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# 1 Using proteins to study how microbes contribute to soil ecosystem services: The current state and

2 future perspectives of soil metaproteomics

3 Robert Starke<sup>1</sup>, Nico Jehmlich<sup>2</sup>, Felipe Bastida<sup>3</sup>

<sup>1</sup>Laboratory of Environmental Microbiology, Institute of Microbiology of the CAS, Vídeňská 1083, 14220
 Praha 4, Czech Republic; <sup>2</sup>Helmholtz-Centre for Environmental Research – UFZ, Department of
 Molecular Systems Biology, Permoserstraße 15, 04318 Leipzig, Germany; <sup>3</sup>CEBAS-CSIC, Campus
 Universitario de Espinardo, Murcia, E-30100, Spain

# 8 Abstract

9 Metaproteomics was established to analyse both the structure and the function of microbial 10 communities and, particularly in soils, their contribution to ecosystem services. In this review, we 11 provide an overview on how the study of the soil metaproteome can provide fundamental information 12 on the role of microbial communities in soil ecosystem services. We further discuss the strengths and 13 weaknesses of soil metaproteomics in comparison to other culture-independent OMIC techniques. We 14 critically review its bottlenecks but also provide strategies to mitigate and possible directions for future research as the direct link of structure and function is advantageous and complementary to 15 16 metagenomics, metatranscriptomics and metametabolomics.

# 17 Soil ecosystem services provided by microbes

Soil, a key component of the terrestrial ecosystem, operates at the interface of the atmosphere, 18 19 biosphere, hydrosphere and lithosphere [1], and provides many ecosystem services - acting as a habitat 20 for soil organisms, regulator of water quality, modifier of the atmospheric composition, medium for plant growth and recycling system for nutrients and organic wastes, etc. [2]. Even though soil is one of 21 22 the most important carbon (C) reservoirs within the Earth's ecosystems [3,4], most of its microbes have 23 not yet been described [5,6]. But, one gram of soil can contain billions of organisms belonging to 24 thousands of different species [7]. In particular, the mean prokaryotic density of 10,000,000 organisms per gram of soil [8] is at least one magnitude of order higher than that found per millilitre of water in the 25 ocean [9]. This implies the importance of understanding microbially-driven soil processes such as C-26 27 cycling and soil fertility that ultimately lead to human wellbeing. Surprisingly, microbial biomass typically 28 accounts for only 1% of soil organic carbon (SOC) [10,11] but its contribution to C cycling is extensive, 29 through its implication in the mineralisation and stabilisation of the soil organic matter (SOM) derived from plants [12-14]. In addition, soil microbial community has been shown to be responsive against 30 climate change factors [15–21]. Hence, understanding the structure and the function of the soil microbial 31

community is crucial to the determination of its contribution to soil ecosystem services and how global
 change and land use intensification will affect this.

The standard metric of quantification in soil microbial ecology studies is the reading of nucleic acid 34 35 counts. The data derived from genomic studies permit: i) the identification and quantification of the 36 relative abundance of microbial populations through amplicon-based approaches (i.e., 16S rRNA genes; ITS fragments, etc.); ii) the quantification and the estimation of the diversity of functional genes (e.g., 37 38 genes related to N cycling, such as amoA, nirS and nirK; and, more recently, genes involved in P cycling, such as PhoD and PhoC [22]); iii) the inference of their functional role through statistical methods which 39 40 relate nucleic acids to functional parameters (i.e., extracellular enzyme activities, basal respiration, etc.); and iv) metagenomic approaches, which aim to decipher the whole repertoire of genes in a soil 41 42 microbial community which attempts to potentially connect phylogeny and functionality.

