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A multi-lake comparative analysis of the General Lake Model (GLM): Stress-testing across a global observatory network

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- The General Lake Model (GLM) is stress tested against 32 globally distributed lakes.
- There was low correlation between input data uncertainty and model performance.
- Model performance related to lake-morphometry, light extinction and flow regime; deep, clear lakes with high residence times had the lowest model error.
- Predictions of temperature were less sensitive to model parameters than thermocline depth and Schmidt stability.

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Abstract

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186 The modelling community has identified challenges for the integration and assessment of
187 lake models due to the diversity of modelling approaches and lakes. In this study, we
188 develop and assess a one-dimensional lake model and apply it to 32 lakes from a global
189 observatory network. The data set included lakes over broad ranges in latitude, climatic
190 zones, size, residence time, mixing regime and trophic level. Model performance was
191 evaluated using several error assessment metrics, and a sensitivity analysis was
192 conducted for nine parameters that governed the surface heat exchange and mixing
193 efficiency. There was low correlation between input data uncertainty and model
194 performance and predictions of temperature were less sensitive to model parameters
195 than prediction of thermocline depth and Schmidt stability. The study provides guidance
196 to where the general model approach and associated assumptions work, and cases where
197 adjustments to model parameterisations and/or structure are required.
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206 **Key Words:** lake model, stratification, GLM, model assessment, global observatory data,
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1 Introduction

246 Vörösmarty et al. (2000) urged the international “water sciences community” to work
247 together in the collation and dissemination of hydrological data and modelling techniques
248 to improve our understanding of freshwater ecosystems and “secure a more complete
249 picture of future water vulnerabilities”. Lakes, in particular, are highly valued ecosystems
250 as they provide important water and food resources, and numerous other ecosystem
251 services (Wilson and Carpenter 1999). Human activities such as fresh water diversion and
252 increased nutrient loading, in addition to indirect pressures from climate change, have
253 led to an increased vulnerability of lakes on a global scale (Folke et al. 2004). These
254 challenges have given rise to international networks of scientists such as the Global Lake
255 Ecological Observatory Network (GLEON: gleon.org). Collaborative networks can take
256 advantage of shared data, techniques, and expertise to enable scientists to address the
257 ecological challenges facing lakes globally (Eigenbrode et al. 2007; Adams 2012; Goring
258 et al. 2014). GLEON was initiated in 2005 as a grassroots science community with a vision
259 to observe, understand and predict freshwater systems at a global scale (Weathers et al.
260 2013).

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271 Collaboration between scientists and synthesis of data collected through international
272 networks has led to advances in our understanding of how lake ecosystems respond to
273 external changes and contribute to effective lake management on a local (Gal et al. 2009),
274 regional (Read et al. 2014; Trolle et al. 2015) and global scale (O'Reilly et al. 2015).
275 Analyses based on data from a broad spectrum of lakes across the globe have provided
276 insight into metabolism and carbon cycling in lakes (Hanson et al. 2011; Solomon et al.
277 2013), the role of wind and heat exchange in lake physics (Read et al. 2012), the impact
278 of climate change (Adrian et al. 2009), response and recovery of lakes to extreme events
279 (Jennings et al. 2012; Klug et al. 2012), incorporation of high frequency data for model
280 validation (Hamilton et al. 2015) and assisted in development of models (Staehr et al.
281 2010; Read et al. 2011; Kara et al. 2012; Hipsey et al. 2017). Further interrogation of the
282 emerging multi-lake datasets offers the potential to advance our understanding of how
283 lakes respond to pressures such as climate or land use change from the individual to
284 global scales.

The collaborative network also creates opportunities for developing and testing modelling tools. Aquatic ecosystem models are recognised as essential instruments to improve understanding of processes, analyse relationships, test hypotheses and predict the state of a system (Trolle et al. 2012). These models have evolved since the first attempts in the early 1920s, with a recent review of aquatic ecosystem models revealing the diversity of existing models from simple 0-D to complex 3-D coupled hydrodynamic-biogeochemical models (Janssen et al. 2015). This diversity creates challenges for integration and synthesis of model approaches (Mooij et al. 2010). The Aquatic Ecosystem Modelling Network (AEMON: <https://sites.google.com/site/aquaticmodelling/home>) originated to foster collaboration and improve model development, predictability, transparency and reliability. One of the major challenges facing modellers is how to develop generic models that can capture the diversity of ecosystems while allowing prediction with confidence of the processes of each system. In order to undertake analytical synthesis across multiple sites, there is a need to assess the transferability of the underlying model and standardise its structure, parameterisation, development and examination. While the need to develop a set of standards for model assessment and reporting is widely recognized (Bennett et al. 2013; Grimm et al. 2014), the ability to test these standards across multiple systems and highlight both strengths and limitations of a particular model remains a challenge.

For lakes and reservoirs in particular, one-dimensional (1-D) models that resolve vertical profiles of temperature and density have found widespread use due to their computational efficiency and minimal calibration requirements. The reduced complexity of 1-D models is advantageous whenever greater computational efficiency is needed, e.g., in ensemble modelling (Trolle et al. 2014), model inter-comparison projects such as LakeMIP (<http://www.unige.ch/climate/lakemip>) (Stepanenko et al. 2010; Thiery et al. 2014), probabilistic studies (Schlabing et al. 2014), long-term scenario analysis (Gilboa et al. 2014) or when linking lake models to global climate models (Balsamo et al. 2012) or catchment models (Hipsey et al. 2015). Moreover, lake managers and reservoir operators prefer models having a simpler application and often rely on 1-D models for this reason (Kerimoglu and Rinke 2013; Weber et al. 2017).

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363 Here we introduce the Multi-Lake Comparison Project (MLCP) undertaken within
364 AEMON. The MLCP is a community driven project, where teams of modellers simulate
365 lakes using common approaches for model setup, assessment and analysis. The
366 underlying purpose of the project was to bring together an international network of
367 scientists and modellers with diverse experience in order to improve our ability to predict
368 how lake ecosystems respond to external drivers. In the first stage, the MLCP took
369 advantage of GLEON and AEMON member data from numerous, diverse lakes to stress
370 test the recently developed General Lake Model (GLM) (Hipsey et al. 2017). GLM is a 1-D
371 hydrodynamic model for use in a broad spectrum of enclosed aquatic ecosystems such as
372 lakes, reservoirs and wetlands. The model is simple in nature and is based on assumptions
373 that are common to previous model applications (Imberger and Patterson 1989; Hamilton
374 and Schladow 1997; Coats et al. 2006). The model conducts a lake mass and energy
375 balance to compute vertical profiles of temperature, salinity and density while accounting
376 for the effect of inflows and outflows, surface heating and cooling, mixing and ice cover on
377 the lake. GLM can be coupled with biogeochemical models to explore the impact of
378 temperature, stratification, and vertical mixing on the dynamics of lake ecology (e.g.
379 Snortheim et al. 2017).
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382 This paper summarises the first phase of the MLCP to develop and stress-test GLM. The
383 stress-test involved applying a single standardised procedure for model set-up,
384 simulation, performance testing and analysis to 32 lakes from across the global network.
385 The main objective of this study was to undertake comparative analysis of model
386 performance using an unprecedented diversity of lake types in order to advance our
387 understanding of limnology and contemporary modelling practices. The specific aims of
388 the study were to:
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- 400 1. ascertain levels of model performance and relate it to model input uncertainty;
- 401 2. identify lake attributes (e.g. depth, inflows, and climate) that correspond with high
402 (or low) prediction accuracy;
- 403 3. relate sensitivity of model output variables to changes in surface exchange, heating
404 and mixing parameters that characterise 1-D lake models;
- 405 4. document the transferability of the model without recalibration of individual
406 parameters among lakes, even where these lakes may strongly differ in their
407 properties; and

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5. provide guidance to lake modellers as to how to focus data collation and model application efforts to improve predictions for lake ecosystems.

433 To ease readability, this main section of the paper includes all text as well as tables and
434 figures relevant to the major methodology and results from the study. Additional data
435 have been provided in the following four appendices as supplementary material to the
436 main study:

- 437 A, describing uncertainty error associated with the model set up;
438 B, extended results describing model performance;
439 C, extended results of the sensitivity analysis; and
440 D, a summary of acknowledgements for each lake.

441 2 Methods

442 2.1 Study site selection

443 Lakes were not chosen a priori based on their attributes, but rather AEMON and GLEON
444 members were invited to participate in the MLCP by volunteering details of their
445 candidate lake to the group (shared via open access spreadsheet). The requirement for
446 inclusion of a lake was based on the following three conditions:
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- 448 1. sufficient temperature data were available for validation (at least 2 years of
449 monthly/regular thermistor chain and/or profile data);
450 2. high-resolution meteorological forcing data from an on-lake buoy or local
451 terrestrial based station were available; and
452 3. gauged or well-estimated inflows and outflows were available over the simulation
453 period to form a reliable lake water balance.
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455 Participants were also required to have a basic knowledge of lake modelling. Instructions
456 as to how to set-up the GLM test cases, and a common binary executable (GLM v2.2.0)
457 were made available for download from the Aquatic EcoDynamics (AED) website
458 (<https://github.com/AquaticEcoDynamics/GLM>). Pre- and post-processing MATLAB
459 scripts were provided to all participants to ensure a common model setup and assessment
460 approach (<https://github.com/AquaticEcoDynamics/GLMm>), and all GLM lake setups
461 were available to other members via a cloud-based, shared folder.
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A total of 32 lakes was chosen for the analysis, with an alphabetic listing of the lakes and their physical characteristics in Table 1. Each lake is associated with a two letter abbreviated code, and for brevity when presenting model results, the lakes are frequently referred to by this code. To illustrate the range of sizes in the lakes included in this study, lake outlines have been drawn to scale in Figure 1. With the exception of lakes Geneva and Kinneret, all lake simulations were run for two years, with the start year and date indicated in Table A3. For Lake Geneva and Lake Kinneret, analyses were performed separately for two alternative 2-year time periods with significant differences in climate and inflows. For Lake Geneva, 2003 to 2004 had higher than average summer air temperatures, precipitation and inflows as well as an uncharacteristically high winter inflow in early 2004. In contrast, 2001 to 2002 experienced closer to the “normal” seasonal cycles of climate and inflows (Anneville et al. 2010). These simulations are referred to as Geneva03 and Geneva01 respectively. For Lake Kinneret, 1997 to 1998 had generally average climatic conditions (Bruce et al. 2006). In contrast, 2003 to 2004 had a rainy winter (Feb-Mar 2003, Jan-Feb 2004), large changes to lake level and lower than normal water temperatures (Berger and Telzch 2005). These simulations are referred to as Kinneret97 and Kinneret03, respectively.

Lake depths ranged from 2.4 to 440 m, and lake surface areas from 104,000 m² to 579,000,000 m² (Table 1). A comparative plot of the hypsographic curves for each of the 32 lakes shows diversity in lake size and bed slope (Figure A1). Annual average inflows ranged from 0 to 3.3×10^7 m³ d⁻¹ and residence times from 1 month to 67 years (Table A3). Lake elevation ranged from 209 m below to 4718 m above sea level (Table 1). Annual average air temperature ranged from below freezing (-9.1°C) to 22.4°C (Table A3). While the majority of the lakes in the MLCP are mid-latitude (both northern and southern hemisphere), two lakes are located in the Arctic (Emaiksoun and Toolik).

2.2 GLM set-up

GLM has several configuration options for simulating surface heating, mixing and inflow and outflow (Hipsey et al. 2017). For this assessment, model set-ups were configured based on the site-specific conditions (e.g., hypsographic curve and number of inflows and outflows), but all simulations adopted the same model algorithms and parameters for mixing, surface heat fluxes, and ice cover. Default parameters adopted are summarised in Table 2.

All simulations were run for 2 years or 730 days starting with initial conditions in the winter or when the lake was most nearly well mixed. For the northern hemisphere lakes the start date was the 1st of January and for lakes located in the southern hemisphere the start date was set at 1st July. The initial conditions were taken from the closest field profile measurements to the start date. The standardised start date was chosen to simplify cross lake comparisons. For the majority of the lakes in the MLCP, mid-winter is also associated with complete mixing thus reducing error associated with uncertainty in initial profiles. A spin up period of 28 days was eliminated from model analysis to further reduce error associated with uncertainty in initial conditions.

Box plots are used to present monthly means and range of input data across all 34 simulations (Figure 2). For input data for each lake, refer to references listed in Table 1 and/or the institutions listed in Table D1. Inflows and outflows are also plotted as monthly averages based on time from the beginning of the simulation (Figure 3a&b). There are no seasonal patterns apparent in the monthly inflows and outflows averaged over the MLCP lakes due to the large variation in peak flow months.

While an effort was made to use lakes with high quality input data, lakes where input data had to be estimated were still selected for the MLCP in order to ensure a sufficient variation in lake characteristics. For seven lakes either inflow, outflow or both were estimated (Bourget, Emaiksoun, Feeagh, Mendota, NamCo, Stechlin and Woods) and the parameter of light attenuation (K_w) was estimated for three lakes (Alexandrina, Muggelsee and Woods). Meteorological data for short wave radiation, air temperature, relative humidity, wind speed and precipitation were supplied either from an on lake station or the closest meteorological station to the lake. Long wave radiation was either measured directly (net or incident) or calculated by GLM using cloud cover data.

In an attempt to assess the errors associated with input data limitations, a qualitative weighting system was used to assess each input variable or constant, where a minimum score is associated with the best available input or observation data (Table A1). Table A2a lists the method of determining the hypsographic curve, distance from lake and frequency of meteorological data and observed data and method of determining inflow, outflow and

extinction coefficient for each lake in the MLCP. This information is used to determine the relative error scale associated with boundary forcing and observed data for each lake (Table A2b), where low refers to low uncertainty in forcing data and high indicates a higher level of error associated with model input. Input error associated with the determination of long wave radiation was not included in the error scaling method.

2.3 Model assessment approach

Measures of model fit used to evaluate model performance included five alternatives listed below. This set of measures of model fit enabled us to standardise comparisons among lakes, track trends in deviations from observed data (Bennett et al. 2013) and to compare with similar lake modelling studies previously published (e.g. Rigosi et al. 2010).

Measures of model fit were calculated as:

- 1) Root mean square error (*RMSE*):

$$RMSE = \sqrt{\frac{\sum_{i=1}^N (P_i - O_i)^2}{N}} \quad (2-1)$$

- 2) Model Efficiency (*MEFF*; Murphy, 1988; Nash and Sutcliffe, 1970):

$$MEFF = 1 - \frac{\sum_{i=1}^N (P_i - O_i)^2}{\sum_{i=1}^N (O_i - \bar{O})^2} \quad (2-2)$$

- 3) Correlation coefficient (*r*):

$$r = \frac{\sum_{i=1}^N (P_i - \bar{P})(O_i - \bar{O})}{[\sum_{i=1}^N (P_i - \bar{P})^2 \sum_{i=1}^N (O_i - \bar{O})^2]^{1/2}} \quad (2-3)$$

- 4) Percent relative error (*PRE*) :

$$PRE = \frac{\sum_{i=1}^N (P_i - O_i)/O_i}{N} * 100 \quad (2-4)$$

- 5) Normalised mean absolute error (*NMAE*) :

$$NMAE = \frac{\sum_{i=1}^N |(P_i - O_i)/O_i|}{N} \quad (2-5)$$

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663 where N is the number of observations, O_i and P_i , the “ i^{th} ” observed and model predicted
664 data and \bar{O} and \bar{P} the mean observed and model predicted data, respectively.
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667 A further advantage of calculating alternative measures of model fit is that different
668 methods of model evaluation highlight different aspects of model performance (Bennett
669 et al. 2013). $RMSE$ is a standard measure of the average deviation of simulated values from
670 observations with values near zero indicating a close match and units that correspond to
671 those of the variable. $MEFF$ is the square of the deviation of simulated values from
672 observations, normalized to the standard deviation of the observed data, such that one
673 indicates perfect fit and zero indicates that the model provides equal predictive skill as
674 the mean of the observed data. The correlation coefficient r gives an indication of the
675 linear relationship between observed and predicted data and is the most common
676 measure for assessing aquatic models (Arhonditsis and Brett 2004). PRE is a measure of
677 the relative deviation of simulated from observed values and can be used to determine
678 the bias in predictions (Bennett et al. 2013). Finally, $NMAE$ is both normalised to the mean,
679 enabling like comparisons between variables and is absolute so that under and over
680 estimations do not cancel each other out.
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683 Initial manual calibration focused on refining input data by adjusting the wind scaling
684 factor and river inflow slope parameters for each lake (the river slope is indicated as ϕ_{inf}
685 in Hipsey et al. (2017), and they are denoted as `wind_factor` and `strmbd_slope` in the
686 configuration file, respectively). Wind factor adjustment was required where wind
687 stations were located some distance from the lake and/or to account for wind sheltering
688 effects (Markfort et al. 2010). River inflow slope was adjusted to correct the magnitude of
689 momentum and entrainment associated with plunging inflows. For lakes where few or no
690 light attenuation or Secchi depth readings were available, K_w was also adjusted until
691 simulated thermocline depth matched that of observed data. Initial calibration was
692 carried out until an $RMSE$ (calculated for all observed temperature data over the
693 simulation period) of less than 2°C was achieved.
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696 We chose a range of thermal metrics to assess model performance at each site: observed
697 full profile temperature data; epilimnion temperature; hypolimnion temperature;
698 thermocline depth and Schmidt Stability (Idso 1973). Schmidt Stability (S_T) and
699 thermocline depth (*thermD*) were calculated for both model output and observed
700 data.
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thermistor data using Lake Analyzer (<http://lakeanalyzer.gleon.org/>), an open source software tool that computes indices of mixing and stratification for lakes and reservoirs (Read et al. 2011). The comparison of *thermD* calculations was included in the analysis as it is a simple, widely-used metric of mixed layer depth, while acknowledging the calculation of *thermD* can be challenging for weakly stratified and polymictic lakes. Also, the approach used in Lake Analyzer identifies the strongest thermal gradient, and may miss important thermal structure. S_T represents resistance to mechanical mixing due to the potential energy inherent in the stratification of the water column, calculated as:

$$S_T = \frac{g}{A_s} \int_0^{z_D} (z - z_v) \rho_z A_z dz \quad (2-6)$$

where g is the acceleration due to gravity, A_s is the surface area of the lake, A_z is the area of the lake at depth z , z_D is the maximum depth of the lake, and z_v is the depth to the centre of volume of the lake, and ρ_z is the water density at depth z . While not used as a direct gauge of model performance, the daily Lake Number (L_N) output as a GLM diagnostic parameter was also used in the cross lake comparison analysis as a measure of the validity of the one-dimensional assumption of the model. L_N balances the strength of stratification to wind induced mixing across the thermocline and is a measure of the potential for mixing across the thermocline (Imberger and Patterson 1989).

$$L_N = \frac{S_T(z_e + z_h)}{2\rho_h u_*^2 A_s^{1/2} z_v} \quad (2-7)$$

where z_e and z_h are the depths to the top and bottom of the metalimnion, respectively, ρ_h is the average density of the hypolimnion and u_* is the surface friction velocity.

2.4 Sensitivity analysis

Sensitivity of model output to nine parameters of mixing and heat exchange was evaluated for each lake. Three of the parameters influence surface heat and momentum exchange: bulk aerodynamic coefficient for sensible heat transfer (C_H), bulk aerodynamic coefficient for latent heat transfer (C_E) and coefficient of wind drag (C_D). The remaining six parameters control surface and hypolimnetic mixing: mixing efficiency for convective

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783 overturn (C_C), mixing efficiency of wind stirring (C_W), mixing efficiency of shear
784 production (C_S), mixing efficiency of unsteady turbulence (C_T), mixing efficiency of Kelvin-
785 Helmholtz turbulent billows (C_{KH}), and mixing efficiency of hypolimnetic turbulence
786 (C_{HYP}), (Table 2). To gauge a response to parameter change, the one-at-a-time (OAT)
787 method (Bruce et al. 2008) was adopted for the first stage of the MLCP, where the model
788 was first run with the model default value to each parameter and then run again
789 increasing and decreasing parameter values by 20%.

