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1 Estimation of critical stress ranges to preserve soil functions for differently textured soils

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19 **Research highlights**

20 Mechanical precompression stress can be similar for different textured soils near field capacity.

21 Texture does not primarily determine the compaction sensitivity of soils.

22 X-ray CT provide valuable additional information about the effect of mechanical stresses.

23 **Abstract**

24 The use of heavy agricultural equipment often produces significant changes in soil physical proper-
25 ties through compaction. Soil compaction is one of the environmental factors in agriculture that
26 adversely act on soil functions such as the provision of air, water, nutrients and pore space for root
27 growth affecting crop yields.

28 In this study, critical stress values are defined as the values at which the effects of compaction
29 result in a limitation of soil functions. Soil functions such as water storage, habitat and plant pro-
30 duction are influenced to a different degree by soil compaction, so we assume there is not some
31 fixed critical stress threshold for soil, but a range of critical stress values, hereafter called critical
32 stress range. It is investigated, if there are differences in the critical stress values with respect to
33 soil functions leading to critical loads that should not be exceeded to prevent negative effects, and
34 if the critical stress ranges differ for different textures.

35 Using column experiments in a greenhouse classic soil mechanical parameters (dry bulk density,
36 mechanical precompression stress), morphometric parameters obtained with X-ray tomography
37 (macroporosity, pore connectivity) from undisturbed soil samples as well as biological parameters
38 (earthworm activity) and crop factors (grain and straw yield) were investigated for soils with differ-
39 ent topsoil textures (top soil 0-15 cm: loam, silty clay loam, silt loam, sandy loam) near field capaci-
40 ty.

41 We found that critical stress values for various parameters and critical stress ranges did not de-
42 pend on topsoil texture at this matric potential. Studies, which recommend maximum values for
43 mechanical loads to prevent harmful soil compaction solely based on texture, should be treated
44 with caution. Although the soil textures at the four sites were quite different, the middle of the criti-
45 cal stress ranges were similar and concurred with the values of the mechanical precompression
46 stresses which were similar at all four sites, too. The agronomic critical stress values of grain and
47 straw yield mostly were impaired at the lower limit of the critical stress ranges and were, therefore,
48 the most sensitive parameters.

49 **Keywords:** pre-compression stress; X-ray CT; soil compaction

50 **Introduction**

51 Plants use a soil as a growth medium, which supplies water, oxygen and nutrients. Agricultural
52 soils are therefore conditioned to create a soil structure that allows optimal plant and root growth,
53 optimal biological activity to release nutrients, and to facilitate the supply of water and oxygen
54 (Carter, 1986). As a result of excessive machine loads as well as improper air pressure in the tires
55 and tillage equipment, harmful soil compaction can occur (Pagliai et al., 2000; Rücknagel et al.,
56 2012). This is especially true for unfavorable water contents during agronomic operations. Soil
57 compaction is an undesirable change in soil structure (Ishaq et al., 2001) which influences not only
58 pore functions such as air and water movement (Lipiec & Hatano, 2003), but also biological activity
59 (Lipiec & Simota, 1994).

60 It is important to evaluate the changes of soil properties caused by soil compaction from differ-
61 ent angles, (Horn & Rostek, 2000; Lipiec & Hatano, 2003), as will be outlined below.
62 Lipiec & Hatano (2003) refer to soil properties including soil strength, aeration, water, and structur-
63 al characteristics that make it easier to characterize soil quality after compaction. A key criterion for
64 the soil's stability when subjected to mechanical loads is precompression stress, a property that is
65 often used in soil mechanics (Horn & Rostek, 2000). Once it is exceeded irreversible changes in
66 soil functions occur (Rücknagel et al., 2017). Macropores, typically defined for pore diameters
67 $>50\ \mu\text{m}$, constitute a comparatively small fraction of the total pore space, but can contribute sub-
68 stantially to total flow, especially at high water content and at high precipitation rates where they
69 cause the phenomena of preferential flow. This is why physical disturbance of macropores caused
70 by soil loading may lead to a significant reduction of hydraulic conductivity and air permeability as
71 described by McKenzie et al. (2009). With a decreasing macroporosity, it can be assumed that
72 larger air-filled pores become discontinuous causing limitations in oxygen supply. In a recent re-
73 view, Rabot et al. (2019) highlight the relevance of the pore structure for the multitude of soil func-
74 tions. Hence, it seems to be promising to complement classical methods by non-destructive imag-

75 ing tools such as computed tomography to better characterize the structure of the pore network in
76 quantitative terms (Jarvis et al., 2017; Pihlap et al., 2019; Pöhlitz et al., 2019). In agricultural sys-
77 tems, biological properties as well as physical or mechanical properties are important components
78 of the soil ecosystem as pointed out by Carter (1986). He also advises that indices of both soil
79 structure and biological conditions are important for understanding the behaviour of soil functions
80 and, thus, the ability to avoid or recover from soil compaction. McKenzie et al. (2009) emphasize
81 that earthworms are the major component of the soil fauna in temperate agro-ecosystems. Also
82 Guei & Tondoh (2012) explain the important role of earthworms as soil “ecosystem engineers”
83 brought by feeding, burrowing and forming habits in and between casts which maintain soil fertility
84 and soil conservation (organic matter and macroaggregates). Soil compaction can gradually affect
85 earthworm activity in soil (Ruiz et al., 2015). They can be sensitive to tillage techniques and can
86 therefore be used as bio-indicators of soil conditions (Lemtiri et al., 2014). Besides soil fauna also
87 plant growth can be a sensitive indicator for soil compaction. Czyz & Tomaszewska (2001) found
88 that grain yield was linearly related to root mass, showing the importance of good soil physical
89 conditions for root growth for optimal yields. Crop yields in compacted soils are usually associated
90 with the extent and function of the root system (Lipiec & Hatano, 2003) which is impaired by de-
91 creasing root penetration due to excessive mechanical resistance (Lipiec & Simota, 1994), reduced
92 infiltration and insufficient aeration (Czyz & Tomaszewska, 2001).

93 To evaluate the impact of soil compaction on soil functions it is of particular interest to know at
94 what state of compaction soil properties such as dry bulk density, porosity or pore connectivity be-
95 come critical. The reduction of a soil property with increasing load might be sharp or gradual. In the
96 latter case the definition of a critical compaction status is difficult and typically set to empirically
97 determined values based on the soil properties of interest like crop growth. Thereby, critical
98 thresholds are expected to depend on climatic conditions, soil type and crop species (Rashid &
99 Sheikh, 1977). Several studies examined dry bulk density (BD) as suitable indicator for critical soil
100 compactions since high BD has been considered limiting for soil aeration and rooting (O’Connell,
101 1975). More specifically, an air capacity, i.e. air-filled porosity at field capacity, of less than 10 %

102 was found to be limiting for crop growth (Lebert et al., 2004; O'Connell, 1975; Reichert et al., 2009;
103 Werner & Paul, 1999) and the corresponding bulk density at which this occurs was determined by
104 regression equations that also take organic matter, particle density and combined silt and clay con-
105 tent into account. Kaufmann et al. (2010) list optimal and limiting values for dry bulk density derived
106 from parabolic relationships between dry bulk density and crop yield which show a pronounced
107 maximum depending on soil conditions, crop species and climate.

