3.31×10^7 and 5.15×10^7 212 m³ yr⁻¹ at the SILB and HAUS stations, respectively). The spatial contribution of mean annual NO_3^- – 213 N load generally followed that of flow volume (e.g., 68 t yr⁻¹ at the SILB station was 43% of the total 214 at the HAUS station). Differences in runoff volume and $NO_3^- - N$ load among subcatchments were 215 due to spatial variability in $NO_3^- - N$ concentrations. From the SILB station to the HAUS station, 216 median $NO_3^- - N$ concentration increased significantly (from 1.01 to 2.68 mg l⁻¹, Wilcoxon signed 217 218 rank test), while median specific runoff decreased significantly (from 169 to 67 mm yr⁻¹, Wilcoxon 219 signed rank test).

220 Despite the high spatial variability, discharge and $NO_3^- - N$ concentration showed strong seasonal patterns (Figure 2). Based on the general hydro-climatic cycling (Figure 2a), we categorized the 221 222 hydrological year into four periods: wetting (October-December, with a continuous increase in mean monthly discharge from 0.90 to 1.80 m³ s⁻¹ at the HAUS station), wet (January-March, with high mean 223 224 discharge of 2.96 m³ s⁻¹), drying (April-June, with mean monthly discharge decreasing from 2.42 to 0.99 m³ s⁻¹) and dry (July-September, with consistently low mean discharge of 0.58 m³ s⁻¹). The 225 226 spatial distribution of flow varied greatly among the hydrological periods. The uppermost 227 subcatchment generated most of the catchment's runoff volume during the wet period but much less 228 during the dry period (e.g., mean discharge in February and September was 70% and 53% that at HAUS, respectively) (Figure 2a). The seasonal pattern of $NO_3^- - N$ was similar to that of discharge at 229 the SILB and MEIS stations, with high concentrations (e.g., $> 3 \text{ mg } l^{-1}$) during the wet period that 230 231 gradually decreased during the drying period, and much lower concentrations (e.g., $< 1 \text{ mg } l^{-1}$) during the dry period (Figure 2b). However, $NO_3^- - N$ at the HAUS station had consistently high 232 concentrations throughout the hydrological cycle (i.e., generally > 2 mg l⁻¹). Mean $NO_3^- - N$ 233 concentration during the dry period was significantly higher at HAUS (2.60 mg l⁻¹) than at SILB (0.58 234 mg l⁻¹). The high mean concentrations at HAUS, much higher than those at MEIS and SILB, were 235 236 likely caused by urban point-source contributions before 2002 (Yang et al., 2018). Nonetheless, the mean monthly concentration decreased slightly from the beginning of the drying period but remained 237 238 much higher than those at MEIS and SILB.



Figure 2. Boxplots of monthly (a) discharge and (b) nitrate-N concentrations at the SILB, MEIS and
HAUS stations from 1994-2017. Outliers were omitted, and red diamond markers represent mean
values. Whiskers represent 1.5 times the interquartile range.

243 **3.2. Event detection**

244 We analyzed 81, 72, and 70 detected events at the SILB, MEIS and HAUS stations, respectively, from 245 2012-2017 (Figure 3, Supplementary Table S1). Events were evenly distributed, while their magnitude and duration varied greatly during the four hydrological periods among the three stations 246 247 (Table 1). Event durations were generally much longer during wetting/wet periods than during 248 drying/dry periods (e.g., mean durations at the SILB station were 11.79 and 4.07 days during the wet and dry periods, respectively). Similarly, event-scale mean discharge and $NO_3^- - N$ had the highest 249 250 values during the wet period and the lowest values during the dry period at the three stations (Table 251 1).





Figure 3. Discharge (Q) and $NO_3^- - N$ concentrations at 15-minute intervals from 2012-2017 at the (a) SILB, (b) MEIS and (c) HAUS stations. A total of 81, 72, and 70 storm events (shaded areas) remained after manual adjustment, respectively.

Table 1. Number of storm events, mean duration, mean discharge (Q) and mean $NO_3^- - N$

concentration during each hydrological period at the SILB, MEIS and HAUS stations.

