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1 **Changes in Terrestrial Water Storage in the Three-North region of China over**
2 **2003-2021: Assessing the Roles of Climate and Vegetation restoration**

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11 **Abstract**

12 Large-scale vegetation restoration has markedly enhanced ecosystem services in the
13 Three-North (TN) region, concurrently exacerbating the pre-existing water resource
14 crisis stemming from climate change in this arid domain. However, the specific
15 contributions of climate and vegetation changes to the evolution of water resources in
16 the TN region have yet to be elucidated. This study leverages data from the Gravity
17 Recovery and Climate Experiment (GRACE) and its Follow-on mission (GRACE-
18 FO), specifically the terrestrial water storage anomaly (TWSA) from 2003 to 2021.
19 We identify two TWSA rising (*TWSA_rI* and *TWSA_rII*) and two declining (*TWSA_dI*
20 and *TWSA_dII*) hotspots in the TN region. By exploring the relative importance of
21 climate drivers and vegetation dynamics in influencing the temporal dynamics of
22 TWSA, we found that vegetation restoration assumes a more pivotal role than climate
23 factors in 45.2% of the total area within the identified hotspots, predominantly evident
24 in TWSA declining hotspots. The heightened TWSA in *TWSA_rI* and *TWSA_rII*
25 primarily stems from precipitation recharge, with precipitation-dominated areas
26 constituting 19.3% of the area within the identified hotspots. Potential
27 evapotranspiration dominates TWSA changes in only 4.6% of the region within the
28 identified hotspots, primarily scattered in *TWSA_dI* and *TWSA_dII*. Furthermore, a
29 comparative assessment of soil moisture and groundwater responses to environmental
30 factors indicates that water storage at shallower layers is more responsive to climate
31 factors, while the impacts of vegetation changes on TWSA are more discernibly
32 reflected in water storage at deeper layers. These findings furnish scientific guidance
33 for future ecological restoration planning and sustainable water resources
34 management.

35 **Keywords:** Vegetation restoration, Terrestrial water storage, Three-North region of
36 China, GRACE

37 **1 Introduction**

38 The dynamic shifts in climate conditions, marked by increased precipitation
39 variability and elevated temperatures, exert profound effects on both global and
40 regional hydrological cycles (Schewe et al., 2014; Chang et al., 2020; Patterson et al.,
41 2022). Besides climate change, anthropogenic activities contribute significantly to the
42 modification of hydrological cycles too, encompassing alterations to land cover,
43 exploitation of water resources, and the implementation of ecological restoration
44 projects, among other factors (Wada et al., 2017; Huang et al., 2023). Globally,
45 hydrological systems in arid regions generally exhibit heightened sensitivity to
46 external environmental influences in contrast to their humid counterparts (Molnar,
47 2001; Xie et al., 2015). The Three-North (TN) region in China, encompassing its
48 northern, northeastern, and northwestern sectors, stands as a substantial arid expanse
49 grappling with a dual challenge: a water resource crisis induced by climate change
50 and additional water resource depletion resulting from large-scale vegetation
51 restoration initiatives (Jia et al., 2017; Lv et al., 2019; Liu et al., 2023a; Zhou et al.,
52 2023). These challenges impede the sustainable development and utilization of
53 regional water resources in the TN region.

54 Vegetation restoration endeavors typically manifest in heightened evapotranspiration
55 rates and reduced runoff volumes, potentially influencing precipitation patterns by
56 reinforcing local moisture recycling (Bai et al., 2020; Zhang et al., 2022; Yang et al.,
57 2023). Given the diverse impact of these alterations in individual hydrological flux on
58 the land water budget, changes in terrestrial water storage emerge as an intuitive
59 indicator for assessing the overarching influence of vegetation restoration on total
60 water resources (Zhao et al., 2021). Terrestrial Water Storage Anomalies (TWSA) data
61 derived from the Gravity Recovery and Climate Experiment (GRACE) satellites
62 indicate recent TWSA depletion attributed to vegetation restoration in various
63 revegetation areas in China, such as the Mu Us Sandyland (Zhao et al., 2021), the
64 Yellow River Basin (Cao et al., 2022), and the Loess Plateau (Liu et al., 2023a).

65 Terrestrial water storage changes represent a complex interplay of influences

66 stemming from both climate change and human activities. Numerous studies have
67 endeavored to discern the relative contributions of distinct factors shaping TWSA
68 dynamics in prominent basins across China. For instance, Li et al. (2020) identified
69 increased vegetation as the primary driver behind the decline in TWSA within the
70 Yellow River basin. In the northeastern basins, Li et al. (2023) elucidated a more
71 substantial impact of vegetation on TWSA trends, contrasting with basins in southern
72 China where precipitation emerged as the predominant factor. However, the specific
73 delineation of the relative contributions of climatic factors and vegetation restoration
74 to TWSA changes within the TN region remains unexplored.

75 Furthermore, soil moisture and groundwater, integral components of terrestrial water
76 storage, have frequently served as pivotal metrics for evaluating shifts in regional
77 water resources prompted by vegetation restoration initiatives. For instance, Jia et al.
78 (2017) observed a diminution in soil moisture associated with afforestation, while
79 Gao et al. (2015) documented notable declines in groundwater levels resulting from
80 large-scale revegetation endeavors. In contrast, Han et al. (2020) reported an increase
81 in soil moisture in the Loess Plateau of China, and Ellison et al. (2017) posited that
82 reforestation could enhance infiltration capacity and groundwater recharge. These
83 disparities underscore the necessity of comprehensive investigations into the
84 responses of soil moisture and groundwater to the combined influences of climate
85 change and vegetation restoration. Such insights are imperative for a nuanced
86 understanding of water resource dynamics in the TN region, where knowledge in this
87 domain remains limited.

