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1 **Wellbore stability in high-temperature granite under true triaxial stress**

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22 **Abbreviations:** AE, acoustic emissions

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26 **Abstract**

27 Supercritical/superhot geothermal success depends on successful drilling. However, wellbore
28 stability in supercritical environments has not been investigated because imposing sufficient
29 stress on high-temperature granite with true triaxial loading is difficult. We conducted wellbore
30 failure experiments on 200–450 °C granite under true triaxial stress, including deformation and
31 acoustic emission measurements, and post-experiment thin section observations. Wellbore
32 failure initiated at stress states consistent with existing brittle failure criterion at the studied
33 temperatures. Using breakout geometry for *in situ* stress estimation may be difficult as shear
34 failure propagation is suppressed at high temperatures; however, boreholes may be inherently
35 stable in high-temperature environments.

36 **Keywords:** wellbore failure, granite, high temperature, supercritical geothermal environment,
37 superhot geothermal environment, true triaxial stress

40 **1. Introduction**

41 Recent attention in geothermal power generation has focused on methods for harvest
42 geothermal energy from unconventional supercritical (or superhot) geothermal environments
43 that exceed the critical temperature of water (374 °C for pure water and 406 °C for seawater)
44 at drillable depths of approximately 2–4 km, as found in Iceland (Friðleifsson et al., 2014;
45 Friðleifsson and Elders, 2017; Steingrímsson et al., 1990), Italy (Baron and Ungemach, 1981;
46 Batini et al., 1983; Ruggieri and Gianelli, 1995), Japan (Kato et al., 1998), Mexico (Espinosa-
47 Paredes and Garcia-Gutierrez, 2003), and the United States (Fournier, 1991; Garcia et al.,
48 2016). Supercritical geothermal environments can provide superheated steam or supercritical
49 water with an extremely high specific enthalpy exceeding 2 MJ/kg (Elders et al., 2014;
50 Friðleifsson and Elders, 2005; Friðleifsson et al. 2007; Smith and Shaw, 1975, 1979; Tester et

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51 al., 2006). Therefore, harvesting supercritical geothermal energy is expected to increase the
52 productivity and sustainability of geothermal energy generation. However, in supercritical
53 geothermal environments on the continental crust, the formation of permeable fracture
54 networks (Tsuchiya et al., 2016; Watanabe et al., 2017a; Weis et al., 2012) can potentially be
55 suppressed by the increased efficiency of mineral crystal plasticity (Parisio et al., 2019, 2020;
56 Tullis and Yund, 1977; Violay et al., 2017), the retrograde solubility of silica minerals
57 (Fournier, 1991; Saishu et al., 2014; Tsuchiya and Hirano, 2007; Watanabe et al., 2021), and
58 the enhanced rates of fracture healing/sealing caused by water–rock reactions (Moore et al.,
59 1994; Morrow et al., 2001; Watanabe et al., 2020). Therefore, it is necessary to develop
60 enhanced geothermal system technologies to artificially create or recreate permeable fracture
61 networks in these supercritical geothermal environments.

62 Recent studies have suggested that the injection of water at or near its critical temperature
63 into granite is likely to create a dense network of microfractures suitable for geothermal energy
64 extraction, a so-called cloud-fracture network, through continuous infiltration and the
65 stimulation of pre-existing microfractures by the low-viscosity water (Goto et al. 2021;
66 Watanabe et al., 2017b, 2019). Additionally, it has been suggested that the pore pressure
67 required to initiate this flow-induced microfracturing may be well predicted by the Griffith
68 failure criterion, which assumes the initiation of brittle failure from pre-existing fractures and
69 is a function of the maximum and minimum principal stresses and tensile strength of rocks
70 (Cox, 2010; Griffith, 1924; Jaeger and Cook, 1979; Secor, 1965). However, the characteristics
71 of wellbore failure (Ewy and Cook, 1990; Morita and Nagano, 2016), and the possibility of
72 utilizing these failures (i.e., induced fractures and breakout) for in situ stress estimation (Brudy
73 and Zoback, 1999; Kato et al., 1998; Zoback et al., 1985) to achieve safe and effective
74 fracturing at the minimum required injection pressure, remain unclear for supercritical
75 geothermal environments. This is because wellbore failure experiments under true triaxial

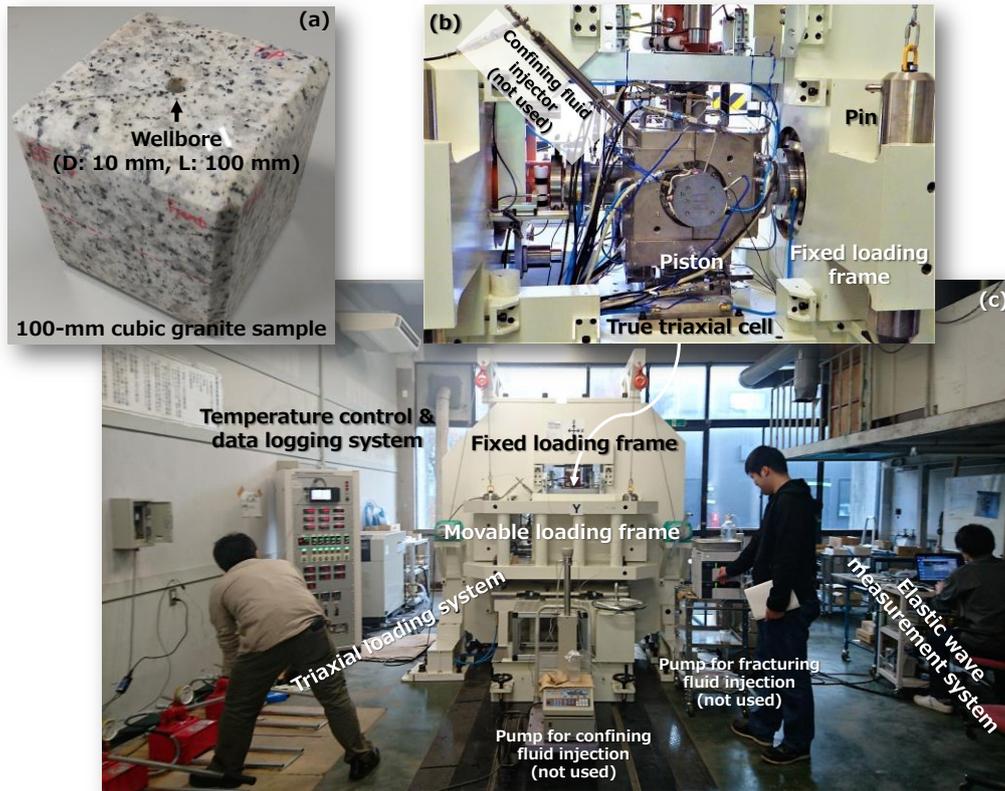
1 76 stress (Lee and Haimson, 1993; Song and Haimson, 1997) have never been conducted on high-
2 77 temperature granite. To the best of our knowledge, wellbore failure experiments under
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4 78 hydrostatic stress have only been conducted on granite at temperatures up to 600 °C (Zhao et
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7 79 al., 2015).

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9 80 The aim of this study is to experimentally simulate wellbore failure in granite under true
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11 81 triaxial stress in supercritical geothermal conditions. We first introduce a novel experimental
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13 82 method for assessing high-temperature wellbore failure in relatively large (10-cm cubic) rock
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15 83 samples under true triaxial stress. Subsequently, we present the results of experiments
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17 84 conducted on granite at 200–450 °C, which show non-catastrophic wellbore failure. Finally,
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19 85 we suggest the predictability of the initiation of wellbore failure using an existing failure
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21 86 criterion for rocks and discuss the possibility of estimating in situ stress from wellbore failure.
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29 88 **2. Experimental methods**

30 31 89 **2.1. Granite and experimental system**

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33 90 Cubes (100 × 100 × 100 mm) of Inada granite from Ibaraki prefecture, Japan (Figure 1),
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35 91 with a single wellbore (diameter: 10 mm, length: 100 mm) at the center, were prepared for the
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37 92 high-temperature wellbore failure experiment under true triaxial stress. Inada granite, of which
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39 93 the porosity at room temperature and atmospheric pressure is 0.5-0.8%, has been used in
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41 94 previous studies on supercritical hydraulic fracturing (Goto et al., 2021; Watanabe et al., 2017b,
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43 95 2019).
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97 Figure 1. Granite sample (a), high-temperature true triaxial cell (b), and experimental
98 (c) in the wellbore failure experiment, adapted from Watanabe et al. (2019).
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100 We used the experimental system (Figure 1) originally developed for supercritical
101 hydraulic fracturing by Watanabe et al. (2019) and applied in Goto et al. (2021). The system
102 consisted of a true triaxial cell, a triaxial loading system, a pump for injecting fracturing fluid
103 into the sample, a pump for injecting confining fluid (high-viscosity plastic melt in this system)
104 along the sample edges via an injector, an elastic wave measurement system, and a temperature
105 control and data logging system. In the present experiment, the elastic wave measurement
106 system was an acoustic emission (AE) measurement device (Physical Acoustics Cooperation's
107 two-channel data acquisition and digital signal processing AE system, PCI-2), and the pumps
108 and injector were not used because no fracturing or confining fluids were required for the
109 experiment.

110 The true triaxial cell (Figure 2) consisted of a pressure vessel with a cubic skeleton,
111 six pistons to apply a compressive load to the $100 \times 100 \times 100$ mm cubic rock sample via a
112 stainless-steel plate (for better mechanical coupling between the piston loading faces and the
113 sample), and thermal insulators used in conjunction with heaters for the pressure vessel. The
114 pressure vessel had six cylindrical holes to allow the pistons to be inserted into the vessel, with
115 graphite packing lubricating the sliding portions. The edges of the sample were chamfered so
116 that the loading face of the sample had 90 mm sides that corresponded to the shape (i.e., $90 \times$
117 90 mm) of both the loading face of the piston and the stainless-steel plate. Each piston was
118 equipped with four cartridge heaters and a pipe through which a thermocouple could reach the
119 vicinity of the sample surface, and had an elastic wave guide bar on the opposite side of the
120 loading face. An AE sensor (R15 α , 150-kHz resonant frequency sensor, Physical Acoustics
121 Corporation) was attached to the face of the elastic wave guide bar, while two AE sensors were
122 used for one pair of the opposed horizontal pistons, where these pistons were used to apply the
123 far-field maximum horizontal stress described in section (2.2). The temperature of the bar was
124 maintained near ambient conditions using a cooling jacket through which water from a chiller
125 circulated.

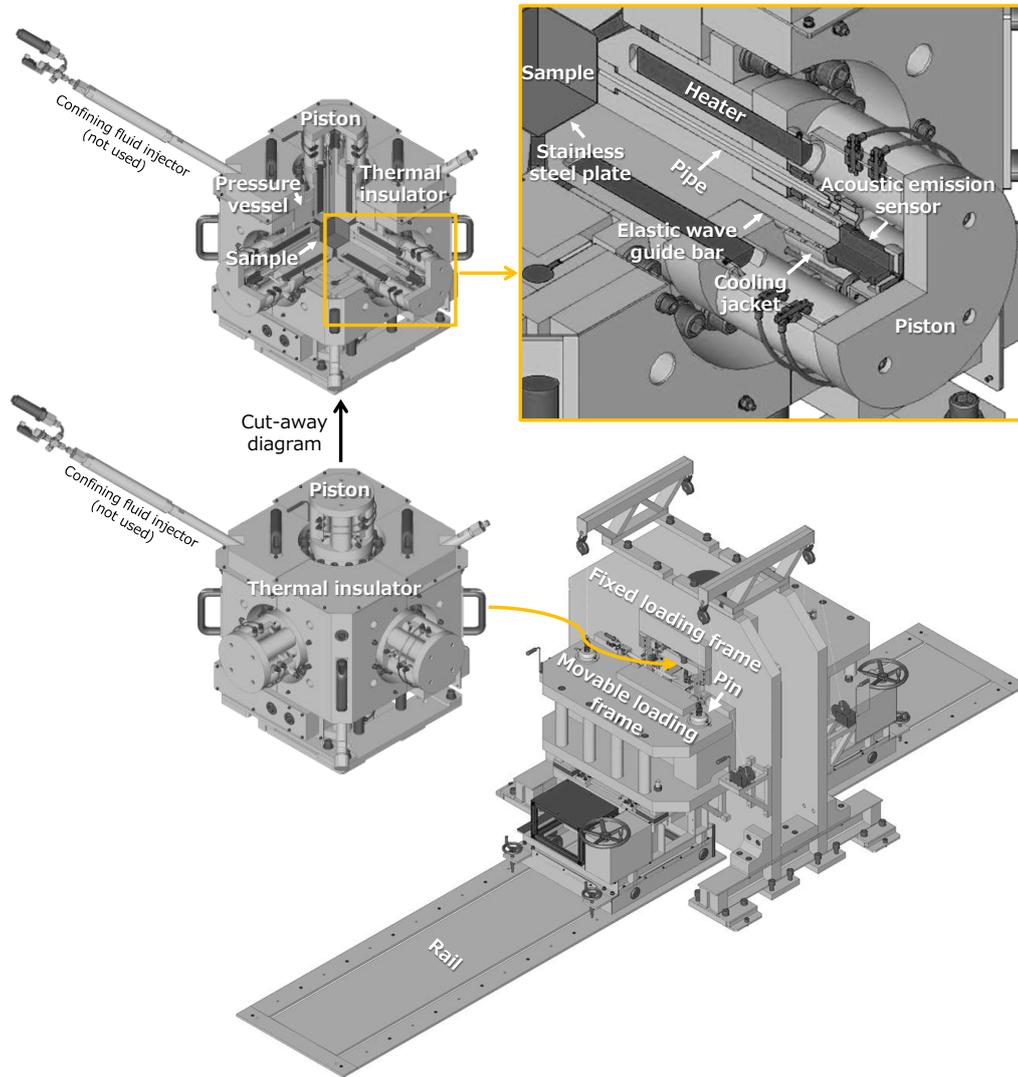


Figure 2. Three-dimensional design of the high-temperature true triaxial cell and triaxial loading frame, adapted from Watanabe et al. (2019).