108 Focusing on only one soil property is problematic because it is possible that one soil property is
109 optimal, while another already shows critical values. The questions arise, whether the soil functions
110 mentioned above react differently to soil compaction, i.e. become limiting at the same applied
111 stress, and whether there are texture dependent differences. In this study, critical stress values are
112 defined as the values at which the effects of compaction result in a limitation of soil functions. Soil
113 functions and their parameters are influenced to a different degree by soil compaction, so we as-
114 sume there is no fixed critical stress threshold for a given soil, but a range of critical stress values,
115 hereafter called critical stress range. We investigate the influence of soil texture on changes in
116 some key properties which are deemed important for different soil functions. Classic soil mechani-
117 cal parameters and macropore characteristics derived from X-ray CT (undisturbed samples) as
118 well as biological and plant parameters (number of biopores, grain and straw yield) are linked. The
119 aim is to explore if there are differences in the critical stress values for the different parameters,
120 and if the critical stress ranges differ for different textures in topsoil near field capacity. We chose
121 field capacity because this corresponds to the field conditions when mechanical loading typically
122 occurs, namely high water content in the spring after the winter melt for sowing and in late autumn
123 for harvesting. We are aware that the strength of a soil can increase with increasing drought (Pöh-
124 litz et al., 2019; Rücknagel et al., 2012). This study should provide extended insights into the top-
125 soil compaction process and structural characteristics of different soil textures near field capacity,
126 and contribute to a suitable management of soils so that soil physical, morphological, biological
127 and plant parameters are not adversely affected by soil compaction.

128 **2. Material and Methods**

129 **2.1. Data acquisition**

130 In autumn 2016 soil samples were taken from the topsoil (0-20 cm) of four differently textured
131 sites (loam, silty clay loam, silt loam, sandy loam) which were specifically selected to represent a
132 wide range of soil textures. The clay content varies between locations from 70 g kg⁻¹ to 280 g kg⁻¹.
133 The sand contents are between 40 g kg⁻¹ and 530 g kg⁻¹. At all sites, long-term reduced tillage
134 (FAO, 1993) took place with cultivator. An overview and description of the locations is given in Ta-
135 ble 1.

136 **2.2. Soil physical measurements**

137 For the soil physical investigations undisturbed samples (volume = 250 cm³, height = 6.1 cm,
138 internal diameter = 7.2 cm) were taken in three repetitions at five places per site (3 x 5 x 4 = 60).

139 The saturated hydraulic conductivity (K_s , cm d⁻¹) of the soil samples (volume = 250 cm³,
140 height = 6 cm) was measured by means of a stationary system (Klute & Dirksen, 1986) with a flow
141 duration of 4 h. The dry bulk density (BD, Mg m⁻³) of the same samples was subsequently deter-
142 mined from the dry weight after drying at 105°C for 48 h (Blake & Hartge, 1986).

143 **2.3. Soil mechanical measurements**

144 For soil mechanical measurements undisturbed soil samples (volume = 220 cm³,
145 height = 2.8 cm, diameter = 10 cm) were taken at four places per site (4 x 4 = 16). The shallow
146 sample geometry was dictated by subsequent oedometer measurements (see below). For the
147 compression tests eight load steps (5, 10, 25, 50, 100, 200, 350 and 550 kPa) were applied to
148 each soil sample.

149 The soil samples were first slowly saturated by capillary rise before being drained for at least
150 seven days in a sandbox with a hanging water column to a matric potential of -6 kPa (Klute, 1986)
151 and then weighed.

152 Fully automated oedometers and the associated software (WINBOD32, Wille Geotechnik, APS
153 Antriebs-, Prüf- und Steuertechnik GmbH, Göttingen-Rosdorf, Germany) were used to determine
154 the stress - strain relationships under drained conditions. Load application was uniaxial. Each load
155 step was applied with a load time of 120 min and a subsequent relaxation time of 15 min with a
156 2 kPa load. The oedometer records settlement with an accuracy of 0.01 mm.

157 After the compression tests, the soil samples were dried at 105°C for 48 h and then weighed
158 (Blake & Hartge, 1986). The dry mass was then divided by the initial sample volume to compute
159 the dry bulk density prior to the compression tests (BD_0). Using the settlement (s), the initial height
160 of the soil sample (h_0), and BD_0 the resulting BD after each load application (BD_{xi}) was calculated
161 as follows:

$$162 \quad BD_{xi} = BD_0 \cdot \frac{h_0}{h_0 - s} \quad (3.1)$$

163 A semi-logarithmic stress - BD_{xi} curve was then created. The mechanical precompression
164 stress was determined based on these curves using the graphical method of Casagrande (1936). It
165 was applied by a number of experimenters to minimize subjectivity (Rücknagel et al., 2010).

166 **2.4. CT examinations**

167 An X-ray CT scan was acquired after each load step (App. 1). Soil samples from the compression
168 tests were scanned with an energy of 150 kV and a beam current of 550 μ A using an industrial X-
169 ray scanner (X-Tek XTH225, Nikon Metrology). One scan comprised 2480 projections with an ex-
170 posure time of 1.41 s (scanning time was $2480 \cdot 1.42s = 3500$ s). A CCD detector panel with
171 2000×1750 diodes recorded the projections. Beam hardening artifacts were reduced with a 0.1 mm
172 copper filter. The CT scans were reconstructed with a spatial resolution of 60 μ m and an 8-bit
173 greyscale resolution using the X-Tek CT Pro software package (Nikon Metrology). This is the max-
174 imum resolution, which allows to scan the entire sample (10 cm in diameter) and to get representa-
175 tive CT images. The results therefore pertain to pore sizes larger than 60 μ m, but can only detect
176 pores larger than two-three voxels faithfully (Vogel et al., 2010). Image processing was performed

177 with the Java software ImageJ 1.50e (Rasband, 1997-2015). To reduce scatter and noise the CT
178 scans were filtered using the “Non-local Means Denoising” plugin in Fiji (Buades et al., 2005).

179 In order to exclude artefacts at the edges of the sample and reduce the data volume a cylindri-
180 cal region of interest (ROI) with a diameter of 90 mm was used in the middle of the reconstructed
181 CT scan. The vertical extent of the ROI was adjusted to the reduction in sample height after each
182 load step. This was based on the positions of small and identifiable features, e.g. stones, at the
183 upper and lower ends of the sample. Regardless of compaction status it was thus always possible
184 to locate the original soil volume again after each consecutive load application.

185 Automatic segmentation was then applied to the scan of the ROI to separate the image into
186 pores and soil matrix. This was carried out using the thresholding method by Otsu (1979).
187 Macroporosity (here pore diameter $> 60 \mu\text{m}$) was quantified as the ratio of the number of pore
188 voxels to the total number of voxels within the ROI. The ImageJ analysis “Particle Analyzer” (Fer-
189 reira & Rasband, 2010-2012) was employed to calculate pore connectivity, which represents the
190 connection probability between two arbitrarily chosen pore voxels, i.e. the chance to belong to
191 same pore cluster. This dimensionless number is also denoted as the Γ indicator (Renard & Allard,
192 2013; Schlüter et al., 2014) and has a value between 0 and 1, where the latter indicates that the
193 soil pores are perfectly connected.