88 As a result, the primary objectives of the current study are (1) to systematically
89 characterize the temporal evolution of TWSA in the TN region of China over the
90 preceding two decades (2003–2021); (2) to attribute the observed TWSA changes to
91 pertinent climate-related variables and the impacts of artificial revegetation; and (3) to
92 discern and systematically compare the nuanced responses of soil moisture and
93 groundwater to variations in environmental factors. The precise identification and
94 attribution of alterations in terrestrial water storage, coupled with a thorough

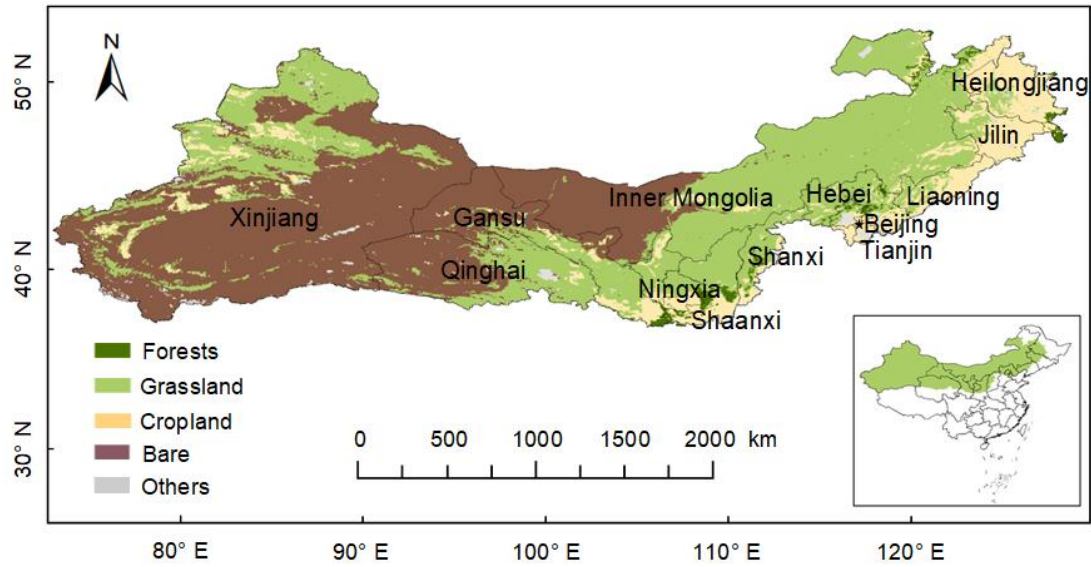
95 examination of soil moisture and groundwater dynamics within areas subjected to
96 revegetation initiatives, are essential for the comprehensive evaluation of ecological
97 restoration endeavors. Furthermore, the insights derived from this study carry
98 substantial implications for guiding future planning initiatives in the realm of
99 sustainable water resource management.

100 **2 Materials and methods**

101 **2.1 Study region**

102 The TN region, encompassing 4 million square kilometers and constituting 42% of
103 China's land area, spans across 13 provinces, autonomous regions, and municipalities
104 as illustrated in Fig. 1. Characterized as a typical arid and semi-arid expanse, a
105 majority of its regions receive an average annual precipitation of less than 400 mm.
106 The TN region predominantly features bare and grassland, with relatively low forest
107 coverage, rendering it susceptible to land degradation, desertification, and dust
108 storms. In response to these challenges, the Chinese government initiated an extensive
109 vegetation restoration effort since 1978 in the TN region, with plans for continuation
110 until 2050.

111 Over the past few decades, a noteworthy surge in vegetation coverage has been
112 observed in the TN region, particularly evident in the forest coverage rate, which
113 escalated from 5.05% in 1978 to 13.57% in 2018 (State Forestry and Grassland
114 Administration, 2018). These initiatives have yielded significant ecological benefits
115 for the region, including soil stabilization, a reduction in dust storm occurrences, and
116 climate regulation. However, they have introduced a paradoxical challenge by
117 exacerbating water scarcity within the region. These features make the TN region
118 stands as an ideal research area for detecting and attributing changes in terrestrial
119 water storage induced by both climate-related factors and vegetation restoration.



120

121 **Figure 1.** Geographical delineation of the study area.

122 **2.2 Datasets**

123 **2.2.1 Water storage**

124 Monthly TWSA datasets employed in this study were derived from both the GRACE
 125 and its successor, the GRACE Follow-on (GRACE-FO) satellites. We utilized
 126 GRACE/GRACE-FO mascon solutions from three institutions: Jet Propulsion
 127 Laboratory (JPL) (<http://grace.jpl.nasa.gov>) (Landerer and Swenson, 2012), Center
 128 for Space Research (CSR) (<http://www2.csr.utexas.edu/grace>) (Save et al., 2016), and
 129 Goddard Space Flight Center (GSFC) ([https://earth.gsfc.nasa.gov/geo/data/grace-](https://earth.gsfc.nasa.gov/geo/data/grace-mascons)
 130 [mascons](https://earth.gsfc.nasa.gov/geo/data/grace-mascons)) (Loomis et al., 2019). The JPL mascon product was employed with
 131 specified gain factors, which were obtained from <http://grace.jpl.nasa.gov>. These
 132 mascon series utilize regularization techniques in data processing, enabling the direct
 133 computation of variations in terrestrial water storage from Level 1B data (Save et al.,
 134 2016). This approach eliminates the need for filtering, smoothing, and other empirical
 135 scaling techniques, thereby reducing errors and providing enhanced signal-to-noise
 136 ratios and spatial resolutions (Watkins et al., 2015; Save et al., 2016). An overlapping
 137 period of JPL, CSR, and GSFC TWSA datasets from 2003 to 2021 was chosen as our
 138 study period. An ensemble mean of these datasets was adopted to mitigate
 139 uncertainties associated with data processing. CSR TWSA data, initially provided at a

140 spatial resolution of $0.25^\circ \times 0.25^\circ$, were resampled to $0.5^\circ \times 0.5^\circ$ to maintain
141 consistency with the resolutions of JPL and GSFC TWSA data. Linear interpolation of
142 adjacent months or averaging monthly TWSA values from adjacent years for a
143 specific month was applied to address missing values (Andrew et al., 2017; Chen et
144 al., 2019).

145 To gain further insights into changes in TWSA, soil moisture (0–1 m, including the
146 surface layer and the root zone) and groundwater data were obtained from Global
147 Land Data Assimilation System Catchment Land Surface Model (GLDAS_CLSM)
148 outputs. The model is driven by meteorological datasets from the European Center for
149 Medium-Range Weather Forecasts (ECMWF) (Li et al., 2019). Groundwater
150 calculation involves subtracting the sum of canopy water, snow water equivalent, and
151 soil moisture from total water storage. The GLDAS_CLSM_2.2 provides daily-scale
152 soil moisture and groundwater data at a spatial resolution of $0.25^\circ \times 0.25^\circ$
153 ([https://disc.gsfc.nasa.gov/datasets/GLDAS_CLSM025_DA1_D_2.2/summary?keyw](https://disc.gsfc.nasa.gov/datasets/GLDAS_CLSM025_DA1_D_2.2/summary?keywords=GLDAS)
154 [ords=GLDAS](https://disc.gsfc.nasa.gov/datasets/GLDAS_CLSM025_DA1_D_2.2/summary?keywords=GLDAS)), which were converted into monthly time series at a resolution of 0.5°
155 $\times 0.5^\circ$ for this study to facilitate analysis.