The triaxial loading system (Figure 2) comprised a fixed loading frame and a movable loading frame. The fixed frame had two hydraulic rams that were placed vertically and horizontally. The movable loading frame contained a horizontal hydraulic ram. Each hydraulic ram had a capacity of 3 MN and was actuated using a manual oil pump. The true triaxial cell was placed on the loading platform of the fixed frame extremely carefully to prevent any loading eccentricity. It was then connected to the AE measurement system and the temperature

136 control and data logging system. When a compressive load was applied, the movable loading
137 frame engaged with the fixed loading frame via four cylindrical pins.

138 The triaxial loading system independently applied compressive loads in three
139 orthogonal directions, using a single hydraulic ram with a fixed loading platen on the opposite
140 side. The hydraulic ram, equipped with a load cell, pushed the piston of the true triaxial cell
141 via a spherical seated platen. The displacement of the piston was ascertained as a proxy for
142 sample deformation, using a linear variable differential transformer displacement transducer.
143 The displacement transducer was attached to a cantilever attached to the fixed loading platen
144 side, so that the measured displacement excluded any expansion between opposing platens
145 caused by loading frame deformation in response to large loads.

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147 **2.2. Experimental procedures and conditions**

148 The first procedure involved increasing the temperature of the pressure vessel and pistons
149 to a prescribed value (200, 350, or 450 °C) while a hydrostatic load (approximately 1 MPa)
150 was applied to a sample with a vertical wellbore. Once the temperature reached the prescribed
151 value, the sample was subjected to an initial stress state with a hydrostatic pressure of 53 MPa.
152 To induce wellbore failure, one of the horizontal stresses, referred to as the far-field minimum
153 horizontal stress (σ_h) in this study, was then decreased continuously and gradually to
154 approximately 5 MPa by returning oil to the manual oil pump, while the far-field maximum
155 horizontal stress (σ_H) and the vertical stress (σ_v) remained constant at the initial level. The
156 manual oil pump was equipped with a flow control valve so that the load could be continuously
157 decreased by returning oil to the pump using a constant valve opening.

158 For post-mortem analysis, the sample was impregnated with a thermosetting acrylic resin
159 under vacuum conditions and then heated to 50 °C under atmospheric pressure, so that the
160 cured resin preserved structural changes within the sample. The resin-impregnated sample was

161 cut horizontally across the middle region containing the wellbore, and a thin section of the
 162 inner region was prepared to assess wellbore failure via optical microscopy.

163 Based on Kirsch's equations (Gholami et al., 2014; Hiramatsu and Oka, 1968; Kirsch,
 164 1898), the state of stress around the wellbore in isotropic elastic homogenous rocks in a
 165 cylindrical coordinate system (r, θ, z) , for which r (from the center of the wellbore), θ (from
 166 the σ_H direction), and z correspond to the radial, tangential, and axial (vertical) directions of
 167 the wellbore, respectively, may be estimated as follows:

$$168 \quad \sigma_r = \frac{1}{2}(\sigma_H + \sigma_h) \left(1 - \frac{R^2}{r^2}\right) + \frac{1}{2}(\sigma_H - \sigma_h) \left(1 - \frac{4R^2}{r^2} + \frac{3R^4}{r^4}\right) \cos(2\theta) + P_w \frac{R^2}{r^2}, \quad (1)$$

$$169 \quad \sigma_\theta = \frac{1}{2}(\sigma_H + \sigma_h) \left(1 + \frac{R^2}{r^2}\right) - \frac{1}{2}(\sigma_H - \sigma_h) \left(1 + \frac{3R^4}{r^4}\right) \cos(2\theta) - P_w \frac{R^2}{r^2}, \text{ and} \quad (2)$$

$$170 \quad \sigma_z = \sigma_V - 2\nu(\sigma_H - \sigma_h) \cos(2\theta), \quad (3)$$

171 where σ_r , σ_θ , and σ_z are the radial, tangential, and axial stresses, respectively, R is the radius of
 172 the wellbore, P_w is the internal wellbore pressure, and ν is the Poisson's ratio, where P_w is
 173 assumed to be zero for the experiments under dry condition in this study. Notably, the
 174 applicability of Kirsch's equations on a wellbore in the granite used in the present study has
 175 been previously confirmed via a hydraulic fracturing experiment on the same granite (diameter:
 176 30 mm, length: 25 mm) with a wellbore (diameter: 1.5 mm, length: 10 mm) (Watanabe et al.,
 177 2017b).

178 Based on the Hoek–Brown failure criterion, a well-known empirical wellbore failure
 179 surface (Hoek and Brown, 1980; Pariseau, 2009) is represented by

$$180 \quad (\sigma_1 - \sigma_3)^2 - (A\sigma_3 + B^2) = 0, \quad (4)$$

$$181 \quad A = \frac{\sigma_c^2 - \sigma_t^2}{\sigma_t}, \text{ and} \quad (5)$$

$$182 \quad B = \sigma_c, \quad (6)$$

183 where σ_1 , σ_2 , and σ_3 are the maximum, intermediate, and minimum principal stresses,
 184 respectively, and σ_c and σ_t are the unconfined compressive and tensile strengths, respectively.

185 According to the literature, the σ_c of the granite used in this study is approximately 150 MPa
186 at 200 °C, 130 MPa at 350 °C, and 120 MPa at 450 °C (Sakai, 1987), while the σ_r and ν of the
187 same granite remain relatively constant across the range of experimental temperatures,
188 approximately 6 MPa and 0.2, respectively (Kinoshita et al., 1997). Using these values, the
189 Hoek–Brown failure criterion provides the following predictions regarding the initiation of
190 tensile and shear failures at $\theta = 0^\circ$ and 90° for the wellbore (Figure 3). Tensile failure, which
191 may produce tensile fractures, by increasing the σ_θ under tension at $\theta = 0^\circ$ starts at the same σ_h
192 level at all temperatures because σ_r is temperature-independent. In contrast, shear failure, which
193 may cause breakout, by increasing the σ_θ under compression at $\theta = 90^\circ$ is initiated at higher σ_h
194 values (i.e., smaller σ_θ) at higher temperatures, because of the smaller σ_c . Table 1 lists these
195 estimated and predicted values. According to our computations, we expect non-catastrophic
196 failure producing millimeter-scale breakout and/or tensile fractures, because the significantly
197 increased σ_θ in both the compression and tension is contained to within 10 mm of the wellbore
198 under the final state of far-field stress (Eq. 2).

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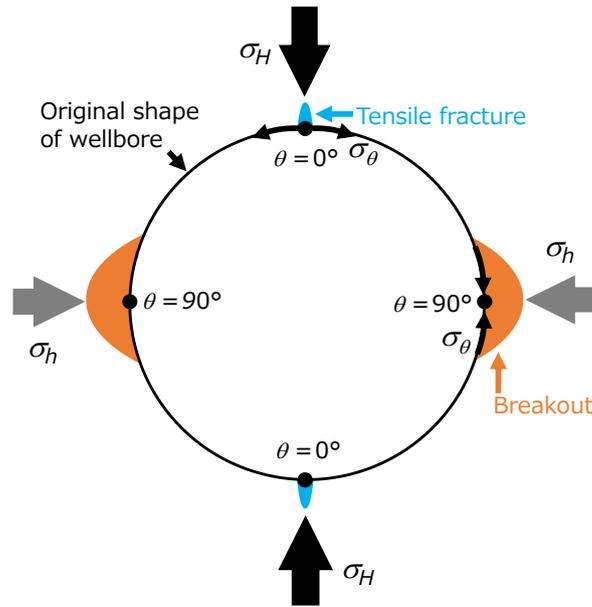


Figure 3. Conceptual illustration of breakout and tensile fractures in the wellbore wall, caused by shear and tensile failures, respectively.

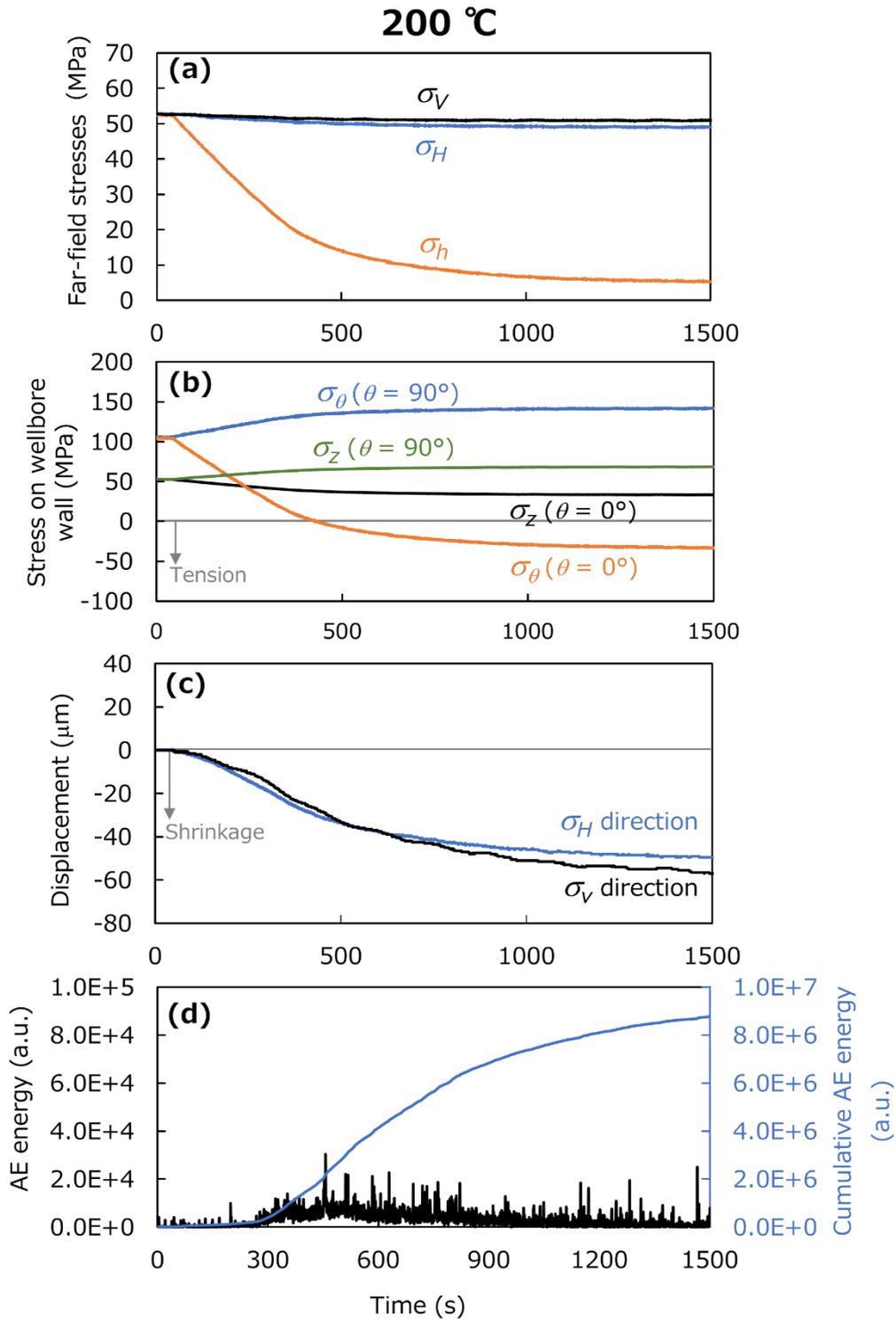
Table 1. Estimated unconfined compressive and tensile strengths (σ_c and σ_t) and predicted far-field minimum horizontal stress (σ_h) for shear and tensile failures at each experimental temperature.

Temperature (°C)	σ_c (MPa)	σ_t (MPa)	σ_h at failure (MPa)	
			Shear failure	Tensile failure
200	150	6	6	15
350	130	6	26	15
450	120	6	36	15

211 **3. Results and discussion**

212 **3.1. Temporal changes in stress, resultant behaviors of displacement and AE**

213 Figures 4–6 show temporal changes in the far-field stresses (σ_h , σ_H , and σ_V) and
214 corresponding changes in the estimated stresses on the wellbore (σ_θ and σ_z) using Eqs. 2 and
215 3, the displacement of the piston (a proxy of sample deformation) in the σ_H and σ_V directions,
216 and the AE energy with its cumulative value, for the three experiments at 200, 350, and 450 °C.
217 In the AE measurement device used in this study, AE energy is calculated as the time integral
218 of the absolute signal voltage and reported in arbitrary units (a.u.). AE energy shown herein is
219 the sum of the values obtained from the two AE sensors; we conducted the experiment at
220 450 °C, twice, and confirmed the absence of significant large AE energy in the second
221 experiment (i.e., after failure in the first experiment) as illustrated in Supplementary Figure S1.



222
 223 Figure 4. Temporal changes in the far-field stresses (a) and corresponding changes in the
 224 estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic
 225 emission (AE) energy and its cumulative value (d), in the experiment at 200 °C.