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795 Sensitivity to changes in parameter values for each of the five lake variables used in the
796 model assessment described above (temperature of the full water column, epilimnion,
797 hypolimnion, $thermD$ and S_T) was analysed. Normalised sensitivity coefficients (S_{ij}) to
798 assess the relative sensitivity of variable i to parameter j were calculated according to:
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$$S_{ij} = \frac{\Delta C_i / C_{is}}{\Delta \beta_j / \beta_{js}} \quad (2-8)$$

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803 where ΔC_i is the change in output variable i , averaged over the simulation period, from
804 the standard or reference value C_{is} (Table 2) and $\Delta \beta_{js}$ is the change in parameter j from the
805 reference value β_j (Fasham et al. 1990).

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808 Sensitivity coefficients were then compared relative to ten characteristics describing the
809 morphometry, climatic conditions and trophic state of the lakes. These properties were
810 the maximum depth, lake volume, ratio of area to maximum depth, ratio of length to
811 width, annual average inflow, residence time, mean air temperature, mean short wave
812 radiation, mean wind speed and extinction coefficient (Table A3).

813 814 3 Results

815 3.1 Model Performance

816 Using the simulated results from running GLM with the standard set of parameters, five
817 model fit metrics ($RMSE$, $MEFF$, r , PRE and $NMAE$) were calculated for five data sets (full
818 profile, epilimnion, hypolimnion temperature, $thermD$ and S_T) for each lake. The full set of
819 results is provided in Appendix B (Table B1) with $NMAE$ results given in Table 3. A
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comprehensive description of model performance for each lake can be found in the plots of modelled versus observed temperature data included in Appendix B.

An analysis of model performance in the prediction of temperature profiles (full profile) demonstrated a robust fit for GLM across the selected metrics, with an average *RMSE* of 1.34°C, *MEFF* of 0.88, *r* of 0.96, *PRE* of -0.16% and *NMAE* of 0.11 (Table B1). The lakes with the lowest *RMSE* included Feagh, Tarawera and Emaiksoun. The highest *RMSE* values were calculated for Ravn, Ammersee and Woods. Ammersee also recorded the lowest values for *MEFF* along with NamCo and Toolik. All values of *r* were > 0.9, with the exception of Toolik. The *PRE* values ranged from +18% for NamCo to -15% for Rassnitzersee. Because lakes had both positive and negative *PRE* (representing a temperature bias, warm and cold respectively) the mean *PRE* was -0.16%. The lowest absolute *PRE* was for GrosseDhuenn (0.33%) which also performed well on all five measures of model fit.

In general, the model performance predicting the epilimnion temperatures was of similar magnitude to the full-profile temperatures (*RMSE* mean = 1.62°C). By analysing the *PRE*, it is clear that the GLM tended to produce both warm and cold temperature biases in the epilimnion, slightly favouring a cold bias (mean *PRE* = -0.84%). For most lakes, model performance metrics were similar for the epilimnion as the full profile with the exception of Windermere and Zurich which performed worse and Oneida which performed better in the computation of epilimnion temperatures.

For the hypolimnetic temperature simulations, average *RMSE* and *NMAE* values were relatively low, 1.31°C and 0.14 respectively. Typically small seasonal variation across all lakes led to greater percentage error between model and simulated data with both warm and cold temperature biases and a tendency to a warm bias (mean *PRE* = 1.97%). The mean *r* value of 0.73 was the lowest of the three temperature-associated properties. Lakes with the highest model performance for hypolimnion temperature included Geneva01, Geneva03 and Como with the lowest being Rassnitzersee, Esthwaite and Blelham. Model efficiency values for the calculation of hypolimnion temperatures were poor with less than a third greater than 0.5 and 44% of lakes recording a value of less than zero.

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903 Thermocline depth (*thermD*) was a difficult parameter to model with the poorest *PRE* and
904 *NMAE* values (Tables 3 & B1). Measures of model performance comparing calculations of
905 observed and simulated *thermD* ranged in value across the lakes with *PRE* values from -
906 16% to +52% and *NMAE* ranging from 0.10 to 0.76 (Tables 3 & B1). The *PRE* values
907 indicate a bias towards over prediction of *thermD* by the model compared to the observed
908 data. This was most apparent in Lake Geneva over the winter months when GLM
909 predicted full mixing (i.e. *thermD* = lake depth) and the field data recorded a shallow
910 *thermD* (<5m). As the lake depth was >300m this resulted in large relative error of greater
911 than 6000%, leading to unfavourable mean measures of fit.
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914 The *NMAE* values for calculation of S_T were generally low. The higher values of *NMAE*
915 were associated with lakes such as Ammersee, Oneida and Pusiano which all had
916 relatively low S_T during the simulated period. The mean *MEFF* and r were both quite high
917 (0.83 and 0.96, respectively) indicating that the general seasonal patterns for S_T
918 prediction across the majority of lakes were well simulated by the model.
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920 Analysis of the relationship between indices of model fit and input quality showed some
921 correlation for the prediction of full profile, epilimnion and hypolimnion temperatures
922 and *thermD* (Table B2). Analysis of measures of *PRE* indicated a cold bias in prediction of
923 both full profile and hypolimnion temperatures when input uncertainty is greatest
924 (Figure 4b). In addition, for lakes where the meteorological measurement station was
925 near or at the lake edge, there was a warm bias and for lakes where meteorological input
926 was sourced from further away, there was a cold bias (Figure 4a). Similarly, there was a
927 warm bias for the prediction of hypolimnetic temperatures for lakes with high frequency
928 meteorological data and a cold bias for lakes with daily meteorological data (Figure 4c).
929 Lakes with lowest input uncertainty associated with the estimation of K_w corresponded
930 with lowest values of r with respect to the prediction of full-profile temperatures (Figure
931 4d) and similarly lakes that had close to ideal ranking of overall input uncertainty scored
932 the lowest values of r for epilimnion temperatures (Figure 4e). This would be attributed
933 to the use of K_w as a calibration parameter for lakes where there were no measurements
934 for light attenuation. High frequency observed data also correlated with high *NMAE*
935 scores for the prediction of hypolimnion temperatures (Figure 4f).
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Analysis of model performance revealed a number of significant correlations linking model performance to lake characteristics (Table B3). For comparison of absolute model performance, the *RMSE* metric was used for temperatures and *MEFF* for *thermD* and *S_T*. Whilst measurements of *PRE* can be a deceptive measure of model performance for lake variables where under and over-prediction occurs in equal measure, they are useful to observe patterns of bias in model prediction. A number of significant correlations between lake characteristics and model error are illustrated in Figure 5 and Figure 6 and described below.

The *RMSE* error associated with the prediction of both full profile and hypolimnion temperatures was generally higher for lakes with high light extinction ($K_w > 0.8 \text{ m}^{-1}$) and lower for clear lakes ($K_w < 0.3 \text{ m}^{-1}$) (Figure 5a&b). A correlation was observed between the *RMSE* associated with the prediction of hypolimnion temperatures and lake depth (Figure 5c), with deep lakes (>100 m) having the lowest values of *RMSE* (<1°C). In terms of relative measures of model performance, for lakes with both low inflows ($< 10^5 \text{ m}^3\text{s}^{-1}$) and low levels of incident short wave radiation averaged over the entire simulation period ($<120 \text{ Wm}^{-2}$) there was a cold bias in prediction of full profile and epilimnion temperatures, respectively (Figure 5c&d). Whilst correlation was relatively low, there was some indication that for lakes with low residence time there was a cold bias in the GLM-predicted hypolimnetic temperatures (Figure 5f).

For prediction of *S_T*, the lake depth, residence time and extinction coefficient all had a significant impact on model performance (Figure 6a, b & c). Generally, clear deep lakes (>100 m), with residence times > 2 years recorded the lowest values of *NMAE*. A reverse pattern of correlation was observed for the prediction of *thermD*, with deep lakes having the highest values of *NMAE* and shallow lakes (<40m) showing highest levels of *thermD* predictive accuracy (Figure 6d). There was a small but significant trend where GLM over estimated *S_T* in lakes with high incident short wave radiation ($>200 \text{ Wm}^{-2}$) (Figure 6e). For prediction of *thermD*, GLM tended towards over-prediction which was more pronounced in colder lakes (air temperature < 10°C) (Figure 6f).

Model performance for the prediction of *thermD* and *S_T* was better for lakes when mean $L_N > 10$, while these lakes tended to record reduced measures of model fit for the

prediction of epilimnion and hypolimnion temperatures (Figure 7a,c,e,g). Conversely, for the small number of lakes with a significant proportion of the stratification period under a regime of $L_N < 1$, prediction of epilimnion and hypolimnion temperatures improved but $thermD$ and S_T decreased (Figure 7b,d,f,h).

3.2 Sensitivity Analysis

The sensitivity analysis (SA) on each of the nine surface exchange and mixing parameters highlighted differences both between lakes and thermal properties (Figure 8a-e). For all three temperature metrics (full profile, epilimnion or hypolimnion) there was little sensitivity to perturbations in physical parameters, when the SA was averaged over the 2 year simulation period. There was some degree of sensitivity to changes in C_d in the calculation of hypolimnion temperatures and to C_e in the calculation of epilimnion temperatures. Sensitivity index (SI) for prediction of both $thermD$ and S_T , were significant (>1) across a broader range of lakes (Figure 8d-e). While there was some variability across the lakes and parameters, model output for both $thermD$ and S_T had greatest sensitivity to perturbations of C_d . Additionally, for S_T there was a consistent level of sensitivity to perturbations of C_e .

The sensitivity of each parameter was compared to a gradient of physical and climate lake properties (Table C1-5) and a number of significant correlations were observed. For each thermal metric, the three most significant correlations to lake characteristics were compared (Figure 9). A common significant ($p<0.05$) trend was recorded for maximum lake depth (Figure 9e, Figure 9m). For the prediction of full profile and epilimnion temperatures, deeper and larger lakes were more sensitive to changes in C_{KH} than small, shallow lakes (Figure 9e). Similarly, for the prediction of $thermD$, deeper lakes were more sensitive to changes in C_c , C_w and C_{KH} than shallow lakes (Figure 9).

A significant correlation with air temperature indicated that lakes with low air temperatures were more sensitive to changes in C_h , C_s and C_{KH} than lakes in warm climates (mean air temperature $> 10^\circ\text{C}$) for the prediction of full-profile temperature (Figure 9c), epilimnion (Figure 9d) and hypolimnion temperatures (Figure 9g). Lakes with low inflows were more sensitive to changes in C_h for the prediction of hypolimnion temperatures than those with larger inflows (Figure 9i). Finally, lakes with highest wind speed recorded greatest SI to C_e in the prediction of S_T (Figure 9m).

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1084 **4 Discussion**
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1087 Historically, lake modellers have adopted simple methods to justify model performance
1088 and suitability, rarely reporting statistical measures of model fit (Arhonditsis and Brett
1089 2004; Arhonditsis et al. 2006). For individual lake applications, these have been adequate
1090 to undertake scenario simulations and further our understanding of site specific
1091 dynamics. However, a common approach to model assessment, both in terms of metrics
1092 that should be applied and identification of a commonly agreed level of model
1093 performance, is necessary to further enhance model development (Bennett et al. 2013).
1094 Undertaking a standardized method of assessment of the community lake model, GLM,
1095 over a diversity of lakes has led to an improved level of understanding of the strengths
1096 and weaknesses in the predictive capacity of simple 1-D lake models. By first ascertaining
1097 an acceptable model error, we were able to elucidate the relation between model
1098 performance and data input uncertainty or lake characteristics (Figure 4; Figure 5).
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1101 The quality of input data was not as significantly related to model performance as
1102 expected. Lakes modelled using daily meteorological input, rather than hourly, did have
1103 the largest values of *NMAE* in the prediction of full profile temperature and *thermD*
1104 (Figure 4), which is not surprising given the importance of diurnal forcing in 1-D model
1105 predictive capability. The greater the meteorological observation distance to the lake
1106 tended to result in both cold-biased temperatures and under prediction of *S_T* (Figure 4).
1107 The cause of warm-biased temperatures and over-prediction of lake stability when
1108 meteorological observations were obtained near or on-lake requires further investigation
1109 (Figure 4). The strong correlation between accuracy of *K_w* measurements and model
1110 performance in the prediction of both full profile temperature and *thermD* (Figure 4)
1111 emphasises both the importance of light extinction in the determination of thermocline
1112 depth and the need to include measurements of *K_w* in routine lake monitoring. The GLM
1113 can be coupled to water quality models such as the Aquatic EcoDynamics Model (AED:
1114 Hipsey et al. 2013) such that seasonal changes in *K_w* would feedback in the model to
1115 potentially improve model prediction particularly in relation to *thermD*; this link is
1116 expected to further improve model accuracy in most circumstances.
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The 1-D nature of the model implicitly assumes that the mixing within the lake can be constrained by processes acting in the vertical and that processes which vary in the horizontal, such as the degree of upwelling of the thermocline, have minimal impact on vertical transport. This assumption is quantified by computation of the Lake Number (Imberger and Patterson, 1989; eq. 2.7). As the L_N is a relative measure of the strength of stratification to surface wind energy, the 1-D model assumption is said to hold true for $L_N \gg 1$ (Imberger and Patterson 1989; Yeates and Imberger 2003). Over the past three decades, the 1-D model approach has been applied to a wide diversity of sites due to its simplicity and tractability relative to 3-D models. However, given that L_N can be highly variable, it has remained unclear what significance the 1-D assumption has on model prediction error for various lake attributes and under what conditions this assumption would no longer hold. The strong correlation ($r^2=[0.70,0.82]$) between the percent of time $L_N < 1$ during the stratified period and the model performance of both *thermD* and *S_T* endorses the use of L_N as an indicator of the validity of the 1-D model assumption, and should be considered when modellers are deciding on model suitability.

A comparison of *PRE* against L_N for the calculation of simulated versus observed *S_T* indicated that lakes with mean $L_N < 1$ tended to underestimate *S_T*. For these lakes, the 1-D assumption as defined by L_N does not hold. One would expect mixing to be underestimated and *S_T* to be higher, unless the resulting warmer near surface temperatures led to greater heat losses by evaporation. Yeates and Imberger (2003) demonstrated that for lakes where deep mixing is important, a 1-D lake model mixing scheme similar to that used in GLM tended to overmix the water column and thus underestimate lake stability and therefore *S_T*. A solution put forward by Yeates and Imberger (2003) included a pseudo two-dimensional algorithm in the 1-D model DYRESM to parameterise internal and boundary fluxes. Similarly Gaudard et al. (2016) proposed a method of adding a seasonal component in the parameterisation of internal seiches that led to improved accuracy in the prediction of deep mixing in the 1-D model SIMSTRAT. Whilst compromising computational efficiency, lake modellers could consider a similar approach when conditions for improved deep mixing accuracy are necessary. For example, this approach could be valuable where upwelling or internal nutrient loading is deemed important or when specific distribution phenomena such as deep chlorophyll maxima are the focus of the modelling study.

Further exploration of how individual lake properties relate to measures of model performance indicated the strongest correlations against K_w and lake depth (Figure 5; Figure 6). Lakes with high K_w (> 0.5), recorded greatest error in the prediction of lake temperatures particularly in the hypolimnion. While there was no significant correlation between the accuracy in prediction of epilimnion temperatures and lake depth, there was a strong positive correlation for measures of model performance in prediction of hypolimnion temperatures and depth (Figure 5). That is, for deeper lakes (>40 m) where surface mixing dynamics have less influence on hypolimnion temperatures, GLM predicts hypolimnion temperatures with greater accuracy. This suggests that while the surface thermodynamics are better represented by the model, prediction of rates of mixing across the metalimnion requires attention and further development to enable more confident prediction across the diversity of lake types. Relatively shallow, well-mixed lakes, such as Feeagh and Emaiksuon, had the highest overall model performance. These lakes are dominated by surface exchange with no thermocline and associated deepening.

The prediction of the lake thermocline depth proved harder to achieve than the lake temperatures. Particularly in moderately deep lakes, small relative deviations in predictions can result in large changes to error magnitude. As the GLM-predicted *thermD* was both deeper and shallower than the observed *thermD* in different lakes, there does not appear to be a consistent bias in the mixing algorithms, and rather, it may be driven by high sensitivity to input parameter uncertainty and require site specific calibration. The positive correlation between *NMAE* of thermocline prediction and lake depth was significant with best fit occurring for lakes less than 50-80 m deep (Figure 6). A tendency to over-predict thermocline depth in the majority of lakes could be attributed to an over-prediction of penetrative heat and may be related to both the application of a standard minimum layer thickness for all lakes and the use of a single average K_w value over 2 annual seasonal cycles. The positive correlation with K_w indicates that a single K_w for all seasonal conditions is not appropriate, particularly for lakes with high mean or seasonally variable K_w values. A consideration for using a K_w weighted towards the summer stratified period could be a solution or coupling to a water quality model with explicit light extinction feedback properties could improve thermocline prediction particularly in lakes with high light extinction ($K_w > 0.5$) (Shatwell et al. 2016).

The absence of strong sensitivity to parameterisation of surface exchange and mixing algorithms in the prediction of temperature profiles (Figure 8) is indicative of the dominance of surface boundary conditions in the thermal budget of individual lakes and negative feedbacks in the surface heating sub-model. In contrast, the prediction of thermocline depth and Schmidt Stability were more sensitive to changes in parameterisation. In particular, the model was sensitive to the shear mixing efficiency and wind drag coefficient parameters. Both parameters are directly related to the transfer of wind energy to mixing. The errors in computing these terms again points to the need for more effort in parameterizing the processes operative when L_N is low and shear increases across the thermocline. Additionally, wind increases in magnitude as it flows across a lake. This effect is important for small and large lakes and is not included when wind is modelled with bulk drag coefficients. Care should be taken in both the accuracy of wind speed measurements as well as the parameterization and classification of these parameters in relation to lake characteristics to improve model performance across a wide variety of lake properties.

In general, simulations of deep lakes with large volumes and residence times were most sensitive to changes in mixing efficiency parameters (as measured by changes in *thermD* and *St*) (Figure 9), which was expected since larger lakes require greater efficiency in transfer of surface momentum input to thermocline deepening and subsequent mixing. Lakes with low K_w were most sensitive to changes in surface exchange parameters. This sensitivity is logical given that in lakes with low K_w , light will penetrate deeper causing a deeper thermocline. Processes which moderate depth of mixing in the epilimnion, such as convection, become important. Being able to model changing dynamics of lakes as K_w changes with modified hydrology and altered loading of chromophoric dissolved organic matter is critical for quantifying the changes associated with climate variability (Snucins and Gunn 2000).

An appealing alternative to the minimal calibration presented here (i.e., input data refinement, wind factor and river inflow slope adjustment) will be the relaxation of the assumption of globally common parameter values for the core hydrodynamic parameters and the adoption of a Bayesian hierarchical calibration framework that reflects the more

realistic notion that each lake (or group of lakes) is peculiar but shares some commonality of behavior with other lakes (Zhang and Arhonditsis 2009; Cheng et al. 2010; Shimoda and Arhonditsis 2015). The proposed approach represents a pragmatic compromise between system- or group-specific and globally common parameter estimates and may be a conceptually sound strategy to accommodate within- and among-lake variability in the context of model application within the global observatory network (Figure 10). Recent work has shown that the delineation of more homogeneous subsets of lakes with respect to their morphological characteristics/hydraulic regimes and their subsequent integration with hierarchical frameworks may give models with better predictive capacity (Cheng et al. 2010; Shimoda and Arhonditsis 2015). In particular, sensitivity analysis patterns identified in this study could be used to identify groups with similarities in behavior (e.g., deep versus shallow lakes, high versus low water transparency) as well as to identify the candidate parameters for the calibration exercise. The prior distributions of the hyper-parameters (or global priors) can be easily formulated on the basis of existing knowledge (e.g., field observations, laboratory studies, and information from the modeling literature) of the relative plausibility of their values. Moreover, the proposed incorporation of mathematical models into Bayesian hierarchical frameworks can also assist the effective modeling of systems with limited knowledge by enabling the transfer of information across systems. With the hierarchical model configuration, we can potentially overcome problems of insufficient local data by “borrowing strength” from well-studied lakes on the basis of distributions that connect systems in space (Zhang and Arhonditsis 2009). Another advantage of a Bayesian calibration configuration will be the ability to express the input uncertainty in the form of probability density functions which can then be propagated through the model structure and may ultimately shape the moments of the posterior predictive distributions.