43 Through sequencing approaches, several studies have revealed the core composition of soil microbial 44 communities at the global scale. For instance, Delgado-Baquerizo and colleagues found that only 2% of bacterial phylotypes, which amounted to 511 individual phylotypes, consistently account for almost half 45 46 of the soil bacterial communities worldwide [23]. This "most wanted" list included Proteobacteria, 47 Actinobacteria, Acidobacteria, Planctomycetes, Chloroflexi, Verrucomicrobia, Bacteroidetes, 48 Gemmatimonadetes, Firmicutes and Armatimonadetes in decreasing order. Complementary, Tedersoo 49 and colleagues used DNA metabarcoding to describe that fungal richness is decoupled from plant diversity but fungi show similar latitudinal diversity gradients to other organisms [24]. Similar to 50 51 bacteria, the abundant phylotypes globally, comprising of Agaricomycetes (50.1%), Sordariomycetes 52 (8.0%), Eurotiomycetes (7.7%), Leotiomycetes (7.1%) and Mortierellomycotina (6.3%), describe more 53 than half of the fungal communities worldwide. When comparing metaproteomics and 16S rRNA gene sequencing in samples from North-America and South-Europe, the patterns of bacterial composition at 54 55 phylum level resembled similar patterns, but the relative abundance of Proteobacteria and Planctomycetes was greater in metaproteomics than metagenomics [25]. Nevertheless, although the 56 57 composition of microbial communities at the phylum level can be partially predicted (either with metagenomics and metaproteomics), the composition of microbial communities at finer taxonomic 58 59 levels could probably arise in greater variations across soil types and climates.

However, sequencing harbors limitations such as the quantitative accuracy [26,27] which is why the actual functionality of microbes should be deciphered by analysis of the catalysts of soil ecosystem services – the proteins, which provide both functional and phylogenetic information [28–30]. This is fundamental since it has been discovered recently that a large part (up to 40%) of the soil DNA can belong to dead cells (relic DNA) which are not active anymore [31], whereas extracellular proteins in soil

65 (e.g., extracellular enzymes) can persist and remain active through stabilisation by humic substances and

66 clays [32].

67 The significance of soil metaproteomics

In 2009, Bastida and colleagues <u>published a review that</u><u>focusedevaluated the initial potentials of soil</u> <u>metaproteomics</u>. This review was focused <u>-</u>on the methodology of soil metaproteomics with the emphasis on the different protein extraction techniques and peptide identification by mass spectrometric analysis as the identification of soil proteins may provide information about the biogeochemical potential of soils and pollutant degradation and act as an indicator of soil quality, identifying which proteins and microorganisms are affected by a degradation process [33]. In their review from 2016, Keiblinger and colleagues discussed the opportunities and limitations of selected

protein extraction techniques in soil and leaf litter metaproteomics, including a step-by-step guideline on application, sampling, sample preparation, extraction and data evaluation strategies, and used recent application to discuss how linking phylogeny and taxonomy can help to gain deeper insights in terrestrial microbial ecology [34]. In this review, we focus on recent applications of metaproteomics but also highlight the bottlenecks and offer a case-study where major limitations are explored (e.g. biomass and soil organic C content). Further, we establish a set of research questions that can be potentially

81 investigated in the coming years with the use of metaproteomics.

82 Among the culture-independent OMIC techniques deployed to gain deeper insights into the structure and function of microbial communities [35], metaproteomics has gained more and more interest from 83 the scientific community as a central element in microbial ecology studies [30], since it deciphers the 84 85 functional relationships between community members [36], particularly in soils [15,36–53]. Thus far, 49 articles on soil metaproteomics have been published, in comparison to 54 on soil metatranscriptomics 86 and 749 on soil metagenomics (PubMed as of July 16<sup>th</sup>, 2018). Genomics is still the method of choice 87 used by many soil scientists to evaluate the diversity and composition of the soil microbial community, 88 89 due to easier handling, lower costs and commercially-available sequencing provided by companies (for 90 instance, Agilent Technologies Inc. and Illumina Inc.) or institutes (for instance, the Joint Genome 91 Institute or Argonne National Laboratory). In comparison, proteomics involves complex sample-92 extraction protocols (which still need to be improved) and cost-intense and sensitive equipment to 93 which only a few laboratories in the world have access. However, the ability of metaproteomics to link the phylogenetic structure to the function is a clear advantage over nucleic acid and metabolomic 94 95 approaches in soil microbial ecology.