156 **2.2.2 Meteorological variables**

157 Monthly precipitation (P) data were sourced from the Multi-Source Weighted-
158 Ensemble Precipitation (MSWEP, <https://www.gloh2o.org/mswep/>) dataset (Beck et
159 al., 2019). This dataset, offering an optimal ensemble of gauge observations, satellite
160 retrievals, and climate reanalysis, provides precipitation information at a spatial
161 resolution of $0.1^\circ \times 0.1^\circ$. Monthly potential evapotranspiration (PET) was obtained
162 from the Global Land Evaporation Amsterdam Model (GLEAM v3.7b,
163 <https://www.gleam.eu/>) at a spatial resolution of $0.25^\circ \times 0.25^\circ$ (Miralles et al., 2011;
164 Martens et al., 2017), where PET is estimated from the Priestley-Taylor model. Both
165 P and PET data were resampled to a spatial resolution of $0.5^\circ \times 0.5^\circ$ to be consistent
166 with the TWSA data.

167 **2.2.3 Vegetation index**

168 Vegetation changes were evaluated using the Normalized Difference Vegetation Index
169 (NDVI). Monthly NDVI data obtained from the Moderate Resolution Imaging
170 Spectroradiometer (MODIS,
171 https://neo.gsfc.nasa.gov/view.php?datasetId=MOD_NDVI_M) were employed in this
172 study (Didan et al., 2015). These data were acquired at a spatial resolution of $0.5^\circ \times$
173 0.5° and were derived from MODIS/Terra Vegetation Indices L3 Global 0.05 Deg
174 CMG V006 (MOD13C1 and MOD13C2).

175 **2.3 Attribution analyses**

176 To attribute the temporal changes of monthly TWSA, we first employed a stepwise
177 multiple linear regression (Draper and Smith, 1998; Clow, 2010) to identify
178 significant predictor variables, assessed at a 5% significance level through an F-test,
179 explaining the temporal variability of TWSA for each grid cell. Subsequently, we
180 applied the dominance analysis approach (Azen and Budescu, 2003) to assess the
181 relative importance of the identified variables. In contrast to conventional methods
182 providing an overall evaluation of variable importance, dominance analysis explores
183 the unique contribution of each predictor by considering all possible subsets of
184 variables. This method entails constructing multiple regression models for each subset
185 and quantifying the extent to which each predictor dominates in explaining the
186 variance of the dependent variable. The dominance weight, computed as the
187 difference in the average coefficient of determination (R^2) between models including
188 and excluding a particular predictor, serves to quantify the relative importance of each
189 variable. This approach accommodates interactions and collinearity among predictors,
190 providing a nuanced understanding of their individual contributions (Vize et al.,
191 2019). The resulting dominance weights facilitate the ranking of predictors based on
192 their average relative importance, contributing to a relatively comprehensive
193 interpretation of variable significance within the entire model context.

194 **3 Results**

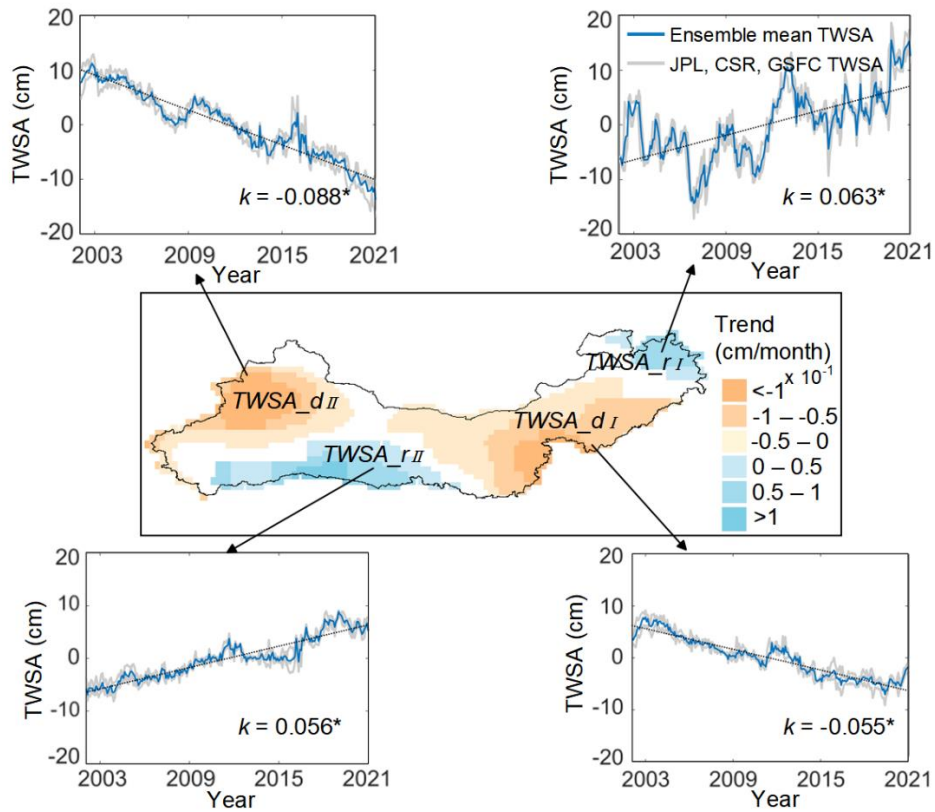
195 **3.1 Hotspots of TWSA changes in the TN region**

196 The linear trend of TWSA was assessed for each grid cell in the TN region spanning

197 from 2003 to 2021, utilizing TWSA datasets provided by JPL, CSR, and GSFC.
198 Hotspots exhibiting consistent and significant increasing or decreasing trends across
199 the three datasets were identified (Supplementary Fig. S1), resulting in the delineation
200 of two TWSA rising (*TWSA_rI* and *TWSA_rII*) and two TWSA declining (*TWSA_dI*
201 and *TWSA_dII*) regions (Fig. 2). Specifically, *TWSA_rI*, predominantly located in the
202 northeastern part of the study area (primarily within Heilongjiang province), covered
203 7.0% of the total grid cells within the identified hotspots. *TWSA_rII*, situated in the
204 southwestern part (mainly within Qinghai province), encompassed 19.6% of the total
205 grid cells within the four hotspots. *TWSA_dI*, spanning portions of Beijing, Tianjin,
206 and the provinces of Hebei, Shanxi, Liaoning, Shaanxi, as well as the Inner Mongolia
207 and Ningxia Hui Autonomous regions, exhibited the largest coverage (43.4%) among
208 the identified hotspots. *TWSA_dII* (geographically associated with the Xinjiang Uygur
209 Autonomous Region) represented 30.0% of the total grid cells within the four
210 hotspots. The areal average TWSA (ensemble mean of JPL, CSR, and GSFC TWSA
211 datasets) over these hotspots is also presented in Fig. 2. Both TWSA rising (declining)
212 hotspots displayed significant increasing (decreasing) trends, with all trends being
213 statistically significant at the 0.05 significance level.