Through international collaboration, this work allowed us to test and to improve the process and performance of a 1-D open source model by simulating thermal structure in lakes with varying physical and climatic characteristics. Initial efforts in setting up a collaborative network of lake modellers were rewarded with improved user support and feedback, refinements and testing to the development team. From its initiation as v1.0 in the MLCP, using feedback and re-coding by network members, the GLM evolved through numerous improvements to the current v2.2 described in this study. The study also

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1383 identified the most sensitive parameters related to surface exchange and mixing that
1384 affect model prediction and therefore performance for each individual lake. These
1385 sensitivities could then be correlated to lake characteristics such as residence time,
1386 meteorological conditions and trophic status. Additionally, this work opens a new
1387 challenge for the community of limnologists involved in ecosystem modelling. Indeed the
1388 next step would be cross lake comparison projects including biogeochemical processes
1389 simulation using a similar open source community biogeochemical model such as the
1390 Framework for Aquatic Biogeochemical Models (FABM: Bruggeman and Bolding, 2014)
1391 and/or AED (Hipsey et al. 2013). The establishment of well-defined standards for
1392 modelling techniques (set up, output analysis), and a diversity of lakes and scientists
1393 provides enormous opportunity for further advances by aquatic ecosystem modellers.
1394 The significance of the MLCP resides in a common and collaborative approach to
1395 answering globally relevant lake science questions, and providing a benchmark for model
1396 performance and an associated parameter set that future applications can refer to.
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1406 networks as well as discussions and working groups held during AEMON workshops and
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2161 Table 1 - Lakes included in the Multi-Lake Comparison Project Stage 1, abbreviation, maximum depth, surface area at maximum depth, crest elevation latitude (°N)
 2162 and longitude (°E).

Lake Name	Abv.	Maximum Depth (m)	Surface Area at Crest (m ²)	Crest Elevation (m)	Latitude	Longitude	Reference
Lake Alexandrina	AL	9.4	655,755,315	3.4	-35.4	139.1	(Hipsey et al. 2014b)
Ammersee	AM	83.7	47,250,000	533.5	48.0	11.1	(Weinberger and Vetter 2014; Bueche et al. 2017)
Blelham	BL	14.5	104,000	14.0	54.4	-3.0	(Woolway et al. 2015)
Lake Bourget	BO	146.0	42,575,000	230.5	45.4	5.9	(Vinçon-Leite et al. 1989, 2014; Kerimoglu et al. 2016)
Cannonsville Reservoir	CA	52.0	19,000,000	351.0	42.1	-75.3	(Samal et al. 2012)
Lake Como	CO	440.0	147,012,649	410.0	46.0	9.3	(Laborde et al. 2010; Copetti et al. 2013; Guyennon et al. 2014)
Lake Constance	CN	253.3	472,650,000	395.0	47.6	9.4	(Wessels 1998; Frassl et al. 2014)
El Gergal	EG	55.0	4,732,669	50.0	37.0	-2.5	(Rigosi et al. 2011)
Emaiksoun	EM	2.4	1,860,000	2.4	71.2	-156.8	(Potter 2011)
Esthwaite	ES	15.5	1,000,000	15.5	54.4	-3.0	(Woolway et al. 2015)
Feeagh	FE	43.0	3,942,266	9.0	53.4	-9.6	(Dalton et al. 2014)
Lake Geneva 2001-2	G1	309.0	578,560,865	371.4	46.4	6.1	(Anneville et al. 2010)
Lake Geneva 2003-4	G3	309.0	578,560,865	371.4	46.4	6.1	(Anneville et al. 2015)
Grosse Dhuenn	GD	48.5	3,750,100	177.5	51.1	7.2	(Weber et al. 2017)
Harp Lake	HA	37.5	713,800	327.0	45.4	-79.1	(Yao et al. 2014)
Lake Iseo	IS	256.0	60,880,350	185.2	45.7	10.1	(Pilotti et al. 2013, 2014; Valerio et al. 2015)
Lake Kinneret 2003-4	K3	44.0	173,000,000	-208.9	32.0	35.6	(Gal et al. 2009)
Lake Kinneret 1997-8	K7	44.0	173,000,000	-208.9	32.0	35.6	(Bruce et al. 2006)
Lake Mendota	ME	25.0	39,581,170	259.0	43.0	-89.4	(Magnuson et al. 2006)
Mount Bold Reservoir	MB	45.4	3,080,000	246.9	-35.1	138.7	(van der Linden and Burch 2016) Rigosi et al. 2015
Muggelsee	MG	8.0	7,318,000	32.4	52.0	13.6	(Huber et al. 2008)
Lake Nam Co	NM	98.9	2,018,230,000	4718.0	30.7	90.6	(Wang et al. 2009)
Oneida	ON	17.0	207,100,000	112.0	43.0	-75.9	(Hetherington et al. 2015)
Lake Pusiano	PU	30.9	8,123,699	27.0	45.8	9.3	(Copetti et al. 2006, 2013; Carraro et al. 2012)

2201	Rappbode	RP	85.6	4,344,724	423.6	51.7	10.9	(Bocaniov et al. 2014)
2202	Rassnitzersee	RS	40.0	3,033,057	85.0	51.3	12.0	(Böhrer et al. 1998; Boehrer et al. 2014)
2203	Ravn	RV	33.0	1,820,000	33.0	56.0	4.8	(Trolle et al. 2008a; b)
2204	Rotorua	RO	22.0	79,722,140	280.0	-38.0	176.3	(Burger et al. 2008)
2205	Stechlin	ST	69.5	4,231,549	60.0	53.2	13.0	(Kirillin et al. 2013)
2206	Tarawera	TA	88.0	40,996,000	297.8	-38.2	176.4	(Hamilton et al. 2006, 2010)
2207	Toolik	TO	24.0	940,119	740.0	68.6	-149.6	(MacIntyre et al. 2009)
2208	Windermere	WI	66.8	14,779,600	66.8	54.4	-3.0	(Woolway et al. 2015)
2209	Woods Lake	WO	10.4	15,000,000	738.2	-42.0	147.0	(Hydro Tasmania 2003)
2210	Lower Lake Zurich	ZU	136.0	66,600,000	406.0	47.3	8.8	(Peeters et al. 2002; Schmid and Köster 2016)

Table 2 - Description, symbols and initial values of the parameters used in the sensitivity analysis.

Symbol	Description	Reference	Initial value
Surface Heat Exchange			
C_h	Bulk aerodynamic coefficient for sensible heat transfer	(Fischer et al. 1979)	0.0013
C_e	Bulk aerodynamic coefficient for latent heat transfer	(Fischer et al. 1979)	0.0013
C_d	Bulk aerodynamic momentum transfer coefficient	(Fischer et al. 1979)	0.0013
Mixing			
C_c	Mixing efficiency - convective overturn	(Yeates and Imberger 2003)	0.2
C_w	Mixing efficiency - wind stirring	(Spigel et al. 1986)	0.23
C_t	Mixing efficiency - unsteady turbulence (acceleration)	(Sherman et al. 1978)	0.3
C_s	Mixing efficiency - shear production	(Sherman et al. 1978)	0.51
C_{KH}	Mixing efficiency - Kelvin-Helmholtz turbulent billows	(Sherman et al. 1978)	0.3
C_{hyp}	Mixing efficiency of hypolimnetic turbulence	(Weinstock 1981)	0.5

Table 3 - *NMAE* for base simulations using standard parameter set against full profile temperature (Full Prof. Temp.) [°C], epilimnion temperature (Epi. Temp.) [°C], Hypolimnion temperature (Hyp. Temp.) [°C], thermocline depth (*thermD*) [m] and Schmidt Stability (*S_t*). Note that for fully mixed lakes or for lakes where temperature profiles were shallower than the thermocline depth, *NMAE* values are listed as not applicable (N/A). N refers to the number of profiles used in the calculation of model performance.

Lake	Full Prof. Temp. (°C)	Epi. Temp. (°C)	Hyp. Temp. (°C)	thermD (m)	ST	N
Alexandrina	0.07	0.07	N/A	N/A	N/A	
Ammersee	0.19	0.20	0.13	0.40	0.17	
Blelham	0.12	0.13	0.31	0.18	0.45	
Bourget	0.08	0.11	0.07	0.32	0.09	
Cannonsville	0.10	0.05	0.15	0.39	0.12	
Como	0.10	0.17	0.06	0.64	0.19	
Constance	0.08	0.09	0.07	0.11	0.16	
ElGergal	0.08	0.06	0.07	0.30	0.27	
Emaiksoun	0.08	0.08	N/A	N/A	N/A	
Esthwaite	0.13	0.11	0.35	0.15	0.24	
Feeagh	0.06	0.04	0.09	0.14	0.30	
Geneva01	0.09	0.11	0.04	0.41	0.22	
Geneva03	0.08	0.05	0.04	0.52	0.20	
GrosseDhunn	0.07	0.05	0.09	0.37	0.09	
Harp	0.18	0.12	0.27	0.68	0.19	
Iseo	0.08	0.10	0.07	0.76	0.16	
Kinneret03	0.07	0.07	0.07	0.28	0.20	
Kinneret97	0.05	0.06	0.05	0.15	0.21	
Mendota	0.11	0.10	0.11	0.30	0.23	
MtBold	0.08	0.08	0.06	0.25	0.43	
Muggelsee	0.07	0.06	N/A	N/A	N/A	
NamCo	0.23	0.17	0.22	0.28	0.35	
Oneida	0.04	0.03	0.06	0.19	0.86	
Pusiano	0.14	0.11	0.26	0.24	0.19	
Rappbode	0.14	0.08	0.12	0.23	0.16	
Rassnitzersee	0.17	0.15	0.23	0.15	0.17	
Ravn	0.19	0.14	0.21	0.27	0.34	
Rotorua	0.07	0.08	0.08	0.09	0.43	
Stechlin	0.13	0.11	0.11	0.33	0.14	
Tarawera	0.04	0.04	0.03	0.27	0.10	
Toolik	0.25	0.26	0.25	0.61	0.43	
Windermere	0.14	0.23	0.26	0.22	0.21	
Woods	0.17	0.17	N/A	N/A	N/A	
Zurich	0.12	0.09	0.16	0.42	0.17	
Mean	0.11	0.10	0.14	0.32	0.25	

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2362	
2363	
2364	Median
2365	0.09
2366	0.09
2367	0.10
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2470	observatory network. Future applications can improve parameter accuracy	
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2484 best relates 53
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Figure 1 – Lake outlines to scale for all lakes in the current MLCP GLM assessment.

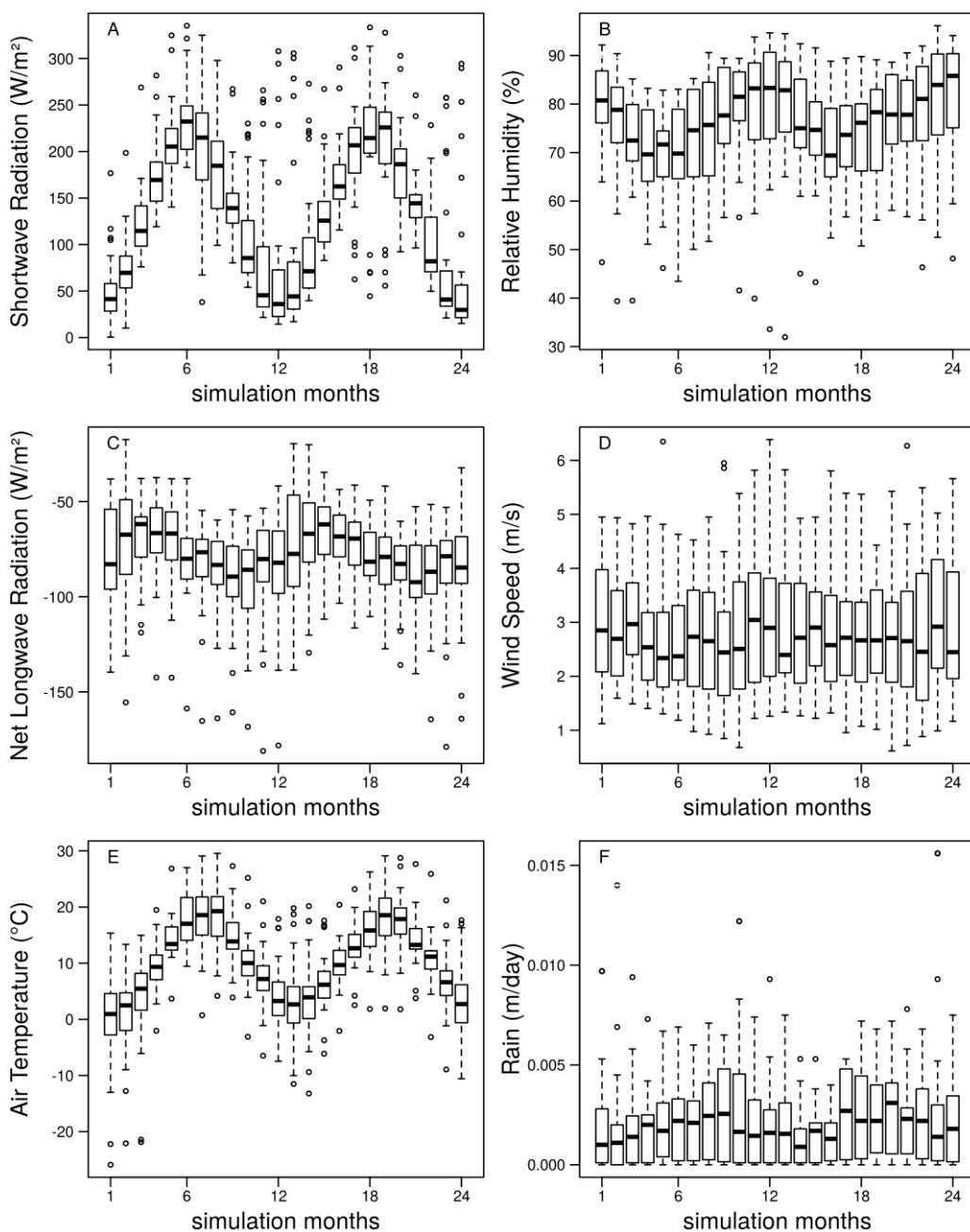


Figure 2 - Time series of monthly mean values across all lakes for (a) short wave radiation, (b) relative humidity, (c) net longwave radiation, (d) wind speed, (e) air temperature and (f) precipitation. For each box, horizontal lines represent median, 25th and 75th percentile, whiskers < 1.5 times the interquartile, and outliers (○) values > 1.5 times the interquartile range. Note that lakes from the Southern Hemisphere start with a shift of 6 months relative to the Northern Hemisphere lakes.

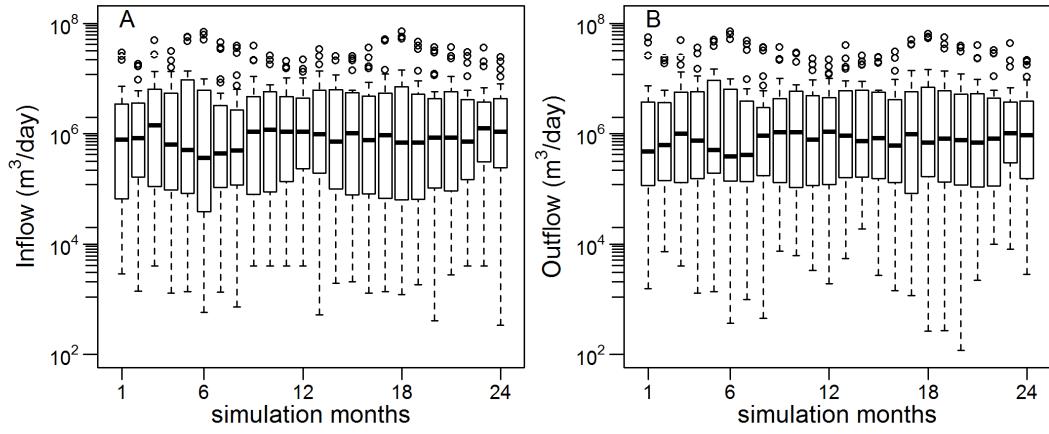


Figure 3 – Time series of monthly mean values across all lakes for (a) inflows and (b) outflows. For each box, horizontal lines represent median, 25th and 75th percentile, whiskers < 1.5 times the interquartile, and outliers (\circ) values > 1.5 times the interquartile range.

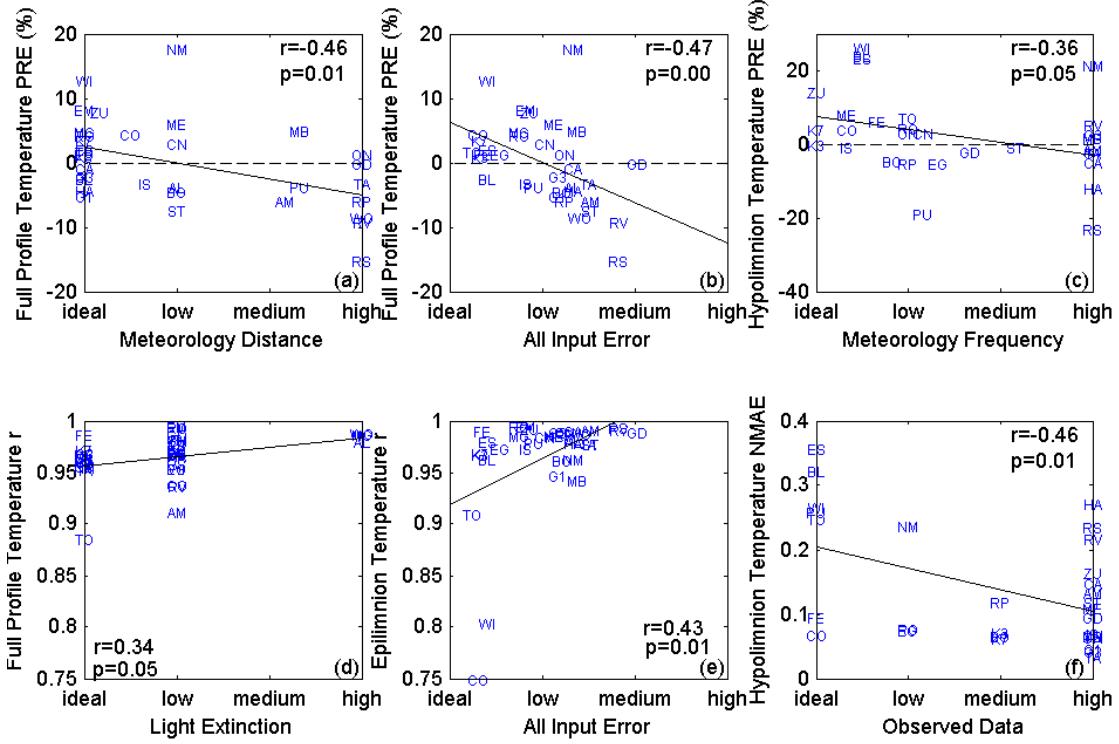


Figure 4 – Correlation between GLM model performance metrics PRE (a-c), r (d-e) and NMAE (f) for prediction of full profile temperatures (a, b & d), epilimnion temperatures (e) and hypolimnion temperatures (c & f) against rankings of input data uncertainty where 0-ideal, 1-low, 2-medium and 3-high level of uncertainty. Refer to Table 1 for lake acronyms and Table A1 for details of input uncertainty ranking system.