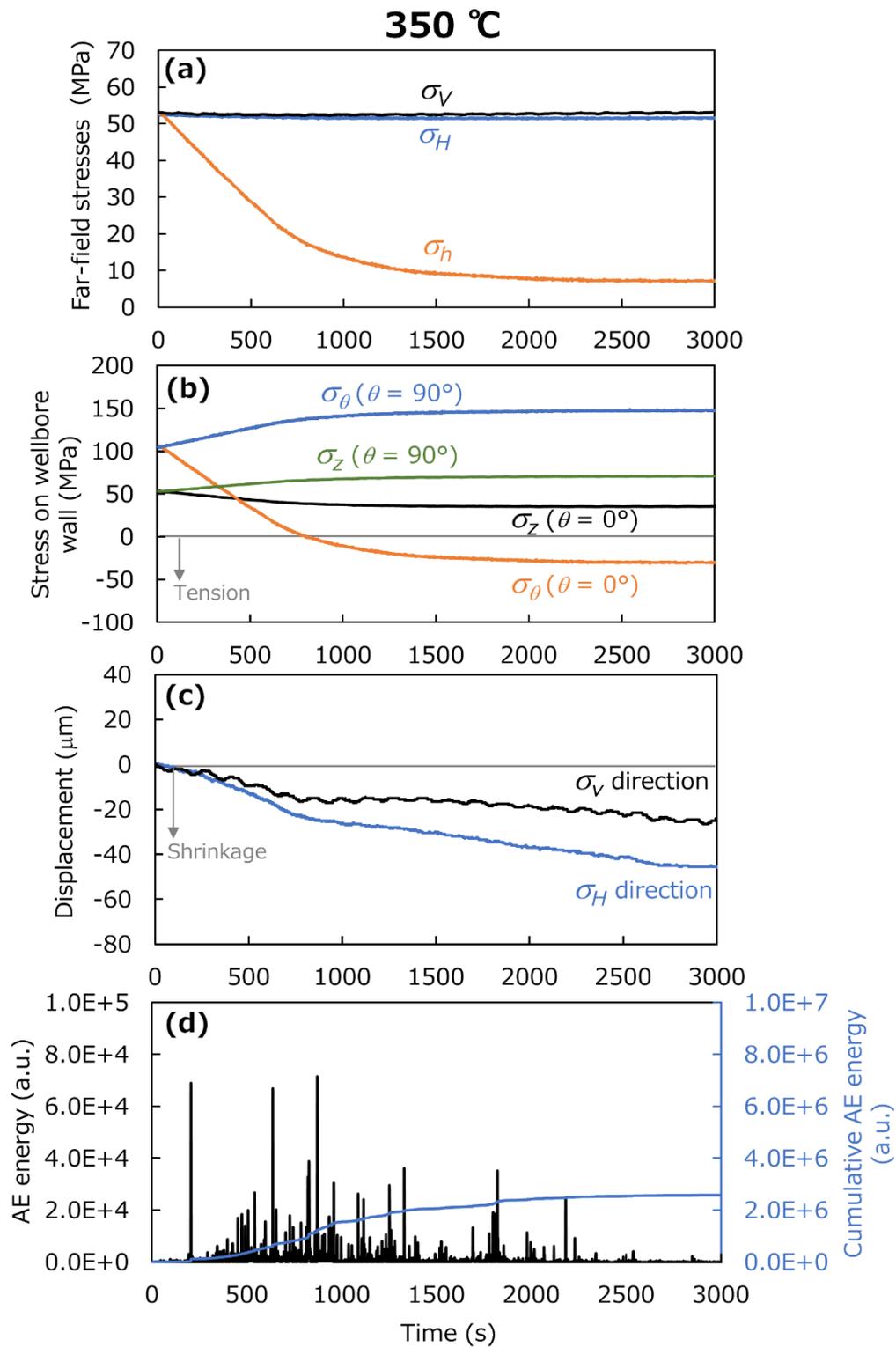


Figure 5. Temporal changes in the far-field stresses (a) and corresponding changes in the estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic emission (AE) energy and its cumulative value (d), in the experiment at 350 °C.

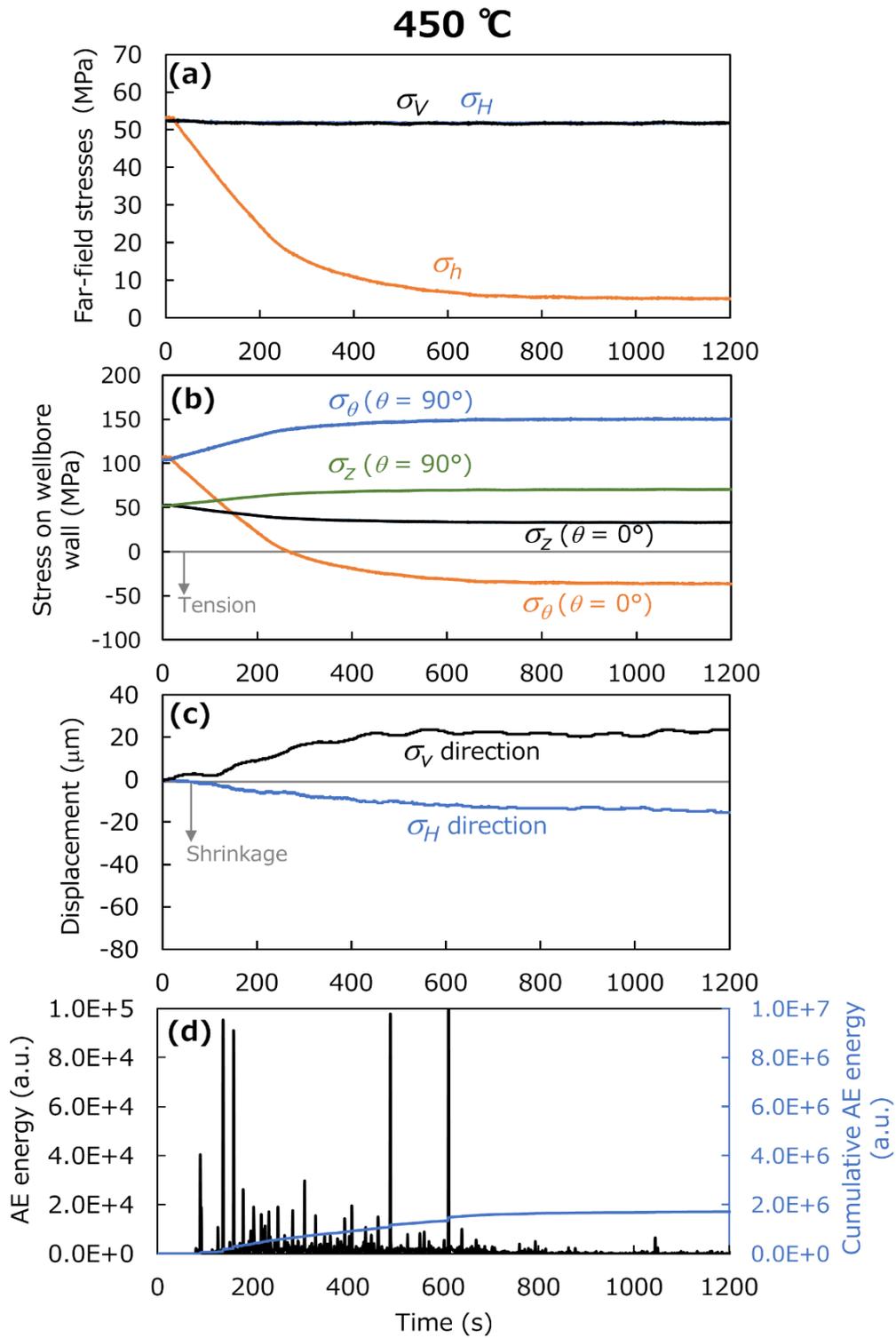


Figure 6. Temporal changes in the far-field stresses (a) and corresponding changes in the estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic emission (AE) energy and its cumulative value (d), in the experiment at 450 °C.

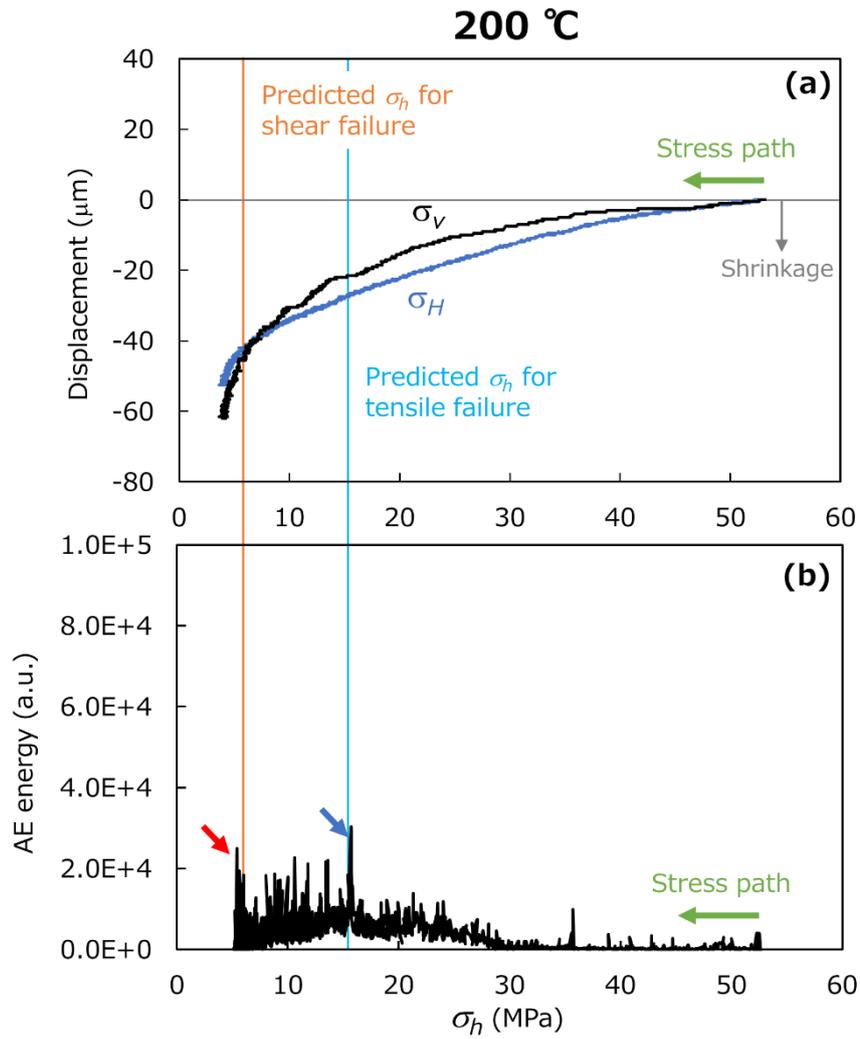
237 In all the experiments, we were able to decrease σ_h smoothly up to approximately 5
238 MPa by transferring the oil back to the manual oil pump, while σ_V and σ_H remained almost
239 constant. The rate of σ_h change was faster for higher σ_h (oil pressure) because the flow control
240 valve opening was kept constant (Figures 4a, 5a, and 6a). The rate of σ_h decline in the
241 experiment at 350 °C was slower than those in the other two experiments because it was
242 conducted at a lower room temperature in winter, which affected the oil viscosity. The final σ_h
243 level in each experiment exceeded the predicted criteria for shear failure (6–36 MPa at 450–
244 200 °C) and/or tensile failure (15 MPa at all temperatures) of the wellbore (Table 1). σ_θ and
245 σ_z also changed nonlinearly in response to nonlinear changes in σ_h (Figures 4b, 5b, and 6b). At
246 $\theta = 0^\circ$ (θ as defined in Figure 3), σ_θ decreased to –30 to –40 MPa (i.e., 30–40 MPa in tension),
247 whereas at $\theta = 90^\circ$, σ_θ increased to approximately 140–150 MPa.

248 Although the final σ_h in each experiment exceeded the predicted criteria for the shear
249 failure and/or tensile failure of the wellbore, the changes in displacement (Figures 4c, 5c, and
250 6c) and cumulative AE energy (Figures 4d, 5d, and 6d) in all experiments were continuous
251 before reaching a plateau, suggesting non-catastrophic shear and tensile failures, as illustrated
252 in Figure 3. The displacement data at 200 °C indicate that the sample shrank similarly in both
253 the σ_H and σ_V directions with decreasing σ_h . With increasing temperature, the sample shrank
254 less in the σ_V and/or σ_H directions, and it dilated in the σ_V direction at 450 °C. Additionally,
255 we observed that the AE energy behavior changed with increasing temperature. For low
256 temperatures, we recorded smaller AE energies frequently. In contrast, at high temperatures,
257 higher AE energies were recorded infrequently, resulting in a lower cumulative AE energy.
258 These changes in displacement and AE behavior with increasing temperature may have been
259 caused by the predicted shift in the predominant failure mode of the wellbore from tensile to
260 shear failure (Table 1).

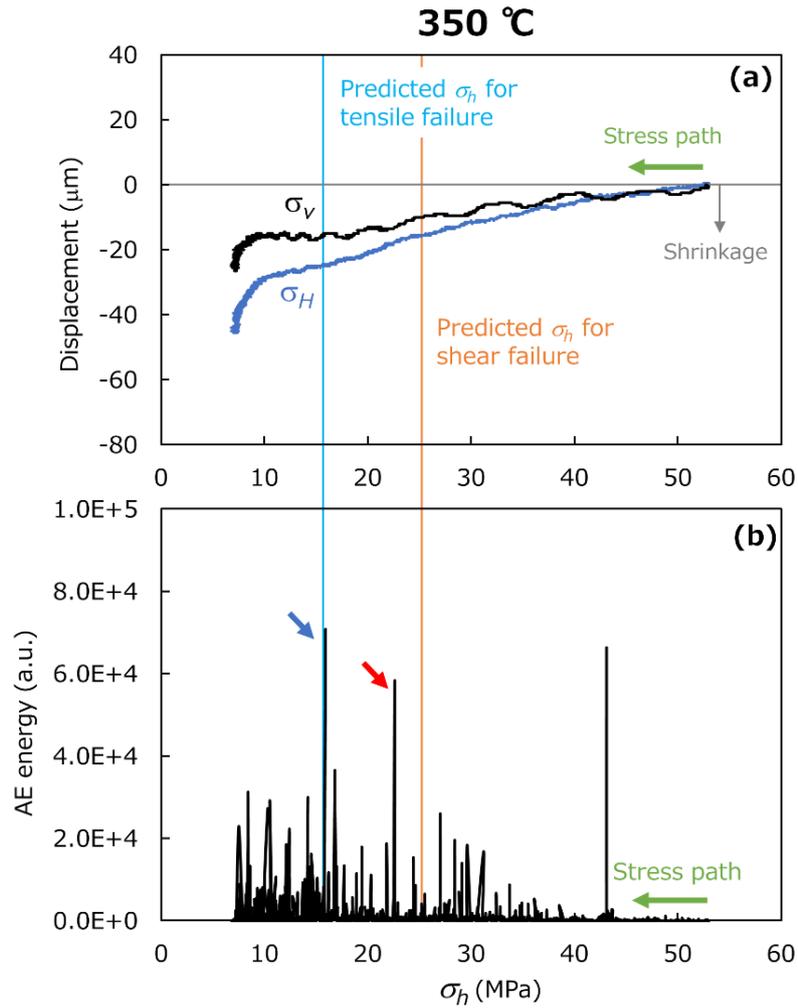
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3.2. Wellbore failure prediction

Figures 7–9 show the changes in the displacement of the piston and AE energy as a function of σ_h in each experiment, together with the predicted σ_h for the shear and tensile failure of the wellbore. Although the sample shrank continuously in both the σ_H and σ_V directions throughout the experiment at 200 °C, the shrinkage accelerated when the loading reached the predicted tensile failure threshold (the stress path is indicated by the green arrow in Figure 7a). Additionally, we recorded large AE energy events, as indicated by the blue arrow, when σ_h reached the predicted tensile failure threshold (Figure 7b). This suggests that the tensile failure of the wellbore started around the expected σ_h level and tensile fractures were generated in the σ_V – σ_H plane, as illustrated in Figure 3, meaning that the sample deformed more easily in the σ_V and σ_H directions. We recorded another set of large AE energy events, as indicated by the red arrow, around the predicted shear failure threshold. However, after the presumed tensile failure, the stress field was altered and the initial shear failure prediction may no longer have been valid. Indeed, the displacement behavior did not change, despite the second set of large AE energy events.