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#### 96 The applications of soil metaproteomics

Soil metaproteomics has been used to describe the soil microbial community in relation to ecosystem services, such as agriculture [38,41,45,54], bioremediation [36,43,52] and C cycling [40,42,47–49,51], and climate change factors (i.e., temperature) [15,37,39,53]. The results obtained from nucleic acids and proteomics, at least at the phylogenetic level, have shown good correlation [15,48,53]. Hence, the use of proteins to describe ecosystem functioning in soils is of interest because it captures both phylogenetic and functional information [15,41,42,48].

# 103 Carbon (C) and nitrogen (N) cycles

104 Carbon cycling is fundamental to the Earth's ecosystem since C is the main component of all biological 105 compounds. The greenhouse gases carbon dioxide (CO<sub>2</sub>) and methane, two forms of C, absorb and 106 retain heat in the Earth's atmosphere and are thus partially responsible for the greenhouse effect [3], 107 which is why an understanding of C cycling is crucial to the comprehension of climate change. Indeed, 108 the concentration of atmospheric CO<sub>2</sub> recorded in summer 2018 was the highest since records began, which underlines the need for a better understanding of the biotic factors (e.g., microbes) controlling 109 110 the mineralisation of SOM, which is driven by heterotrophic soil microbes [55]. Within the C-cycle, the 111 restoration of C-limited soils, such as those frequently found under arid and semi-arid climates, by the 112 application of organic amendments (i.e., sludges, composts, etc.) is a major task that is necessary to 113 maintain soil fertility. Through a metaproteomic approach, it has been found that organic amendments 114 strongly impact soil ecosystem processes, at both the cellular and extracellular level, depending on the type of organic amendment [47]. Further, metaproteomics revealed that not only the nutrient 115 116 concentrations but also the C:N:P stoichiometry influenced both the structure and the activity of the 117 microbial community during beech litter decomposition [40], whereas deforestation induced a longterm loss of bacterial biomass and enzyme activity but increased bacterial diversity [49]. Moreover, the 118 119 active diversity (measured by metaproteomics), but not the total microbial diversity (through amplicon-120 based genomics), was shown to be related to the availability of organic C [48] in a semi-arid ecosystem. 121 Lastly, metaproteomic approaches have been utilised to track the N-cycle. The degradation of plant-122 derived polymeric organic C (i.e., lignin, cellulose, etc.) is dominated by fungi [56,57], but bacteria were 123 recently shown to dominate the short-term assimilation of plant-derived N, presumably in easily-124 available sources, by tracing <sup>15</sup>N incorporation into proteins of the microbial community [41].

### 125 Pollution, bioremediation and restoration

126 The growth of the human population is increasing the pollution of the environment and hence a detailed 127 understanding of the microbial populations associated with the biodegradation of pollutants is required, 128 as well as of the main biochemical pathways of biodegradation, which have biotechnological 129 applications. In fact, metaproteomics revealed that petroleum pollution induced an increase in the 130 relative abundance of Proteobacteria but a decrease in Rhizobiales - which ultimately led to promotion 131 of the removal of polycyclic aromatic hydrocarbons and alkanes [36]; this work also provided parallel 132 information on the biochemical pathways of the biodegradation. Further, the rhizosphere community of 133 constructed wetlands, generally used to treat contaminated water, was shown to have diurnal 134 polyhydroxyalkanoate metabolism, which is aligned with the diurnal cycle in plants [58]. Similarly, 135 naphthalene, an aromatic contaminant, resulted in the separation of a bio-stimulated community - able to utilise naphthalene - from a non-stimulated community exhibiting a greater metabolic window than 136 137 previously predicted [52], suggesting a core community for specific functionality. Such an active 138 community was found in permafrost-affected cryosols, for atmospheric-methane-oxidising bacteria [37].