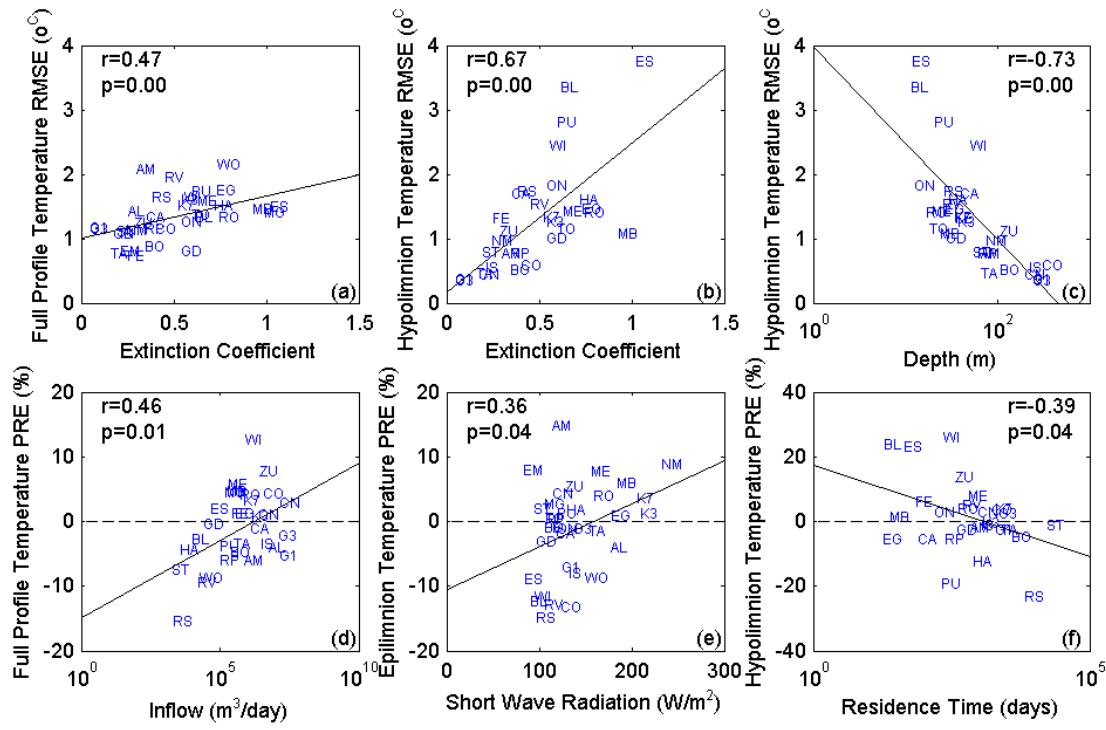


Figure 5 - GLM model performance metrics for prediction of full profile temperature (a&d), epilimnion temperature (e) and hypolimnion temperature (b,c&f) against lake characteristics. Refer to Table 1 for lake acronyms.

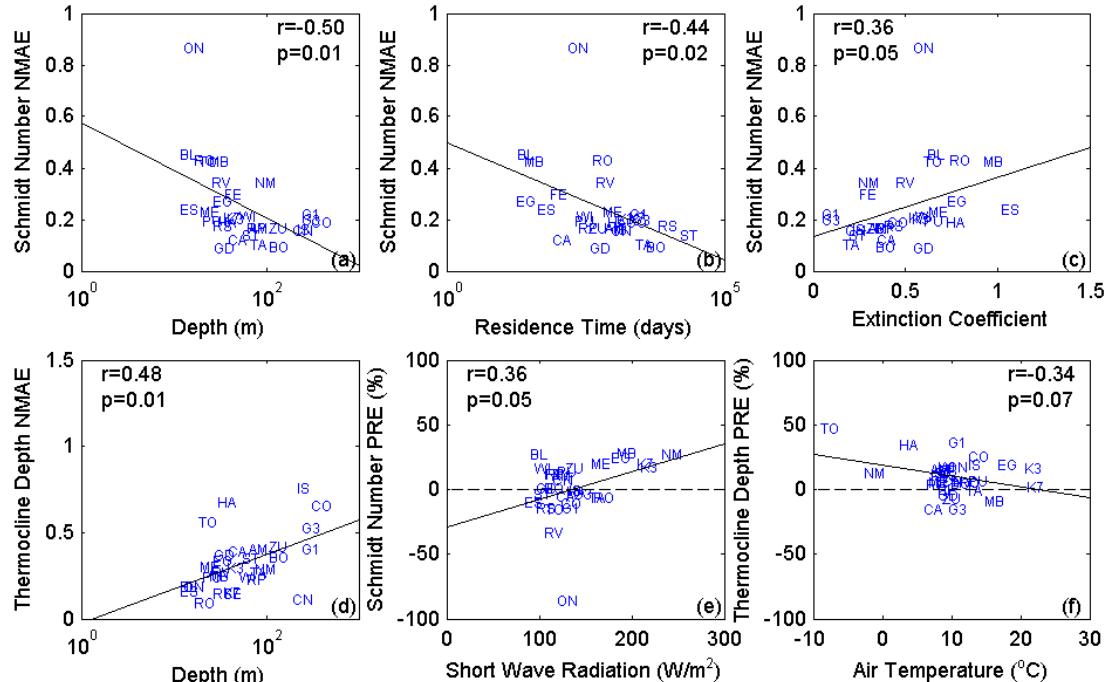


Figure 6 - GLM model performance metrics for prediction of thermocline depth (d,f) and Schmidt stability (a,b,c&e) against lake characteristics. Refer to Table 1 for lake acronyms.

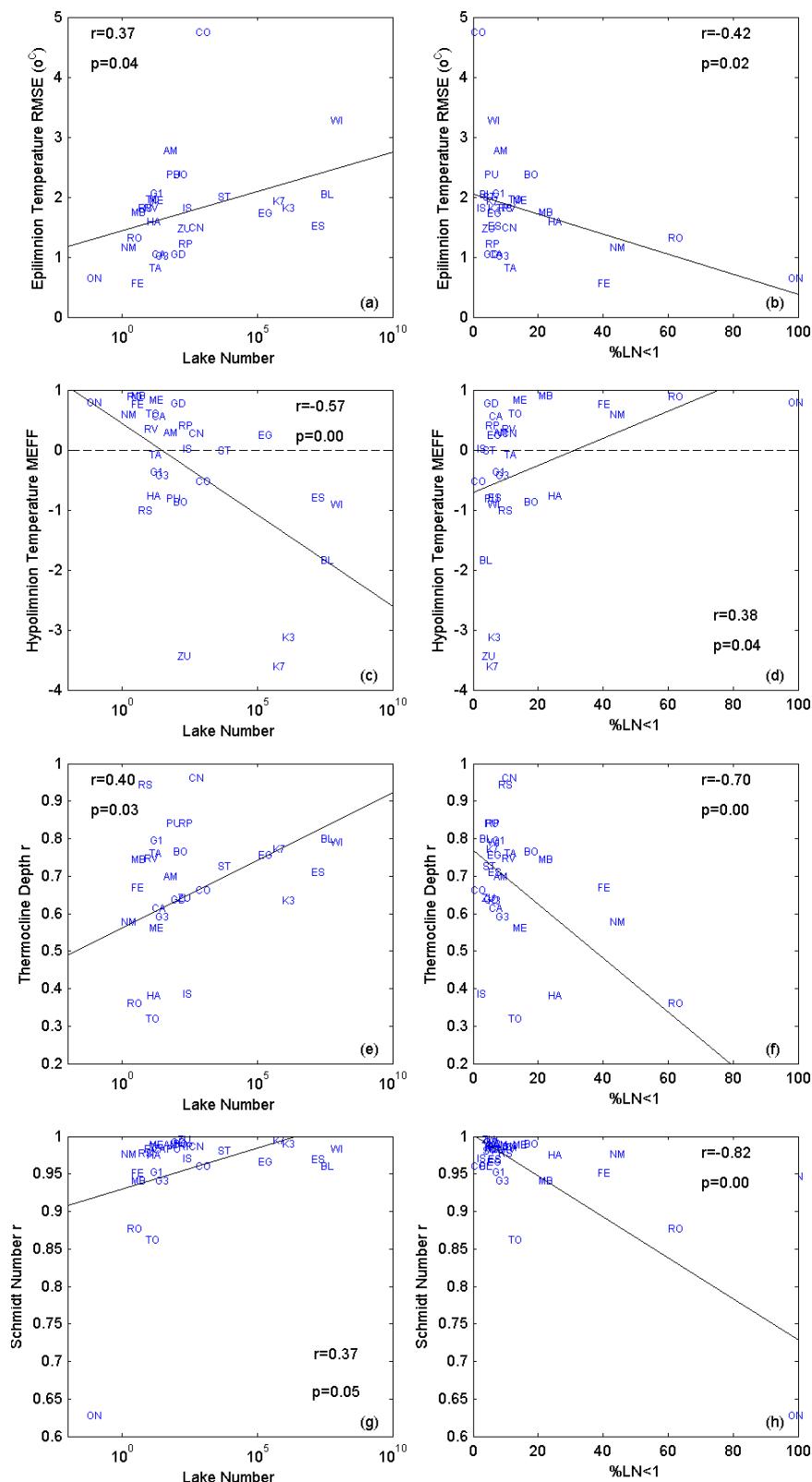
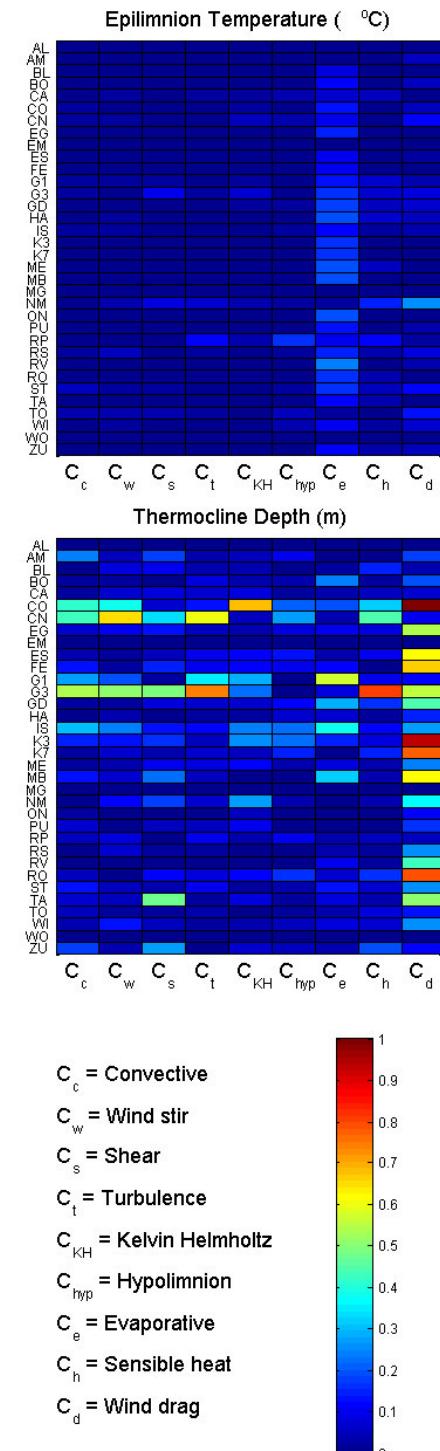
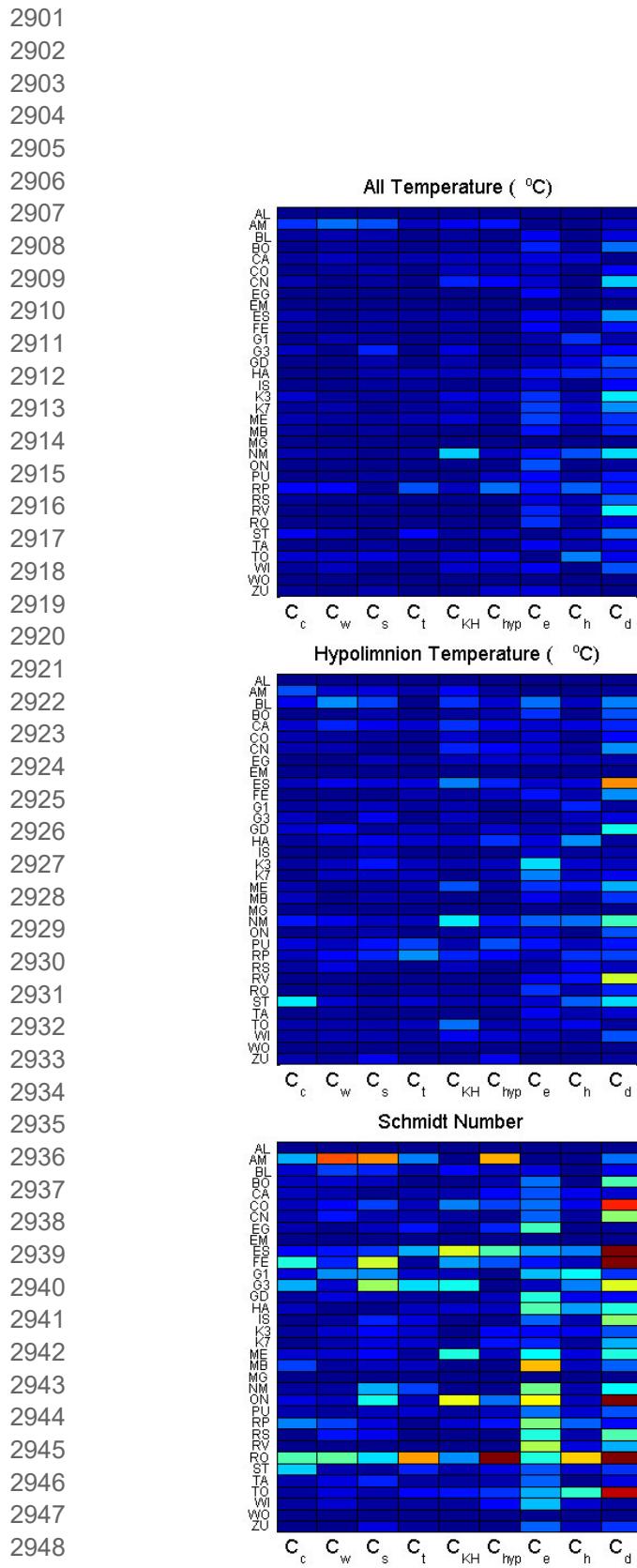
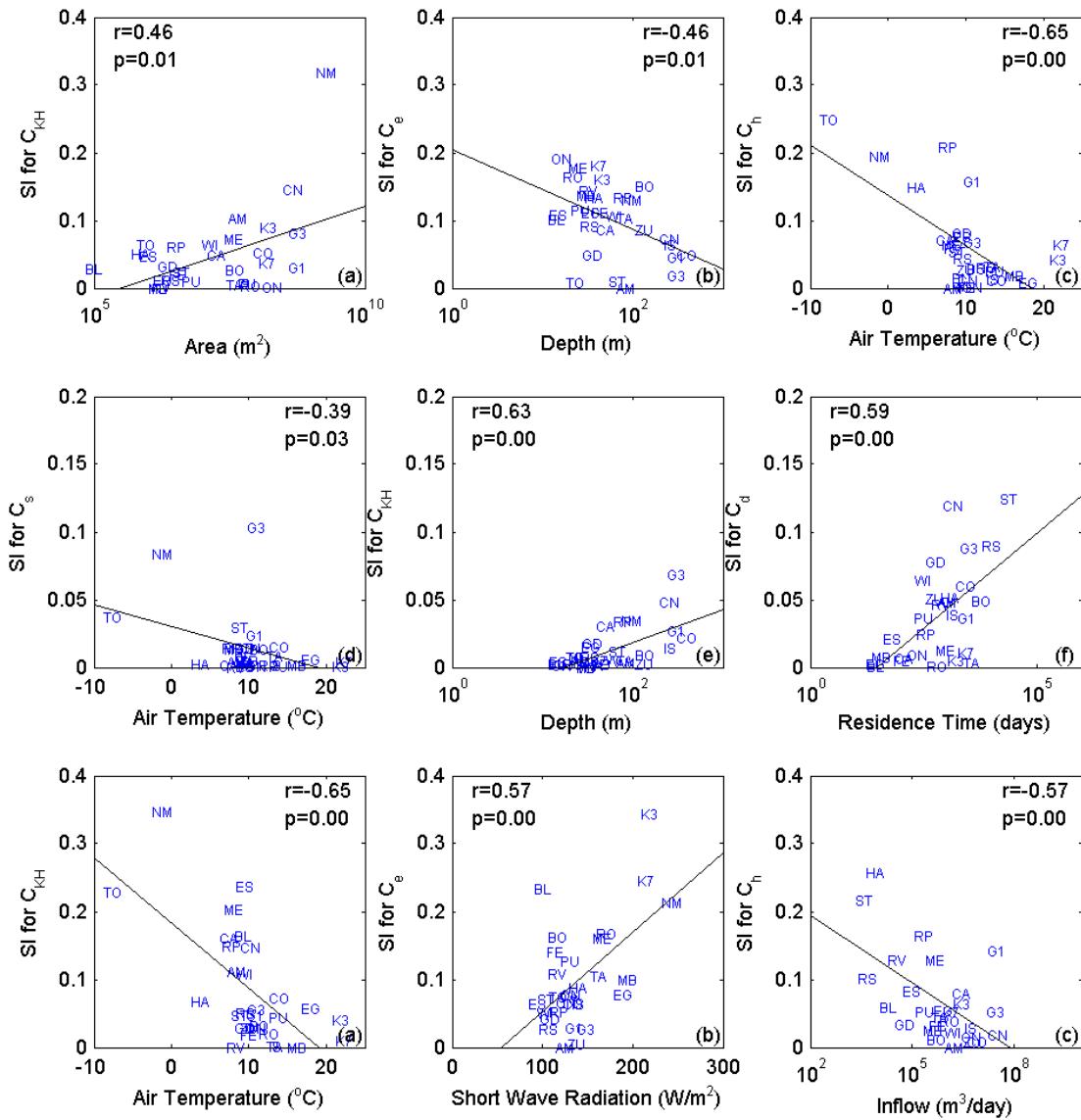


Figure 7 - GLM model performance metrics for prediction of epilimnion temperature (a,b), hypolimnion temperature (c,d), thermocline depth (e,f) and Schmidt stability (gh) against Lake Number and %LN<1. Refer to Table 1 for lake acronyms