278
 279 Figure 7. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)
 280 as a function of σ_h in the experiment at 200 °C, together with the predicted σ_h for the shear and
 281 tensile failure of the wellbore wall.



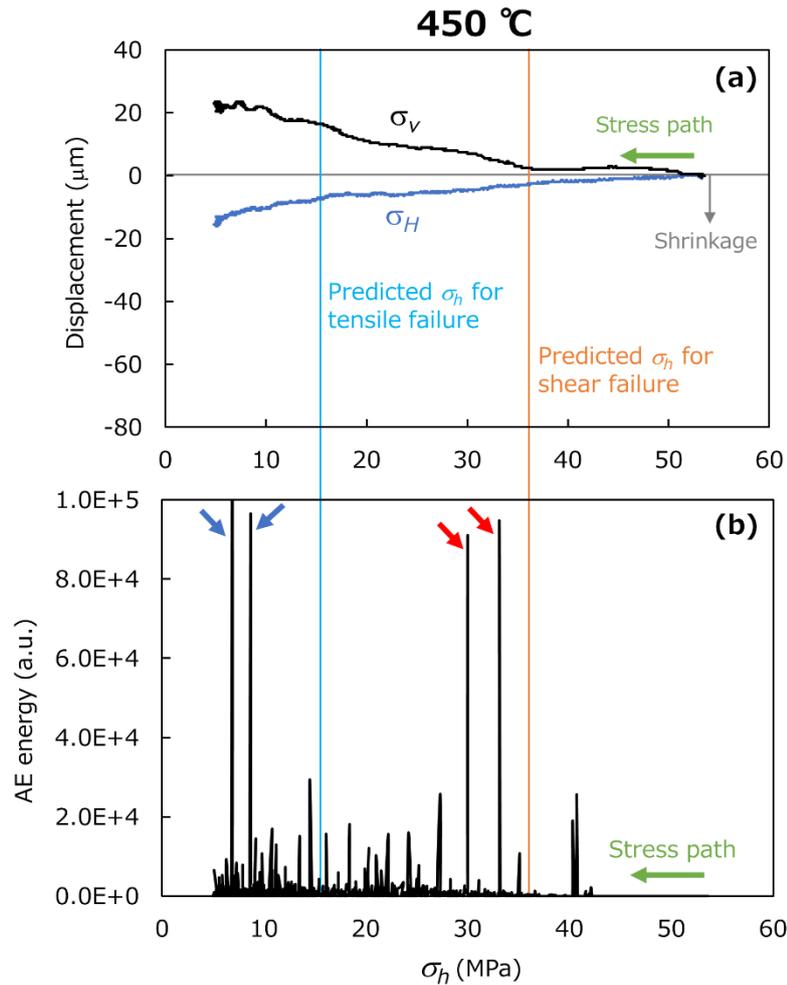
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284 Figure 8. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)

285 as a function of σ_h in the experiment at 350 °C, together with the predicted σ_h for the shear and

286 tensile failure of the wellbore wall.

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288
 289 Figure 9. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)
 290 as a function of σ_h in the experiment at 450 °C, together with the predicted σ_h for the shear and
 291 tensile failure of the wellbore wall.

292
 293 In contrast, at 450 °C, significant dilation in the σ_v direction started after the predicted
 294 shear failure loading (Figure 9a) and large energy AE events occurred, as indicated by the red
 295 arrows in Figure 9b. This suggests the occurrence of wellbore shear failure that caused dilation
 296 in the σ_v direction, as Zhao et al. (2015) observed shear failure and increased axial deformation
 297 in a wellbore in granite under hydrostatic stress. After the predicted tensile failure, remarkably
 298 large energy AE events occurred again, as indicated by the blue arrows in Figure 9b. However,
 299 the second set of large energy AE events occurred at a much smaller σ_h than the predicted value

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300 for tensile failure, and there was no clear indication of tensile failure in the deformation
301 behavior. This suggests that shear failure dominated the wellbore failure throughout the
302 experiment at 450 °C.

303 In the 350 °C experiment (Figure 8), after the sample reached the predicted shear failure,
304 the σ_h -dependence of the shrinkage in the σ_V direction decreased, flattening the displacement
305 curve at an σ_h of approximately 10–20 MPa. Based on the discussion of the 450 °C experiment,
306 this change in displacement behavior can be attributed to shear failure inducing dilation in the
307 σ_V direction. Subsequently, at $\sigma_h < 10$ MPa, the σ_h -dependence of the shrinkage in both the σ_V
308 and σ_H directions increased. Based on the discussion of the 200 °C experiment, this can be
309 attributed to tensile failure inducing shrinkage in both the σ_V and σ_H directions. Indeed, we
310 recorded large AE energies, as indicated by the red and blue arrows, around the predicted shear
311 and tensile failures. This suggests that, at 350 °C, both shear and tensile failure significantly
312 contributed to wellbore failure, owing to their relatively similar σ_h levels.

313 The above discussion suggests that the wellbore failure in the present experiments
314 occurred at σ_h levels that were largely consistent with the predicted values (Table 1). Therefore,
315 it is possible to predict wellbore failure in supercritical geothermal environments based on
316 existing rock failure criterion, such as the Hoek–Brown criterion used in this study.

317 318 **3.3. Possibility of using wellbore failure for in situ stress estimation**

319 Figure 10 shows optical microphotographs of the thin sections prepared from the
320 samples after the experiment (Figure 10a–c) and from an intact sample, for comparison, which
321 was not used in any experiment (Figure 10d). In these microphotographs, the profile of the
322 wellbore is traced in red, whereas the yellow line indicates a perfect circle

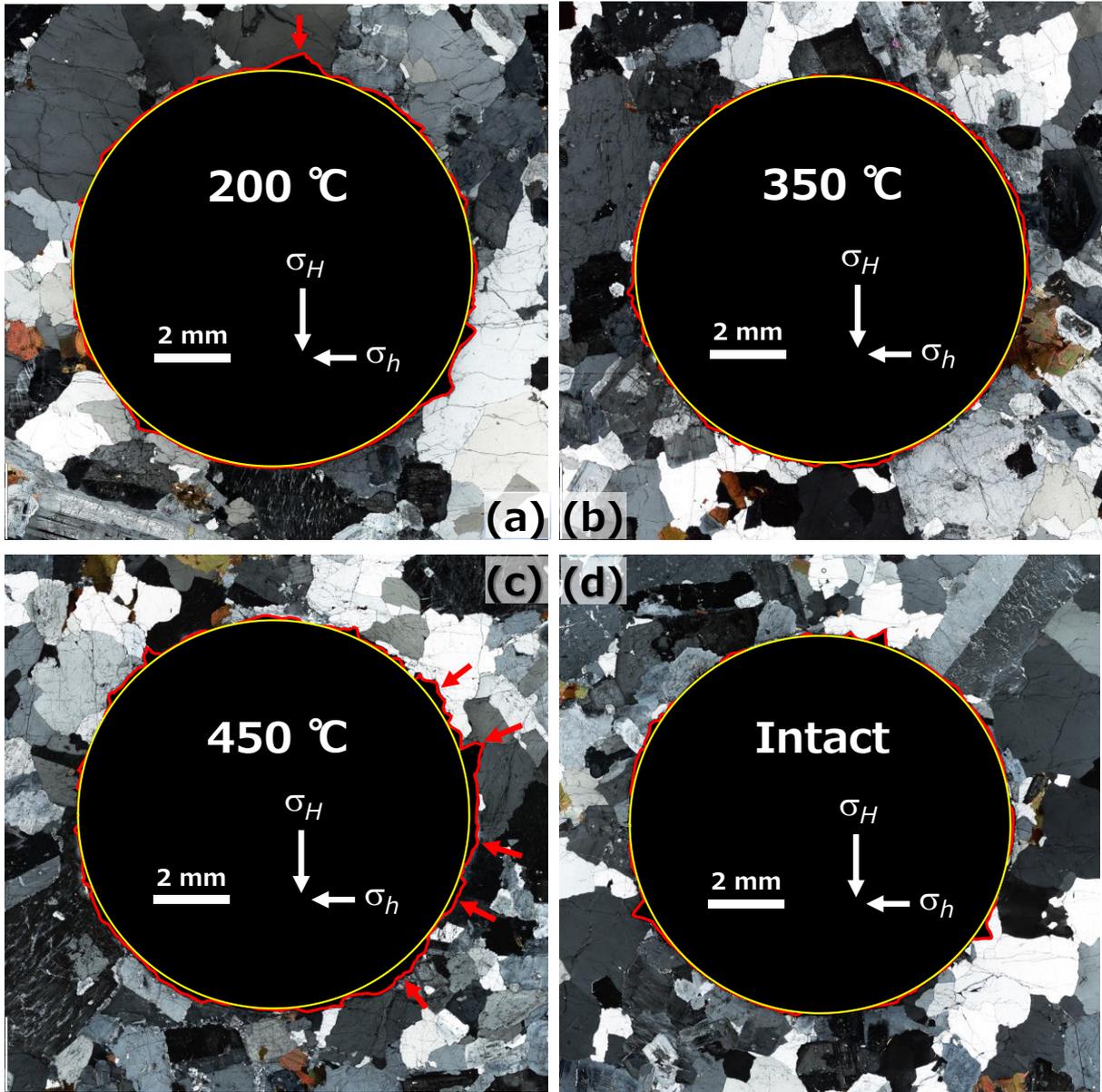


Figure 10. Optical microphotographs of thin sections prepared from the samples after the experiments at 200 (a), 350 (b), and 450 °C (c), and from an intact sample for comparison (d).

Even the intact sample profile does not completely match the perfect circle because of the relatively small defects with acute triangular shapes caused by boring (Figure 10d). However, the 200 °C profile deviates more from the perfect circle owing to the presence of larger defects in the wellbore, as indicated by the red arrow in Figure 10a. The largest defect appears to be a tensile fracture produced by the dominant failure at 200 °C (i.e., tensile failure

333 at $\theta = 0^\circ$ in Figure 3). The 450 °C profile also deviated significantly from the perfect circle
334 owing to the presence of much broader defects, as indicated by the red arrows (Figure 10c).
335 The defects appear to be breakouts produced by the dominant failure at 450 °C (i.e., shear
336 failure at $\theta = 90^\circ$ in Figure 3). The changes in the wellbore shape are consistent with our
337 speculation that the dominant failure mode shifts from tensile (200 °C) to shear (450 °C). In
338 contrast, the 350 °C profile did not significantly deviate from the perfect circle, although small
339 deviations appear around the entire wellbore wall (Figure 10b). It would therefore be
340 reasonable to assume no clear defects at specific locations if, as suggested in the previous
341 section (3.2), both shear and tensile failure significantly contributed to the wellbore failure at
342 350 °C.

343 The present experimental results therefore suggest that breakout and tensile fractures
344 may occur in wellbores in supercritical geothermal environments. However, at either 350 °C
345 or 450 °C, we did not observe the textbook breakout shape (Figure 3), such as the shape
346 observed in previous studies on granite at room temperature (Lee and Haimson, 1993; Song
347 and Haimson, 1997). This suggests that high temperatures suppress the propagation of shear
348 failure and its formation of clear breakout patterns. This may have been related to the
349 intermittent large energy AE events at 350 and 450 °C (Figures 5d and 6d), because such
350 intermittent large AE events were not observed in the previous study (Lee and Haimson, 1993).
351 This, in turn, suggests that in situ stress estimation from the shape of the breakout would be
352 difficult in supercritical geothermal environments because of the suppression of failure
353 propagation at high temperatures. However, if shear failure propagation is suppressed and
354 wellbore failure does not propagate further into the formation at high temperatures, this would
355 greatly improve wellbore stability and suggests the possibility of underbalanced drilling, which
356 would substantially reduce the risk of circulation loss during drilling. If our conclusions are
357 correct, they would have tremendous implications for drilling practices in supercritical

358 geothermal development. The suppressed failure propagation from the wellbore wall into the
359 rock may have been caused by a rapid relaxation of stress concentration around the wellbore
360 since several points on the circumference of the wellbore wall simultaneously reached failure,
361 or became close, due to the relatively small unconfined compressive strength. However, this
362 was not confirmed, hence warrants further investigation. Thus, extensive future studies are
363 encouraged to examine the suppression of failure propagation at high temperatures and to
364 clarify the mechanisms involved. Ideally, such studies should address wellbore failure in high-
365 temperature granite along the temperature and stress paths that may occur during drilling, with
366 considerations on influences of fluid pressure and thermal stress as well as *in situ* preexisting
367 weakness planes.

369 4. Conclusions

370 For the successful development of enhanced geothermal systems in supercritical
371 geothermal environments through hydraulic fracturing, it is essential to establish wellbore
372 stability and understand *in situ* stresses. The results of the first experiments on wellbore failure
373 in 200–450 °C granite under true triaxial stress suggest that non-catastrophic tensile and shear
374 failure occurs in the wellbore at such temperatures, and their initiation is predictable based on
375 existing brittle failure criterion for rocks. The results also suggest that the resultant shape of
376 the wellbore after failure, particularly after breakout by shear failure, is not ideal for *in situ*
377 stress estimation, because of the suppressed propagation of shear failure at high temperatures.
378 In other words, there is a possibility that the high temperatures in supercritical geothermal
379 environments contribute to wellbore stability. Therefore, extensive future studies are
380 encouraged to examine the suppression of failure propagation at high temperatures and to
381 clarify the underlying mechanisms.

382

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27 394 system.

31 396 **Data availability**

33
34 397 The data supporting the findings of this study are available from the corresponding
35
36 398 author upon reasonable request.