139 Climate change

140 Climate change is expected to induce a loss of SOC by increasing soil respiration rates [55,59], which will 141 ultimately exacerbate the greenhouse effect [60]. Recent evidence indicates that warming is accelerating the activity of soil microorganisms at the global scale, with the subsequent increase in CO<sub>2</sub> 142 release from soil to atmosphere [61]. As soil respiration is mainly performed by heterotrophic soil 143 144 microorganisms [55], understanding the structure and function of the microbial community is vital to 145 comprehend responses to climate change. Changes in the phylogenetic composition of the microbial 146 community and in its functional potential and activity were shown to represent different states of thaw 147 in soils [53]. In contrast, warming resulted only in subtle differences at the phylum and class levels of the community whereas functionality increased, particularly energy production and conversion [39]. 148 Similarly, the structure and diversity of the total bacterial community were unaffected by drought even 149 150 though microbial biomass and ecosystem multifunctionality decreased, in a semi-arid ecosystem [15]. In 151 fact, climate was reported to be a fundamental driver of the protein abundance of various phyla in a 152 wide range of soils from different continents [25].

#### 153 Soil fertility and agriculture

154 Agricultural productivity does not depend only on the physical and chemical quality of soil (e.g., texture, 155 density, nutrients and water availability, etc.) or the absence of plant pathogens. The soil microbial community plays a critical role in agricultural productivity through its cycling of nutrients, making them 156 157 available for plant growth. Indeed, the soil microbial diversity is fundamental to the maintenance of soil 158 productivity and the resulting economic benefits [62]. Soil metaproteomics could help to dig deeper into 159 the functions and role of microbial populations in plant-soil interactions in an agricultural context. In this sense, crop yields and the optimal growth conditions have been evaluated through metaproteomic 160 approaches. For instance, the mere presence of plants influences the metaproteome and soil bacterial 161 162 functions [63]. Further, a yield decline of crops (i.e., sugarcane) was shown to be related to the soil 163 metaproteome, as well as to enzyme activities and the catabolic diversity of the microbial community 164 [38]. Moreover, changes in microbial abundance in the rhizosphere have been related to the protection 165 of tomato plants from wilt disease, presumably by beneficial microbial proteins in healthy soil [64], and to ecological and functional adaptations to varying water management in semi-arid ecosystems [41]. 166

#### 167 The bottlenecks of soil metaproteomics

# 168 *Protein extraction, sample processing and measurement*

169 Protein extraction from soil is time-consuming as it comprises protein extraction (2 days) followed by gel 170 electrophoresis (1 day), tryptic digestion for at least 16 hours or overnight and desalting (2 days) (Figure 171 1). The commonly-used extractants are sodium dodecyl sulphate and trichloroacetic acid (SDS-TCA) [50] 172 and SDS-phenol [65]. The extracted proteins are then analysed by 2D-SDS-PAGE gel electrophoresis for 173 imaging and MALDI-TOF analysis [66] or, more commonly, by 1D-SDS-PAGE for LC-MS/MS analysis. The 1D approach comprises tryptic digestion of the proteins in-solution or in-gel, desalting via ZipTip 174 175 columns or SCX spin columns, measurement with a coupled system of liquid chromatography and tandem mass spectrometry (LC-MS/MS), database searches with Proteome Discoverer (using Mascot or 176 177 Sequest as the algorithm) against every known protein-coding sequence or a metagenome, the use of 178 Percolator to calculate false discovery rates (FDR) [67], phylogenetic and functional assignment of protein groups with the "Proteomics results Pruning & Homology group Annotation engine" 179 180 (PROPHANE) [68], calculation of protein abundances - as either the normalised spectral abundance 181 factor (NSAF) [69] or the area under the curve (AUC) - and, finally, data upload to ProteomeXchange via 182 PRIDE [70] (Please refer to the reviews by Muth and colleagues [71], and Heyer and colleagues [72] for 183 further details on the bioinformatic data analysis of metaproteomes). However, slight differences in the

extraction method, sample preparation, instrumentation and data analysis were found among soil metaproteomic studies. Each of these steps has been optimised but requires further validation to obtain high-quality proteomics data. However, among these steps, protein extraction and data analysis (see next section) can influence the outcome the most.