	SILB					ME	EIS		HAUS				
Periods	Wetting	Wet	Drying	Dry	Wetting	Wet	Drying	Dry	Wetting	Wet	Drying	Dry	
Number	25	16	20	20	17	18	20	17	20	15	18	17	
Duration (days)	9.5	11.8	5.1	4.1	11.0	11.0	4.7	5.4	9.2	11.6	5.4	4.8	
$Q(m^3 s^{-1})$	1.12	1.95	1.42	0.56	1.73	2.56	2.42	0.85	1.93	3.29	2.69	1.21	
$NO_3^{-}-N (\text{mg } l^{-1})$	2.29	3.38	1.30	0.85	2.42	3.15	1.24	0.98	2.49	3.33	2.16	2.04	

259 **3.3. Hysteresis pattern analysis**

260 The patterns of HI and CI showed both spatial and seasonal variations. Negative HI (i.e., counterclockwise hysteresis) dominated at the SILB, MEIS and HAUS stations (i.e., ca. 78%, 88% and 63% 261 of all events, respectively) (Figure 4a). Most events with positive HI (i.e., clockwise hysteresis) 262 263 occurred during drying/dry periods (78%, 67% and 88% at the SILB, MEIS and HAUS stations, respectively). The median HI was significantly different among four hydrological periods at each 264 265 station (Kruskal-Wallis test). The trend for CI differed from that of HI. Positive CI (i.e., accretion 266 effect) dominated in the uppermost and middle subcatchments (i.e., 89% and 79% of all events at the SILB and MEIS stations, respectively) (Figure 4b). However, 66% of events at the HAUS station had 267 268 negative CI (i.e., dilution effect). The median CI between SILB and HAUS, and between MEIS and HAUS were significantly different (Wilcoxon signed rank test). Mean CI was positive during all four 269 270 hydrological periods at the SILB and MEIS stations but was positive only during the wet period at the HAUS station (Figure 4b). 271





Figure 4. Boxplots of (a) hysteresis index (*HI*) and (b) concentration-change index (*CI*) of storm
events during four hydrological periods at the SILB, MEIS and HAUS stations. Red points represent
means. Whiskers represent 1.5 times the interquartile range.

276 We combined the two hysteresis indices and categorized all events into four categories (Figure 5).

277 Event-scale hysteresis patterns varied greatly among the hydrological periods and landscape

- 278 conditions (Figure 6). At the SILB station, the general hysteresis pattern was negative HI combined
- with positive CI (ca. 72%, Figure 6). The events in this category had the longest duration and highest

total precipitation (Table 2). The pattern of positive *HI* combined with positive *CI* accounted for 17%

of events at SILB, most of which occurred during drying/dry periods. Among events with negative CI,

those combined with negative HI occurred during wetting/wet periods (Figure 6). In contrast, events

283 with positive HI occurred during drying/dry periods and had the shortest duration, lowest total

- 284 precipitation and lowest discharge and $NO_3^- N$ concentration (Table 2).
- At the MEIS station, the general hysteresis pattern was the same as at the SILB station (i.e., 71% of

events had negative *HI* with positive *CI*), but the percentage of positive *HI* with positive *CI* decreased

to 8%, indicating an overall lower accretion effect (Table 2). Patterns of negative HI with negative CI

increased to 17% of events at MEIS, with a relatively short duration and low discharge (Table 2). The

pattern of positive *HI* with negative *CI* was least common and was evenly distributed among the

wetting, drying and dry periods (Figure 6). Similarly, this pattern had the shortest duration and lowest total precipitation, discharge and $NO_3^- - N$ concentration (Table 2).

The hysteresis pattern at the HAUS station differed strongly from those at the two upstream stations
(Figure 6). The percentage of events with negative *HI* combined with positive *CI*, which dominated at
upstream stations, decreased to only 27% at HAUS and had the longest duration, highest total

precipitation and highest discharge and $NO_3^- - N$ concentration (Table 2). The general pattern was

negative *HI* with negative *CI*, which accounted for ca. 36% of all events. The pattern of positive *HI*

with negative CI increased from < 5% at the upper two stations to 30% at the HAUS station. The

pattern of positive *HI* with positive *CI* occurred only during drying/dry periods (Figure 6). Notably,

regardless of CI, patterns with positive HI usually occurred during drying/dry periods (ca. 88%), with

300 a short duration and low discharge and $NO_3^- - N$ concentration, while patterns with negative HI

- 301 occurred more during wetting/wet periods (ca. 73%), with a long duration and high discharge and
- 302 $NO_3^- N$ concentration (Table 2).