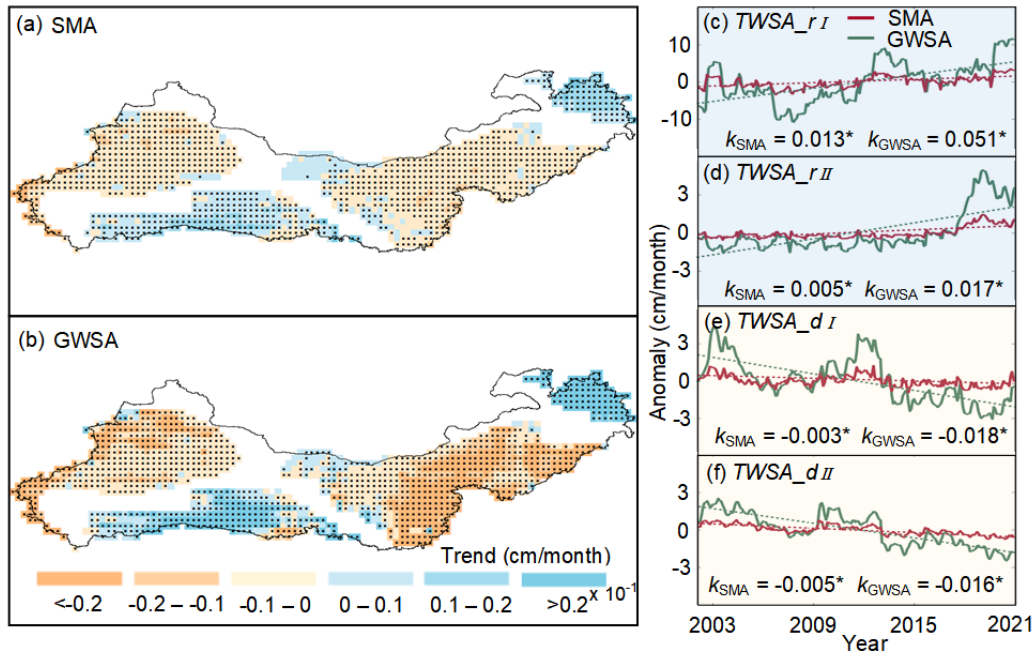
214 Spatial trends in GLDAS-derived soil moisture anomalies (SMA) and groundwater
215 anomalies (GWSA) in the identified hotspots are illustrated in Figs. 3a and b,
216 respectively. Generally, both SMA and GWSA demonstrated increasing trends in
217 *TWSA_rI* and *TWSA_rII* and decreasing trends in *TWSA_dI* and *TWSA_dII* from 2003
218 to 2021. However, the trends of SMA were generally weaker than those of GWSA,
219 attributable to the GLDAS_CLSM's definition of soil moisture as the water storage in
220 the top 1 meter of the soil, containing a comparatively smaller amount of water
221 compared to groundwater (Supplementary Fig. S2). In addition, the areal average time
222 series of SMA and GWSA in the four hotspots of TWSA changes are illustrated in
223 Figs. 3c–f. Similar to TWSA, both SMA and GWSA exhibited significant increasing
224 trends in *TWSA_rI* and *TWSA_rII*, with corresponding significant decreasing trends in
225 *TWSA_dI* and *TWSA_dII*. Notably, the amplitudes of groundwater were larger than

226 those of soil moisture in the identified hotspots.



227

228 **Figure 2.** Spatial distribution of TWSA hotspots exhibiting significant rising trends
229 (*TWSA_rI* and *TWSA_rII*) and declining trends (*TWSA_dI* and *TWSA_dII*), alongside
230 the interannual variations in areal average TWSA (ensemble mean derived from JPL,
231 CSR, and GSFC TWSA datasets) during the period 2003–2021 within the specified
232 hotspots. The symbol “*” denotes instances where the linear trend (*k*) is statistically
233 significant at the 5% significance level.