The extraction of extracellular proteins remains a challenge due to their physicochemical interactions 188 189 with SOM and soil mineral particles [33], which were reported to lower the extraction efficiencies and 190 final peptide yields [46,73]. For example, an increasing content of montmorillonite, an expansible 2:1 191 clay, in soil was found to decrease the amount of extracted proteins and to hinder protein identification, 192 likely due to conformational changes or degradation [74]. In addition to the interaction with clay, it was 193 demonstrated that the contact between proteins and soil-borne humic substances, potentially due to 194 hydrophobic interactions, led to protein modifications that affected the identification by MS; this was 195 especially prominent in proteins of lower molecular weight and of less-complex structure [75]. 196 Previously hypothesised were hydrophobic interactions caused by dehydration, which could destabilise 197 and partially disassemble humic substances in contact with the proteins [76] or supramolecular aggregation of low molecular weight organic molecules held together by weak molecular surface forces 198 199 [77]. However, it remains unknown which modifications specifically lead to lower identification rates.

The relationship between the concentration of SOC or microbial biomass (both as *S*) and the number of identified peptides (v) using equation 1 as previously described [78] are illustrated in Figure 2, with  $V_{max}$ as the maximum number of identified peptides and  $K_m$  the concentration of SOC or microbial biomass at which half of the maximum number of peptides are identified.

204 Eq. 1:  $v = \frac{V_{max}[S]}{K_m + [S]}$ 

205 These results were obtained from a gradient of semi-arid soils that covered a range of SOC 206 concentrations. The SOC concentration has a negative impact on the number of spectra and, 207 consequently, on the number of identified peptides (Figure 2a). The relationship between the SOC concentration and the number of identified peptides reveals a plateau of extracted proteins at a 208 209 breaking point of 1.5% SOC, despite the greater microbial biomass in soils richer in SOC that should logically lead to higher protein content. In fact, the efficiency of protein extraction is lowered not only 210 by increasing SOC but also by increasing microbial biomass content (Figure 2b). The number of peptide 211 212 identifications stagnated when the microbial biomass reached a breaking point of 16.8 nmol fatty acids 213 per gram of soil, which follows previously-reported shielding from disruption treatments at high cell 214 densities, due to aggregation [79]. Hence, (i) more-efficient cell lysis techniques and (ii) more-efficient 215 protein extraction techniques must be investigated to assure the quality and quantity of proteomics

216 data. Further, the extraction techniques and buffers commonly used for the measurement of 217 extracellular enzyme activities were shown to have low extraction efficiencies, which points to strong 218 adsorption of proteins on humic substances and clays [80]. In addition, a definitive quantification of the 219 extracted proteins was still lacking in many studies [46,80] as the scientific community has commonly 220 relied on colorimetric methods - for instance, the Bradford assay - that are biased by the interference of 221 humic compounds [81,82]. However, correct quantification is crucial, to evaluate both the efficiency of 222 protein extraction and the full contribution of the protein repertoire to soil ecosystem services. In this 223 regard, the quantification of the amino acids from the extracted proteins [46] or the amido black assay 224 [83] could provide the most-straightforward approaches to estimating the amount of extracted proteins. 225 Otherwise, total ion chromatogram counts from liquid chromatography could be used but accurate 226 quantification requires internal standards and calibration for each protein of interest. However, the 227 intensity cannot be directly related to abundance due to differences in ionisation efficiencies between 228 peptides and discrepancies between ionisation and detection since only one out of every 100 to 10,000 229 analyte ions generated is detected [84-86].