194 **2.5. Soil biological and agronomic measurements**

195 ***Column preparation***

196 Sufficiently large soil volumes are required for soil biological and agronomic measurements to min-
197 imize the impact of wall artifacts. Undisturbed soil cores of such size could not be extracted in the
198 field. Therefore, a total of approximately 750 kg of disturbed soil from 0-20 cm depth was taken
199 from each site ($\sim 750 \text{ kg} \times 4 = 3000 \text{ kg}$) to set up greenhouse experiments with repacked soil at
200 different compaction levels. For the column experiments with earthworms (*Lumbricus terrestris*)
201 and for the column experiments with spring barley (*Hordeum vulgare*) six dry bulk densities were
202 produced in five repetitions for each site ($6 \times 5 \times 4 = 120$ columns). For both column experiments,

203 the columns were positioned in a randomized order. The two column experiments took place inde-
204 pendently of each other.

205 Opaque polyvinyl chloride (PVC) pipes were used for the column experiments (19 cm inside
206 diameter, 283.52 cm² surface area, 30 cm height). For each site it was tested in advance how far
207 the soil was compactable which depends on soil texture. This was done by compacting the loose
208 soil piece by piece in a 5 cm layer. In the end, we used the soil mass to calculate the final density.
209 This final density corresponds to the highest density we were able to produce in the laboratory.
210 From this density, which was different for each test site, we produced 5 further densities in steps of
211 0.07 Mg m⁻³. Also, we compared these highest densities with the values that resulted from the
212 stress strain tests. These corresponded to each other. Thus, for the four sites dry bulk density
213 ranges of 1.42-1.77 Mg m⁻³ (loam), 1.28-1.63 Mg m⁻³ (silty clay loam), 1.21-1.56 Mg m⁻³ (silt loam)
214 and 0.72-1.07 Mg m⁻³ (sandy loam) were produced.

215 The soil was manually compacted at a water content near field capacity directly after soil sam-
216 pling using metal plates. To get the first layer a pre-weighed amount of soil was filled into a column
217 and beaten until its volume was reduced to the required extent (5 cm). The second layer was then
218 filled onto the compacted first layer in the column and treated in the same manner, and so on, until
219 all the pre-weighed soil was filled into the column. In the columns representing a given site the
220 same soil mass was used for all six dry bulk densities. This resulted in decreasing filling heights
221 with increasing dry bulk density. The soil mass was chosen to yield a maximum filling height of
222 25 cm at the lowest dry bulk density. Using the same soil mass entails the same amount of nutri-
223 ents in each column representing the same field site.

224 On the inside of the columns a duct tape which reached about 2.5 cm thickness into the col-
225 umn was attached after the second (10 cm), third (15 cm) and fourth (20 cm) layer counting from
226 the top. This was done to prevent earthworms from crawling and plant roots from growing preferen-
227 tially along the column wall.

228 **Biological measurements**

229 The earthworm genus *Lumbricus terrestris* L. was used for column experiments. In each col-
230 umn, six earthworms were placed on the soil surface. To prevent soil drying while permitting gas
231 exchange (O₂), 30 g of wheat and oat straw about 5 cm in length was mixed together and placed
232 on each surface. The upper end of the columns was covered with gauze and the lower one with a
233 fleece to prevent the earthworms from escaping. The experimental conditions were constant dark-
234 ness at 20°C. The total burrowing period of *L. terrestris* was 18 days.

235 After the experiment, the straw was carefully removed and collected by hand, together with the
236 earthworm casts on the soil, to make the earthworm burrows visible on the surface. The biopores
237 were counted at the top (0 cm), after the second (10 cm), third (15 cm), fourth (20 cm) and bottom
238 layer (25 cm) of the columns, summed up, divided by the number of counted layers and converted
239 to number per square meters representing average biopore activity of the first 25 cm topsoil in
240 field.

241 **Agronomic measurements**

242 For the agronomic investigations, summer barley (*Hordeum vulgare* L.) of the variety Avalon
243 was used. In each column, 15 plants were sown, covered with a roughly 3 cm thick soil layer of the
244 same soil, and thinned to ten plants after emergence. For nitrogen, a target value of 90 kg N ha⁻¹
245 was adjusted. N_{min} was determined (Tab. 1) and subtracted from the target value to obtain the
246 amount of nitrogen to be applied to the surface of the columns in the form of calcium ammonium
247 nitrate parallel to sowing.

248 The bottom of each column was covered with a fleece to prevent roots from growing out of the
249 column. The columns were weighed regularly every few days to monitor water loss by evapotran-
250 spiration. This loss was then compensated by watering so that lack of water did not restrict plant
251 growth.

252 The experiments were carried out in a greenhouse with the climatic conditions regulated ac-
253 cording to the BBCH stages of cereals (Witzenberger et al., 1989) as follows: (i) germination, leaf

254 development and tillering: constant 15°C with a 12 h photoperiod; (ii) stem elongation and booting:
255 20°C during the day and 15°C at night with a 14 h photoperiod; (iii) inflorescence emergence,
256 heading, flowering, anthesis, development of fruit, ripening and senescence: 25°C during the day
257 and 20°C at night with a 15 h photoperiod. The total growing period of the crop was 130 days.

258 At the senescence stage, the plants were harvested 1 cm above the soil. Straw and grain were
259 dried separately at 105°C to a constant weight, which was then converted into a yield (g m^{-2}).

260 **2.6. Derivation of critical stress values and critical stress ranges**

261 *Critical stress values and critical stress ranges*

262 With precompression stress and CT-derived parameters we target the critical stresses were
263 transitions in the properties are strongest or reach absolute thresholds, whereas for biopores and
264 yield we target optimum values, i.e. where the onset of detrimental effects by compaction occur.
265 We think that this is justified by the fact that the former are soil inherent properties, whereas the
266 later reflect emerging biological behavior constrained by the present soil structure. The respective
267 minimum and maximum spread have been added to the critical stress values. The critical stress
268 range is between the lowest and highest critical stress value of the examined parameters for each
269 soil.

270 *Mechanical precompression stress*

271 The mechanical precompression stress is widely viewed as the most important measure to as-
272 sess harmful soil compaction. Hence, it is used here as a critical stress value for the mechanical
273 component of a soil (section 2.3.).

274 *Macroporosity*

275 According to Werner & Paul (1999) an air capacity of ≥ 8 Vol.-% at pF1.8 (pores $> 50 \mu\text{m}$) is
276 necessary to maintain the ecological functionality of cohesive soils. Air capacity, when measured at
277 this matric potential, can be considered equivalent to macroporosity (Drewry et al., 2008). Here,
278 macroporosity was determined with CT quantitative image analysis at a similar resolution (pore
279 size). Following Werner & Paul (1999) a macroporosity ≥ 8 % was considered to be the minimum

280 required. So, as soon as a macroporosity of 0.08 was reached, the corresponding stress value was
281 considered to be the critical stress value.