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1 **Wellbore stability in high-temperature granite under true triaxial stress**

2

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22 **Abbreviations:** AE, acoustic emissions

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26 **Abstract**

27 Supercritical/superhot geothermal success depends on successful drilling. However, wellbore
28 stability in supercritical environments has not been investigated because imposing sufficient
29 stress on high-temperature granite with true triaxial loading is difficult. We conducted wellbore
30 failure experiments on 200–450 °C granite under true triaxial stress, including deformation and
31 acoustic emission measurements, and post-experiment thin section observations. Wellbore
32 failure initiated at stress states consistent with existing brittle failure criterion at the studied
33 temperatures. Using breakout geometry for *in situ* stress estimation may be difficult as shear
34 failure propagation is suppressed at high temperatures; however, boreholes may be inherently
35 stable in high-temperature environments.

36 **Keywords:** wellbore failure, granite, high temperature, supercritical geothermal environment,
37 superhot geothermal environment, true triaxial stress

38

39

40 **1. Introduction**

41 Recent attention in geothermal power generation has focused on methods for harvest
42 geothermal energy from unconventional supercritical (or superhot) geothermal environments
43 that exceed the critical temperature of water (374 °C for pure water and 406 °C for seawater)
44 at drillable depths of approximately 2–4 km, as found in Iceland (Friðleifsson et al., 2014;
45 Friðleifsson and Elders, 2017; Steingrímsson et al., 1990), Italy (Baron and Ungemach, 1981;
46 Batini et al., 1983; Ruggieri and Gianelli, 1995), Japan (Kato et al., 1998), Mexico (Espinosa-
47 Paredes and Garcia-Gutierrez, 2003), and the United States (Fournier, 1991; Garcia et al.,
48 2016). Supercritical geothermal environments can provide superheated steam or supercritical
49 water with an extremely high specific enthalpy exceeding 2 MJ/kg (Elders et al., 2014;
50 Friðleifsson and Elders, 2005; Friðleifsson et al. 2007; Smith and Shaw, 1975, 1979; Tester et

51 al., 2006). Therefore, harvesting supercritical geothermal energy is expected to increase the
52 productivity and sustainability of geothermal energy generation. However, in supercritical
53 geothermal environments on the continental crust, the formation of permeable fracture
54 networks (Tsuchiya et al., 2016; Watanabe et al., 2017a; Weis et al., 2012) can potentially be
55 suppressed by the increased efficiency of mineral crystal plasticity (Parisio et al., 2019, 2020;
56 Tullis and Yund, 1977; Violay et al., 2017), the retrograde solubility of silica minerals
57 (Fournier, 1991; Saishu et al., 2014; Tsuchiya and Hirano, 2007; Watanabe et al., 2021), and
58 the enhanced rates of fracture healing/sealing caused by water–rock reactions (Moore et al.,
59 1994; Morrow et al., 2001; Watanabe et al., 2020). Therefore, it is necessary to develop
60 enhanced geothermal system technologies to artificially create or recreate permeable fracture
61 networks in these supercritical geothermal environments.

62 Recent studies have suggested that the injection of water at or near its critical temperature
63 into granite is likely to create a dense network of microfractures suitable for geothermal energy
64 extraction, a so-called cloud-fracture network, through continuous infiltration and the
65 stimulation of pre-existing microfractures by the low-viscosity water (Goto et al. 2021;
66 Watanabe et al., 2017b, 2019). Additionally, it has been suggested that the pore pressure
67 required to initiate this flow-induced microfracturing may be well predicted by the Griffith
68 failure criterion, which assumes the initiation of brittle failure from pre-existing fractures and
69 is a function of the maximum and minimum principal stresses and tensile strength of rocks
70 (Cox, 2010; Griffith, 1924; Jaeger and Cook, 1979; Secor, 1965). However, the characteristics
71 of wellbore failure (Ewy and Cook, 1990; Morita and Nagano, 2016), and the possibility of
72 utilizing these failures (i.e., induced fractures and breakout) for in situ stress estimation (Brudy
73 and Zoback, 1999; Kato et al., 1998; Zoback et al., 1985) to achieve safe and effective
74 fracturing at the minimum required injection pressure, remain unclear for supercritical
75 geothermal environments. This is because wellbore failure experiments under true triaxial

76 stress (Lee and Haimson, 1993; Song and Haimson, 1997) have never been conducted on high-
77 temperature granite. To the best of our knowledge, wellbore failure experiments under
78 hydrostatic stress have only been conducted on granite at temperatures up to 600 °C (Zhao et
79 al., 2015).

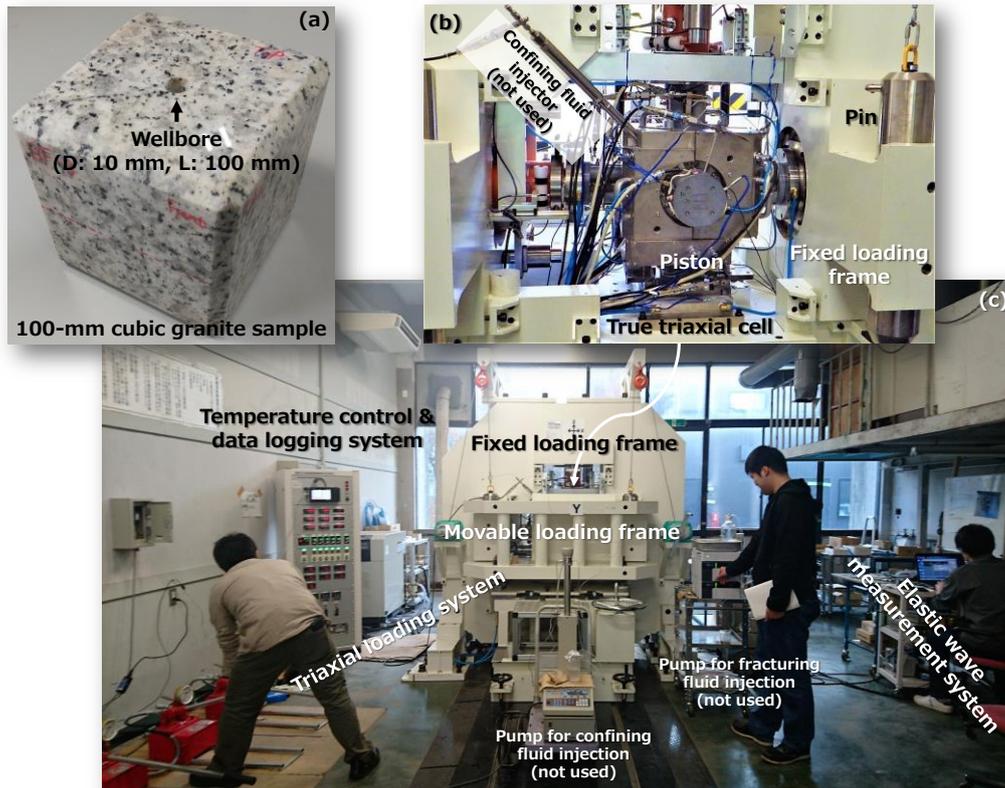
80 The aim of this study is to experimentally simulate wellbore failure in granite under true
81 triaxial stress in supercritical geothermal conditions. We first introduce a novel experimental
82 method for assessing high-temperature wellbore failure in relatively large (10-cm cubic) rock
83 samples under true triaxial stress. Subsequently, we present the results of experiments
84 conducted on granite at 200–450 °C, which show non-catastrophic wellbore failure. Finally,
85 we suggest the predictability of the initiation of wellbore failure using an existing failure
86 criterion for rocks and discuss the possibility of estimating in situ stress from wellbore failure.

87

88 **2. Experimental methods**

89 **2.1. Granite and experimental system**

90 Cubes (100 × 100 × 100 mm) of Inada granite from Ibaraki prefecture, Japan (Figure 1),
91 with a single wellbore (diameter: 10 mm, length: 100 mm) at the center, were prepared for the
92 high-temperature wellbore failure experiment under true triaxial stress. Inada granite, of which
93 the porosity at room temperature and atmospheric pressure is 0.5-0.8%, has been used in
94 previous studies on supercritical hydraulic fracturing (Goto et al., 2021; Watanabe et al., 2017b,
95 2019).



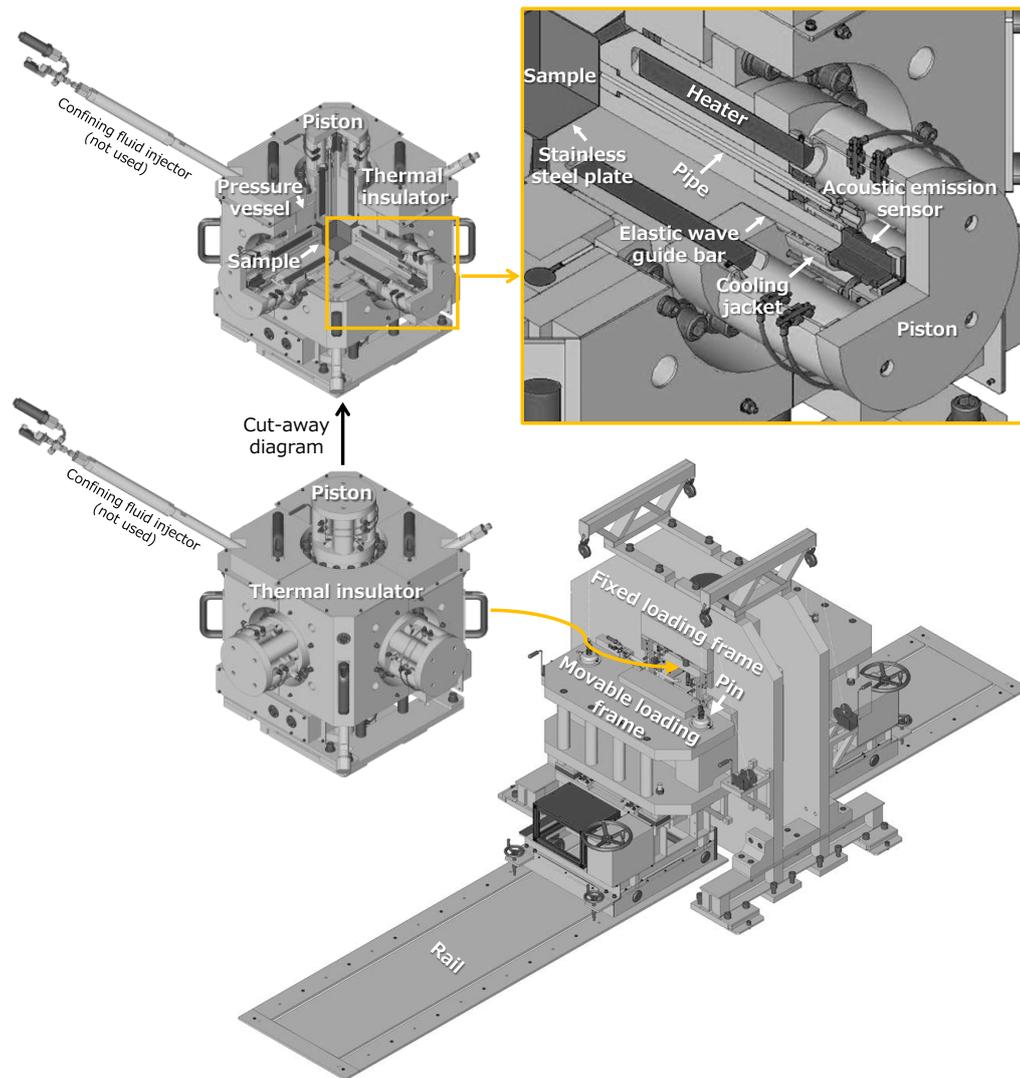
96

97 Figure 1. Granite sample (a), high-temperature true triaxial cell (b), and experimental
 98 (c) in the wellbore failure experiment, adapted from Watanabe et al. (2019).

99

100 We used the experimental system (Figure 1) originally developed for supercritical
 101 hydraulic fracturing by Watanabe et al. (2019) and applied in Goto et al. (2021). The system
 102 consisted of a true triaxial cell, a triaxial loading system, a pump for injecting fracturing fluid
 103 into the sample, a pump for injecting confining fluid (high-viscosity plastic melt in this system)
 104 along the sample edges via an injector, an elastic wave measurement system, and a temperature
 105 control and data logging system. In the present experiment, the elastic wave measurement
 106 system was an acoustic emission (AE) measurement device (Physical Acoustics Cooperation's
 107 two-channel data acquisition and digital signal processing AE system, PCI-2), and the pumps
 108 and injector were not used because no fracturing or confining fluids were required for the
 109 experiment.

110 The true triaxial cell (Figure 2) consisted of a pressure vessel with a cubic skeleton,
111 six pistons to apply a compressive load to the $100 \times 100 \times 100$ mm cubic rock sample via a
112 stainless-steel plate (for better mechanical coupling between the piston loading faces and the
113 sample), and thermal insulators used in conjunction with heaters for the pressure vessel. The
114 pressure vessel had six cylindrical holes to allow the pistons to be inserted into the vessel, with
115 graphite packing lubricating the sliding portions. The edges of the sample were chamfered so
116 that the loading face of the sample had 90 mm sides that corresponded to the shape (i.e., $90 \times$
117 90 mm) of both the loading face of the piston and the stainless-steel plate. Each piston was
118 equipped with four cartridge heaters and a pipe through which a thermocouple could reach the
119 vicinity of the sample surface, and had an elastic wave guide bar on the opposite side of the
120 loading face. An AE sensor (R15 α , 150-kHz resonant frequency sensor, Physical Acoustics
121 Corporation) was attached to the face of the elastic wave guide bar, while two AE sensors were
122 used for one pair of the opposed horizontal pistons, where these pistons were used to apply the
123 far-field maximum horizontal stress described in section (2.2). The temperature of the bar was
124 maintained near ambient conditions using a cooling jacket through which water from a chiller
125 circulated.



126

127 Figure 2. Three-dimensional design of the high-temperature true triaxial cell and triaxial
 128 loading frame, adapted from Watanabe et al. (2019).

129

130 The triaxial loading system (Figure 2) comprised a fixed loading frame and a movable
 131 loading frame. The fixed frame had two hydraulic rams that were placed vertically and
 132 horizontally. The movable loading frame contained a horizontal hydraulic ram. Each hydraulic
 133 ram had a capacity of 3 MN and was actuated using a manual oil pump. The true triaxial cell
 134 was placed on the loading platform of the fixed frame extremely carefully to prevent any
 135 loading eccentricity. It was then connected to the AE measurement system and the temperature

136 control and data logging system. When a compressive load was applied, the movable loading
137 frame engaged with the fixed loading frame via four cylindrical pins.