# 230 Data analysis and functional assignments

231 The backbone of proteomic analysis is the database, as the measured peptides from the sample are 232 matched, using data analysis programs such as MASCOT [87,88] or SEQUEST [89], with the database 233 sequences. Hence, knowledge of what is in the sample is as important as the quantity and quality of the 234 proteins within the database. Unsurprisingly, since most organisms in soil are unknown [5,6], the database searches usually involve the utilisation of all known protein sequences - which comprise 235 79,565,724 TrEMBL and 335,933 Swiss-Prot entries for bacteria, 8,114,426 TrEMBL and 33,876 Swiss-236 237 Prot entries for fungi, 3,591,467 TrEMBL and 20,986 Swiss-Prot entries for archaea and 629,133 TrEMBL 238 and 4,636 Swiss-Prot entries for nematodes. These are weighted in favor of intensively-investigated organisms such as the bacterial E. coli K12 (4,497), the fungal S. cerevisiae (6,729) or the nematode C. 239 elegans (26,846) (Uniprot as of July 7<sup>th</sup>, 2018 [90]). Instead of using all known sequences, one could 240 241 provide a database by using shotgun sequencing of DNA [91,92], which is both costly and time-242 consuming and requires additional programs and knowledge. A possible solution could be to make 243 sequences publicly available as a community effort like the Human Genome Project [93], which has 244 already been started for proteomic datasets with PRIDE [70,94–98] but, hereafter, soil studies focusing 245 on all clades of the ecosystem should be emphasised.

Recent studies have focused on the community composition to assess ecosystem functioning but lackdirect functional classification even though the peptides identified with MS were assigned to proteins

248 with both taxonomic and functional information. Further, in complex communities containing thousands 249 of different species per gram of soil [7], not many proteins will be assigned to individual species of the 250 community as one measurement generally yields a few thousand proteins. Moreover, these microbial 251 communities are not homogenous, generally having a few highly-abundant species and many species of 252 low abundance. Worse still, proteins with ubiquitous functions such as ATP synthases or chaperones -253 which every cell contains, regardless of its type - are more abundant than specialised enzymes such as 254 benzoyl-CoA reductases and the many proteins whose functions are still unknown. These functional 255 proteins are generally less abundant than "housekeeping" proteins (e.g., proteins involved in the Krebs 256 cycle, cellular respiration, lipid metabolism, DNA replication, etc.), which makes them hard to detect 257 among the most-abundant proteins. Hence, while the taxonomic affiliation generates profound data on 258 what is there, the functional information is biased by ubiquitous enzymes and, depending on the degree 259 of heterogeneity, by specialised enzymes of the most-abundant organisms. Furthermore, the 260 identification of extracellular soil enzymes (i.e., glucosidases, phosphatases, cellulases, etc.), which are 261 deeply involved in nutrient cycles [32], is hampered by the inefficient extraction of the extracellular 262 metaproteome of soil because proteins are frequently linked to humic substances or clays [80]; hence, the genomic database is still poor in this respect. To circumvent these issues and directly assess 263 functionality, targeted approaches, instead of the commonly-used untargeted approaches [99,100], 264 265 have to be considered; this implies the selection of specific proteins for the measurements, by targeting the masses of peptides from proteins. Briefly, to achieve consistent quantification of sets of proteins 266 267 across a variety of different samples and replicates, targeted MS - such as selected reaction monitoring 268 [101,102] - could provide specific assays for the detection and quantification of proteins over the whole 269 range of cellular concentrations [103].