282 *Pore connectivity*

283 There are no critical values given in the literature regarding pore connectivity, i.e. for the col-
284 lapse of a well-connected pore network into many isolated pores. In this study, the first significant
285 change in connectivity with increasing load application is therefore considered the critical stress
286 value.

287 *Biopores, Grain & Straw yield*

288 For the relationship between dry bulk density and yield an optimal dry bulk density at maximum
289 yields is frequently reported (Czyz & Tomaszewska., 2001, Czyz, 2004). Since we noted that the
290 volume of biopores is directly related to bulk density, we assumed, just as described with the yield,
291 that an optimal dry bulk density also leads to maximum biopore numbers. BD_{opt} is then used to
292 derive the critical stress value with the help of the stress - BD_{xi} diagram as follows (Fig. 1): BD_{opt} is
293 targeted on the abscissa of the BD – Biopores or BD – yield diagram. Then the critical stress value
294 at which BD_{opt} value is reached can be read of the ordinate of the stress - BD_{xi} diagram. At lower
295 BD there is little burrowing activity, because there is less need to dig to obtain food and shelter or
296 there is bad soil root contact (Kemper et al., 1988; Shah et al., 2017; Stovold et al., 2004). At high-
297 er BD burrowing is mechanically restricted or there is a high mechanical resistance to root penetra-
298 tion and reduced availability of oxygen, water and nutrients (Daddow & Warrington, 1983;
299 Håkansson, 1989; Saqib et al., 2004).

300 **2.7. Statistical analyses**

301 The statistical analyses were carried out with the statistics program 'R Studio' (version
302 0.99.893, R Foundation for Statistical Computing).

303 The arithmetic mean values for BD_{xi} , macroporosity, pore connectivity, biopores density and
304 crop yield were calculated separately for the repetitions of each site. The means of the log-
305 normally distributed saturated hydraulic conductivity and precompression stress values were calcu-

306 lated based on the logarithmized values. For the variance analyses, all parameters were tested for
307 normal distribution (Shapiro-Wilk test) and variance homogeneity (Levene's test).

308 A one-way analysis of variance was performed for all parameters between sites within each
309 load step, and between the eight load steps of the compression tests and between the six dry bulk
310 densities from the column experiments for each location separately. The Tukey honestly significant
311 difference test was applied to determine differences in group means and was considered signifi-
312 cant at a significance level of $p \leq 0.05$.

313 **3. Results**

314 **3.1. Soil mechanical parameters and accompanying critical stress values**

315 The stress - BD_{xi} diagrams (Fig. 2) show classic compaction curves for each of the four sites.
316 The BD_{xi} curves for the four sites run more or less parallel, but are shifted up or down with respect
317 to their initial densities, where the loam site had the highest and the sandy loam site the lowest
318 density. The values of the mechanical precompression stress ($\sigma_{P\ BD_{xi}}$) determined from the stress -
319 BD_{xi} diagrams do therefore not differ significantly between the sites. They are in the range of log
320 $\sigma_P = 1.65$ to 1.85 (45 to 71 kPa).

321 The increase in load results in significant decreases in macroporosity (Fig. 3) and pore connec-
322 tivity (Fig. 4), regardless of the site. Only at the lowest load steps, did the loam and silt loam sites
323 on the one hand, and the silty clay loam and sandy loam sites on the other differ significantly from
324 each other in terms of macroporosity and pore connectivity (App. 2). Because of different initial dry
325 bulk densities, the loam, the silt loam and the sandy loam site have lower macroporosity and pore
326 connectivity values than the silty clay loam site.

327 The stress - macroporosity curves (Fig. 3) show the most significant decrease in macroporosity
328 in the load range between 25 and 200 kPa. At the highest load steps (350 to 550 kPa) macropores
329 disappeared almost completely. This is true for all four sites. The macroporosity at the loam site
330 can already be regarded as critical (<0.08) at a load of 12 kPa (Fig. 3A). At the silty clay loam site

331 (Fig. 3B) the critical stress value for macroporosity is reached at a much higher load of 155 kPa.
332 The silt loam (Fig. 3C) and sandy loam (Fig. 3D) sites have a critical stress value for macroporosity
333 of 52 and 72 kPa, respectively, which is near the value for the mechanical precompression stress
334 for these two sites.

335 The shape of the stress - pore connectivity curves (Fig. 4) for all sites is similar to that of the
336 stress - macroporosity curves, with the difference that pore connectivity remains more constant
337 with increasing load for a longer time before decreasing steeply in the load range of 50-200 kPa. In
338 all soils a significant reduction in connectivity happened at a load of 100 kPa except for the silt
339 loam site (Fig. 4C) where this ad-hoc definition of a critical stress value for pore connectivity oc-
340 curred at 200 kPa. This is in contrast to other soils where the collapse of connectivity happened
341 already at a load of 100 kPa.

342 **3.2. Soil biological parameters and accompanying critical stress values**

343 *L. terrestris* dug successfully into the soil in all columns, even at the highest dry bulk densities.
344 Mortality was negligible and was not related to BD. *L. terrestris* formed permanent continuous bur-
345 rows with little branching. Several individuals of *L. terrestris* in a column sometimes used the same
346 burrow system. This means not every earthworm dug a biopore. The shape of the biopores was
347 affected by the compaction procedure which led to a series of layers with a slightly higher bulk
348 density in the upper part as compared to the lower. Interestingly, *L. terrestris* tended to follow the
349 less compacted zone on top of each layer so that the burrows in those areas were horizontal. This
350 was especially observed at the maximum burrowing depths. But this peculiarity did not affect the
351 estimated number for biopore density. There were biopores even at the lower end of some lower
352 density columns.

353 For all soils, except for the loam, the density of biopores slightly increased with increasing bulk
354 density before it declined above some optimal bulk density (Fig. 5B-D). For the loam soil no effect
355 of BD on the number of biopores was observed (Fig. 5A, App. 3). In this case the critical stress
356 value of 11 kPa was set at $BD_{xi} = 1.42 \text{ Mg m}^{-3}$ where the mean biopore density was obtained.

357 To determine an optimal BD with respect to the burrowing activity of earthworms a polynomial
358 could be fitted reasonably well to the measures biopore densities except for the loam (Fig. 5B-D).
359 We used the maximum value of the fitted polynomials to determine the optimal BD and critical load
360 which led to 1.40 Mg m⁻³/80 kPa, 1.35 Mg m⁻³/43 kPa and 0.82 Mg m⁻³/20 kPa for the silt clay loam,
361 silt loam and sandy loam respectively.