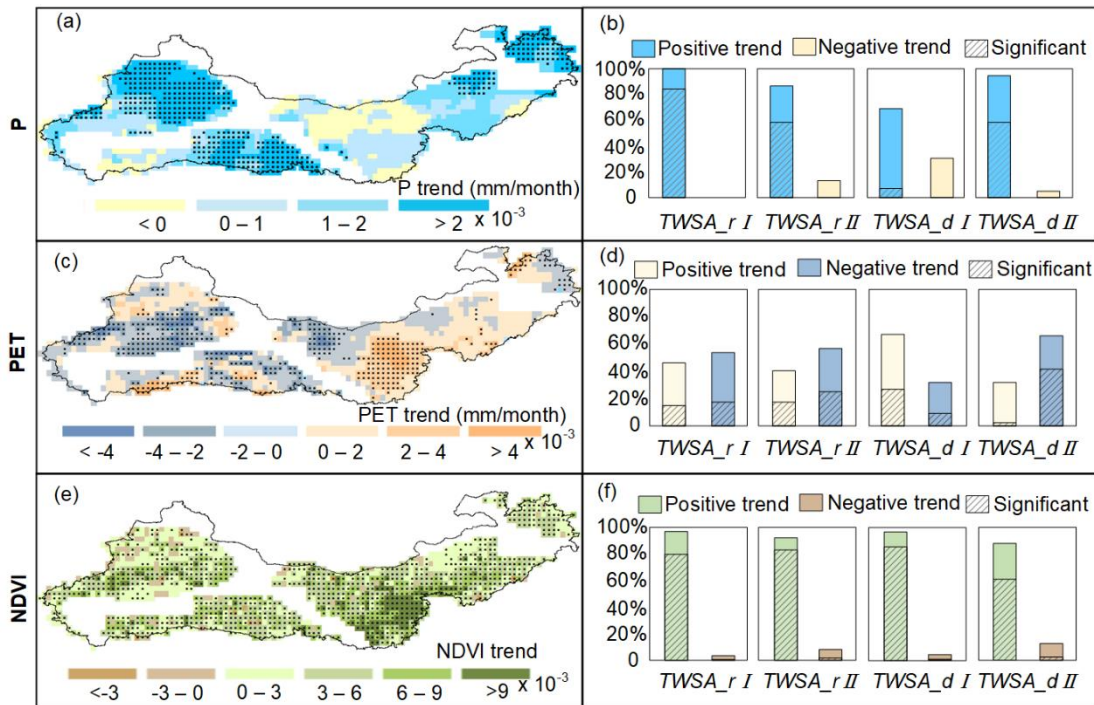
234

235 **Figure 3.** Spatial trend maps depicting (a) soil moisture anomalies (SMA) and (b)
 236 groundwater anomalies (GWSA) in the four identified hotspots during the period
 237 2003–2021. Interannual variations of SMA and GWSA are presented for (c) *TWSA_rI*,
 238 (d) *TWSA_rII*, (e) *TWSA_dI*, and (f) *TWSA_dII*. The symbol “*” denotes instances
 239 where the linear trend (*k*) is statistically significant at the 5% significance level.

240 3.2 Attribution of TWSA changes

241 In this study, the atmospheric water supply (*P*) and demand (PET) are regarded as
 242 crucial climate drivers influencing TWSA, and their impacts are compared with those
 243 of artificial vegetation restoration, as indicated by changes in the NDVI. The linear
 244 trends of these three environmental factors (*P*, PET, and NDVI) across the four
 245 identified hotspots are presented in Fig. 4. Precipitation has exhibited an increase in
 246 81.0% of the total grid cells within the four hotspots, with 46.3% of them being
 247 statistically significant at the 0.05 significance level (Fig. 4a). Grid cells displaying
 248 decreasing precipitation were observed in *TWSA_rII*, *TWSA_dI*, and *TWSA_dII*,
 249 although none of these trends passed the significance test (Fig. 4b). PET showed a
 250 decreasing trend in 50.4% of the total grid cells within the four hotspots, and the
 251 significantly decreasing PET (29.9%) was predominantly observed in *TWSA_rII*,
 252 *TWSA_dIII*, and the western part of *TWSA_dI* (Figs. 4c–d). Conversely, PET displayed

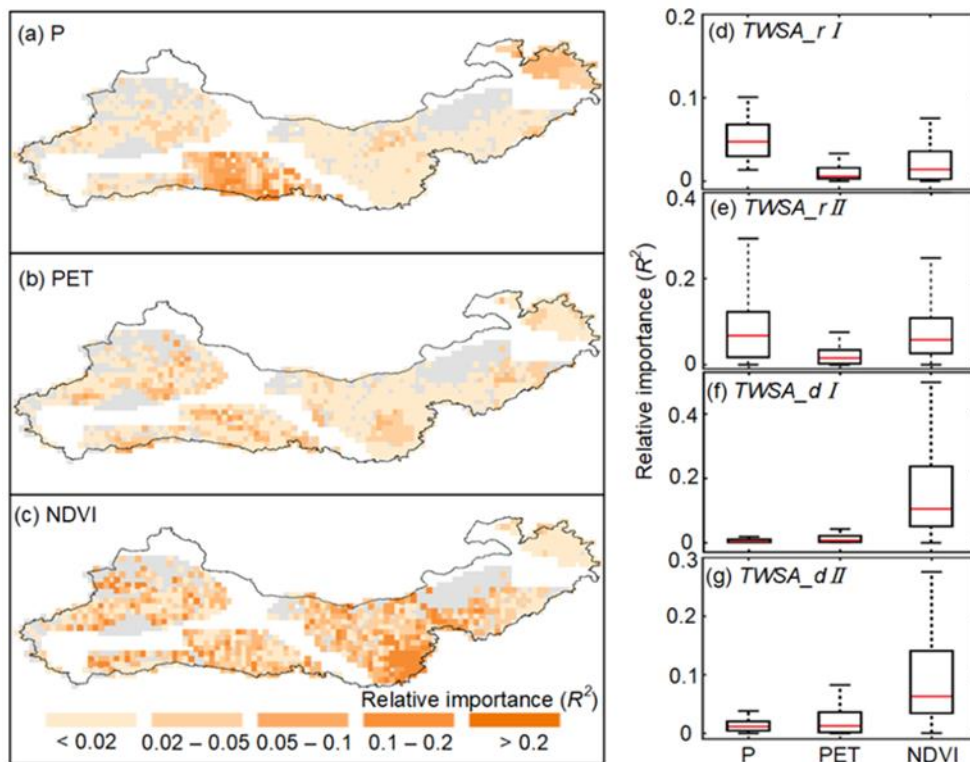
253 an increasing trend in 49.6% of the grid cells within the four hotspots, with 46.9% of
 254 them passing the significance test. Significant increases in PET were mainly observed
 255 in the Loess Plateau region of China, situated in *TWSA_dI* (Fig. 4c). Additionally,
 256 NDVI exhibited an increasing trend in 92.6% of the grid cells within the four
 257 hotspots, with 83.0% of them being statistically significant, indicating a notable
 258 greening of vegetation in the TN region over the past two decades (Figs. 4e–f). The
 259 most conspicuous greening area was identified in the Loess Plateau region, where the
 260 Grain to Green Program (GTGP), recognized as the largest ecological restoration plan
 261 ever in a developing country (Wang et al., 2016), has been implemented since 1999.



262
 263 **Figure 4.** Spatial distribution and histograms illustrating the trends of (a–b)
 264 precipitation (*P*), (c–d) potential evapotranspiration (PET), and (e–f) vegetation index
 265 (NDVI) across the four identified hotspots during the period 2003–2021.