# 270 Stable isotope probing (SIP) of proteins

271 Metaproteomics can provide information on both phylogeny and functionality [36] in microbial ecology 272 studies [30], but only the use of SIP can directly link microbial populations within the whole community 273 to a specific function [104]. Possible labelling targets in proteins are C, hydrogen, N, oxygen and sulphur 274 [105]. All of which can be now analysed by the tool MetaProSIP [106] that uses the identification of the 275 peptide on MS<sup>2</sup>-level together with detecting and describing the incorporation pattern on MS<sup>1</sup>-level by 276 relative isotope abundance (RIA) and labelling ratio (LR). Of these, C is the one used most frequently in environmental SIP studies [107,108] due to its high abundance in amino acids and since it is a main 277 278 component of the bacterial biomass [109]. Otherwise, hydrogen has the second-highest abundance in 279 proteins, but its use is limited by the rapid HD-exchange [110] which will occur in living cells and during 280 sample preparation. In addition, the toxicity of highly-deuterated water has been described for 281 eukaryotes; in particular, cell division is hampered during the formation of the mitotic spindle [111]. By 282 contrast, prokaryotes were reported to grow in artificial conditions with a high deuterium content [112], 283 up to 98% D<sub>2</sub>O [113]. However, the chromatographic properties of deuterated compounds - originating 284 from the higher hydrophilicity of CD bonds, compared to CH bonds [114] - result in significantly-different 285 retention times [115]. Nitrogen, on the other hand, has a lower abundance in amino acids, resulting in 286 smaller mass shifts and reduced sensitivity [116]. Oxygen has not been used in protein-SIP experiments, 287 perhaps because of the hampering of direct incorporation from the substrate into biomass due to the 288 exchange of 18-oxygen for 16-oxygen water through enzymatic reactions [117]. Lastly, sulphur is only 289 present in cysteine and methionine, which is beneficial for the tracking of general metabolic activity 290 [118] but, due to the low amount of sulphur in proteins, according to the average model, the utilisation 291 of labelled sulphur in protein-SIP appears inauspicious. As of now, the only reported isotopic labelling of 292 soils has been done with 15-nitrogen, using labelled plant material (tobacco) to investigate its 293 assimilation by the soil microbial community, through protein-SIP [42]. In this study, between 11 and 26 294 labelled peptides were identified per time point over the course of 14 days; this showed that bacteria 295 dominated the short-term assimilation of plant N. However, if one relates this number of labelled peptides to the existence of 10,000,000 organisms per gram of soil [8], then the validity of the data and 296 297 their value regarding the determination of ecosystem functioning are challenged. Metabolic labelling of 298 energy or nutrient sources leads to dynamic and unpredictable mass shifts [109] and, hence, small 299 amounts of label indicate a low number of active key players within the microbial community with 300 respect to specific substrates such as N. This is further complicated by the fact that microbial biomass. 301 even though its derived products significantly contribute to SOM [12–14], typically accounts for only 1% 302 of SOC [10,11] and many of its species are thought to be dormant [16,31,119,120]. To achieve higher 303 labelling efficiency, the utilisation of heavy water, labelled with either deuterium or 18-oxygen, is 304 promising as it assesses the global activity regardless of the specific substrate. It was successfully demonstrated in acid-mine-drainage biofilms using  $D_2O$  [121] and in soil communities using  $H_2^{18}O$  in 305 DNA [122-124] and RNA [125,126]. In fact, SIP of fully-18-oxygen-labelled water in RNA was used 306 307 recently to elucidate that 94% of soil taxa produced new rRNA and therefore were metabolically active 308 [127] - which is the opposite of the common-acceptedly state of high dormancy of soil microbes. 309 However, this approach has yet to be demonstrated for proteins and in soils. In particular, 18-oxygen 310 seems more promising than deuterium as (i) the abiotic HD-exchange of acidic hydrogens once the 311 proteins are in contact with unlabelled water [110] could cause depletion of the label and (ii) the cell

division of eukaryotes is hampered due to the negative effect of deuterium on the mitotic spindle formation [111], even though it has been reported that bacteria survive in 98% D<sub>2</sub>O [113]. Hence, deuterated water could negatively impact the activity of eukaryotes such as fungi and nematodes, which would falsely indicate high bacterial activity.