Figure 5. Examples of the four hysteresis types detected at the HAUS station: (a) an event in Aug. 2014 with a positive hysteresis index (*HI*) and positive concentration-change index (*CI*); (b) event in Jul. 2016 with a positive *HI* and negative *CI*; (c) event in Jan. 2012 with a negative *HI* and positive *CI*; and (d) event in Nov. 2017 with a negative *HI* and negative *CI*. Blue and red lines represent discharge and $NO_3^- - N$ concentration, respectively. Inset plots show the corresponding hysteresis loops (from blue to red), in which the x-axis and y-axis are normalized values of discharge and $NO_3^- - N$ concentration, respectively.



Figure 6. Hysteresis index (*HI*) and concentration-change index (*CI*) patterns at the SILB, MEIS and
HAUS stations during the wetting, wet, drying and dry periods.

Table 2. Statistics of events of four hysteresis types based on sign of the hysteresis index (*HI*) and concentration-change index (*CI*) at the SILB, MEIS and HAUS stations (including number of storm events, mean duration, mean total precipitation (P), mean discharge (Q) and mean $NO_3^- - N$ concentration.

	SILB				MEIS				HAUS				
(HI/CI)	(+/+)	(+/-)	(-/+)	(-/-)	(+/+)	(+/-)	(-/+)	(-/-)	(+/+)	(+/-)	(-/+)	(-/-)	
Number	14	4	58	5	6	3	51	12	5	21	19	25	
Duration (day)	4.6	2.5	8.7	6.7	14.0	4.4	8.0	5.4	2.1	4.8	10.2	9.3	
Total P (mm)	18.70	4.17	27.66	20.00	27.60	7.52	23.78	20.26	12.32	20.45	22.64	19.28	
$Q (m^3 s^{-1})$	0.90	0.63	1.28	1.85	2.60	0.89	1.89	1.62	0.88	1.19	3.66	2.32	
$NO_3^{-1}-N \text{ (mg l}^{-1}\text{)}$	1.23	0.78	2.05	3.01	2.07	1.52	1.92	1.74	1.99	2.11	2.99	2.49	

319 **3.4. Shared events analysis**

To investigate the influence of the landscape on seasonal flow and nitrate dynamics, we analyzed 24 catchment-wide shared events. Most shared events had negative *HI* at the SILB and MEIS stations, with only one event with positive *HI* at the MEIS station (Supplementary Table S2). Three shared events had positive *HI* during the dry period at the HAUS station, with a high contribution of NO_3^- – *N* load from the lowermost subcatchment (Table S2). The number of shared events with negative *CI* increased from upstream to downstream (i.e., 2, 3 and 12 at the SILB, MEIS and HAUS stations, respectively) (Table S2).

327 Runoff volume (R_V) and nitrate-N load (N_L) contributions from each subcatchment had strong

seasonal variations at the event scale (Figure 7). During wetting/wet periods, $R_{V,UP}$: $R_{V,ALL}$ was

significantly higher than $R_{V,MID}$: $R_{V,ALL}$ and $R_{V,LOW}$: $R_{V,ALL}$ (Kruskal-Wallis test). $R_{V,MID}$: $R_{V,ALL}$ and

330 $R_{V,LOW}$: $R_{V,ALL}$ varied during the wetting period, but the former ratio increased and the latter one

decreased during the wet period. During the drying period, $R_{V,UP}$: $R_{V,ALL}$ decreased quickly to a

332 proportion similar to $R_{V,MID}$: $R_{V,ALL}$, while $R_{V,LOW}$: $R_{V,ALL}$ increased but remained lower than

333 $R_{V,UP}$: $R_{V,ALL}$ and $R_{V,MID}$: $R_{V,ALL}$. During the dry period, contributions from the three subcatchments

varied within a similar range, with slightly higher $R_{V,UP}$: $R_{V,ALL}$ than $R_{V,MID}$: $R_{V,ALL}$ and

335 $R_{V,LOW}$: $R_{V,ALL}$.

336 Nitrate-N load contributions had a different seasonal pattern from runoff volume contributions (Figure

337 7b). During wetting/wet periods, nitrate-N load contributions from the three subcatchments generally

followed runoff volume contributions (i.e., that from the uppermost subcatchment was significantly

higher than that form the middle and lowermost subcatchments, Kruskal-Wallis test). Likewise,

340 $N_{L,MID}$: $N_{L,ALL}$ and $N_{L,LOW}$: $N_{L,ALL}$ varied, with the former ratio increasing and the latter one

- decreasing. During the drying period, however, $N_{L,MID}$: $N_{L,ALL}$ increased quickly and ultimately
- exceeded $N_{L,UP}$: $N_{L,ALL}$, while both $N_{L,UP}$: $N_{L,ALL}$ and $N_{L,MID}$: $N_{L,ALL}$ decreased to a similar low
- 343 proportion. During the dry period, $N_{L,LOW}$: $N_{L,ALL}$ decreased. Thus, the subcatchment that contributed
- the most nitrate-N load varied among the four hydrological periods.