362 **3.3. Agronomic parameters and accompanying critical stress values**

363 For the silt loam and sandy loam soil grain yields increased with increasing bulk density (Fig. 6,
364 App. 3) before it decreased again after an optimal bulk density was exceeded. For the loam and
365 silty clay loam soils, grain yields decrease already with the first investigated reduction in bulk den-
366 sity. The loam (Fig. 6A) and silty clay loam (Fig. 6B) have the largest amplitudes in grain yields.
367 Correspondingly, the graphs of the polynomial functions are steeper than those for the silt loam
368 (Fig. 6C) and sandy loam (Fig. 6D) sites. For the latter two soils, no significant differences in grain
369 yield with increasing density are found.

370 To determine an optimal BD with respect to the grain yield a polynomial could be fitted reason-
371 ably well to the measures grain yield densities except for the silt loam (Fig. 6A-D). We used the
372 maximum value of the fitted polynomials to determine the optimal BD and critical load which led to
373 1.39 Mg m⁻³/7 kPa, 1.30 Mg m⁻³/45 kPa, 1.36 Mg m⁻³/50 kPa and 0.89 Mg m⁻³/56 kPa for the loam,
374 silt clay loam, silt loam and sandy loam respectively.

375 In principle, the results for grain yield are reflected in the results for straw yield, with the differ-
376 ence that more straw was formed than grain, which led to shifts in the optimum density values
377 (Fig. 7).

378 The maximum straw yield occurs at a density of 1.47 Mg m⁻³ for loam (Fig. 7A), 1.31 Mg m⁻³ for
379 silty clay loam (Fig. 7B), 1.35 Mg m⁻³ for silt loam (Fig. 7C), and 0.87 Mg m⁻³ for sandy loam
380 (Fig. 7D). This corresponds to critical stress values of 30 kPa for loam, 48 kPa for silty clay loam,
381 43 kPa for silt loam, and 50 kPa for sandy loam. Again, for the last three sites the derived critical
382 stress values for straw yield are in the range of the mechanical precompression stresses.

383 **3.4. Critical stress ranges**

384 The critical stress values given in chapters 3.1 to 3.3. are summarized in Figure 8. The mini-
385 mum and maximum stress values for each soil determine the critical stress range, indicated by
386 dotted vertical lines. The critical stress ranges differ between the soils. While the stress range is
387 rather large for the loam, it is similar and much narrower for the silty clay loam, silt loam and sandy
388 loam.

389 For the relations between individual critical stress values some general trends can be derived.
390 The spread of the mechanical precompression stress are very small for all four soils. For most of
391 the soils the critical stress value for macroporosity is smaller than the one for pore connectivity.
392 The critical stress values for the biopores, grain and straw yield are usually in the lower part of the
393 critical stress ranges.

394 **4. Discussion**

395 **4.1. General remarks**

396 Before looking at the critical stress values and critical stress ranges resulting from soil compac-
397 tion the effects of compaction on the examined soil functions observed here should be briefly dis-
398 cussed.

- 399 • The initial BD's for the four sites were much lower than site-specific plant root limiting BD's
400 according to Kaufmann et al. (2010). Furthermore, the K_s values were much higher than the
401 minimum rate recommended by Werner & Paul (1999). Hence, the soil structures were in-
402 tact at all sites.
- 403 • Despite different initial densities, all stress - BD_{xi} curves showed a similar shape. The indi-
404 vidual curves were merely shifted somewhat along the ordinate. Therefore, the values for
405 the mechanical precompression stress were similar.

- 406 • For soils with low initial macroporosity the critical stress value for macroporosity is smaller
407 than the one for pore connectivity, whereas the values are similar for soils with high initial
408 macroporosity. This discrepancy is caused by the different criteria of how critical stress val-
409 ues are determined for both properties. For macroporosity absolute values are considered
410 (threshold at 0.08), whereas for pore connectivity relative differences mattered (significant
411 changes between consecutive load steps). This also restricted critical stress values to ac-
412 tual measured load steps as no interpolation is possible. In the future the method could be
413 changed to an absolute threshold of 0.5. It has been shown frequently that pore connectivi-
414 ty drops in a narrow macroporosity range of 5-10% (Jarvis et al., 2017; Pihlap et al., 2019;
415 Pöhlitz et al., 2019). This also explains the empirical air capacity limit of about 8 % by vol-
416 ume in Werner & Paul (1999), because then the air-filled pore space becomes discontinu-
417 ous and the oxygen supply is disturbed. Even if the mechanical precompression stress is
418 exceeded, the macropores are still connected.
- 419 • Several individuals of *L. terrestris* shared biopores. This is in agreement with the studies of
420 Jegou et al. (1998) and Joschko et al. (1989).
- 421 • The burrowing activity of *L. terrestris* varied with dry bulk density, following a parabolic
422 curve except for the loam site. At low BD there was little burrowing activity, because there
423 is less need to dig to obtain food and shelter. Up to a certain point (optimal BD) the number
424 of burrows increased with increasing soil density. Beyond this point the burrowing activity
425 decreased with BD, because *L. terrestris* was mechanically restricted, i.e. soil strength
426 seemed to be a limiting factor at higher densities. A lower burrowing rate at higher densities
427 was also found by Kemper et al. (1988) and Stovold et al. (2004). Schrader et al. (2007)
428 and Kemper et al. (1988) observed that *L. terrestris* failed to penetrate a silt loam soil with a
429 BD of
430 1.60-1.70 Mg m⁻³. Similarly, Horn (1999) named 1.67 Mg m⁻³ as the BD limit for *L. terrestris*.
- 431 • We observed that at high soil densities, *L. terrestris* tended to remain in a particular loca-
432 tion. Perreault & Whalen (2006) also reported such a behaviour, if *L. terrestris* was exposed

433 to unfavorably cold and wet soil conditions. Recall that in this study, the soil was at 20°C
434 and moist, i.e. the conditions were favorable.

435 • Very loose (bad soil root contact) as well as heavily compacted soil (high mechanical re-
436 sistance to root penetration, reduced availability of oxygen, water and nutrients) leads to
437 reduced yields, as found by Daddow & Warrington (1983), Håkansson (1989), Saqib et al.
438 (2004) and Shah et al. (2017), too.

439 • The parabolic relationship between dry bulk density and burrowing activity of *L. terrestris* as
440 well as grain and straw yield of *H. vulgare* depends on texture. The optimal BD and the
441 steepness of the curve differ.

442 • It seems that an increased content of organic matter like at the sandy loam site (51 g kg⁻¹)
443 only has an effect on the dry bulk density of a soil, but not on the other parameters studied
444 here.

445 • The mechanical precompression stress reflects critical stress value for yield or critical
446 stress ranges in general.

447 **4.2. Critical stress ranges**

448 In this study, critical stress values for various parameters in topsoil near field capacity were
449 presented. Single values were taken from fitted curves. However, in some cases, the curves were
450 fairly flat and the data had a high standard deviation. Consequently, it is possible to pick a number
451 of plausible critical stress values for a given parameter. For some the possible critical stress values
452 have a rather big spread, for example the critical stress value for biopores at the loam site. This is
453 because biopore density is insensitive to dry bulk density. Overall, for grain and straw yield the
454 spread among soils is widest.

455 The values for the mechanical precompression stress at the four sites are fairly similar. The
456 spread is the lowest for all parameters. In general, the mechanical precompression stress lies in

457 the middle of the critical stress ranges for the other five investigated parameters and covers them
458 reasonably well.