266 Figures 5a–c respectively present the relative importance, as indicated by the R^2
 267 values derived from the dominance analysis, of *P*, PET, and NDVI in elucidating the
 268 temporal variability of monthly TWSA in the four identified hotspots. 70% of the grid
 269 cells within the identified hotspots were selectively displayed based on the criterion
 270 that the total explanatory power of the three influencing factors (*P*, PET, and NDVI)

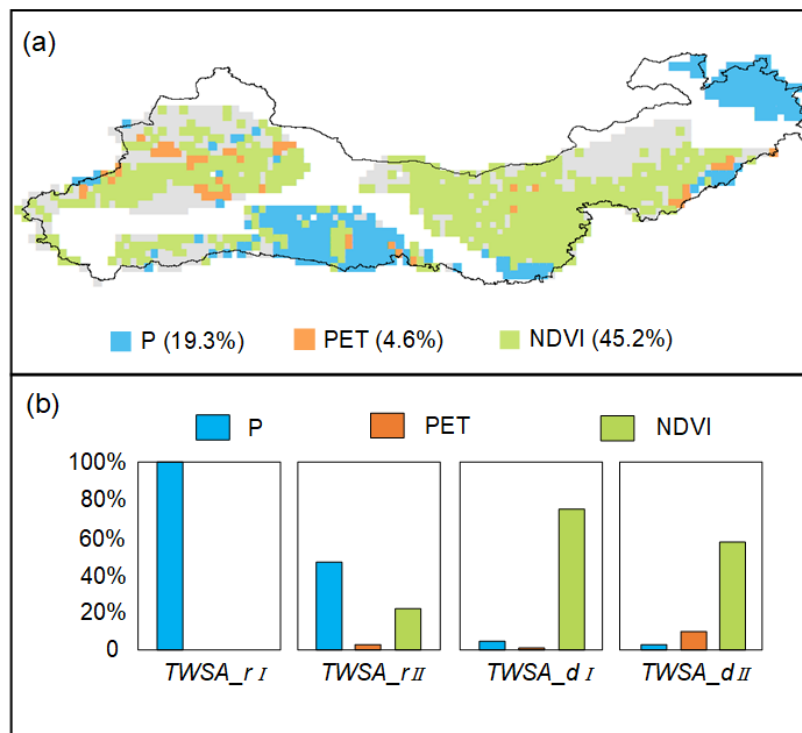
271 surpassed the 5% significance level, represented by graded colors. TWSA changes in
 272 gray areas were not significantly associated with either the two climate-related factors
 273 or vegetation restoration during the study period from 2003 to 2021. In general, *P*
 274 played a more pivotal role than PET and NDVI in influencing TWSA changes in
 275 *TWSA_rI* and *TWSA_rII* (Fig. 5d–e). Conversely, NDVI exhibited greater relative
 276 importance than climate-related factors in explaining the temporal variability of
 277 TWSA in *TWSA_dI* and *TWSA_dII* (Figs. 5 f–g).



278
 279 **Figure 5.** Spatial distribution of the relative importance (R^2) of (a) *P*, (b) PET and (c)
 280 NDVI in explaining TWSA temporal variability (areas in gray means none of the
 281 environmental factor has been selected as significant predictor variable in explaining
 282 TWSA temporal variability). The comparisons of the relative importance of each
 283 environmental factor in the four hotspots: (d) *TWSA_rI*, (e) *TWSA_rII*, (f) *TWSA_dI*
 284 and (g) *TWSA_dII*.

285 Subsequently, the dominant driver influencing TWSA was identified through
 286 comparing the relative importance of *P*, PET, and NDVI at each grid cell. As
 287 illustrated in Fig. 6a, nearly half of the area (45.2%) within the four identified

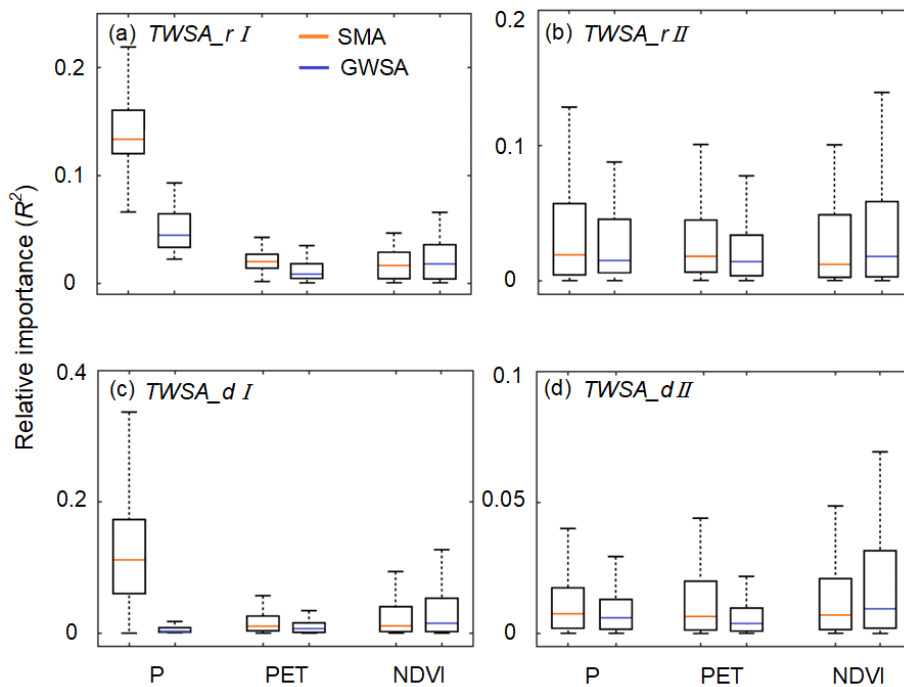
288 hotspots exhibited TWSA changes primarily dominated by NDVI. Areas dominated
 289 by P and PET accounted for 19.3% and 4.6% of the grid cells within the hotspots,
 290 respectively. Specifically, P -dominant areas were predominantly situated in regions
 291 where TWSA exhibited significant increasing trends (i.e., $TWSA_{rI}$ and $TWSA_{rII}$).
 292 Conversely, in regions with TWSA displaying significant decreasing trends (i.e.,
 293 $TWSA_{dI}$ and $TWSA_{dII}$), NDVI predominantly drove the observed TWSA changes.
 294 While the NDVI in $TWSA_{dII}$ generally exhibited a weaker increasing trend
 295 compared to that in $TWSA_{dI}$, it's important to note the relatively shallow
 296 groundwater table in $TWSA_{dII}$ (Fan and Miguez-Macho, 2013), where vegetation
 297 heavily relies on groundwater. Consequently, even a modest increase in NDVI within
 298 $TWSA_{dII}$ could potentially result in a decline in TWSA. These outcomes underscore
 299 the replenishing effects of precipitation on terrestrial water storage and the impact of
 300 vegetation restoration in consuming terrestrial water storage.