2952 Figure 8 - Sensitivity indices for a) full profile temperature, b) epilimnion temperature, c)
2953 hypolimnion temperature, d) thermocline depth and e) Schmidt stability. The colour bar
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2955 (indicating the percent response in thermodynamic metric is greater than the change in
2956 physical parameter) has been highlighted.
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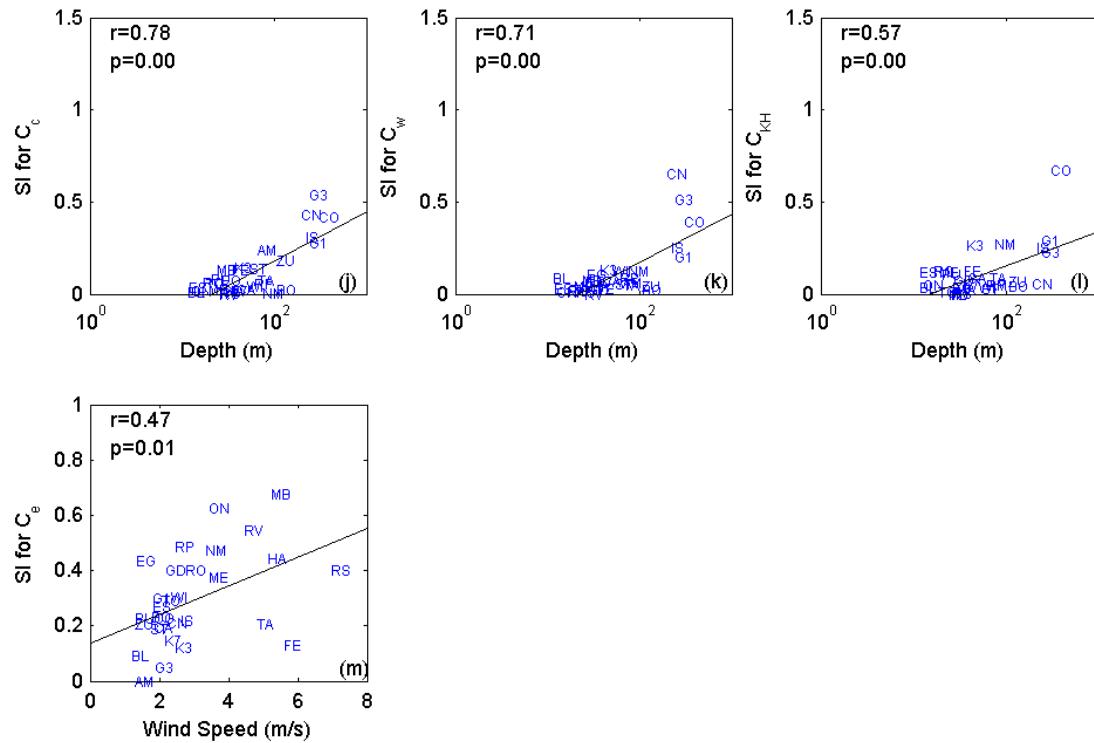
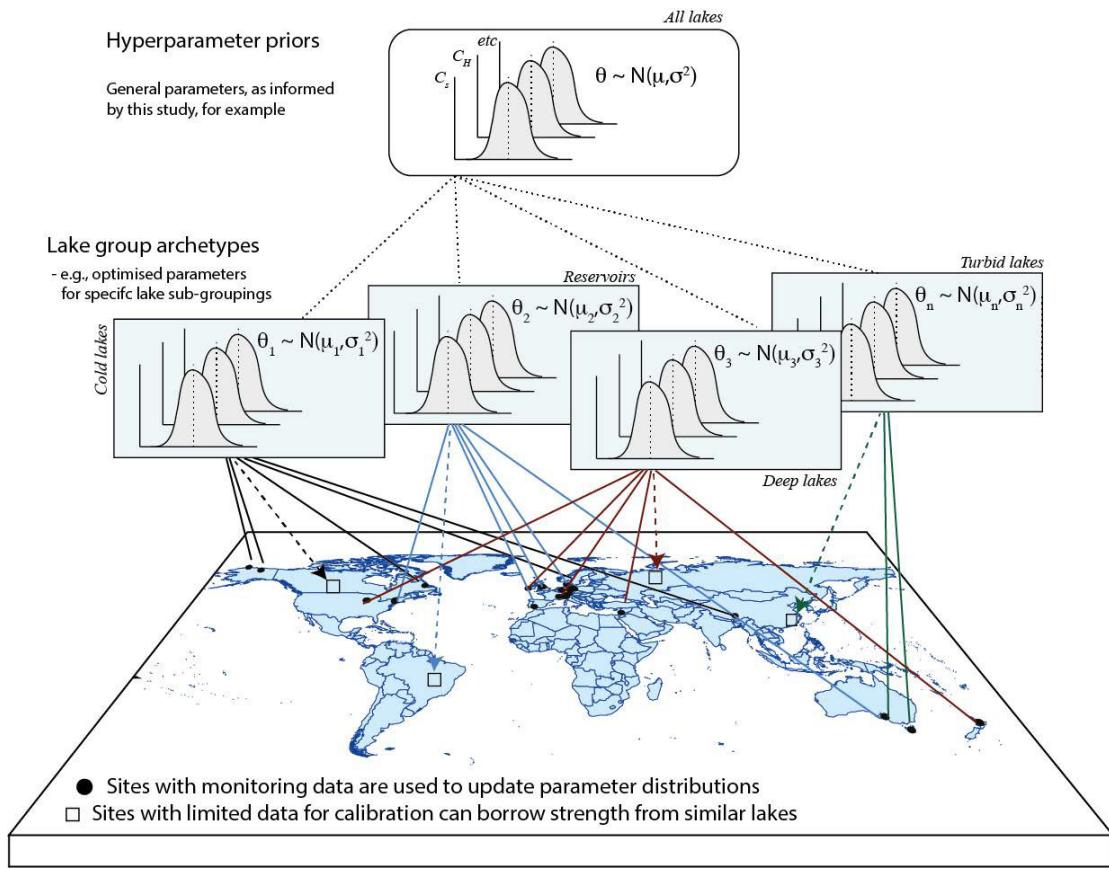


Figure 9 - Significant correlation between sensitivity indices of GLM physical parameters for the prediction of: full profile temperatures and (a) surface area, (b) lake depth and (c) air temperature; epilimnion temperature and (d) air temperature, (e) lake depth and (f) residence time; hypolimnion temperatures and (g) air temperature, (h) short wave radiation and (i) inflow; thermocline depth as a function of lake depth (j-l); and Schmidt stability as a function of wind speed (m).



3109 Figure 10 - A conceptual overview of future lake modelling applications to best
3110 integrate model applications with the increasing volumes of sensor data. In this
3111 study no parameter fitting was undertaken for GLM and parameters presented
3112 herein could be used as the hyperparameter prior for all lakes within the
3113 observatory network. Future applications can improve parameter accuracy
3114 within a Bayesian hierarchical framework based on suitable groupings of lakes
3115 into distinct archetypes. Other lakes with limited data for robust calibration, can
3116 adopt standard model parameters depending on the lake archetype, to which it
3117 best relates.
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 3142
 3143 **5 Appendix A – Model Input**
 3144

3145 Table A1 – Input uncertainty ranking system

Rank	0 - ideal	1 - high	2 - medium	3 - low
Morphometry	digitised		estimated from topographic drawing	estimated
Meteorology - Distance	on lake (< 1km)	<5km from lake	<10km from lake	>10km from lake/ estimated
Meteorology - Frequency	sub-hourly	hourly	sub-daily	daily
Flow	gauged	modelled		estimated
Kw	mean from light measurements	Secchi depth mean from > 12 measurements/year	Secchi depth mean from < 12 measurements/year	estimated
Frequency of Observed Data	>=daily	>= weekly	>=monthly	< monthly

3181 Table A2 – Input data quality for each lake, (a) measurement, (b) rank. A value of 9999 indicates that input data has been estimated.

3182 Lake Name	3183 Morphome try	3184 Distance (km)						3185 Sampling interval (hours)								3186 Number of		
		3187 Short Wave Rad.	3188 Long Wave Rad.	3189 Air Temp.	3190 Rel. Hum.	3191 Wind speed	3192 Precipita tion	3193 Short Wave Rad.	3194 Long Wave Rad.	3195 Air Temp.	3196 Rel. Hum.	3197 Wind speed	3198 Precipitation	3199 Inflow	3200 Outflow	3201 Kw	3202 Obs Data	
Alexandrina	digital	1.2	1.2	1.2	1.2	1.2	1.2	0.25	0.25	0.25	0.25	0.25	0.25	model	model	estimated	14	
Ammersee	digital	10	12	10	10	10	10	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	24	
Blelham	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2920
Bourget	contour	2	2	2	2	2	2	1.00	1.00	1.00	1.00	1.00	1.00	0.10	gauge	estimated	secchi	74
Cannonsville	contour	0	0	0	0	0	0	24.00	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	31
Como	digital	0	50	0	0	0	0	1	0.02	1.00	0.02	0.02	0.02	1.00	gauge	gauge	secchi	2916
Constance	digital	1.2	1.2	1.2	1.2	1.2	1.2	1.00	1.00	1.00	1.00	1.00	1.00	6.00	gauge	gauge	secchi	31
ElGergal	digital	0.3	0.3	0.3	0.3	0.3	0.3	1.00	1.00	1.00	1.00	1.00	1.00	24.00	gauge	gauge	secchi	124
Emaiksoun	digital	0	0	0	0	0	0	0.00	0.00	0.00	0.00	0.00	0.00	0.00	estimated	estimated	secchi	467
Esthwaite	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2799
Feeagh	digital	0.7	0.7	0.35	0.35	0.35	0.35	0.03	24.00	0.00	0.00	0.00	0.00	1.00	gauge	estimated	light	2913
Geneva03	contour	1	1	1	1	1	1	24.00	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	39
Geneva05	contour	1	1	1	1	1	1	24.00	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	30
GrosseDhuenn	estimated	37	9999	22.5	22.5	22.5	9999	1.00	9999.00	1.00	1.00	1.00	9999.00	model	gauge	secchi	17	
Harp	contour	0.5	0.5	0.5	0.5	0.5	0.5	24.00	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	20
Iseo	digital	0	20	0	0	0	0	3	0.03	1.00	0.03	0.03	0.03	1.00	gauge	gauge	secchi	24
Kinneret03	digital	0	0	0	0	0	0	0.02	0.02	0.02	0.02	0.02	0.02	0.02	gauge	gauge	light	93
Kinneret97	digital	0	0	0	0	0	0	0.02	0.02	0.02	0.02	0.02	0.02	0.02	gauge	gauge	light	78
Mendota	contour	2.5	2.5	2.5	2.5	2.5	2.5	<0.01	1.00	<0.01	<0.01	<0.01	<0.01	1.00	gauge	estimated	light	30
MtBold	digital	12	22	10	10	12	5	24.00	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	46
Muggelsee	digital	0	0	0	0	0	0	0.08	24.00	0.08	0.08	0.08	0.08	0.08	model	model	estimated	1747
NamCo	digital	1.6	1.6	1.6	1.6	1.6	1.6	24.00	24.00	24.00	24.00	24.00	24.00	24.00	estimated	estimated	light	731
Oneida	digital	21	21	21	21	21	21	1.00	1.00	1.00	1.00	1.00	1.00	1.00	model	gauge	light	40
Pusiano	digital	11	9999	4	17	11	4	1.00	3.00	1.00	1.00	1.00	1.00	1.00	model	model	secchi	2320
Rappbode	digital	16	16	16	16	16	16	1.00	1.00	1.00	1.00	1.00	1.00	1.00	model	gauge	secchi	58

3221	Rassnitzersee	digital	12	12	12	12	12	12	24.00	24.00	24.00	24.00	24.00	24.00	model	model	secchi	16
3222	Ravn	contour	50	50	50	50	50	50	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	secchi	44
3223	Rotorua	digital	0	0	0	0	0	0	1.00	1.00	1.00	1.00	1.00	1.00	gauge	estimated	secchi	677
3224	Stechlin	digital	5	5	5	5	5	5	3.00	3.00	3.00	3.00	3.00	24.00	estimated	estimated	light	41
3225	Tarawera	digital	15	15	15	15	15	15	24.00	24.00	24.00	24.00	24.00	24.00	gauge	gauge	light	21
3226	Toolik	digital	0.05	0.05	0.05	0.05	0.05	0.05	1.00	1.00	1.00	1.00	1.00	1.00	gauge	model	light	2920
3227	Windermere	contour	0	0	0	0	0	0	0.03	24.00	0.03	0.03	0.03	0.03	gauge	gauge	secchi	2920
3228	Woods	digital	9999	9999	33.6	33.6	33.6	33.6	1.00	1.00	1.00	1.00	1.00	1.00	estimated	gauge	estimated	731
3229	Zurich	contour	0.5	2.0	0.5	0.5	0.5	0.5	0.17	0.17	0.17	0.17	0.17	0.17	gauge	gauge	secchi	24
3230																		
3231																		
3232																		
3233																		

3234	Lake Name	Morphometry	Distance		Sampling interval					Number of									
			Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation	Short Wave Rad.	Long Wave Rad.	Air Temp.	Rel. Hum.	Wind speed	Precipitation	Inflow	Outflow	Kw	Obs Data	Mean
3235	Alexandrina	0	1	1	1	1	1	1	0	0	0	0	0	0	1	1	3	3	1.33
3236	Ammersee	0	2	3	2	2	2	2	3	3	3	3	3	3	0	0	1	3	1.53
3237	Blelham	1	0	0	0	0	0	0	0	3	0	0	0	0	0	0	1	0	0.42
3238	Bourget	1	1	1	1	1	1	1	1	1	1	1	1	0	0	3	1	2	1.22
3239	Cannonsville	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	1	3	1.33
3240	Como	0	0	3	0	0	0	0	0	1	0	0	0	1	0	0	1	0	0.31
3241	Constance	0	1	1	1	1	1	1	1	1	1	1	1	2	0	0	1	3	1.03
3242	ElGergal	0	0	0	0	0	0	0	1	1	1	1	1	3	0	0	1	1	0.56
3243	Emaiksoun	0	0	0	0	0	0	0	0	0	0	0	0	0	3	3	1	1	0.83
3244	Esthwaite	1	0	0	0	0	0	0	0	3	0	0	0	0	0	0	1	0	0.42
3245	Feeagh	0	0	0	0	0	0	0	0	3	0	0	0	1	0	3	0	0	0.36
3246	Geneva03	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	0	3	1.17
3247	Geneva05	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	0	3	1.17
3248	GrosseDhuenn	3	3	3	3	3	3	3	1	3	1	1	1	3	1	0	1	3	2.03
3249	Harp	1	0	0	0	0	0	0	3	3	3	3	3	3	0	0	1	3	1.33
3250	Iseo	0	0	3	0	0	0	1	0	1	0	0	0	1	0	0	1	3	0.83

3261	Kinneret03	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	2	0.33
3262	Kinneret97	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	0	2	0.33
3263	Mendota	1	1	1	1	1	1	1	0	1	0	0	0	0	1	0	3	0	3	1.14
3264	MtBold	0	3	3	2	2	3	1	3	3	3	3	3	3	0	0	0	0	3	1.39
3265	Muggelsee	0	0	0	0	0	0	0	0	3	0	0	0	0	1	1	1	3	0	0.75
3266	NamCo	0	1	1	1	1	1	1	3	3	3	3	3	3	3	3	3	0	1	1.33
3267	Oneida	0	3	3	3	3	3	3	1	1	1	1	1	1	1	0	0	3	1.25	
3268	Pusiano	0	3	3	1	3	3	1	1	2	1	1	1	1	1	1	1	1	0	0.92
3269	Rappbode	0	3	3	3	3	3	3	1	1	1	1	1	1	1	0	1	2	1.25	
3270	Rassnitzersee	0	3	3	3	3	3	3	3	3	3	3	3	3	1	1	1	3	1.83	
3271	Ravn	1	3	3	3	3	3	3	3	3	3	3	3	3	3	0	0	1	3	1.83
3272	Rotorua	0	0	0	0	0	0	0	1	1	1	1	1	1	0	3	1	1	1	0.75
3273	Stechlin	0	1	1	1	1	1	1	2	2	2	2	2	3	3	3	0	3	1.53	
3274	Tarawera	0	3	3	3	3	3	3	3	3	3	3	3	3	0	0	0	3	1.50	
3275	Toolik	0	0	0	0	0	0	0	1	1	1	1	1	1	0	1	0	0	0	0.25
3276	Windermere	1	0	0	0	0	0	0	0	3	0	0	0	0	0	0	0	1	0	0.42
3277	Woods	0	3	3	3	3	3	3	1	1	1	1	1	1	3	0	3	0	1.42	
3278	Zurich	1	0	1	0	0	0	0	0	0	0	0	0	0	0	0	0	1	3	0.83

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3301 Table A3 – Input lake characteristics used for comparative analysis. Values for lake depth, surface area, volume and meteorological data are averaged
 3302 over the simulation period. Area: Depth Ratio, is the ratio between average lake surface area and lake depth.
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3304 Abbrv.	3305 Lake Name	3306 Depth (m)	3307 Surface Area (m ²)	3308 Area: Depth Ratio	3309 Volume (m ³)	3310 Length: Width Ratio	3311 Inflow (m ³ /d)	3312 Res. Time (days)	3313 Short Wave Radiation (W/m ²)	3314 Air Temp. (°C)	3315 Wind Speed (m/s)	3316 K _w (m ⁻¹)	3317 Start Date
AL	Alexandrina	6.1	587,158,666	9.66E+07	1.11E+09	1.14	1.12E+07	98.96	186.98	14.88	3.81	0.30	1-Jul-10
AM	Ammersee	83.8	46,427,854	5.64E+05	1.81E+09	3.00	1.70E+06	1061.06	125.18	8.40	1.58	0.35	1-Jan-09
BL	Blelham	14.7	103,802	7.07E+03	5.78E+05	3.75	2.05E+04	28.14	100.78	9.41	1.46	0.67	1-Jan-08
BO	Bourget	140.5	40,379,890	3.03E+05	3.31E+09	5.83	5.48E+05	6035.65	117.27	11.48	2.08	0.40	1-Jan-09
CA	Cannonsville	49.7	18,014,685	3.83E+05	3.28E+08	27.40	2.80E+06	116.90	128.97	7.36	2.11	0.40	1-Jan-03
CO	Como	400.2	133,890,923	3.67E+05	2.21E+10	15.00	8.42E+06	2618.03	134.28	13.96	2.14	0.46	1-Jan-05
CN	Constance	253.4	465,703,202	1.87E+06	4.70E+10	4.00	3.26E+07	1443.46	126.05	10.26	2.51	0.23	1-Jan-94
EG	El Gergal	34.9	1,951,384	1.35E+05	2.47E+07	19.38	7.40E+05	33.43	189.57	18.14	1.64	0.79	1-Jan-01
EM	Emaiksoun	2.2	1,871,484	8.77E+05	2.89E+06	2.30	0.00E+00	N/A	93.80	-9.09	4.10	0.27	1-Jan-12
ES	Esthwaite	15.0	1,003,558	6.69E+04	6.38E+06	7.67	9.79E+04	65.16	94.79	9.50	2.09	1.07	1-Jan-08
FE	Feeagh	44.9	3,620,712	8.78E+04	6.49E+07	3.86	6.02E+05	107.72	115.20	10.11	5.87	0.30	1-Jan-11
G1	Geneva01	308.9	577,950,600	1.87E+06	9.12E+10	5.21	3.13E+07	2914.51	135.43	10.91	2.10	0.10	1-Jan-03
G3	Geneva03	308.8	577,404,068	1.87E+06	9.12E+10	5.21	2.77E+07	3292.98	148.04	10.95	2.14	0.10	1-Jan-01
GD	GrosseDhuenn	33.8	2,214,418	1.11E+05	3.26E+07	8.36	1.05E+05	309.21	108.20	9.51	2.50	0.60	1-Jan-96
HA	Harp	37.2	706,244	1.92E+04	9.31E+06	1.50	7.96E+03	1170.14	139.26	3.73	5.39	0.77	1-Jan-92
IS	Iseo	260.5	60,829,158	2.34E+05	7.91E+09	8.33	5.48E+06	1443.16	140.28	13.55	2.83	0.25	1-Jan-10
K3	Kinneret03	47.5	171,786,293	3.64E+06	4.96E+09	1.62	2.93E+06	1690.75	219.52	22.02	2.74	0.59	1-Jan-03
K7	Kinneret97	43.6	166,425,534	3.97E+06	4.29E+09	1.62	1.62E+06	2643.24	215.43	22.37	2.41	0.57	1-Jan-97
ME	Mendota	25.0	39,229,728	1.58E+06	4.96E+08	2.00	4.96E+05	1000.35	167.57	8.23	3.74	0.69	1-Jan-09
MB	MtBold	31.7	1,703,574	1.07E+05	1.99E+07	7.33	1.70E+05	116.81	195.67	16.36	5.55	0.98	1-Jul-03
MG	Muggelsee	7.8	7,173,756	9.43E+05	3.38E+07	1.69	3.65E+05	92.55	117.60	10.26	3.85	1.05	1-Jan-04
NM	NamCo	98.9	1,942,514,246	2.04E+07	1.00E+11	2.38	0.00E+00	N/A	244.03	-1.13	3.64	0.30	1-Jan-12
ON	Oneida	16.4	199,785,009	1.26E+07	1.35E+09	3.79	5.68E+06	236.69	131.55	10.96	3.73	0.60	1-Jan-11
PU	Pusiano	26.8	6,537,710	3.03E+05	7.72E+07	1.97	2.39E+05	322.45	131.21	13.91	1.62	0.66	1-Jan-02