Figure 7. Contribution of (a) runoff volume (R_V) and (b) nitrate-N load (N_L) from each subcatchment to the catchment outlet for shared events throughout the year. Dashed lines separate the four hydrological periods. Solid lines indicate high-order polynomial regressions.

349 4. Discussion

Terrestrial nitrate transport at the catchment scale is strongly related to flow dynamics. In the Selke 350 catchment, the land-to-stream transport has been identified as a key driven factor for surface water 351 dynamics (Dupas et al., 2017). Due to the complex combination of meteorological, hydrological, 352 353 geographical and pedological characteristics of each subcatchment, nitrate dynamics varied from upstream to downstream in the nested Selke catchment. Based on the hysteresis indices, four 354 355 hysteresis patterns can be conceptualized (Figure 8). The driving factors of their occurrences in the 356 Selke catchment (Figure 6) depend on the combinations of landscape features and hydrological 357 conditions at the seasonal scale.

358 4.1. Characteristic patterns of flow and nitrate dynamics

359 The four hysteresis patterns varied strongly spatially and temporally. In the uppermost and forest-

- 360 dominated middle subcatchments, the most common hysteresis pattern across different seasons was
- 361 counter-clockwise hysteresis with an accretion effect during the rising limb (i.e., negative *HI* with
- positive CI, c.a. 70% at the SILB and MEIS stations). At the event scale, nitrate sources near stream

363 reaches were flushed out quickly, which resulted in an accretion effect during the rising limb of the hydrograph, given the low ambient nitrate concentrations (Figure 8d). Although the uppermost 364 mountainous regions contain considerable areas of agricultural land, nitrate cannot accumulate in the 365 deeper subsurface due to the shallow impermeable bedrock and the consequently flashier flow 366 367 pathways (Dupas et al., 2017; Yang et al., 2019). This feature results in higher nitrate concentrations in the interflow than in the baseflow. Consequently, a synchronous seasonal pattern of discharge and 368 369 nitrate concentration was observed (i.e., generally high values during wetting/wet interflow-370 dominated periods and low values during drying/dry baseflow-dominated periods, Figure 2). Due to 371 sufficient precipitation during events and well-established hydrological connectivity, interflow can 372 transport distal terrestrial nitrate sources to the stream, which further increases surface water nitrate 373 concentrations. The time lag between hydrological celerity and solute transport velocity makes nitrate 374 concentration normally peak after discharge do, which results in counter-clockwise hysteresis. However, this pattern decreased at the HAUS station, accounting for only 27 % and occurred more 375 376 frequently during wetting/wet periods. Under high-flow and low-temperature conditions, the pattern at 377 the catchment outlet was controlled more by the two upstream subcatchments due to their large 378 contributions to both runoff volume and nitrate-N load (Figure 7) and low in-stream nitrate uptake 379 (Rode et al., 2016a). Therefore, during wetting/wet periods, the hysteresis pattern at the HAUS station 380 can depend more on upstream features than on those of the lowermost subcatchment.

381 The propagation effects also influenced the pattern of counter-clockwise hysteresis with a dilution 382 effect at the three stations (Figure 8c). The uppermost and middle subcatchments had more saturation overland flow with lower nitrate concentration in forest areas (Zimmermann et al., 2006). Overland 383 384 flow near streams can be generated quickly during wetting/wet periods with high discharge, causing a dilution effect at the beginning of events. The dominant runoff component of overland flow during the 385 386 rising limb can be replaced by interflow quickly, with higher nitrate concentration, resulting in 387 counter-clockwise hysteresis. This pattern was the most common pattern at the HAUS station. During 388 wetting/wet periods, this hysteresis pattern at HAUS can be affected by the two upstream 389 subcatchments as mentioned. Moreover, stream water routed from upstream subcatchments can also