138 The triaxial loading system independently applied compressive loads in three
139 orthogonal directions, using a single hydraulic ram with a fixed loading platen on the opposite
140 side. The hydraulic ram, equipped with a load cell, pushed the piston of the true triaxial cell
141 via a spherical seated platen. The displacement of the piston was ascertained as a proxy for
142 sample deformation, using a linear variable differential transformer displacement transducer.
143 The displacement transducer was attached to a cantilever attached to the fixed loading platen
144 side, so that the measured displacement excluded any expansion between opposing platens
145 caused by loading frame deformation in response to large loads.

146

147 **2.2. Experimental procedures and conditions**

148 The first procedure involved increasing the temperature of the pressure vessel and pistons
149 to a prescribed value (200, 350, or 450 °C) while a hydrostatic load (approximately 1 MPa)
150 was applied to a sample with a vertical wellbore. Once the temperature reached the prescribed
151 value, the sample was subjected to an initial stress state with a hydrostatic pressure of 53 MPa.
152 To induce wellbore failure, one of the horizontal stresses, referred to as the far-field minimum
153 horizontal stress (σ_h) in this study, was then decreased continuously and gradually to
154 approximately 5 MPa by returning oil to the manual oil pump, while the far-field maximum
155 horizontal stress (σ_H) and the vertical stress (σ_V) remained constant at the initial level. The
156 manual oil pump was equipped with a flow control valve so that the load could be continuously
157 decreased by returning oil to the pump using a constant valve opening.

158 For post-mortem analysis, the sample was impregnated with a thermosetting acrylic resin
159 under vacuum conditions and then heated to 50 °C under atmospheric pressure, so that the
160 cured resin preserved structural changes within the sample. The resin-impregnated sample was

161 cut horizontally across the middle region containing the wellbore, and a thin section of the
 162 inner region was prepared to assess wellbore failure via optical microscopy.

163 Based on Kirsch's equations (Gholami et al., 2014; Hiramatsu and Oka, 1968; Kirsch,
 164 1898), the state of stress around the wellbore in isotropic elastic homogenous rocks in a
 165 cylindrical coordinate system (r, θ, z) , for which r (from the center of the wellbore), θ (from
 166 the σ_H direction), and z correspond to the radial, tangential, and axial (vertical) directions of
 167 the wellbore, respectively, may be estimated as follows:

$$168 \quad \sigma_r = \frac{1}{2}(\sigma_H + \sigma_h) \left(1 - \frac{R^2}{r^2}\right) + \frac{1}{2}(\sigma_H - \sigma_h) \left(1 - \frac{4R^2}{r^2} + \frac{3R^4}{r^4}\right) \cos(2\theta) + P_w \frac{R^2}{r^2}, \quad (1)$$

$$169 \quad \sigma_\theta = \frac{1}{2}(\sigma_H + \sigma_h) \left(1 + \frac{R^2}{r^2}\right) - \frac{1}{2}(\sigma_H - \sigma_h) \left(1 + \frac{3R^4}{r^4}\right) \cos(2\theta) - P_w \frac{R^2}{r^2}, \text{ and} \quad (2)$$

$$170 \quad \sigma_z = \sigma_V - 2\nu(\sigma_H - \sigma_h) \cos(2\theta), \quad (3)$$

171 where σ_r , σ_θ , and σ_z are the radial, tangential, and axial stresses, respectively, R is the radius of
 172 the wellbore, P_w is the internal wellbore pressure, and ν is the Poisson's ratio, where P_w is
 173 assumed to be zero for the experiments under dry condition in this study. Notably, the
 174 applicability of Kirsch's equations on a wellbore in the granite used in the present study has
 175 been previously confirmed via a hydraulic fracturing experiment on the same granite (diameter:
 176 30 mm, length: 25 mm) with a wellbore (diameter: 1.5 mm, length: 10 mm) (Watanabe et al.,
 177 2017b).

178 Based on the Hoek–Brown failure criterion, a well-known empirical wellbore failure
 179 surface (Hoek and Brown, 1980; Pariseau, 2009) is represented by

$$180 \quad (\sigma_1 - \sigma_3)^2 - (A\sigma_3 + B^2) = 0, \quad (4)$$

$$181 \quad A = \frac{\sigma_c^2 - \sigma_t^2}{\sigma_t}, \text{ and} \quad (5)$$

$$182 \quad B = \sigma_c, \quad (6)$$

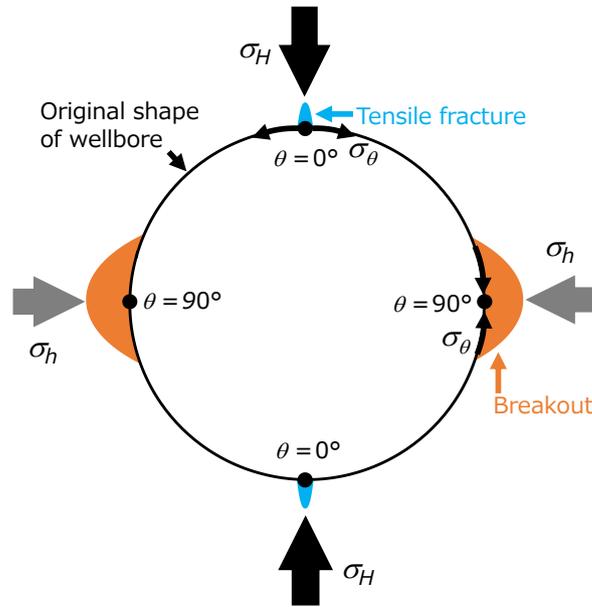
183 where σ_1 , σ_2 , and σ_3 are the maximum, intermediate, and minimum principal stresses,
 184 respectively, and σ_c and σ_t are the unconfined compressive and tensile strengths, respectively.

185 According to the literature, the σ_c of the granite used in this study is approximately 150 MPa
186 at 200 °C, 130 MPa at 350 °C, and 120 MPa at 450 °C (Sakai, 1987), while the σ_t and ν of the
187 same granite remain relatively constant across the range of experimental temperatures,
188 approximately 6 MPa and 0.2, respectively (Kinoshita et al., 1997). Using these values, the
189 Hoek–Brown failure criterion provides the following predictions regarding the initiation of
190 tensile and shear failures at $\theta = 0^\circ$ and 90° for the wellbore (Figure 3). Tensile failure, which
191 may produce tensile fractures, by increasing the σ_θ under tension at $\theta = 0^\circ$ starts at the same σ_h
192 level at all temperatures because σ_t is temperature-independent. In contrast, shear failure, which
193 may cause breakout, by increasing the σ_θ under compression at $\theta = 90^\circ$ is initiated at higher σ_h
194 values (i.e., smaller σ_θ) at higher temperatures, because of the smaller σ_c . Table 1 lists these
195 estimated and predicted values. According to our computations, we expect non-catastrophic
196 failure producing millimeter-scale breakout and/or tensile fractures, because the significantly
197 increased σ_θ in both the compression and tension is contained to within 10 mm of the wellbore
198 under the final state of far-field stress (Eq. 2).

199

200

201



202

203 Figure 3. Conceptual illustration of breakout and tensile fractures in the wellbore wall, caused
 204 by shear and tensile failures, respectively.

205

206 Table 1. Estimated unconfined compressive and tensile strengths (σ_c and σ_t) and predicted far-
 207 field minimum horizontal stress (σ_h) for shear and tensile failures at each experimental
 208 temperature.

Temperature (°C)	σ_c (MPa)	σ_t (MPa)	σ_h at failure (MPa)	
			Shear failure	Tensile failure
200	150	6	6	15
350	130	6	26	15
450	120	6	36	15

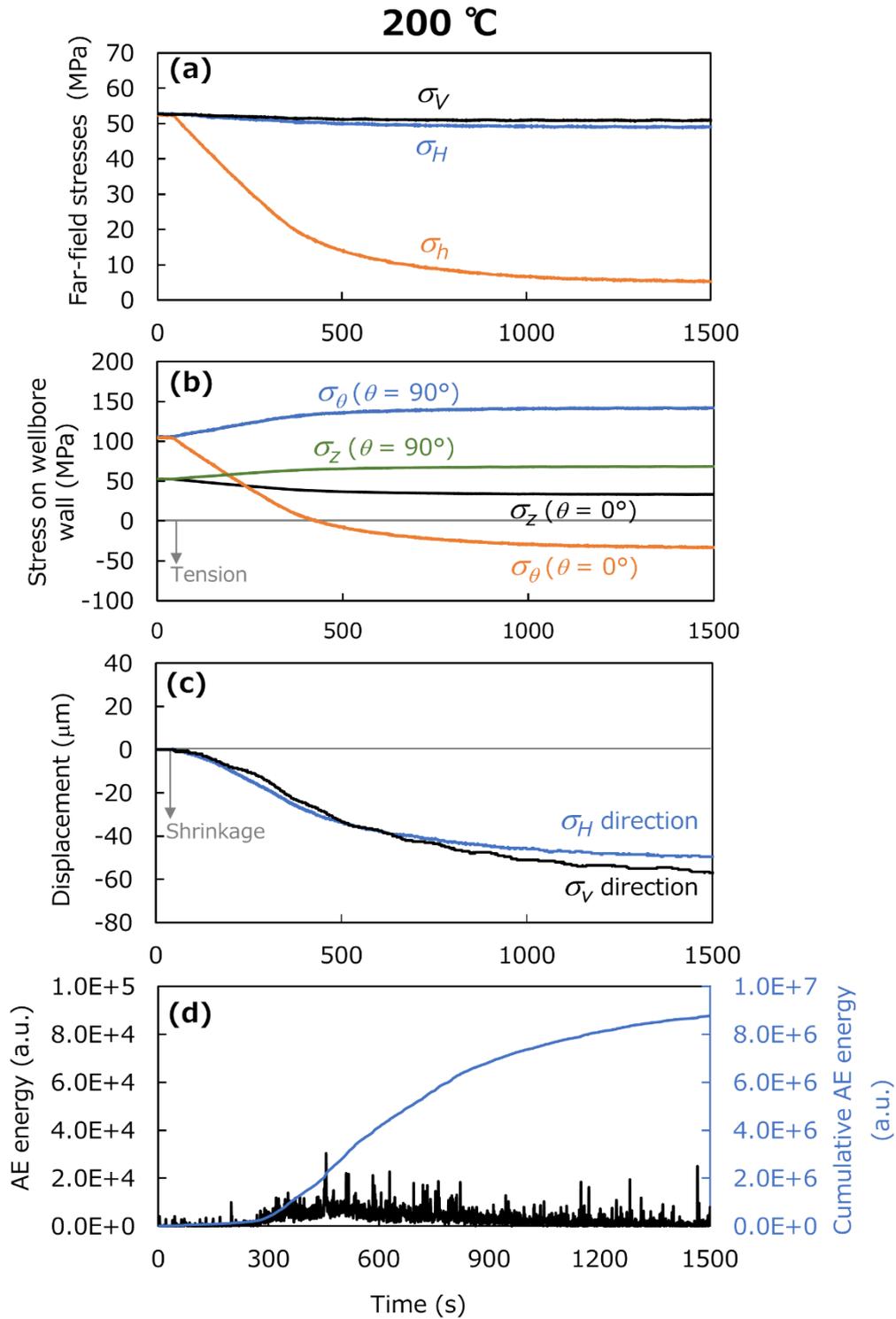
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210

211 **3. Results and discussion**

212 **3.1. Temporal changes in stress, resultant behaviors of displacement and AE**

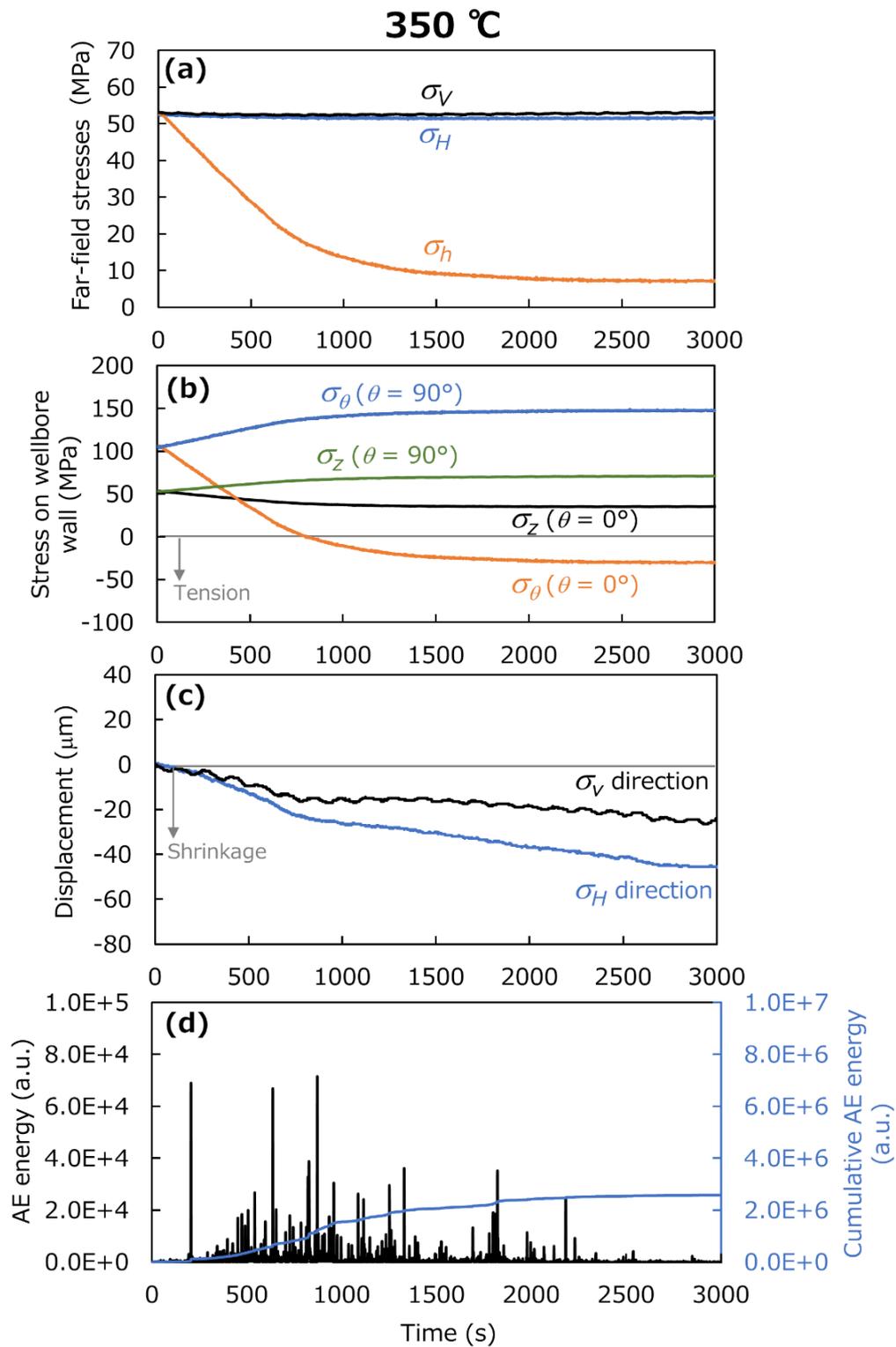
213 Figures 4–6 show temporal changes in the far-field stresses (σ_h , σ_H , and σ_V) and
214 corresponding changes in the estimated stresses on the wellbore (σ_θ and σ_z) using Eqs. 2 and
215 3, the displacement of the piston (a proxy of sample deformation) in the σ_H and σ_V directions,
216 and the AE energy with its cumulative value, for the three experiments at 200, 350, and 450 °C.
217 In the AE measurement device used in this study, AE energy is calculated as the time integral
218 of the absolute signal voltage and reported in arbitrary units (a.u.). AE energy shown herein is
219 the sum of the values obtained from the two AE sensors; we conducted the experiment at
220 450 °C, twice, and confirmed the absence of significant large AE energy in the second
221 experiment (i.e., after failure in the first experiment) as illustrated in Supplementary Figure S1.