3341	RP	Rappbode	79.4	3,453,019	5.47E+04	8.60E+07	16.00	2.31E+05	371.77	118.44	7.82	2.77	0.40	1-Jan-08
3342	RS	Rassnitzersee	34.8	2,714,136	8.71E+04	5.42E+07	1.24	5.14E+03	10551.99	107.95	9.66	4.26	0.44	1-Jan-01
3343	RV	Ravn	32.3	1,748,402	6.00E+04	2.61E+07	1.33	3.57E+04	729.93	116.19	8.37	4.72	0.50	1-Jan-03
3344	RO	Rotorua	21.8	78,779,626	3.66E+06	7.95E+08	1.17	1.23E+06	645.34	170.60	12.64	3.07	0.80	1-Jul-07
3345	ST	Stechlin	69.5	4,230,060	6.11E+04	9.75E+07	0.76	4.00E+03	24364.00	104.49	8.93	2.02	0.25	1-Jan-01
3346	TA	Tarawera	82.8	39,406,707	4.95E+05	2.18E+09	1.17	6.11E+05	3570.44	162.49	13.34	5.07	0.21	1-Jul-02
3347	TO	Toolik	23.5	920,947	3.99E+04	6.18E+06	1.50	1.29E+04	479.13	117.31	-7.54	2.36	0.65	1-Jan-06
3349	WI	Windermere	63.7	14,360,584	2.32E+05	5.15E+08	12.13	1.61E+06	319.72	104.08	9.44	2.60	0.60	1-Jan-08
3351	WO	Woods	9.2	13,451,648	1.63E+06	5.55E+07	1.31	4.93E+04	1125.18	162.37	6.62	4.97	0.80	1-Jul-11
3352	ZU	Zurich	136.0	66,593,085	4.90E+05	3.37E+09	11.20	6.20E+06	543.19	138.46	10.06	1.57	0.34	1-Jan-03
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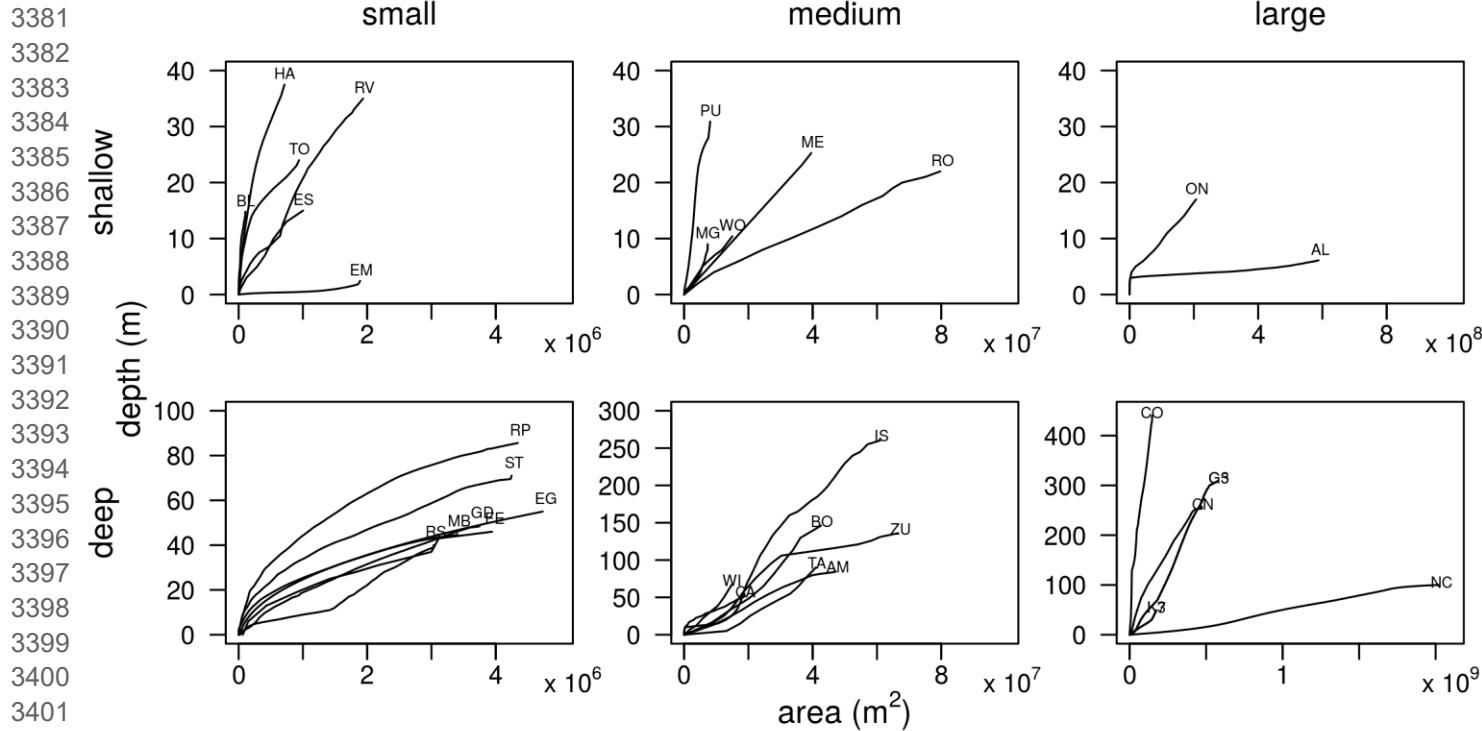
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Figure A1 - Hypsographic curves for (a) small, shallow lakes, (b) medium, shallow lakes, (c) large, shallow lakes, (d) small, deep lakes, (e) medium, deep lakes, (f) large, deep lakes.

6 Appendix B – Analysis of Model Performance

Table B1 – Model performance metrics for base simulations using standard parameter set. Note that for fully mixed lakes or for lakes where temperature profiles were shallower than the thermocline depth, *NMAE* values are listed as not applicable (N/A).

	Full-profile Temperature					Epilimnion Temperature					Hypolimnion Temperature					Thermocline Depth					Schmidt Stability				
	RMS E	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E	RMSE	MEF F	r	PRE	NMA E	RMS E	MEF F	r	PRE	NMA E
Alexandrina	1.43	0.86	0.98	-3.9	0.07	1.44	0.86	0.98	-4.0	0.07	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
Ammersee	2.07	0.79	0.91	-6.0	0.19	2.78	0.84	0.99	14.8	0.20	0.78	0.30	0.59	-1.7	0.13	28.0	0.28	0.70	15.6	0.40	278	0.95	0.99	10.6	0.17
Blelham	1.32	0.89	0.97	-3.0	0.12	2.02	0.85	0.96	-12.0	0.13	3.24	-1.60	0.84	21.6	0.31	2.6	0.54	0.80	-0.2	0.18	24	0.65	0.96	27.0	0.45
Bourget	0.87	0.93	0.97	-4.5	0.08	2.01	0.91	0.97	0.1	0.11	0.53	-0.95	0.40	-4.9	0.07	36.4	0.56	0.79	3.9	0.32	823	0.98	0.99	0.9	0.09
Cannonsville	1.33	0.94	0.97	-1.0	0.10	1.06	0.97	0.99	-1.8	0.05	1.70	0.57	0.79	-5.3	0.15	11.6	0.32	0.62	-15.0	0.39	125	0.96	0.98	-5.7	0.12
Como	1.13	0.86	0.94	4.2	0.10	4.32	0.52	0.78	-10.0	0.17	0.57	-0.49	0.48	3.1	0.06	82.7	-0.43	0.67	26.6	0.64	5498	0.90	0.96	-10.1	0.19
Constance	1.08	0.95	0.98	2.8	0.08	1.49	0.95	0.98	4.4	0.09	0.44	0.25	0.74	2.8	0.07	31.0	0.92	0.96	6.7	0.11	1372	0.95	0.99	7.0	0.16
ElGergal	1.72	0.81	0.95	1.1	0.08	1.54	0.91	0.98	1.4	0.06	1.38	0.38	0.80	-5.6	0.07	8.5	0.55	0.79	16.8	0.30	328	0.74	0.97	24.1	0.27
Emaiksoun	0.80	0.95	0.99	8.0	0.08	0.80	0.95	0.99	8.0	0.08	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A
Esthwaite	1.49	0.89	0.98	1.4	0.13	1.50	0.93	0.98	-8.9	0.11	3.64	-0.66	0.92	21.4	0.35	2.7	0.40	0.71	10.3	0.15	12	0.93	0.97	-9.9	0.24
Feeagh	0.72	0.95	0.99	1.2	0.06	0.53	0.98	0.99	-0.9	0.04	1.30	0.80	0.97	5.8	0.09	9.7	0.41	0.67	6.3	0.14	34	0.90	0.95	11.6	0.30
Geneva01	1.18	0.92	0.96	-5.3	0.09	2.07	0.86	0.95	-7.0	0.11	0.36	-0.34	0.65	-2.6	0.04	94.4	0.46	0.80	36.3	0.41	3350	0.87	0.95	-14.0	0.22
Geneva03	1.16	0.93	0.97	-2.1	0.08	1.02	0.98	0.99	-1.0	0.05	0.34	-0.39	0.67	2.6	0.04	123.0	0.22	0.59	-15.6	0.52	3977	0.88	0.94	-3.6	0.20
GrosseDhuenn	0.81	0.97	0.99	-0.3	0.07	1.05	0.97	0.99	-3.1	0.05	1.02	0.78	0.90	-2.5	0.09	13.5	0.37	0.64	-4.1	0.37	69	0.98	0.99	0.7	0.09
Harp	1.54	0.92	0.96	-4.6	0.18	1.60	0.95	0.98	1.7	0.12	1.63	-0.79	0.70	-12.5	0.27	5.9	-0.48	0.38	34.3	0.68	60	0.94	0.98	-1.9	0.19
Iseo	1.07	0.96	0.98	-3.4	0.08	1.83	0.92	0.97	-8.0	0.10	0.55	0.03	0.56	-1.0	0.07	122.9	-0.33	0.39	19.3	0.76	2620	0.94	0.97	-0.5	0.16
Kinneret03	1.60	0.88	0.96	0.9	0.07	1.76	0.87	0.97	1.4	0.07	1.28	-3.04	0.28	-0.6	0.07	10.9	0.31	0.66	15.4	0.28	527	0.88	0.99	17.3	0.20

3461	Kinneret97	1.49	0.87	0.97	3.1	0.05	1.65	0.89	0.99	4.6	0.06	1.10	-2.16	0.46	3.0	0.05	7.8	0.56	0.79	2.6	0.15	571	0.87	0.99	19.1	0.21
3462	Mendota	1.60	0.92	0.97	5.9	0.11	1.94	0.94	0.98	7.9	0.10	1.42	0.84	0.95	7.8	0.11	7.8	0.15	0.56	5.3	0.30	96	0.88	0.99	20.0	0.23
3463	MtBold	1.47	0.87	0.96	4.8	0.08	1.74	0.80	0.94	6.0	0.08	1.08	0.90	0.96	1.7	0.06	11.4	0.50	0.75	-9.3	0.25	146	0.57	0.94	28.7	0.43
3464	Muggelsee	1.40	0.92	0.99	4.6	0.07	1.24	0.94	0.98	2.7	0.06	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	
3465	NamCo	1.13	0.85	0.95	17.6	0.23	1.04	0.93	0.98	11.1	0.17	0.85	0.70	0.93	19.1	0.22	29.6	0.16	0.58	13.4	0.28	304	0.81	0.98	25.7	0.35
3466	Oneida	1.26	0.91	0.96	1.1	0.04	0.65	0.97	0.99	-1.0	0.03	1.83	0.79	0.91	2.8	0.06	3.0	-1.49	0.10	17.1	0.19	26	-0.48	0.63	-86.3	0.86
3467	Pusiano	1.74	0.90	0.97	-3.8	0.14	2.38	0.90	0.98	1.2	0.11	2.81	-0.81	0.26	-19.3	0.26	5.3	0.69	0.84	6.4	0.24	146	0.88	0.98	13.3	0.19
3468	Rappbode	1.15	0.91	0.97	-6.0	0.14	1.22	0.96	0.99	0.4	0.08	0.77	0.41	0.74	-5.3	0.12	13.3	0.67	0.84	3.6	0.23	254	0.92	0.99	11.5	0.16
3469	Rassnitzersee	1.64	0.80	0.96	-15.3	0.17	1.82	0.90	0.99	-14.9	0.15	1.73	-1.00	0.77	-23.3	0.23	5.2	0.82	0.94	14.4	0.15	116	0.94	0.98	-14.9	0.17
3470	Ravn	1.94	0.85	0.94	-9.3	0.19	1.81	0.91	0.99	-12.9	0.14	1.53	0.36	0.88	5.2	0.21	8.3	0.34	0.75	10.4	0.27	149	0.80	0.98	-33.1	0.34
3471	Rotorua	1.33	0.91	0.99	3.8	0.07	1.33	0.91	0.99	3.9	0.08	1.38	0.89	0.99	3.9	0.08	3.9	0.04	0.36	4.9	0.09	11	0.74	0.88	-6.6	0.43
3472	Stechlin	1.11	0.91	0.96	-7.3	0.13	1.73	0.93	0.99	2.5	0.11	0.77	0.04	0.80	-1.0	0.11	21.2	0.42	0.74	11.9	0.33	159	0.96	0.98	-3.9	0.14
3473	Tarawera	0.77	0.86	0.95	-3.3	0.04	0.82	0.93	0.98	-1.4	0.04	0.47	-0.08	0.60	-2.4	0.03	21.4	0.46	0.76	-1.3	0.27	424	0.97	0.99	-6.6	0.10
3474	Toolik	1.36	0.77	0.88	2.1	0.25	1.94	0.82	0.91	1.4	0.26	1.15	0.61	0.81	7.4	0.25	11.3	-0.89	0.29	52.3	0.61	17	0.74	0.86	-15.7	0.43
3475	Windermere	1.61	0.82	0.95	12.4	0.14	3.21	0.54	0.81	-11.1	0.23	2.39	-0.82	0.85	25.2	0.26	10.2	0.20	0.79	15.8	0.22	271	0.90	0.98	16.2	0.21
3476	Woods	2.14	0.82	0.99	-9.0	0.17	2.13	0.82	0.99	-9.1	0.17	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	N/A	
3477	Zurich	1.24	0.94	0.98	7.6	0.12	1.48	0.96	0.99	5.3	0.09	1.11	-3.05	0.60	13.6	0.16	50.8	0.30	0.64	-7.6	0.42	816	0.94	1.00	16.2	0.17
3478	Mean	1.34	0.89	0.96	-0.16	0.11	1.67	0.89	0.97	-0.84	0.10	1.31	-0.25	0.73	1.97	0.14	26.5	0.23	0.66	9.89	0.32	753	0.83	0.96	1.23	0.25
3479	Median	1.33	0.90	0.97	0.26	0.09	1.62	0.92	0.98	0.23	0.09	1.13	0.03	0.78	2.13	0.10	11.3	0.36	0.71	8.53	0.28	206	0.90	0.98	0.81	0.20

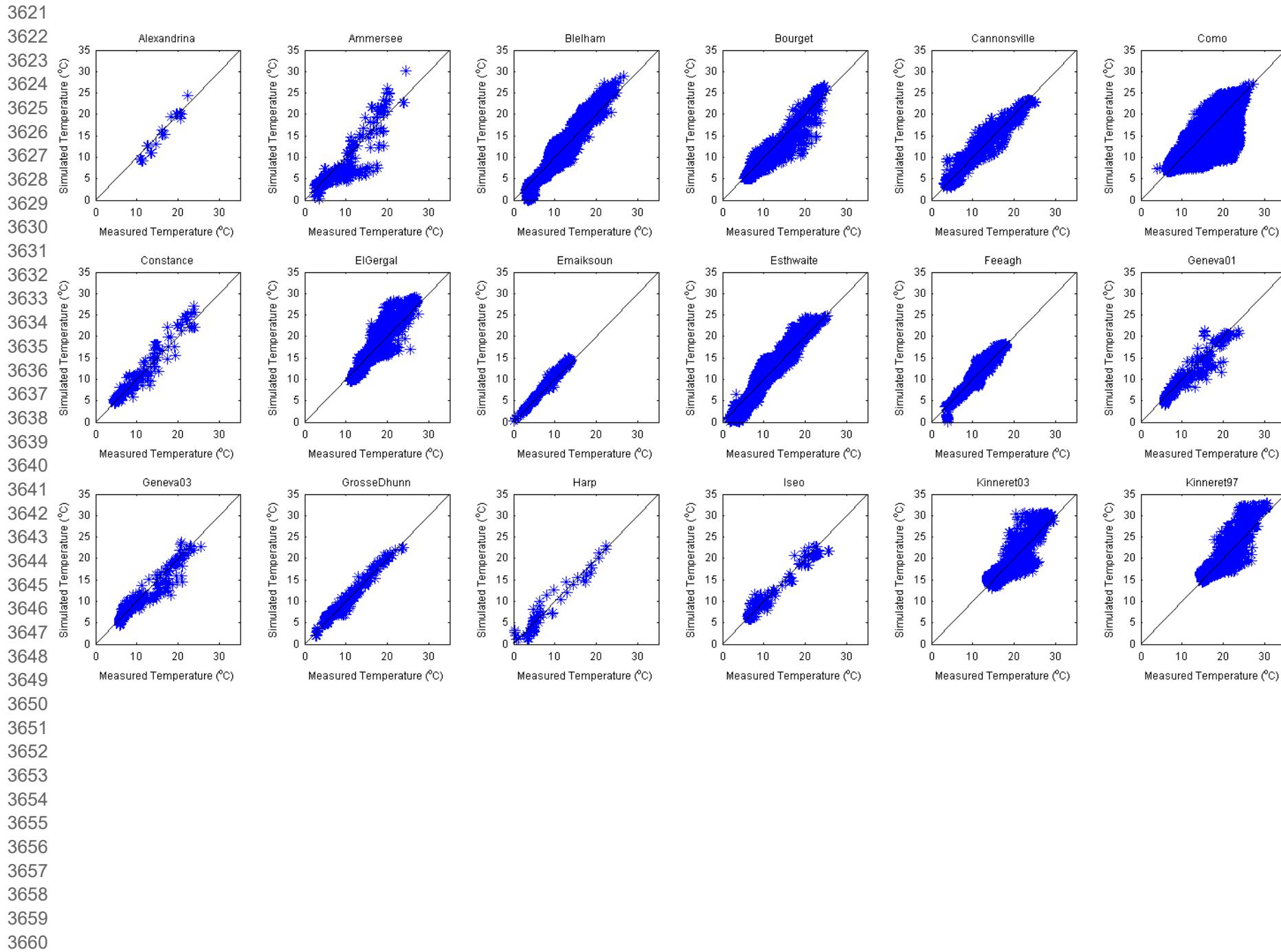
3501 Table B2 – Significance and correlation between model performance metrics and input uncertainty for morphometry (morph), distance to meteorological
 3502 station (distmet), frequency of meteorological data (freqmet), determination of inflow and outflows (flow), light extinction coefficient (Kw), frequency of
 3503 observed temperature profiles (obs) and average error ranking (mean). Significant correlations highlighted in red and corresponding r in yellow.
 3504

		Full Profile Temperature						Epilimnion Temperature						Hypolimnion Temperature						Thermocline Depth						Schmidt Stability			
		RMS	MEF	NMA	RMS	MEF	NMA	RMS	MEF	NMA	RMS	NSE	r	PRE	NMA	RMS	MEF	NMA	RMS	PRE	NMA	RMS	MEF	F	r	PRE	NMA		
		E	F	r	PRE	E	F	r	PRE	E	F	r	PRE	E	E	E	E	E	E	NSE	r	PRE	E	E	F	r	PRE	E	
r	morph	0.05	0.26	0.03	-0.01	0.04	0.05	0.05	-0.10	-0.25	-0.02	0.09	0.12	0.22	0.22	0.14	-0.02	0.18	0.23	0.01	-0.01	-0.06	0.25	0.24	-0.09	-0.09	-0.27		
	distmet	0.27	-0.27	-0.16	-0.44	0.17	0.00	0.04	0.20	-0.18	0.00	0.14	-0.21	-0.22	-0.35	0.09	-0.18	0.09	0.07	-0.13	-0.14	-0.22	-0.15	-0.10	-0.38	0.04			
	freqmet	0.11	-0.19	-0.27	-0.32	0.34	0.00	0.07	0.06	-0.01	0.11	-0.09	-0.11	-0.01	-0.35	0.11	0.24	-0.34	-0.18	0.32	0.50	-0.08	0.11	0.10	-0.21	-0.13			
	flow	-0.21	0.06	0.23	0.17	0.16	-0.24	0.18	0.24	0.29	0.09	-0.05	-0.01	0.07	0.00	0.05	-0.10	-0.19	-0.17	0.14	-0.01	-0.25	-0.08	-0.08	0.16	0.14			
	Kw	0.30	-0.03	0.38	-0.16	0.03	0.15	-0.10	0.08	-0.27	0.06	0.33	-0.17	-0.19	-0.16	0.41	-0.17	0.42	0.40	-0.27	-0.29	0.02	0.24	0.26	0.03	-0.29			
	obs	-0.03	0.21	-0.02	-0.31	-0.16	-0.17	0.30	0.31	0.13	-0.32	-0.33	-0.13	-0.09	-0.29	-0.32	0.34	-0.18	-0.07	0.25	0.40	-0.01	0.07	0.10	-0.31	-0.25			
	mean	0.17	-0.02	0.01	-0.44	0.20	-0.10	0.22	0.32	-0.06	-0.05	-0.06	-0.18	-0.09	-0.37	0.05	0.10	-0.11	0.00	0.17	0.24	-0.20	0.08	0.11	-0.33	-0.21			
p	morph	0.80	0.14	0.88	0.94	0.83	0.77	0.78	0.57	0.15	0.93	0.65	0.51	0.24	0.25	0.47	0.93	0.33	0.22	0.95	0.97	0.76	0.19	0.20	0.64	0.15			
	distmet	0.12	0.13	0.38	0.01	0.32	0.99	0.82	0.27	0.32	0.99	0.47	0.26	0.24	0.06	0.64	0.34	0.65	0.71	0.51	0.44	0.23	0.44	0.60	0.04	0.84			
	freqmet	0.54	0.29	0.12	0.07	0.05	0.98	0.69	0.72	0.94	0.54	0.64	0.55	0.96	0.06	0.57	0.19	0.07	0.33	0.09	0.01	0.69	0.56	0.58	0.26	0.48			
	flow	0.24	0.72	0.19	0.34	0.35	0.17	0.31	0.17	0.09	0.62	0.79	0.94	0.72	0.99	0.79	0.59	0.31	0.36	0.46	0.94	0.18	0.69	0.67	0.41	0.46			
	Kw	0.08	0.85	0.03	0.38	0.89	0.41	0.59	0.65	0.12	0.73	0.07	0.36	0.32	0.39	0.03	0.36	0.02	0.03	0.15	0.12	0.92	0.21	0.16	0.87	0.12			
	obs	0.85	0.23	0.92	0.07	0.37	0.33	0.09	0.07	0.46	0.07	0.08	0.51	0.65	0.12	0.09	0.06	0.33	0.72	0.18	0.03	0.96	0.71	0.61	0.09	0.18			
	mean	0.33	0.90	0.97	0.01	0.25	0.57	0.22	0.06	0.73	0.76	0.77	0.33	0.64	0.05	0.77	0.59	0.55	0.99	0.36	0.20	0.29	0.66	0.55	0.07	0.27			