301
 302 **Figure 6.** (a) Spatial distribution and (b) histograms illustrating the dominant
 303 influencing factor among P , PET and NDVI for TWSA temporal variability in the
 304 four identified hotspots.

305 **3.3 Comparison of SMA and GWSA responses to environmental changes**

306 As water storage at different layers of the soil may respond differently to
 307 environmental changes, we additionally conducted a comparative analysis of the
 308 relative importance of P , PET, and NDVI in elucidating the temporal variability of
 309 SMA and GWSA in the four identified hotspots, as depicted in Fig. 7. Climate-related
 310 factors, specifically P and PET, generally exhibited a larger relative importance in
 311 explaining the temporal variability of SMA than in elucidating changes in GWSA
 312 across the identified hotspots. Additionally, NDVI demonstrated a more pronounced
 313 relative importance in elucidating changes in GWSA compared to its role in
 314 explaining SMA in all identified hotspots. This observation suggests that water
 315 storage in shallower layers is more susceptible to influences from climate-related
 316 factors, while deeper layers exhibit a higher degree of responsiveness to dynamics in
 317 vegetation.



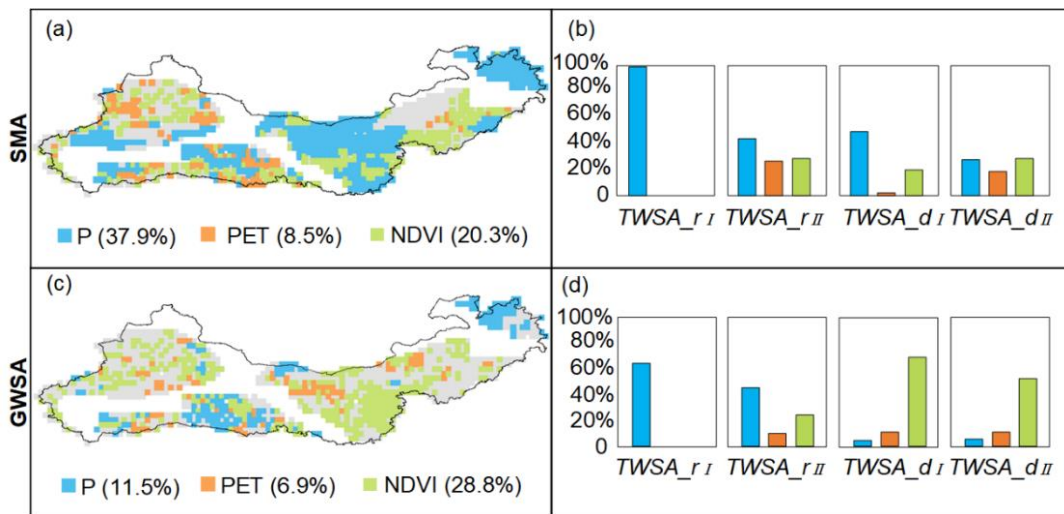
318
 319 **Figure 7.** Relative importance of P , PET and NDVI in explaining SMA and GWSA
 320 temporal variability in (a) $TWSA_{rI}$, (b) $TWSA_{rII}$, (c) $TWSA_{dI}$ and (d) $TWSA_{dII}$.

321 By comparing the relative importance of P , PET and NDVI in explaining SMA and
 322 GWSA, dominant influencing factors for SMA and GWSA changes were identified in
 323 Fig. 8. The temporal variability of SMA and GWSA was primarily governed by

324 precipitation in 37.9% and 11.5% of the grid cells within the four identified hotspots,
 325 as illustrated in Figs. 8a and c. In regions characterized by increased SMA and GWSA
 326 (i.e., *TWSA_rI* and *TWSA_rII*), climate-related factors exerted a more substantial
 327 influence than NDVI on water storage across larger areas. Notably, increased
 328 precipitation emerged as the predominant driver, replenishing water storage at both
 329 shallow and deep layers.

330 In regions displaying decreased SMA and GWSA (i.e., *TWSA_dI* and *TWSA_dII*),
 331 water storage at the shallow layer exhibited greater responsiveness to climate-related
 332 factors (e.g., decreased precipitation and increased PET), leading to a decline in SMA
 333 in approximately 50% of grid cells in *TWSA_dI* and *TWSA_dII* (Fig. 8b).

334 Concurrently, water storage depletion at the deep layer was predominantly driven by
 335 increased NDVI, with NDVI dominating TWSA changes in over 50% of the areas in
 336 *TWSA_dI* and *TWSA_dII* (Fig. 8d).



337
 338 **Figure 8.** Spatial distribution and histograms illustrating the dominant influencing
 339 factor among *P*, *PET* and *NDVI* for (a–b) *SMA* and (c–d) *GWSA* temporal variability
 340 in the four identified hotspots.

341 **4 Discussion**

342 In the context of climate change and vegetation restoration, a comprehensive
 343 understanding of the evolutionary dynamics and underlying causes of *TWSA* in arid
 344 regions is imperative. Here we identified a noteworthy reduction in *TWSA* across

345 nearly three-quarters of the TN region, encompassing extensive areas of the Loess
346 Plateau (identified as *TWSA_dI*) and northwestern China (identified as *TWSA_dII*).
347 Attribution analysis indicates that the primary factor contributing to the declining
348 TWSA in these regions is enhanced water consumption induced by vegetation
349 restoration. This observation suggests a potential over-planting phenomenon in these
350 areas, which is also supported by earlier studies (e.g., Liu et al., 2023a; Liu et al.,
351 2023b). Conversely, we also identified two TWSA rising hotspots within the TN
352 region, specifically *TWSA_rI* in the northeastern sector and *TWSA_rII* in the
353 southwestern region, which was primarily due to precipitation replenishment that
354 surpassed the water consumption associated with vegetation restoration. Additionally,
355 increased TWSA in *TWSA_rII* may be partially ascribed to glacial mass gain in the
356 Karakoram Mountains (Karakoram anomaly) (Forsythe et al., 2017; Farinotti et al.,
357 2020) and the expansion of lakes in the Tibetan Plateau's endorheic basin (Zhang et
358 al., 2017; Li et al., 2022). These findings highlight the intricate dynamics of the
359 impact of vegetation restoration on TWSA, emphasizing that its influence is not
360 universally associated with a decline in TWSA. Precipitation replenishment emerges
361 as a crucial factor capable of substantially or entirely offsetting the augmented water
362 consumption resulting from vegetation restoration. Therefore, the continuation of
363 revegetation initiatives is deemed suitable in regions where the positive effects of
364 precipitation recharge outweigh the negative impact of additional water consumption
365 by extra vegetation. This recommendation aligns with the guidance provided by Xu et
366 al. (2023), who propose a similar strategy for semiarid and subhumid areas in
367 northern China. In these regions, where the benefits stemming from climate change
368 surpass the adverse contributions of vegetation to soil moisture, the ongoing efforts in
369 revegetation are advised. Conversely, in hyperarid and arid subareas, a focus on
370 nature-based recovery is recommended.