459 The critical stress value for biopores, usually appeared at the lower part of the critical stress
460 range. For the loam site, one can assume that the critical stress value is very low due to the high
461 sand content, because the rough surface of the sand particles can impair burrowing activity
462 (McKenzie et al., 2009). At higher dry bulk densities this effect may increase, because the particles
463 move closer together, which further increases friction (Horn et al., 1995). However, with increasing
464 dry bulk density no significant differences of biopore number could be determined. This could be
465 attributed to the lack of aggregation (Beisecker et al., 1994) and the accompanying evenly distrib-
466 uted resistance. That may be why the earthworms were still able to burrow through the soil even
467 after the mechanical precompression stress was exceeded. For the sandy loam site, the high con-
468 tent of organic matter compensates the aforementioned effect of a high sand content, because it
469 ensures a loose and thus easily penetrable packing. In general, the dry bulk densities were low so
470 that there was little burrowing resistance.

471 The critical stress value for grain yield is very close to the one for straw yield. The loam soil
472 seems to be an exception to this but generalizations are difficult due to large spread in the data.
473 Grain and straw yield tend to become critical before the mechanical precompression stress is
474 reached. If one looks at the spread, grain and straw yield seem to be the most sensitive parame-
475 ters and form the lower limit of the critical stress range. This is because they are determined at the
476 bulk density of optimal growth and not at the bulk density of steepest decrease in some property,
477 e.g. pore connectivity. Iler & Stevenson (1991) report that sandy soils have high growth-limiting dry
478 bulk densities of around $1.65\text{-}1.75 \text{ Mg m}^{-3}$, but that plants already show a significant reduction in
479 growth before those densities are reached. In addition, plants differ in their tolerance of soil com-
480 paction affecting germination and growth of some species but not others (Skinner et al., 2009).
481 Ultimately, grain and straw yield are the most important parameters for the farmer.

482 Following these explanations it is not surprising that at different sites the critical stress value for
483 a given parameter may lie at the upper or lower end of the critical stress range for all parameters

484 considered here, and that its value and spread may differ between sites. For example, the critical
485 stress range is widest at the loam site and narrowest at the silty clay loam site. However, if the
486 middle of the critical stress ranges (recall the logarithmic scale) are looked at, it can be seen that
487 they are similar in value (50 - 100 kPa), and also similar to the values of the mechanical precom-
488 pression stresses (45 - 71 kPa). Hence, the mechanical precompression stress turns out to be well
489 suited as a measure of the critical stress values and the critical stress ranges. Nevertheless, useful
490 additional information about the effect of mechanical stresses on the soil structure is gained from
491 considering the other parameters as well.

492 The results show that topsoil texture does not determine the compaction sensitivity near field
493 capacity alone or at least not primarily. Field capacity has an important role because at this point
494 the soil is most susceptible to compaction and corresponds to the field conditions when mechanical
495 loading typically occurs, namely high water content in the spring after the winter melt and in late
496 autumn for sowing and harvesting. Furthermore, this moisture level is the closest state of the ag-
497 gregates for clay soils, which we have at the Buttelstedt site. However, of course the relationship
498 between precompression stress and soil moisture (matric potential) is different for different soil
499 textures – clay is very weak when wet but becomes very hard when dry, while sandy soils will not
500 show as much change with initial matric potential (Schjønning & Lamandé 2018).

501 **5. Conclusions**

502 In this study we found that critical stress with respect to soil functions does not depend on soil
503 texture. This can be concluded from a number of physical and biological soil characteristics (me-
504 chanical precompression stress, macroporosity, pore connectivity, number of biopores, grain and
505 straw yield) which are known to reflect soil functions and which were measured after applying in-
506 creasing mechanical loads for differently textured topsoils near field capacity. This means that
507 studies, which recommend maximum values for mechanical loads to prevent harmful soil compac-
508 tion solely based on texture, should be treated with caution.

509 Based on the different soil characteristics we could identify critical stress ranges which slightly
510 differ between the soils. However, despite the considerably contrast in soil texture, the centers of
511 the determined stress ranges were similar and concurred with the values of the mechanical pre-
512 compression stresses which were similar for all four soils. Hence, the mechanical precompression
513 stress was confirmed to be a suitable measure of the critical stress values and the critical stress
514 ranges. Nevertheless, useful additional information is gained from considering the other character-
515 istics as well. The critical stress values for indicators of plant growth (i.e. grain and straw yield),
516 were mostly found at the lower limit of the critical stress ranges meaning that plant growth is most
517 sensitively affected by soil compaction.

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525 **References**

- 526 Arvidsson, J, Keller, T., 2004. Soil precompression stress: I. A survey of Swedish arable soils. Soil
527 Till. Res. 77, 85-95.
- 528 Beisecker, R., 1994. Einfluss langjährig unterschiedlicher Bodenbearbeitungssysteme auf das Bo-
529 dengefüge, die Wasserinfiltration und die Stoffverlagerung eines Löß- und eines Sandbodens.
530 Bodenökologie und Bodengenese 12.

531 Berli, M., Kirby, J., Springman, S., Schulin, R.. 2003. Modelling compaction of agricultural subsoils
532 by tracked heavy construction machinery under various moisture conditions in Switzerland. *Soil*
533 *Till. Res.* 73, 57–66.

534 Blake, G.R., Hartge, K.H., 1986. Dry bulk density. In: Klute, A. (Ed.), *Methods of soil analysis. Part*
535 *1: Physical and mineralogical methods*, 2nd edition. American Society of Agronomy Monograph
536 No. 28, Madison, Wisconsin, USA, pp. 363–382.

537 Buades, A., Coll, B., Morel, J.M., 2005. A review of image denoising algorithms, with a new one.
538 *Multiscale Model. Simul.* 4, 490-530.

539 Carter, M. R., 1986. Microbial biomass as an index for tillage-induced changes in soil biological
540 properties. *Soil Till. Res.* 7, 29-40.

541 Casagrande, A. 1936. The determination of the pre-consolidation load and its practical signifi-
542 cance. *Proceedings of the International Conference on Soil Mechanics and Foundation Engi-*
543 *neering*, Vol. III, pp. 60–64. Harvard University, Cambridge, Massachusetts, USA.

544 Czyz, E. A., 2004. Effects of traffic on soil aeration, dry bulk density and growth of spring barley.
545 *Soil Till. Res.* 79, 153–166.

546 Czyz, E. A., Tomaszewska, J., Dexter, A. R., 2001. Response of spring barley to changes of com-
547 paction and aeration of sandy soil under model conditions. *Int. Agrophysics* 15, 9–12.

548 Daddow, R.L., Warrington, G.E., 1983. Growth-limiting soil bulk densities as influenced by soil tex-
549 ture. WSDG Report TN-00005. USDA Forest Service, Watershed Systems Development
550 Group, Fort Collins, Colorado, USA.

551 Dewry, J.J., Cameron, K.C., Buchan, G.D., 2008. Pasture yield and soil physical property respons-
552 es to soil compaction from treading and grazing – a review. *Aust. J. Soil Res.* 46, 237–256.