390 dilute nitrate concentration in the lowermost subcatchment, since the latter generally has higher nitrate concentration during low flow conditions (Figure 2), changing the hysteresis pattern from upstream 391 392 accretion to downstream dilution across different seasons (Table S2). Besides influences from upstream subcatchments, features of the lowermost subcatchment can also cause a dilution effect. 393 394 Urban/arable areas have been recognized to influence runoff generation in a catchment during storm events (Bronstert et al., 2002; Niehoff et al., 2002). Agricultural and municipal construction results in 395 396 quick surface flow, which tends to decrease nitrate concentration. Dilution effects caused by surface 397 flow during storm events were also observed in a mountainous agricultural catchment in California 398 (USA) (Goodridge and Melack, 2012), mountainous agricultural catchments in the tropics (Jacobs et 399 al., 2018) and urbanized catchments in North America (Barco et al., 2008). In this case, nitrate 400 concentration decreased quickly at the beginning of the rising limb in the uppermost and lowermost 401 subcatchments (Figure 5c), which both contain urban and arable areas (Figure 1). Interflow can then 402 dominate quickly due to hydrological connectivity during wet periods, or baseflow can dominate 403 again after quick flow during the dry period. In both situations, nitrate concentration increased after 404 quick flow before the discharge peak, resulting in counter-clockwise hysteresis (Figure 8c). Therefore, 405 this pattern dominated at the HAUS station not only due to events that propagated from upstream 406 subcatchments, but also events generated in the lowermost subcatchment.

408 Hydrological connectivity from land to stream can be limited by the higher temperature and lower soil 409 moisture during drying/dry periods. Distal nitrate sources become immobilized and are not exported to the stream at low discharge, and thus resulted in clockwise hysteresis (Figure 5a). Our findings are 410 411 in line with those of Baker and Showers (2019) who concluded that clockwise hysteresis was favored 412 when antecedent soil moisture was low. In the uppermost subcatchment, the accretion effect was 413 caused by plentiful proximal nitrate sources, as mentioned. In the middle subcatchment, atmospheric 414 deposition is the main source of N in forest areas (MacDonald et al., 2002). Nitrate concentration has 415 been reported to be higher in the topsoil than in groundwater due to uptake by deep-rooted vegetation 416 and denitrification in deep soil layers (Chaves et al., 2009). Consequently, fewer distal nitrate sources

It was noticeable that clockwise hysteresis occurred more during drying/dry periods (Figure 6).

407

417 can be transported to the stream during the baseflow-dominated dry period in the middle than in the uppermost subcatchment, which decreased the number of clockwise hysteresis at the MEIS station 418 (Table 2). This pattern was the least common pattern at the HAUS station and occurred only during 419 420 drying/dry periods (Figure 6). None of this pattern at HAUS belonged to shared event (Table S2). 421 This indicated that this pattern could represent nitrate dynamics exclusively in the lowermost 422 subcatchment. The lowermost subcatchment usually stores more nitrate due to mineral fertilizer and 423 manure application. Saturated nitrate sources that are near the stream are easily flushed out when storm events occur, even at low discharge, leading to accretion effects (Figure 8b). Subsequently, 424 425 nitrate concentration decreases quickly due to insufficient precipitation and low hydrological 426 connectivity with distal nitrate sources, resulting in clockwise hysteresis. 427 The features of the lowermost subcatchment caused more clockwise hysteresis with the dilution effect 428 (Figure 8a), which was the second-most common pattern at the HAUS station. This pattern appeared 429 more during drying/dry periods when the lowermost subcatchment contributed considerable nitrate-N 430 load and had substantial influence on nitrate export dynamics (Figure 7b). In this situation, the 431 dilution effect at the HAUS station could be related to quick flow from paved areas in the lowermost 432 subcatchment, as mentioned. Nitrate concentration followed a vertical gradient in the lowermost 433 subcatchment: lower in the topsoil due to plant uptake and higher in deep soil due to legacy nitrate (Outram et al., 2016). Hydrological connectivity was low during drying/dry periods, and interflow 434 435 could not increase enough to transport deep nitrate-rich sources to the stream, which results in the continued decrease in nitrate concentration during the falling limb. Therefore, clockwise hysteresis 436 occurred more at the HAUS station than at the two upstream stations (i.e., only 4 and 3 times at SILB 437 438 and MEIS, respectively).



440 **Figure 8**. Conceptual interpretation of nitrate-discharge relationships during events for the four 441 combinations of hysteresis index (*HI*) and concentration-change index (*CI*). Blue lines indicate flow 442 behavior, and red lines indicate $NO_3^- - N$ behavior. Arrows indicate dominant mechanisms during 443 rising/falling limbs.