371 In addition to examining total water storage changes, this study addresses the
372 knowledge gap regarding water storage variations at different soil depths in the TN
373 region. The results reveal a noteworthy consistency in trends between SMA and

374 GWSA, with GWSA exhibiting a larger magnitude. The reduced GWSA observed in
375 *TWSA_dI* and *TWSA_dII* aligns with findings from previous studies. Specifically, Han
376 et al. (2020) documented GWSA depletion in the Loess Plateau of China
377 (corresponding to *TWSA_dI*) during 2003–2015, while Liu et al. (2023b) reported a
378 decline in GWSA in northwestern China (associated with *TWSA_dII*) from 2003 to
379 2018. The decline in SMA within *TWSA_dI* and *TWSA_dII* also aligns with earlier
380 studies by Jia et al. (2017) and Liu et al. (2023b), respectively. However, it is
381 noteworthy that Han et al. (2020) reported an increased SMA in specific areas within
382 the Loess Plateau, attributing it to the positive impact of the water conservancy
383 system in Ningxia and Hetao irrigation areas. This disparity underscores the
384 complexity of accurately detecting changes in water storage, emphasizing the
385 necessity of considering the influence of various anthropogenic activities.

386 Notably, anthropogenic mass flow signals in GRACE/GRACE-FO observations
387 attributed to coal mining and transportation (Tang et al., 2013; Zhou et al., 2023), as
388 well as domestic and industrial water consumption and irrigation (Felfelani et al.,
389 2017), were not specifically accounted for. The stronger decreasing trends in
390 GRACE-derived TWSA compared to GLDAS model-derived TWSA in both
391 *TWSA_dI* and *TWSA_dII* (Supplementary Fig. S3) suggest potential influences from
392 the aforementioned human activities. Annual water consumption records from the
393 National Bureau of Statistics (<http://www.stats.gov.cn/sj/ndsj/>) for *TWSA_dI* and
394 *TWSA_dII* reveal a notable rise in domestic, industrial, agricultural, and ecological
395 water consumption over the past two decades (Supplementary Figs. S4a and b). These
396 regions, corresponding to areas of groundwater over-exploitation in China (Huang et
397 al., 2023), align with the North China Plain for *TWSA_dI* and the Tarim Basin for
398 *TWSA_dII*. In addition, a decline in TWSA and GWSA was observed during the
399 period 2003 – 2009 (Figs. 2 and 3c), which might also be related to the increased
400 water consumption due to human activities (Fig. S4c). Thus, it is evident that, in
401 addition to artificial vegetation restoration, water consumption resulting from diverse
402 human activities has also contributed to TWSA depletion.

403 In addition, the causal relationship between vegetation and TWSA revealed herein
404 should be viewed with caution. We assume that the observed increase in vegetation
405 over the TN region can largely be attributed to human-led ecological restoration
406 projects, but it should be noted that the survival and subsequent growth of artificially
407 planted vegetation were heavily contingent upon local water conditions (McVicar et
408 al., 2007). Our results showed that increased TWSA in around 20.0% area of
409 *TWSA_rII* was dominated by NDVI (Fig. 6). This might be explained by the enhanced
410 water retention capacity of soil that benefits from increased vegetation. More
411 extensive vegetation coverage and improved root systems could facilitate water
412 infiltration and soil hydrophilicity (Lange et al., 2009; Guan et al., 2010; Wang et al.,
413 2013; Chen et al., 2021), and consequently impede runoff formation and contribute to
414 increases in terrestrial water storage. This issue deserves further exploration in the
415 future.

416 **5 Conclusions**

417 This study utilized TWSA data derived from the GRACE and GRACE-FO missions
418 to examine water storage dynamics within the context of climate change and
419 vegetation restoration in the TN region from 2003 to 2021. Four distinctive hotspots
420 were identified, characterized by either ascending (*TWSA_rI* and *TWSA_rII*) or
421 descending (*TWSA_dI* and *TWSA_dII*) trends in TWSA. These findings were
422 corroborated by soil moisture and groundwater data obtained from the GLDAS
423 model.

424 To attribute the observed regional TWSA changes to climatic drivers (*P* and PET) and
425 artificial vegetation restoration, dominance analysis was applied to assess the relative
426 importance of *P*, PET, and NDVI in influencing the temporal variability of TWSA.
427 Results indicated that vegetation exerted a more pronounced influence than climate-
428 related factors on TWSA changes in descending TWSA hotspots (*TWSA_dI* and
429 *TWSA_dII*), while precipitation predominantly governed TWSA changes in most
430 ascending TWSA hotspots (*TWSA_rI* and *TWSA_rII*) during the study period. Regions
431 dominated by NDVI constituted 45.2% of the area within the identified hotspots,

432 while those dominated by *P* and PET covered 19.3% and 4.6% of the area,
433 respectively. Furthermore, our study revealed that water storage at shallow layers
434 exhibited heightened responsiveness to climate-related factors, whereas the impacts of
435 vegetation changes on TWSA were more discernible in water storage at deeper layers.
436 In summary, the outcomes of this study contribute valuable insights into the evolution
437 and causative factors of TWSA changes in the TN region. Additionally, they offer
438 scientific guidance for sustainable water resource management and inform decision-
439 making processes related to future ecological restoration initiatives.

440 **CRedit authorship contribution statement**

441 Ajiao Chen: Conceptualization, Methodology, Formal analysis, Resources, Writing –
442 original draft, Funding acquisition. Jinghua Xiong: Resources, Writing – review &
443 editing. Shixue Wu: Conceptualization, Methodology. Yuting Yang:
444 Conceptualization, Methodology, Writing – review & editing, Supervision, Funding
445 acquisition.

446 **Declaration of Competing Interest**

447 The authors declare that they have no known competing financial interests or personal
448 relationships that could have appeared to influence the work reported in this paper.

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