553 Diserens, E., 2009. Calculating the contact area of trailer tyres in the field. *Soil Till. Res.* 103, 302–
554 309.

555 FAO, 1993. *Soil tillage in Africa: needs and challenges*. FAO Soils Bulletin No. 69. FAO, Rome,
556 Italy.

557 FAO, 1998. World Reference Base for Soil Resources. ISSS–ISRIC–FAO. World Soil Resources
558 Report No. 84, Rome, Italy.

559 Ferreira, T., Rasband, W.S., 2010-2012. “ImageJ user guide - IJ 1.46”, [im-agej.nih.gov/ij/
560 docs/guide/](http://imagej.nih.gov/ij/docs/guide/)

561 Gee, G.W., J.W. Bauder. 1986. Particle-size analysis. In: Klute, A. (Ed.), Methods of soil analysis.
562 Part 1: Physical and mineralogical methods, 2nd edition. American Society of Agronomy Mono-
563 graph No. 28, Madison, Wisconsin, USA, pp. 383–411.

564 Guéi, A.M., Tondoh, J.E., 2012. Ecological preferences of earthworms for land-use types in semi-
565 deciduous forest areas, Ivory Coast. *Ecol. Ind.* 18, 644-651.

566 Håkansson, I., 1989. Compaction of the plough layer-which degree of compactness is the best?
567 *Swedish Univ. Agric. Sci., Uppsala. Fakta/markväxter* 1, 4.

568 Horn, R., 1999. Verdichtung von Böden – Überlegungen zum Prozess und zur Prognose der me-
569 chanischen Belastbarkeit. *Wasser und Boden* 51, 9–13.

570 Horn, R., Domzal, H., Slowinska-Jurkiewicz, A., van Ouwerkerk, C., 1995. Soil compaction pro-
571 cesses and their effects on the structure of arable soils and the environment. *Soil Till. Res.* 35,
572 23–36.

573 Horn, R., Rostek, J., 2000. Subsoil compaction processes – state of knowledge. *Advances in Ge-
574 oEcology. Reiskirchen, Catena*, 32.

575 Iler G.S., Stevenson C.K., 1991. The effects of soil compaction on the production of processing
576 vegetables and field crops. A review. Ridgetown College of Agricultural Technology, Ontario.

577 Ishaq., M., Hassan, A., Saeed, M., Ibrahim, M., Lal, R., 2001. Subsoil compaction on crops in Pun-
578 jab, Pakistan. I. Soil physical properties and crop yield. *Soil Till. Res.* 59, 57–65.

579 Jarvis, N., Larsbo, M., Koetsel, J., 2017. Connectivity and percolation of structural pore networks in
580 a cultivated silt loam soil quantified by X-ray tomography. *Geoderma* 287, 71–79.

- 581 Jegou, D., Cluzeau, D., Wolf, H. J., Gandon, Y., Trehen, P., 1998. Assessment of the burrow sys-
582 tem of *Lumbricus terrestris*, *Aporrectodea giardi*, and *Aporrectodea caliginosa* using X-ray
583 computed tomography. Biol. Fertil. Soils 26, 116–121.
- 584 Joschko, M., Diestel, H., Larink, O., 1989. Assessment of earthworm burrowing efficiency in com-
585 pacted soil with a combination of morphological and soil physical measurements. Biol. Fert.
586 Soils 8, 191–196.
- 587 Kaufmann, M., Tobias, S., Schulin, R., 2010. Comparison of critical limits for crop plant growth
588 based on different indicators for the state of soil compaction. J. Plant Nutr. Soil Sci. 173, 573–
589 583.
- 590 Kemper, W.D., Jolley, P., Roseneau, R.C., 1988. Soil management to prevent earthworms from
591 riddling irrigation ditch banks. Irrig. Sci. 9, 79–87.
- 592 Klute, A., 1986. Water Retention: Laboratory Methods. In: Klute, A. (Ed.), Methods of soil analysis.
593 Part 1: Physical and mineralogical methods, 2nd edition. American Society of Agronomy Mono-
594 graph No. 28, Madison, Wisconsin, USA, pp. 635–662.
- 595 Klute, A., Dirksen, C., 1986. Hydraulic Conductivity and Diffusivity: Laboratory Methods. In: Klute,
596 A. (Ed.), Methods of soil analysis. Part 1: Physical and mineralogical methods, 2nd edition.
597 American Society of Agronomy Monograph No. 28, Madison, Wisconsin, USA, pp. 687–734.
- 598 Krück, S., Joschko, M., Schultz-Sternberg, R., Kroschewski, B., Tessmann, J., 2006. A classifica-
599 tion scheme for earthworm populations (Lumbricidae) in cultivated agricultural soils in Bran-
600 denburg, Germany. J. Plant Nutr. Soil Sci. 169 (5), 651-660.
- 601 Lebert, M., Brunotte, J, and Sommer, C., 2004. Ableitung von Kriterien zur Charakterisierung einer
602 schädlichen Bodenverdichtung entstandenen durch nutzungsbedingte Verdichtung von Bö-
603 den/Regelungen zur Gefahrenabwehr, Umweltbundesamt, Berlin, 122.
- 604 Lehfeldt, J., 1988. Auswirkungen von Krümenbasisverdichtungen auf die Durchwurzelbarkeit san-
605 diger und lehmiger Bodensubstrate bei Anbau verschiedener. Arch. Acker-Pflanzenb. Bodenk.
606 32, 533–539.

607 Lemitri, A., Gilles, C., Alabi, T., Cluzeau, D., Zirbes, I., Haubruge, E., Franics, F., 2014. Impacts of
608 earthworms on soil components and dynamics. A review. *Biotechnol. Agron. Soc. Environ.* 18,
609 121–133.

610 Lipiec, J., Hatano, R., 2003. Quantification of compaction effects on soil physical properties and
611 crop growth. *Geoderma* 116, 107–136.

612 Lipiec, J., Simota, C., 1994. Role of soil and climate factors in influencing crop responses to soil
613 compaction in central and eastern Europe. *Developm.Agric. Eng.* 11, 365–390.

614 McKenzie, B. M., Kühner, S. MacKenzie, K., Peth, S., Horn, R., 2009. Soil compaction by uniaxial
615 loading and the survival of the earthworm *Aporrectodea caliginosa*. *Soil Till. Res.* 104, 320–
616 323.

617 Obour, P. B., Keller, T., Lamandé, M., Munkholm, L. J., 2019. Pore structure characteristics and
618 soil workability along a clay gradient. *Geoderma* 337, 1186-1195.

619 O’Connell, D.J., 1975. The measurement of apparent specific gravity of soils and its relationship to
620 mechanical composition and plant root growth. In: *Soil physical conditions and crop production*.
621 Ministry of Agriculture, Fisheries, and Food, Technical Bulletin 29. H.M. Stat. Off., London, pp.
622 298–313.

623 Otsu, N., 1979. A threshold selection method from gray-level histograms. *IEEE Trans. Syst. Man,*
624 *Cybern. B, Cybern.* Vol. SMC-9, No. 1, 62-66.