444 **4.2. Implications and limits**

Overall, the hysteresis pattern based on the C-Q relationship varied among landscapes, reflecting 445 446 unique flow pathways and spatial nitrate storage. Due to different human activities (e.g., fertilizer 447 application, drainage construction) and geological conditions, different patterns can dominate for a 448 given dominant land use. For example, Carey et al. (2014) observed more clockwise hysteresis and a 449 dilution effect in coastal catchments in which forest dominated. In contrast, we observed that counter-450 clockwise hysteresis with an accretion effect dominated in the upstream mountainous forest areas. 451 This difference indicates that scientific monitoring should be set up according to meteorological, 452 hydrological, geographical and pedological features rather than based only on land use. This is 453 especially important in a nested catchment, where upstream subcatchment(s) can influence downstream subcatchment(s), and thus conceal nutrient export dynamics of the latter. Furthermore, 454 differences in hydrological connectivity and biogeochemical processes due to seasonal variations can 455

456 cause variable flow and nitrate dynamics (e.g., more clockwise hysteresis during drying/dry periods 457 and more counter-clockwise hysteresis during wetting/wet periods). To disentangle seasonal effects, 458 long-term high-frequency monitoring can provide reliable datasets. Our findings can be used to target 459 mitigation measures when specific *HI/CI* combinations dominate in a given catchment. For example, 460 when the C-Q relationship suggests that proximal nitrate sources dominate, management actions can 461 focus on agricultural land near streams, but when distal nitrate sources dominate, catchment-wide 462 actions are required.

463 However, hysteresis analysis should also consider hydrographs with dual peaks or extremely long 464 recession periods, which may influence the analysis (Lloyd et al., 2016b; Williams et al., 2018). Such 465 hydrographs were rare in our study, but since they can improve understanding of process mechanisms during certain periods, they were not excluded. For example, the only shared event with positive HI at 466 467 the MEIS station showed counter-clockwise hysteresis early in the event but was influenced by a dual 468 peak and longer falling limb, which yielded a low positive value of HI (ca. 0.006). This analysis can 469 improve understanding of the corresponding shared event at the SILB station that had negative HI: multiple distal sources can be consumed after a long period of flushing (e.g., during a large storm 470 event) and cause lower nitrate concentrations during the late falling limb. This case requires 471 472 catchment-wide management actions instead of focus on proximal streams, despite a positive HI at the MEIS station. Thus, events should be assessed carefully to avoid unreliable conclusions from 473 474 statistical results, especially those with low values of HI or CI. Doing so may provide a detailed picture of variations in flow/nitrate dynamics. 475

Our comprehensive analysis of landscape and seasonality effects on flow and nitrate dynamics
focused mainly on high-frequency monitoring data. Therefore, it is difficult to quantify the landscape
effect in the three subcatchments of the Selke catchment, even when considering shared events. The
landscape effect should be explored further using physical-based hydrological water-quality models.
Current modeling is based mainly on daily data (Yang et al., 2018), which may bias detection of
storm events and calculation of the nitrate-N load during storm events. Thus, future studies require

482 high-frequency modeling, which can be used to quantify influential factors that result in different flow483 and nitrate dynamics and provide targeted advice for water management.

484 5. Conclusions

According to C-Q relationships, counter-clockwise hysteresis with an accretion effect dominated the catchment throughout the year; however, hysteresis was clockwise during specific periods in each subcatchment. Clockwise hysteresis occurred more during the dry period, indicating low hydrological connectivity from land to stream for export of distal nitrate sources. Dilution effects dominated in the lowermost catchment, which may have been influenced by flow propagating from upstream subcatchments during the wet period or generated by quick flow from paved areas.

491 When analyzing shared events, the uppermost subcatchment always dominated runoff volume and 492 dominated nitrate-N load during all periods except the dry period, when the lowermost subcatchment 493 dominated nitrate-N load, which indicates the substantial contribution of nitrate export regimes from 494 the lower urban/arable area. At the event scale, this alternation suggests that high nitrate-loaded 495 interflow dominated in the upper mountainous subcatchments, while quick runoff (e.g., surface flow 496 with low nitrate concentration) dominated in the lowermost subcatchment. This difference in nitrate 497 export can increase during dry/hot seasons, when hydrological connectivity and biogeochemical 498 processes change greatly.

These conclusions depend greatly on high-frequency data, which enabled events to be detected and nitrate-N load to be calculated more accurately. Although complex hydrographs may have influenced our results, the interpretation of the fundamental mechanism of variable C-Q relationships remains reliable. Water or agricultural management should be considered in complex conditions in which several mechanisms may coexist. Thus, a continuous and scientific monitoring strategy in a nested catchment is important to capture the nitrate export regime at the seasonal and catchment scale.

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