222

223 Figure 4. Temporal changes in the far-field stresses (a) and corresponding changes in the
 224 estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic
 225 emission (AE) energy and its cumulative value (d), in the experiment at 200 °C.

226



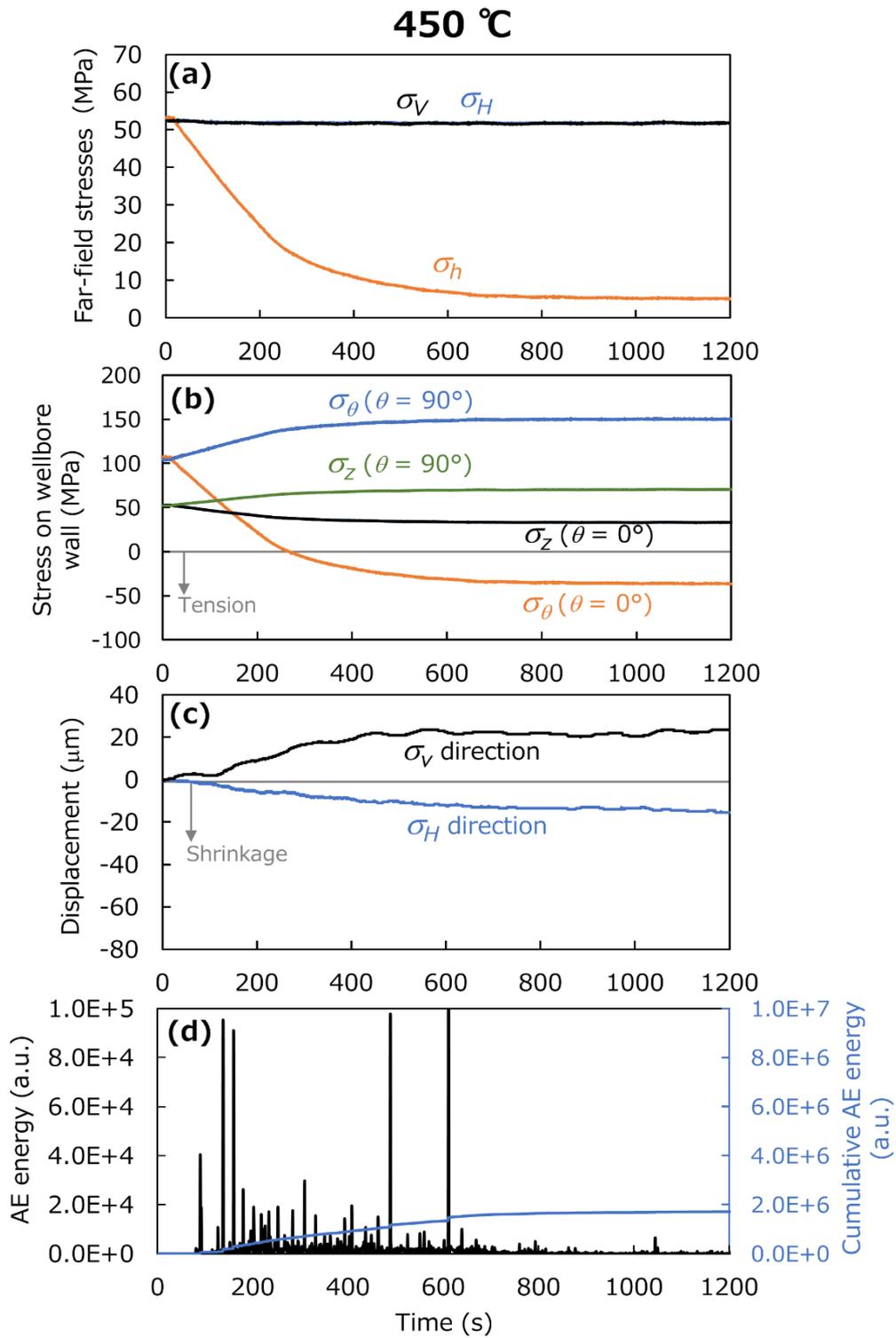
227

228 Figure 5. Temporal changes in the far-field stresses (a) and corresponding changes in the

229 estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic

230 emission (AE) energy and its cumulative value (d), in the experiment at 350 °C.

231



232

233 Figure 6. Temporal changes in the far-field stresses (a) and corresponding changes in the

234 estimated stresses on the wellbore wall (b), displacement of the piston (c), and acoustic

235 emission (AE) energy and its cumulative value (d), in the experiment at 450 °C.

236

237 In all the experiments, we were able to decrease σ_h smoothly up to approximately 5
238 MPa by transferring the oil back to the manual oil pump, while σ_V and σ_H remained almost
239 constant. The rate of σ_h change was faster for higher σ_h (oil pressure) because the flow control
240 valve opening was kept constant (Figures 4a, 5a, and 6a). The rate of σ_h decline in the
241 experiment at 350 °C was slower than those in the other two experiments because it was
242 conducted at a lower room temperature in winter, which affected the oil viscosity. The final σ_h
243 level in each experiment exceeded the predicted criteria for shear failure (6–36 MPa at 450–
244 200 °C) and/or tensile failure (15 MPa at all temperatures) of the wellbore (Table 1). σ_θ and
245 σ_z also changed nonlinearly in response to nonlinear changes in σ_h (Figures 4b, 5b, and 6b). At
246 $\theta = 0^\circ$ (θ as defined in Figure 3), σ_θ decreased to –30 to –40 MPa (i.e., 30–40 MPa in tension),
247 whereas at $\theta = 90^\circ$, σ_θ increased to approximately 140–150 MPa.

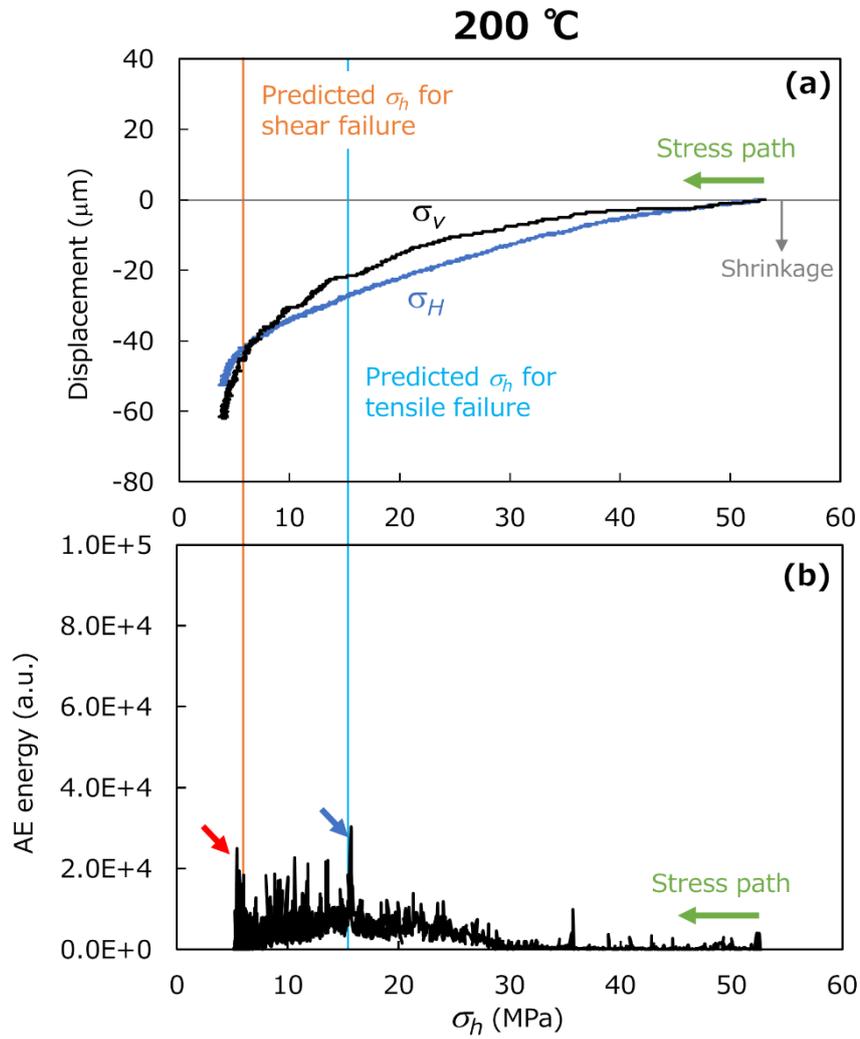
248 Although the final σ_h in each experiment exceeded the predicted criteria for the shear
249 failure and/or tensile failure of the wellbore, the changes in displacement (Figures 4c, 5c, and
250 6c) and cumulative AE energy (Figures 4d, 5d, and 6d) in all experiments were continuous
251 before reaching a plateau, suggesting non-catastrophic shear and tensile failures, as illustrated
252 in Figure 3. The displacement data at 200 °C indicate that the sample shrank similarly in both
253 the σ_H and σ_V directions with decreasing σ_h . With increasing temperature, the sample shrank
254 less in the σ_V and/or σ_H directions, and it dilated in the σ_V direction at 450 °C. Additionally,
255 we observed that the AE energy behavior changed with increasing temperature. For low
256 temperatures, we recorded smaller AE energies frequently. In contrast, at high temperatures,
257 higher AE energies were recorded infrequently, resulting in a lower cumulative AE energy.
258 These changes in displacement and AE behavior with increasing temperature may have been
259 caused by the predicted shift in the predominant failure mode of the wellbore from tensile to
260 shear failure (Table 1).

261

262 3.2. Wellbore failure prediction

263 Figures 7–9 show the changes in the displacement of the piston and AE energy as a
264 function of σ_h in each experiment, together with the predicted σ_h for the shear and tensile failure
265 of the wellbore. Although the sample shrank continuously in both the σ_H and σ_V directions
266 throughout the experiment at 200 °C, the shrinkage accelerated when the loading reached the
267 predicted tensile failure threshold (the stress path is indicated by the green arrow in Figure 7a).
268 Additionally, we recorded large AE energy events, as indicated by the blue arrow, when σ_h
269 reached the predicted tensile failure threshold (Figure 7b). This suggests that the tensile failure
270 of the wellbore started around the expected σ_h level and tensile fractures were generated in the
271 σ_V – σ_H plane, as illustrated in Figure 3, meaning that the sample deformed more easily in the σ_V
272 and σ_H directions. We recorded another set of large AE energy events, as indicated by the red
273 arrow, around the predicted shear failure threshold. However, after the presumed tensile failure,
274 the stress field was altered and the initial shear failure prediction may no longer have been
275 valid. Indeed, the displacement behavior did not change, despite the second set of large AE
276 energy events.

277



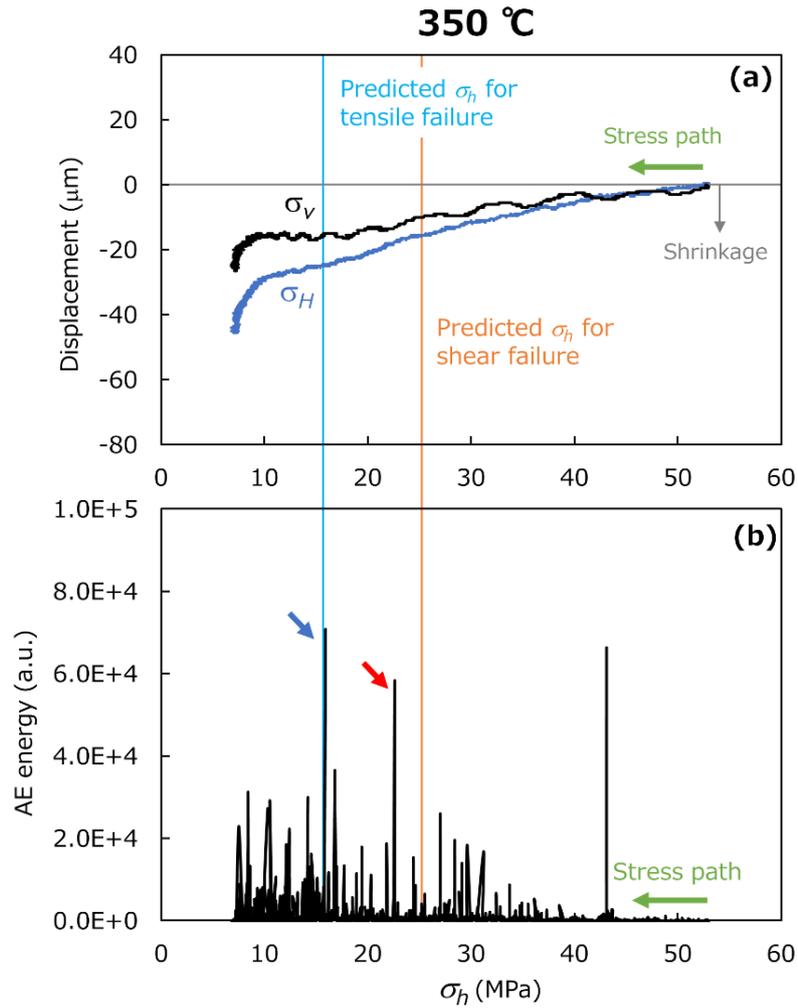
278

279 Figure 7. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)

280 as a function of σ_h in the experiment at 200 °C, together with the predicted σ_h for the shear and

281 tensile failure of the wellbore wall.