# 316 Research questions to be answered by soil metaproteomics

317 To date, the knowledge derived from soil metaproteomics has been rather descriptive, mainly providing 318 information on the community composition and cellular functional attributes. However, the advent of 319 new ecological concepts in the field of soil metaproteomics [128] will provide a functional basis for the 320 development of microbial-ecological theories in soil science. In this respect, we propose several 321 fundamental research questions that could be investigated with the use of metaproteomics (of course, 322 in combination with classical and other OMIC approaches in soil science). For instance, how do land use 323 and associated changes in soil nutrients determine the protein repertoire and its functional repercussion 324 within soil ecosystem services? Further, environments with higher phylogenetic diversity are supposed 325 to be more resilient/resistant against harmful stressors such as drought [15-21,129]. Hence, does the 326 diversity of proteins reflect the resilience/resistance of a soil microbial community? This would be critical 327 with respect to gaining a mechanistic understanding how microbial communities can resist climate 328 change factors such as global warming or aridity. If so, can we predict "functional markers" of climate-329 change resistance in the soil microbial community? Finally, it has been shown recently how microbes 330 interact with each other at the global scale [130] and how climate can affect not only soil microbial 331 communities but also the co-occurrence networks [131]. Consequently, is metaproteomics able to 332 provide functional insights into these interactions? While proteomics has proven to be powerful, with 333 advantages over other culture-independent OMIC tools and culture-dependent techniques, the 334 bottlenecks must be reviewed and actively targeted to sustain high-quality research using proteins from 335 environmental samples. In fact, better integration of techniques and research groups can guide the 336 scientific community to a more-comprehensive understanding of soil ecosystem services since the data 337 are complementary and not exclusive. Besides, more effort must be made to improve methods of 338 protein extraction from soil as well as the identification of low-abundance proteins with functional importance. In conclusion, despite the current limitations, soil metaproteomics has rapidly developed 339 340 since 2007 [132] and there is room for terrific improvements in the near future. Soil metaproteomics is 341 contributing to our understanding of fundamental and global aspects of soil microbial ecology (and will 342 continue to do so), which will ultimately help us to understand the role of microbes in ecosystem

services in the light of the land-use intensification and global changes forecasted for the comingdecades.

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760 761 Figure 1: The workflow of metaproteomics (blue) and protein stable isotope probing (orange) of soil samples. 762 Proteins are commonly extracted with sodium dodecyl sulphate (SDS) [65] or SDS-phenol [50], before in-solution 763 [133] or in-gel [134] tryptic digestion; sulphur bonds of proteins (S-S) are reduced with dithiothreitol (DTT) and 764 alkylated (shown as -AA) with iodoacetamide (IAA) [135]. Tryptic peptides are desalted using strong cation 765 exchange (SCX) columns [136,137] or ZipTip® pipette tips prior to mass spectrometric measurement (LC-MS<sup>2</sup>). MS<sup>2</sup> 766 is used to identify the sequence of the peptide for database matching, with MASCOT [87,88] or SEQUEST [89] as 767 the search algorithm. MS<sup>1</sup> is used for quantification as the normalised spectral abundance factor (NSAF) [69] and stable isotope probing yields the relative isotope abundance (RIA) and labelling ratio (LR), estimated by 768 769 MetaProSIP [106]. Taxonomy and function are assigned via PROPHANE [68]. The raw data are finally uploaded to 770 ProteomeXchange via PRIDE [70].



771Soil organic carbon (%)Microbial biomass (nmol FA g<sup>-1</sup> soil)772Figure 2: The relationship between soil organic carbon and the number of peptides.  $V_{max}$  was 6791.558±404.841773(P<2e-16) and  $K_m$  was 0.496±0.201 (P=0.016) (a). The breaking point was at a  $K_m$  of 1.48695, with 95% confidence774intervals of 0.647 and 3.044. The relationship between microbial biomass and the number of peptides showed a775 $V_{max}$  of 8620.149±1122.251 (P=2.1e-10) and a  $K_m$  of 5.609±2.372 (P=0.0214) (b). The breaking point was at a  $K_m$  of77616.827, with 95% confidence intervals of 7.087 and 37.799. The data were obtained from a gradient of semi-arid777soils that covered a range of soil organic carbon concentration and microbial biomass estimated through the778phospholipid fatty acids (PLFAs) content [48].

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# 783 Congratulations to Journal of Proteomics' 10th Anniversary

The authors of this review would like to congratulate "Journal of Proteomics" on its 10th Anniversary. We are working in soil metaproteomics since a few years ago and it is always convenient to have a forum with recognized expertise where our research can be properly understood. This is particularly true when considering complex biological samples such as soils that require technical and bioinformatic developments. We are grateful to "Journal of Proteomics" for publishing some of our studies that are contributing to understand soil microbial ecology beyond genomics and keep us motivated to work further in the topic.