3541
 3542 Table B3 – Significance (p) and correlation (r) between model performance metrics and mean values of lake volume (V), surface area (Area), depth (D),
 3543 surface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature (T_{air}),
 3544 wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat), Lake Number (LN) and percent time when LN is less than one (%<1). Significant
 3545 correlations highlighted in red and corresponding r in yellow.
 3546
 3547

		Full-profile Temperature					Epilimnion Temperature					Hypolimnion Temperature					Thermocline Depth					Schmidt Stability				
		RMS E	MEF F	r PRE	NMA E	RMS E	MEF F	r PRE	NMA E	RMS E	MEF F	r PRE	NMA E	RMS E	MEF F	r PRE	NMA E	RMS E	MEF F	r PRE	NMA E	RMS E	MEF F	r PRE	NMA E	
r	V	-0.20	0.14	0.00	0.19	-0.22	0.12	-0.09	-0.13	0.18	-0.10	-0.65	-0.11	-0.37	0.02	-0.60	0.65	-0.03	-0.02	-0.05	0.16	0.63	0.07	0.02	-0.05	-0.20
	Area	-0.13	0.12	0.08	0.26	-0.26	0.01	-0.04	-0.03	0.26	-0.16	-0.56	-0.07	-0.30	0.06	-0.58	0.52	-0.10	-0.14	-0.06	0.05	0.50	-0.07	-0.11	-0.10	-0.05
	D	-0.25	0.14	-0.21	-0.03	-0.02	0.31	-0.17	-0.29	-0.04	0.07	-0.74	-0.12	-0.42	-0.08	-0.50	0.84	0.11	0.23	0.02	0.47	0.80	0.38	0.31	0.10	-0.50
	A/D	0.00	0.03	0.21	0.32	-0.30	-0.17	0.05	0.13	0.32	-0.25	-0.35	0.00	-0.15	0.09	-0.51	0.22	-0.15	-0.26	-0.11	-0.18	0.22	-0.29	-0.29	-0.14	0.20
	L/W	-0.09	0.12	0.02	0.14	-0.13	0.12	-0.16	-0.28	-0.14	-0.12	0.01	0.10	0.04	0.01	-0.09	0.12	0.06	0.11	-0.24	0.18	0.19	0.10	0.14	0.12	-0.21
	Inf	-0.23	0.21	0.02	0.46	-0.53	0.03	-0.14	-0.21	0.23	-0.25	-0.43	-0.02	-0.26	0.19	-0.59	0.57	-0.19	-0.18	-0.05	0.17	0.58	-0.14	-0.21	-0.05	0.06
	RT	-0.16	0.01	-0.19	-0.38	0.13	0.20	-0.02	-0.02	0.00	0.19	-0.58	-0.21	-0.43	-0.38	-0.33	0.38	0.03	0.09	0.23	0.26	0.36	0.32	0.18	-0.21	-0.44
	sw	0.18	-0.21	0.00	0.32	-0.12	-0.10	0.01	0.10	0.38	-0.18	-0.29	-0.09	-0.14	0.03	-0.38	0.01	0.01	-0.09	-0.09	-0.06	0.02	-0.08	0.04	0.36	0.12
	T_{air}	0.16	0.05	0.21	-0.18	-0.59	0.16	-0.12	-0.01	-0.16	-0.46	-0.04	-0.40	-0.43	-0.21	-0.52	0.08	0.34	0.29	-0.37	-0.28	0.19	-0.02	0.14	0.18	-0.13
	u_{wind}	0.03	-0.12	0.12	-0.19	0.05	-0.29	0.13	0.18	-0.15	-0.09	-0.04	0.24	0.35	-0.27	0.02	-0.26	-0.03	-0.05	0.03	-0.18	-0.25	-0.11	-0.10	-0.18	0.16
	K_w	0.47	-0.22	0.04	0.14	0.08	0.12	-0.14	-0.06	-0.05	0.08	0.67	0.02	0.27	0.15	0.43	-0.61	-0.14	-0.16	0.06	-0.23	-0.46	-0.26	-0.17	0.09	0.36
	Lat	-0.17	0.05	-0.16	-0.11	0.32	0.04	-0.03	-0.13	-0.22	0.35	0.24	0.12	0.23	0.13	0.48	-0.03	-0.13	0.02	0.32	0.16	-0.05	0.11	-0.05	-0.27	-0.01
	LN	0.21	-0.18	0.09	0.14	-0.09	0.36	-0.39	-0.31	-0.30	0.17	0.46	-0.55	-0.19	0.43	0.31	-0.14	0.31	0.41	-0.02	-0.18	-0.03	0.25	0.37	0.42	-0.22
	%<1	-0.14	0.10	0.10	0.25	-0.16	-0.43	0.24	0.20	0.23	-0.24	0.04	0.39	0.42	0.08	-0.14	-0.25	-0.56	-0.70	0.07	-0.29	-0.25	-0.78	-0.82	-0.53	0.78
p	V	0.25	0.41	0.99	0.27	0.21	0.50	0.60	0.46	0.31	0.57	0.00	0.57	0.04	0.92	0.00	0.00	0.89	0.90	0.79	0.38	0.00	0.70	0.92	0.81	0.30
	Area	0.45	0.52	0.65	0.14	0.14	0.96	0.82	0.86	0.15	0.36	0.00	0.71	0.11	0.76	0.00	0.00	0.60	0.47	0.74	0.78	0.00	0.72	0.55	0.60	0.81
	D	0.15	0.42	0.23	0.87	0.90	0.08	0.35	0.09	0.84	0.68	0.00	0.54	0.02	0.66	0.00	0.00	0.55	0.23	0.91	0.01	0.00	0.04	0.10	0.62	0.01
	A/D	0.98	0.85	0.22	0.06	0.08	0.34	0.79	0.45	0.06	0.15	0.06	0.99	0.43	0.65	0.00	0.24	0.41	0.16	0.56	0.34	0.23	0.12	0.12	0.45	0.30
	L/W	0.61	0.49	0.92	0.42	0.48	0.50	0.36	0.11	0.43	0.51	0.94	0.62	0.83	0.94	0.63	0.51	0.74	0.56	0.19	0.33	0.32	0.60	0.47	0.53	0.25
	Inf	0.22	0.25	0.91	0.01	0.00	0.86	0.45	0.25	0.21	0.18	0.02	0.91	0.18	0.33	0.00	0.00	0.33	0.37	0.79	0.38	0.00	0.49	0.29	0.79	0.78
	RT	0.40	0.95	0.32	0.04	0.47	0.28	0.92	0.94	1.00	0.31	0.00	0.29	0.02	0.05	0.09	0.05	0.90	0.64	0.24	0.18	0.06	0.10	0.35	0.29	0.02

3581	sw	0.30	0.24	1.00	0.06	0.49	0.59	0.98	0.57	0.02	0.30	0.12	0.63	0.45	0.88	0.04	0.96	0.94	0.62	0.63	0.75	0.93	0.66	0.82	0.05	0.54
3582	T _{air}	0.37	0.78	0.23	0.32	0.00	0.37	0.52	0.98	0.38	0.01	0.85	0.03	0.02	0.26	0.00	0.69	0.07	0.12	0.04	0.14	0.31	0.91	0.47	0.35	0.49
3583	u _{wind}	0.86	0.51	0.49	0.27	0.80	0.09	0.48	0.32	0.39	0.63	0.83	0.20	0.06	0.15	0.91	0.16	0.88	0.79	0.89	0.35	0.17	0.57	0.62	0.35	0.40
3584	Kw	0.00	0.20	0.84	0.42	0.66	0.52	0.44	0.72	0.77	0.67	0.00	0.90	0.14	0.42	0.02	0.00	0.47	0.39	0.74	0.22	0.01	0.17	0.38	0.64	0.05
3585	Lat	0.33	0.79	0.36	0.53	0.07	0.82	0.89	0.46	0.22	0.04	0.20	0.52	0.22	0.49	0.01	0.88	0.50	0.93	0.08	0.39	0.77	0.57	0.80	0.16	0.96
3587	LN	0.26	0.35	0.62	0.45	0.63	0.05	0.03	0.10	0.10	0.37	0.01	0.00	0.30	0.02	0.10	0.48	0.10	0.02	0.92	0.35	0.86	0.18	0.04	0.02	0.24
3588	%<1	0.45	0.60	0.61	0.19	0.41	0.02	0.20	0.29	0.23	0.19	0.82	0.03	0.02	0.66	0.45	0.19	0.00	0.00	0.72	0.12	0.18	0.00	0.00	0.00	0.00
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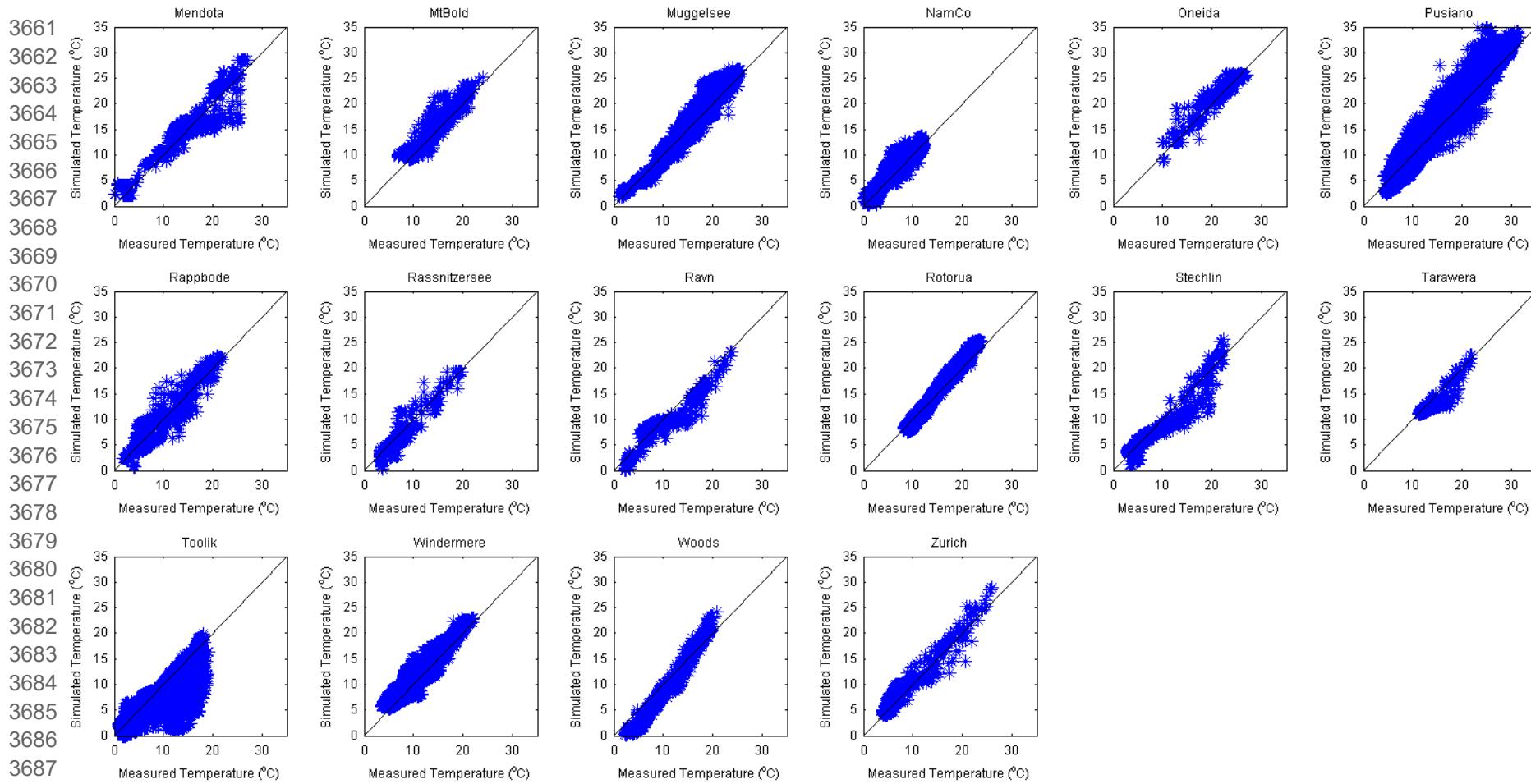


Figure B1 – Plot of modelled vs observed temperature data for each of the MLCP lakes.

3701 7 Appendix C – Sensitivity Analysis

3702

3703 Table C1 - Significance (p) and correlation (r) between sensitivity indices for full profile temperature and mean values of lake volume (V), surface area
 3704 (Area), depth (D), suface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air
 3705 temperature (T_{air}), wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red
 3706 and corresponding r in yellow.

	Attribute	C_c	C_w	C_s	C_t	C_{KH}	C_{hyp}	C_e	C_h	C_d
r	V	0.08	0.04	-0.09	-0.10	0.47	0.04	-0.07	-0.20	0.02
	Area	0.15	0.09	0.02	0.09	0.33	0.16	-0.43	-0.18	-0.06
	D	0.06	0.03	-0.12	-0.13	0.49	0.02	0.06	-0.16	0.03
	A/D	-0.02	-0.03	-0.16	-0.22	0.43	-0.08	0.30	-0.13	0.05
	L/W	-0.08	0.14	-0.10	0.15	-0.05	0.21	-0.15	-0.08	-0.40
	Inf	0.01	0.06	0.00	-0.24	0.43	0.07	-0.08	-0.41	-0.32
	RT	0.21	0.04	-0.01	0.18	0.18	0.06	-0.24	0.09	0.26
	sw	-0.10	-0.11	-0.18	-0.14	0.39	-0.12	0.41	0.12	0.18
	T_{air}	-0.23	-0.24	-0.26	-0.22	-0.40	-0.30	0.34	-0.67	0.03
	u_{wind}	-0.30	-0.32	-0.19	-0.15	-0.14	-0.30	0.33	0.09	0.22
	K_w	-0.25	-0.16	-0.04	-0.16	-0.30	-0.14	0.42	0.12	0.11
	Lat	0.22	0.16	0.24	0.18	-0.24	0.20	-0.49	0.19	-0.06
p	V	0.68	0.85	0.64	0.60	0.01	0.84	0.73	0.30	0.91
	Area	0.42	0.65	0.90	0.65	0.08	0.40	0.02	0.34	0.75
	D	0.75	0.89	0.54	0.48	0.01	0.92	0.77	0.39	0.90
	A/D	0.90	0.87	0.39	0.24	0.02	0.67	0.11	0.49	0.79
	L/W	0.68	0.47	0.60	0.44	0.80	0.27	0.44	0.66	0.03
	Inf	0.95	0.74	0.98	0.23	0.02	0.72	0.68	0.03	0.10
	RT	0.29	0.83	0.97	0.36	0.37	0.76	0.21	0.67	0.18
	sw	0.61	0.56	0.35	0.47	0.04	0.52	0.02	0.54	0.35
	T_{air}	0.22	0.21	0.17	0.25	0.03	0.11	0.07	0.00	0.88
	u_{wind}	0.11	0.08	0.31	0.43	0.48	0.10	0.08	0.65	0.25
	K_w	0.18	0.39	0.84	0.39	0.11	0.45	0.02	0.53	0.57

3741		Lat	0.25	0.41	0.20	0.35	0.20	0.28	0.01	0.31	0.76
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3781 Table C2 -- Significance (p) and correlation (r) between sensitivity indices for epilimnion temperature and mean values of lake volume (V), surface area
 3782 (Area), depth (D), suface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air
 3783 temperature (T_{air}), wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red
 3784 and corresponding r in yellow.

	Attribute	C_c	C_w	C_s	C_t	C_{KH}	C_{hyp}	C_e	C_h	C_d
r	V	0.24	0.19	0.31	0.25	0.56	-0.10	-0.20	0.21	0.31
	Area	0.40	0.29	0.27	0.30	0.63	0.09	-0.25	0.21	0.33
	D	0.18	0.14	0.31	0.22	0.49	-0.13	-0.17	0.20	0.28
	A/D	0.00	0.01	0.24	0.12	0.30	-0.21	-0.05	0.13	0.17
	L/W	-0.07	0.10	-0.13	0.20	0.25	0.41	-0.18	0.14	-0.22
	Inf	0.13	0.09	0.07	0.14	0.51	-0.05	-0.31	-0.16	-0.09
	RT	0.44	0.29	0.40	0.09	0.22	-0.06	0.09	0.21	0.60
	sw	-0.13	-0.13	0.27	0.10	0.03	-0.20	0.09	0.19	0.08
	T_{air}	-0.21	-0.50	-0.54	-0.40	-0.10	-0.34	0.43	-0.46	-0.61
	u_{wind}	-0.18	0.08	-0.11	-0.05	-0.26	-0.10	0.35	0.08	-0.02
	K_w	-0.43	-0.37	-0.33	-0.37	-0.58	-0.14	0.32	-0.31	-0.36
	Lat	0.26	0.23	-0.06	0.03	-0.05	0.26	-0.16	-0.15	0.12
p	V	0.19	0.31	0.10	0.19	0.00	0.62	0.30	0.27	0.10
	Area	0.03	0.12	0.15	0.10	0.00	0.63	0.18	0.26	0.08
	D	0.35	0.45	0.10	0.24	0.01	0.50	0.38	0.29	0.13
	A/D	0.99	0.94	0.20	0.51	0.11	0.27	0.79	0.50	0.37
	L/W	0.71	0.60	0.49	0.28	0.18	0.02	0.34	0.45	0.24
	Inf	0.50	0.64	0.73	0.48	0.01	0.80	0.11	0.40	0.66
	RT	0.02	0.14	0.03	0.67	0.25	0.75	0.66	0.28	0.00
	sw	0.50	0.49	0.14	0.62	0.86	0.30	0.65	0.31	0.67
	T_{air}	0.27	0.00	0.00	0.03	0.58	0.06	0.02	0.01	0.00
	u_{wind}	0.33	0.68	0.56	0.78	0.17	0.61	0.06	0.68	0.94
	K_w	0.02	0.04	0.07	0.05	0.00	0.45	0.09	0.10	0.05
	Lat	0.17	0.22	0.77	0.88	0.80	0.16	0.39	0.43	0.51

3821 Table C3 - - Significance (p) and correlation (r) between sensitivity indices for hypolimnion temperature and mean values of lake volume (V), surface area
 3822 (Area), depth (D), suface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air
 3823 temperature (T_{air}), wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red
 3824 and corresponding r in yellow.