282



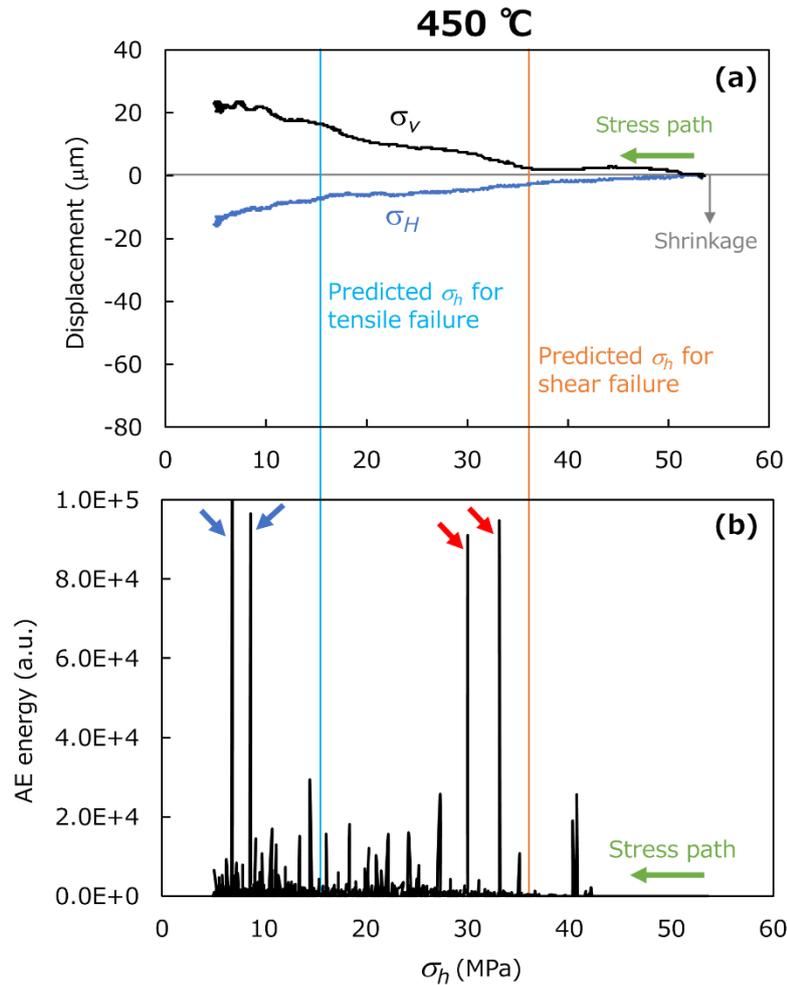
283

284 Figure 8. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)

285 as a function of σ_h in the experiment at 350 °C, together with the predicted σ_h for the shear and

286 tensile failure of the wellbore wall.

287



288

289 Figure 9. Changes in the displacement of the piston (a) and acoustic emission (AE) energy (b)
 290 as a function of σ_h in the experiment at 450 °C, together with the predicted σ_h for the shear and
 291 tensile failure of the wellbore wall.

292

293 In contrast, at 450 °C, significant dilation in the σ_v direction started after the predicted
 294 shear failure loading (Figure 9a) and large energy AE events occurred, as indicated by the red
 295 arrows in Figure 9b. This suggests the occurrence of wellbore shear failure that caused dilation
 296 in the σ_v direction, as Zhao et al. (2015) observed shear failure and increased axial deformation
 297 in a wellbore in granite under hydrostatic stress. After the predicted tensile failure, remarkably
 298 large energy AE events occurred again, as indicated by the blue arrows in Figure 9b. However,
 299 the second set of large energy AE events occurred at a much smaller σ_h than the predicted value

300 for tensile failure, and there was no clear indication of tensile failure in the deformation
301 behavior. This suggests that shear failure dominated the wellbore failure throughout the
302 experiment at 450 °C.

303 In the 350 °C experiment (Figure 8), after the sample reached the predicted shear failure,
304 the σ_h -dependence of the shrinkage in the σ_V direction decreased, flattening the displacement
305 curve at an σ_h of approximately 10–20 MPa. Based on the discussion of the 450 °C experiment,
306 this change in displacement behavior can be attributed to shear failure inducing dilation in the
307 σ_V direction. Subsequently, at $\sigma_h < 10$ MPa, the σ_h -dependence of the shrinkage in both the σ_V
308 and σ_H directions increased. Based on the discussion of the 200 °C experiment, this can be
309 attributed to tensile failure inducing shrinkage in both the σ_V and σ_H directions. Indeed, we
310 recorded large AE energies, as indicated by the red and blue arrows, around the predicted shear
311 and tensile failures. This suggests that, at 350 °C, both shear and tensile failure significantly
312 contributed to wellbore failure, owing to their relatively similar σ_h levels.

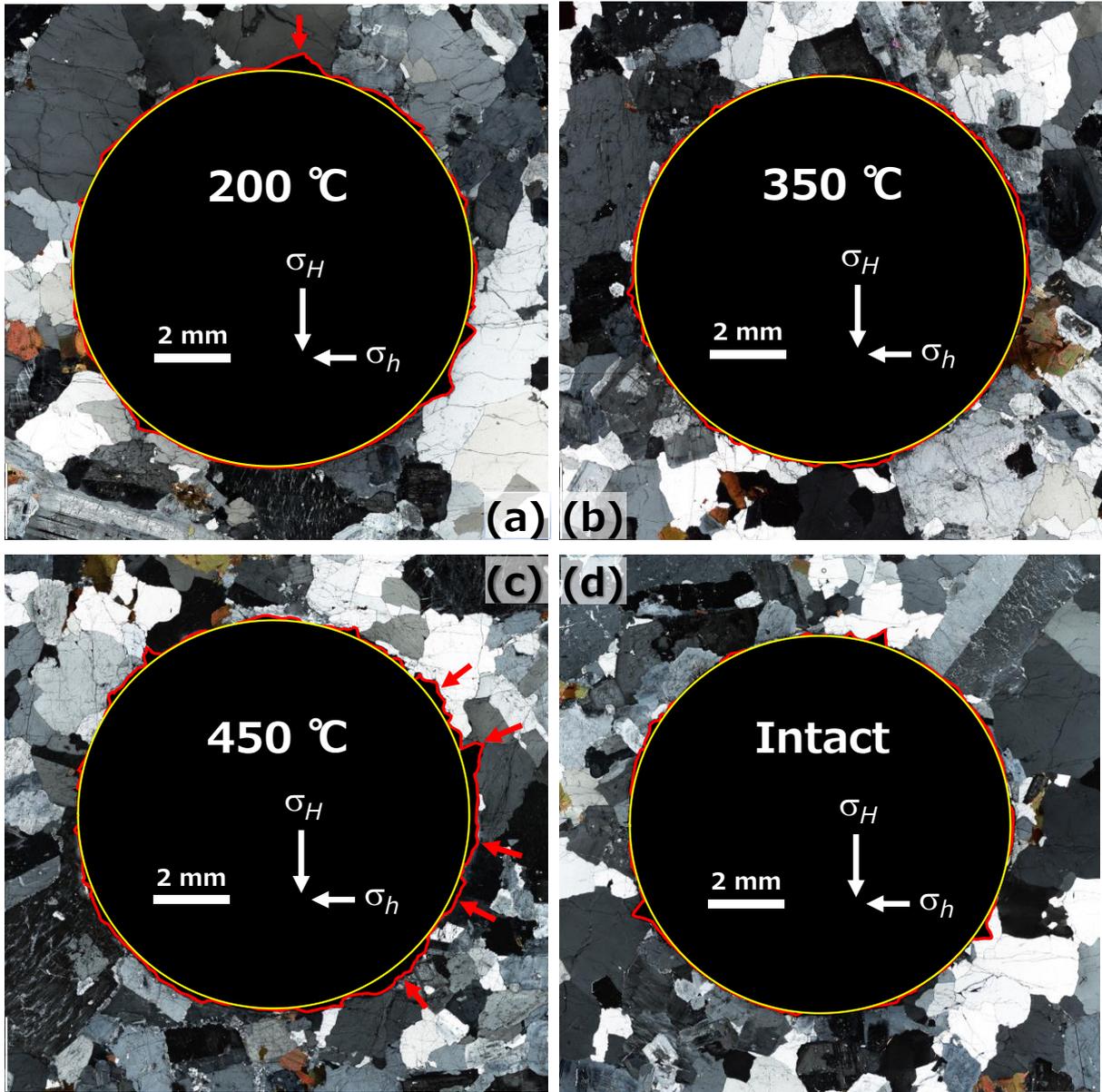
313 The above discussion suggests that the wellbore failure in the present experiments
314 occurred at σ_h levels that were largely consistent with the predicted values (Table 1). Therefore,
315 it is possible to predict wellbore failure in supercritical geothermal environments based on
316 existing rock failure criterion, such as the Hoek–Brown criterion used in this study.

317

318 **3.3. Possibility of using wellbore failure for in situ stress estimation**

319 Figure 10 shows optical microphotographs of the thin sections prepared from the
320 samples after the experiment (Figure 10a–c) and from an intact sample, for comparison, which
321 was not used in any experiment (Figure 10d). In these microphotographs, the profile of the
322 wellbore is traced in red, whereas the yellow line indicates a perfect circle

323



324

325 Figure 10. Optical microphotographs of thin sections prepared from the samples after the
 326 experiments at 200 (a), 350 (b), and 450 °C (c), and from an intact sample for comparison (d).

327

328 Even the intact sample profile does not completely match the perfect circle because of
 329 the relatively small defects with acute triangular shapes caused by boring (Figure 10d).
 330 However, the 200 °C profile deviates more from the perfect circle owing to the presence of
 331 larger defects in the wellbore, as indicated by the red arrow in Figure 10a. The largest defect
 332 appears to be a tensile fracture produced by the dominant failure at 200 °C (i.e., tensile failure

333 at $\theta = 0^\circ$ in Figure 3). The 450 °C profile also deviated significantly from the perfect circle
334 owing to the presence of much broader defects, as indicated by the red arrows (Figure 10c).
335 The defects appear to be breakouts produced by the dominant failure at 450 °C (i.e., shear
336 failure at $\theta = 90^\circ$ in Figure 3). The changes in the wellbore shape are consistent with our
337 speculation that the dominant failure mode shifts from tensile (200 °C) to shear (450 °C). In
338 contrast, the 350 °C profile did not significantly deviate from the perfect circle, although small
339 deviations appear around the entire wellbore wall (Figure 10b). It would therefore be
340 reasonable to assume no clear defects at specific locations if, as suggested in the previous
341 section (3.2), both shear and tensile failure significantly contributed to the wellbore failure at
342 350 °C.

343 The present experimental results therefore suggest that breakout and tensile fractures
344 may occur in wellbores in supercritical geothermal environments. However, at either 350 °C
345 or 450 °C, we did not observe the textbook breakout shape (Figure 3), such as the shape
346 observed in previous studies on granite at room temperature (Lee and Haimson, 1993; Song
347 and Haimson, 1997). This suggests that high temperatures suppress the propagation of shear
348 failure and its formation of clear breakout patterns. This may have been related to the
349 intermittent large energy AE events at 350 and 450 °C (Figures 5d and 6d), because such
350 intermittent large AE events were not observed in the previous study (Lee and Haimson, 1993).
351 This, in turn, suggests that in situ stress estimation from the shape of the breakout would be
352 difficult in supercritical geothermal environments because of the suppression of failure
353 propagation at high temperatures. However, if shear failure propagation is suppressed and
354 wellbore failure does not propagate further into the formation at high temperatures, this would
355 greatly improve wellbore stability and suggests the possibility of underbalanced drilling, which
356 would substantially reduce the risk of circulation loss during drilling. If our conclusions are
357 correct, they would have tremendous implications for drilling practices in supercritical

358 geothermal development. The suppressed failure propagation from the wellbore wall into the
359 rock may have been caused by a rapid relaxation of stress concentration around the wellbore
360 since several points on the circumference of the wellbore wall simultaneously reached failure,
361 or became close, due to the relatively small unconfined compressive strength. However, this
362 was not confirmed, hence warrants further investigation. Thus, extensive future studies are
363 encouraged to examine the suppression of failure propagation at high temperatures and to
364 clarify the mechanisms involved. Ideally, such studies should address wellbore failure in high-
365 temperature granite along the temperature and stress paths that may occur during drilling, with
366 considerations on influences of fluid pressure and thermal stress as well as *in situ* preexisting
367 weakness planes.

368

369 **4. Conclusions**

370 For the successful development of enhanced geothermal systems in supercritical
371 geothermal environments through hydraulic fracturing, it is essential to establish wellbore
372 stability and understand *in situ* stresses. The results of the first experiments on wellbore failure
373 in 200–450 °C granite under true triaxial stress suggest that non-catastrophic tensile and shear
374 failure occurs in the wellbore at such temperatures, and their initiation is predictable based on
375 existing brittle failure criterion for rocks. The results also suggest that the resultant shape of
376 the wellbore after failure, particularly after breakout by shear failure, is not ideal for *in situ*
377 stress estimation, because of the suppressed propagation of shear failure at high temperatures.
378 In other words, there is a possibility that the high temperatures in supercritical geothermal
379 environments contribute to wellbore stability. Therefore, extensive future studies are
380 encouraged to examine the suppression of failure propagation at high temperatures and to
381 clarify the underlying mechanisms.

382

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395

396 **Data availability**

397 The data supporting the findings of this study are available from the corresponding
398 author upon reasonable request.

399

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