	Attribute	C_c	C_w	C_s	C_t	C_{KH}	C_{hyp}	C_e	C_h	C_d
r	V	-0.11	-0.37	-0.33	-0.26	-0.07	-0.22	0.08	-0.27	-0.26
	Area	-0.03	-0.34	-0.29	-0.21	-0.14	-0.20	-0.23	-0.22	-0.32
	D	-0.11	-0.34	-0.31	-0.24	-0.02	-0.18	0.16	-0.25	-0.22
	A/D	-0.14	-0.29	-0.26	-0.19	0.02	-0.14	0.31	-0.22	-0.12
	L/W	-0.10	0.19	0.20	0.17	0.14	0.22	-0.39	-0.15	-0.09
	Inf	-0.39	-0.39	-0.20	-0.25	-0.08	-0.19	-0.02	-0.68	-0.38
	RT	0.18	-0.29	-0.39	-0.13	-0.34	-0.36	-0.03	0.17	-0.24
	sw	-0.14	-0.22	0.01	-0.05	0.03	-0.08	0.59	0.11	-0.22
	T_{air}	-0.22	-0.23	0.01	-0.11	-0.64	-0.22	0.32	-0.50	-0.20
	u_{wind}	-0.29	-0.21	-0.35	-0.17	-0.21	-0.22	0.01	0.27	0.04
	K_w	-0.14	0.15	0.27	0.17	0.13	0.25	0.16	0.07	0.27
	Lat	0.11	0.14	-0.03	0.09	0.15	0.00	-0.48	0.03	0.21
p	V	0.55	0.04	0.07	0.16	0.71	0.25	0.69	0.15	0.17
	Area	0.88	0.07	0.12	0.27	0.46	0.28	0.22	0.24	0.09
	D	0.55	0.06	0.10	0.20	0.92	0.34	0.40	0.19	0.25
	A/D	0.46	0.12	0.17	0.30	0.93	0.46	0.10	0.24	0.52
	L/W	0.61	0.32	0.29	0.36	0.46	0.25	0.04	0.42	0.65
	Inf	0.04	0.04	0.30	0.21	0.68	0.34	0.93	0.00	0.05
	RT	0.37	0.14	0.04	0.50	0.08	0.06	0.88	0.40	0.22
	sw	0.47	0.24	0.96	0.78	0.87	0.66	0.00	0.56	0.25
	T_{air}	0.25	0.22	0.94	0.57	0.00	0.25	0.09	0.01	0.30
	u_{wind}	0.12	0.26	0.06	0.36	0.27	0.25	0.96	0.15	0.85
	K_w	0.46	0.42	0.15	0.38	0.48	0.18	0.38	0.71	0.16
	Lat	0.55	0.46	0.87	0.63	0.44	0.99	0.01	0.86	0.28

3861 Table C4 - - Significance (p) and correlation (r) between sensitivity indices for thermocline depth and mean values of lake volume (V), surface area (Area),
 3862 depth (D), suface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature
 3863 (T_{air}), wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red and
 3864 corresponding r in yellow.

	Attribute	C_c	C_w	C_s	C_t	C_{KH}	C_{hyp}	C_e	C_h	C_d
r	V	0.48	0.43	0.30	0.35	0.58	0.29	0.23	0.31	0.25
	Area	0.57	0.54	0.29	0.42	0.61	0.27	0.40	0.36	0.17
	D	0.41	0.36	0.27	0.31	0.51	0.28	0.14	0.27	0.25
	A/D	0.22	0.19	0.22	0.18	0.33	0.20	0.02	0.15	0.25
	L/W	0.01	0.16	-0.01	0.04	0.04	0.06	0.06	0.15	-0.04
	Inf	0.50	0.43	0.34	0.38	0.53	0.33	0.23	0.40	0.23
	RT	0.25	0.13	-0.06	0.10	0.28	0.08	0.16	0.07	0.10
	sw	0.03	0.08	0.31	-0.06	0.10	0.17	0.05	0.03	0.38
	T_{air}	0.11	0.13	0.19	0.06	0.09	0.34	0.19	0.18	0.44
	u_{wind}	-0.19	-0.13	0.06	-0.18	-0.22	-0.19	0.00	-0.33	0.06
	K_w	-0.42	-0.24	-0.27	-0.26	-0.46	0.03	-0.23	-0.04	0.08
	Lat	-0.01	-0.11	-0.34	0.01	-0.09	-0.19	-0.06	-0.05	-0.32
p	V	0.01	0.02	0.11	0.06	0.00	0.12	0.21	0.10	0.17
	Area	0.00	0.00	0.12	0.02	0.00	0.15	0.03	0.05	0.38
	D	0.03	0.05	0.15	0.10	0.00	0.14	0.45	0.15	0.18
	A/D	0.24	0.31	0.24	0.35	0.08	0.29	0.93	0.42	0.18
	L/W	0.94	0.39	0.94	0.83	0.83	0.74	0.77	0.43	0.84
	Inf	0.01	0.02	0.07	0.04	0.00	0.09	0.25	0.04	0.24
	RT	0.21	0.50	0.77	0.61	0.15	0.70	0.40	0.71	0.60
	sw	0.86	0.68	0.10	0.75	0.60	0.37	0.79	0.88	0.04
	T_{air}	0.55	0.49	0.31	0.75	0.63	0.07	0.30	0.34	0.02
	u_{wind}	0.31	0.49	0.74	0.34	0.25	0.31	1.00	0.07	0.76
	K_w	0.02	0.20	0.14	0.16	0.01	0.89	0.22	0.83	0.68
	Lat	0.95	0.58	0.07	0.96	0.64	0.30	0.77	0.81	0.08

3901 Table C5 -- Significance (p) and correlation (r) between sensitivity indices for Schmidt stability and mean values of lake volume (V), surface area (Area),
 3902 depth (D), suface area divided by depth (A/D), length divided by width (L/W), inflow (Inf), residence time (RT), short wave radiation (sw), air temperature
 3903 (T_{air}), wind speed (u_{wind}), light extinction coefficient (K_w), latitude (Lat) and Lake Number (LN). Significant correlations highlighted in red and
 3904 corresponding r in yellow.

	Attribute	C_c	C_w	C_s	C_t	C_{KH}	C_{hyp}	C_e	C_h	C_d
r	V	0.02	0.13	0.27	0.15	-0.08	-0.06	-0.16	-0.07	0.00
	Area	0.00	0.07	0.10	-0.06	-0.30	-0.28	-0.25	-0.15	-0.31
	D	0.04	0.14	0.31	0.22	0.02	0.04	-0.09	-0.01	0.13
	A/D	0.05	0.12	0.33	0.30	0.15	0.17	0.07	0.03	0.30
	L/W	-0.15	-0.16	-0.21	-0.17	-0.06	-0.08	0.03	-0.12	-0.21
	Inf	0.02	0.17	0.31	0.16	0.13	0.07	-0.16	0.07	0.13
	RT	0.01	0.07	-0.05	0.00	-0.25	-0.14	-0.20	0.02	-0.13
	sw	-0.10	-0.14	0.04	0.16	-0.22	0.00	0.17	-0.04	0.00
	T_{air}	0.00	-0.05	-0.01	0.00	-0.08	0.03	-0.15	-0.29	0.00
	u_{wind}	0.04	-0.18	0.03	-0.21	0.00	-0.15	0.43	-0.05	0.14
	K_w	-0.11	-0.11	-0.25	0.15	0.31	0.32	0.32	0.24	0.26
	Lat	0.06	0.08	-0.07	-0.16	0.16	-0.05	-0.13	0.17	-0.05
p	V	0.93	0.48	0.14	0.43	0.67	0.77	0.41	0.73	0.98
	Area	0.99	0.70	0.62	0.74	0.10	0.13	0.19	0.44	0.10
	D	0.83	0.45	0.09	0.24	0.91	0.85	0.62	0.95	0.49
	A/D	0.80	0.53	0.08	0.11	0.44	0.37	0.70	0.87	0.11
	L/W	0.43	0.40	0.26	0.36	0.77	0.67	0.86	0.54	0.27
	Inf	0.92	0.40	0.11	0.43	0.52	0.74	0.42	0.72	0.52
	RT	0.95	0.72	0.78	0.98	0.20	0.48	0.31	0.94	0.50
	sw	0.60	0.45	0.84	0.40	0.24	1.00	0.38	0.85	0.98
	T_{air}	0.98	0.80	0.95	0.98	0.66	0.88	0.44	0.11	0.99
	u_{wind}	0.85	0.33	0.87	0.26	0.99	0.42	0.02	0.77	0.46
	K_w	0.57	0.55	0.19	0.41	0.09	0.08	0.08	0.20	0.17
	Lat	0.77	0.68	0.70	0.39	0.39	0.80	0.50	0.36	0.79

8 Appendix D – Acknowledgements

Table D1 – Acknowledgements by individual lake. Note people includes additional staff and students who helped set up the GLM not included in the list of authors.

Lake Name	Morphometry	Meteorology	Flow	Field Data	Institutions	Funding	People
Alexandrina	Department of Environment, Water and Natural Resources	Natural Resources SA Murray-Darling Basin	Murray-Darling Basin Authority (MDBA)	Natural Resources SA Murray-Darling Basin	The University of Western Australia	ARC Discovery Grant DP130104078	Alex Perry
Ammersee	Bavarian Environment Agency	German Weather Service,Wielenbach; Bavarian Agency of Agriculture, Rothenfeld + Westerschondorf	Water Agency Weilheim	Water Agency Weilheim		Bavarian Environment Agency; Bavarian State Ministry of the Environment and Consumer Protection	Otfried Baume
Blelham	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
Bourget	Delebecque, 1898 and IFREMER 1992	Météo France	DREAL Rhône-Alpes	© SOERE OLA-IS, INRA Thonon-les-Bains, CISALB, [2012], developed by INRA's ORE Eco-Information system	INRA Thonon les Bains	CISALB, Agence de l'eau Rhône-Méditerranée-Corse, SOERE OLA, Ecole des Ponts ParisTech, INRA	Orlane Anneville
Cannonsville	New York City Department of Environmental Protection	Cannonsville Dam Station	Trout Creek and West Branch Delaware River	New York City Department of Environmental Protection, Kingston, NY	New York City Department of Environmental Protection	New York City Department of Environmental Protection	Karen E.B. Moore
Como	Lombardy Region	Water Research Institute - National Research Council of Italy (IRSA-CNR); Bergamo-Orio al Serio Aiport; Regional Authority for Environmental Protection (ARPA Lombardia)	Consorzio dell'Adda	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Centro Volta Como ; Istituto Nazionale della Montagna	Simulake Project; WAAESs Project	Gianni Tartari (Water Research Institute - National Research Council of Italy (IRSA-CNR)

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3982	Constance	IGKB (Internationale Gewässerschutzkommission für den Bodensee	German meteorological service (DWD)	IGKB (Internationale Gewässerschutzkommission für den Bodensee	IGKB (Internationale Gewässerschutzkommission für den Bodensee	LUBW Landesanstalt für Umwelt, Messungen und Naturschutz Baden-Württemberg; Limnological Institute, University of Konstanz	DFG, grant Ri 2040/1-1	Thomas Wolf
3983								
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3987	El Gergal	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	Empresa Metropolitana de Abastecimiento y Saneamiento de Aguas de Sevilla, S.A.	University of Granada	CGL2005-04070/HID	Carmelo Escot
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3992	Emaiksoun	Kenneth M. Hinkel, University of Cincinnati	Brittany L. Potter, University of Nebraska-Lincoln		Brittany L. Potter, University of Nebraska-Lincoln; Kenneth M. Hinkel, University of Cincinnati	LimnoTech; University of Nebraska-Lincoln; University of Cincinnati	NSF ARC-1107792	Brittany L. Potter, Kenneth M. Hinkel
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3996	Esthwaite	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
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4000	Feeagh	Marine Institute	Met Eireann; Marine Institute	Marine Institute	Marine Institute	Marine Institute	Marine Institute	Eleanor Jennings; Elizabeth Ryder; Mary Dillane; Russell Poole; Burrisheoole field staff
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4002								
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4004								
4005	Geneva03	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL	SOERE OLA-IS, INRA Thonon-les-Bains, CIPEL			Orlane Anneville
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4008	Geneva05	As Above	As Above	As Above	As Above	As Above	As Above	As Above
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4010	GrosseDhuen	Wupperverband (reservoir manager)	German Weather Service stations Luedenscheid, Cologne-Bonn	Wupperverband (reservoir manager)	Wupperverband (reservoir manager), Helmholtz Centre for Environmental Research UFZ	Helmholtz Centre for Environmental Research UFZ	District Council Cologne and Ministry of Environment North Rhine-Westphalia	Karsten Rahn, Martin Wieprecht, Wilfried Scharf
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4013	Harp		Dorset Environmental Science Centre; Environment Canada	Dorset Environmental Science Centre				
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4022	Iseo	Regione Lombardia	Università degli Studi di Brescia; Bergamo-Orio al Serio Aiport; Regional Authority for Environmental Protection (ARPA Lombardia)	Consorzio dell'Oglio	Università degli Studi di Brescia; Prof. Letizia Garibaldi (Università degli Studi di Milano-Bicocca)			
4023	Kinneret03	Kinneret Limnological Laboratory	Isr. Meteorol. Service; Kinneret Limnological Laboratory	Isr. Hydrological Service	Kinneret Limnological Laboratory, IOLR	Israel Water Authority	Israel Water Authority	
4024	Kinneret97	As Above	As Above	As Above	As Above	As Above	As Above	As Above
4025	Mendota	Wisconsin Departement of Natural Resources	SSEC-GAMIS-RIG UW-Madison; US National Climatic Data Center	U.S. Geological Survey	University of Wisconsin - Madison LTER program	University of Wisconsin, Madison	NSF grant DEB-0822700 (North Temperate Lakes Long-Term Ecological Research)	
4026	MtBold	South Australian Water Corporation	Bureau of Meteorology; South Australian Water Corporation, Happy Valley Reservoir	Department of Water and Natural Resources; SA Water Major Systems	South Australian Water Corporation	University of Adelaide, Adelaide, South Australia	Water Research Foundation, Boulder, CO, USA; South Australian Water Corporation	Mike Burch; Rob Daly;
4027	Muggelsee	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Senatsverwaltung für Stadtentwicklung und Umwelt Berlin	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Leibniz Institute of Freshwater Ecology and Inland Fisheries (IGB)	Thomas Hintze (IGB) for operating the Müggelsee Lake Station.
4028	NamCo	Institute of Tibetan Plateau Research, Chinese Academy of Sciences	Institute of Tibetan Plateau Research, Chinese Academy of Sciences			Institute of Tibetan Plateau Research, Chinese Academy of Sciences;Institut für Geographie, Friedrich-Schiller-Universität	the National Basic Research Program of China (2012CB956100), National Natural Science Foundation of China (41071123) and CADY project (TP2:03G0813F) from BMBF, Germany	
4029	Oneida	National Oceanic and Atmospheric Administration; Cornell Biological Field Station	Northeast Regional Climate Center	United States Geological Survey	Cornell Biological Field Station	Cornell University	Cornell University Brown Endowment; New York State Department of Environmental Conservation; United States Department of Agriculture 0226747	Lars Rudstam
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4061	Pusiano	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Regional Authority for Environmental Protection (ARPA Lombardia); National Oceanic and Atmospheric Administration	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Water Research Institute - National Research Council of Italy (IRSA-CNR)	Parco Valle Lambro; Fondazione CARIPLO	Progetto PIROGA	Gianni Tartari and Franco Salerno (Water Research Institute - National Research Council of Italy (IRSA-CNR))
4062	Rappbode	Reservoir authority of the state of Saxony-Anhalt (Talsperrenbetrieb Sachsen-Anhalt)	German Meteorological Service (DWD)	Rappbode Reservoir Authority (Talsperrenbetrieb Sachsen-Anhalt)	Fernwasserversorgung Elbaue Osthartz; Helmholtz Centre for Environmental Research - UFZ; Talsperrenbetrieb Sachsen-Anhalt	Helmholtz Centre for Environmental Research - UFZ		Karsten Rahn (UFZ, SEEFO); Martin Wieprecht (UFZ, SEEFO); Maren Dietze (Talsperrenbetrieb Sachsen-Anhalt); Dieter Noga (DWD); Marco Matthes (Fernwasserversorgung Elbaue Osthartz).
4063	Rassnitzersee	Helmholtz Centre for Environmental Research - UFZ	German Meteorological Service (DWD)	Lausitzer and Mitteldeutsche Braunkohle Verwaltungsgesellschaft - LMBV	Uwe Kiwel(UFZ) and Karsten Rahn (UFZ)	Helmholtz Centre for Environmental Research - UFZ	Helmholtz Centre for Environmental Research - UFZ	Uwe Kiwel(UFZ) and Karsten Rahn (UFZ)
4064	Ravn	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	Danish Meteorological Institute (DMI). Data made available for analyses relating to the National Monitoring Program for Water and Nature.	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	National Monitoring Program for Water and Nature. Data hosted by Aarhus University.	Aarhus University	CLEAR centre of excellence (Villum-Kann Rasmussen Foundation)	
4065	Rotorua	Digitised bathymetry held by University of Waikato, Bay of Plenty Regional Council	Meteorological Service of New Zealand Limited. Data obtained via 'cliflo' database (National Institute of Water and Atmospheric Research (NIWA), New Zealand).	National Institute of Water and Atmospheric Research.	Bay of Plenty Regional Council	University of Waikato-Environmental Research Institute	Bay of Plenty Regional Council Chair in Lake Restoration at University of Waikato	
4066	Stechlin		German Meteorological Service (DWD), Umwelt Bundesamt (UBA), Energiewerke Nord GmbH (Betriebsteil Kernkraftwerk Rheinsberg)		Leibniz-Institute of Freshwater Ecology and Inland Fisheries	Leibniz-Institute of Freshwater Ecology and Inland Fisheries	German Federal Ministry of Research and Education (BMBF Project KLIMZUG-INKABB TP22)	Peter Kasprzak

4101	Tarawera	Digitised bathymetry held by University of Waikato, Bay of Plenty Regional Council	National Institute of Water and Atmospheric Research	Bay of Plenty Regional Council	Bay of Plenty Regional Council	University of Waikato-Environmental Research Institute	Bay of Plenty Regional Council Chair in Lake Restoration at University of Waikato	Andy Bruere
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4105	Toolik	Toolik Field Station - Environmental Data Center (TFS EDC). Jason J. Stuckey (TFS GIS Manager) performed a bathymetrical study of the lake in 2008 using a Garmin GPSMAP 188 Sounder	Meteorological datasets provided by the Toolik Field Station Environmental Data Center are based upon work supported by the U. S. National Science Foundation (NSF) under grants #455541 and #1048361	Arctic Long Term Ecological Research (ARC LTER), funded by NSF, Division of Environmental Biology (DEB) 0423385	(a) Arctic Long Term Ecological Research (ARC LTER); (b) Toolik Field Station - Environmental Data Center (TFS EDC), University of Alaska Fairbanks (UAF)	(a) Department of Ecology, Evolution, and Marine Biology, University of California, Santa Barbara, California; (b) Marine Science Institute, University of California, Santa Barbara, California	NFS Support, from two different grants of Arctic Natural Sciences (ANS) to Sally MacIntyre: #0714085 and #1204267	
4106	Windermere	Ramsbottom, A.E. 1976. Depth charts of the Cumbrian Lakes. Freshwater Biological Association	Centre for Ecology and Hydrology	Environment Agency	Centre for Ecology and Hydrology	University of Reading; Centre for Ecology and Hydrology	UKLEON (NE/I007407/1)	Bernard Tebay
4107	Woods	Hydro Tasmania	Australian Bureau of Meteorology	Hydro Tasmania	Hydro Tasmania	University of Tasmania	ARC Linkage Grant LP130100756	Carolyn Maxwell, Leon Barmuta, Abhijeet Kulkarni, Aditya Singh
4108	Zurich	David M. Livingstone	MeteoSwiss	Federal Office for the Environment (FOEN)	Wasserversorgung der Stadt Zürich (Zurich Water Supply)	Eawag: Swiss Federal Institute of Aquatic Science and Technology	Amt für Abfall, Wasser, Energie und Luft (AWEL) of the Canton of Zurich (Lake Monitoring)	
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