625 Pagliai, M., Pellegrini, S., Vignozzi, N., Rousseva, S., Grasselli, O., 2000. The quantification of the
626 effect of subsoil compaction on soil property and related physical properties under conventional
627 to reduced management practices. *Adv. Geocol.* 32, 305–313.

628 Perreault, J.M., Whalen, J.K., 2006. Earthworm burrowing in laboratory microcosms as influenced
629 by soil temperature and moisture. *Pedobiologia* 50, 397–403.

630 Pihlap, E., Vuko, M., Lucas, M., Steffens, M., Schloter, M., Vetterlein, D., Eendenich, M., Kögel-
631 Knabner, I., 2019. Initial soil formation in an agriculturally reclaimed open-cast mining area -
632 the role of management and loess parent material. *Soil Till. Res.* 191, 224-237.

633 Pöhlitz, J., Rücknagel, J., Koblenz, B., Schlüter, S., Vogel, H.-J., Christen, O., 2019. Computed
634 tomography as an extension of classical methods in the analysis of soil compaction, exempli-
635 fied on samples from two tillage treatments and at two moisture tensions. *Geoderma* 346, 52-
636 62.

637 Rabot, E. Wiesmeier, M., Schlüter, S., Vogel, H.-J., 2018. Soil structure as an indicator of soil func-
638 tions. A review. *Geoderma* 314, 122-137.

639 Rasband W.S., 1997 - 2015. ImageJ. U.S. National Institute of Health, Bethesda, Maryland, USA.
640 <http://imagej.nih.gov/ij/>

641 Rashid, S., Sheikh, K.H., 1977. Response of wheat to different levels of soil compaction. *Phyton*
642 18, 43–56.

643 Reichelt, J. M., Suzuki, L. E. A. S., Reinert, D. J., Horn, R., and Håkansson, I., 2009. Reference
644 bulk density and critical degree-of-compactness for no-till crop production in subtropical highly
645 weathered soils, *Soil & Till. Res.* 102, 242-254.

646 Renard, P., Allard, D., 2013. Connectivity metrics for subsurface flow and transport. *Adv. Water*
647 *Res.* 51, 168-196.

648 Ruiz, S. Or, D., Schymanski, S., 2015. Soil Penetration by Earthworms and Plant Roots–
649 "Mechanical Energetics of Bioturbation of Compacted Soils. *PloS one* 10: e0128914.
650 10.1371/journal.pone.0128914.

651 Rücknagel, J., Hofmann, B., Paul, R., Christen, O., Hülsbergen, K., 2007. Estimating precompres-
652 sion stress of structured soils on the basis of aggregate density and dry bulk density. *Soil Till.*
653 *Res.* 92, 213–220.

654 Rücknagel, J., Brandhuber, R., Hofmann, B., Lebert, M., Marschall, K., Paul, R., Stock, O., Chris-
655 ten, O., 2010. Variance of mechanical precompression stress in graphic estimations using the
656 Casagrande method and derived mathematical models. *Soil Till. Res.* 106, 165–170.

657 Rücknagel., J., Christen, O., Hofmann, B., Ulrich, S., 2012. A simple model to estimate change in
658 precompression stress as a function of water content on the basis of precompression stress at
659 field capacity. *Geoderma* 177-178, 1-7.

660 Rücknagel, J., Rademacher, A., Götze, P., Hofmann, B., Christen, O., 2017. Uniaxial compression
661 behaviour and soil physical quality of topsoils under conventional and conservation tillage. *Ge-
662 oderma* 286, 1–7.

663 Saqib, M., Akhtar, J., Qureshi, R. H., 2004. Pot study on wheat growth in saline and waterlogged
664 compacted soil. I. Grain yield and yield components. *Soil Till. Res.* 77, 169–177.

665 Schjønning, P., Lamandé, M. 2018. Models for prediction of soil precompression stress from readi-
666 ly available soil properties. *Geoderma* 320, 115-125.

667 Schjonning, P. & Rasmussen, K. J., 2000. Soil strength and soil pore characteristics for direct
668 drilled and ploughed soils. *Soil Till. Res.* 57 (1), 68-82.

669 Schjonning, P., van den Akker, J.J.H., Keller, T., Greve, M.H., Lamande, M., Simojoki, A., Stettler,
670 M., Arvidsson, J., Breuning-Madsen, H., 2015. Drive-Pressure-State-Impact-Response
671 (DPSIR) analysis and risk assessment for soil compaction – a European perspective. *Adv. Agr.*
672 133, 183–237.

673 Schlüter, S., Sheppard, A., Brown, K., Wildenschild, D., 2014. Image processing of multiphase
674 images obtained via X-ray microtomography: A review. *Water Resour. Res.* 50, 3615-3639.

675 Schrader, S., Rogasik, H., Onasch, I., Jegou, D., 2007. Assessment of soil structural differentiation
676 around earthworm burrows by means of X-ray computed tomography and scanning electron
677 microscopy. *Geoderma* 137, 378–387.

- 678 Shah, A. N., Tanveer, M., Shahzad, B., Yang, G., Fahad, S., Ali, S., Bukhari, M. A., Tung, S. A.,
679 Hafeez, A., Souliyanonh, B., 2017. Soil compaction effects on soil health and crop productivity:
680 An overview. *Environ. Sci. Pollut. Res.* 24, 10056–10067.
- 681 Skinner, A., Lunt, I., Spooner, P., McIntyre, S., 2009. The effect of soil compaction on germination
682 and early growth of *Eucalyptus albens* and an exotic annual grass. *Austral. Ecol.* 34, 698-704.
- 683 Stovold, R. J., Whally, W. R., Harris, P. J., White, P. W., 2004. Spatial variation in soil compaction,
684 and the burrowing activity of the earthworm *Aporrectodea caliginosa*. *Biol. Fertil. Soil* 39, 360-
685 365.
- 686 Tursic, I., 1982. The effect of soil compaction and mineral fertilization on the yield of spring barley.
687 *Polj. Znan. Smotra-Agr. Consp.* 58, 39-48 (in Croatian with an English summary).
- 688 Vogel, H.-J., Weller, U., and Schlüter, S. 2010: Quantification of soil structure based on Minkoski
689 functions, *Computers & Geo-sciences*, 36, 1236-1245.
- 690 Werner, D., Paul, R., 1999. Kennzeichnung der Verdichtungsgefährdung landwirtschaftlich genutz-
691 ter Böden. *Wasser und Boden* 51, 10-14.
- 692 Witzemberger, A., Hack, H., van den Boom, T., 1989. Erläuterungen zum BBCH-Dezimal-Code für
693 die Entwicklungsstadien des Getreides - mit Abbildungen. *Gesunde Pflanzen* 41, 384-388.

Table 1: Description of the sampled sites.

Site	T (°C)	N (mm)	Taxonomy ^a	Texture		Texture class ^b	TOC (g kg ⁻¹)	pH	Nutrients			N _{min} (kg N ha ⁻¹)	BD (Mg m ⁻³)	K _s (cm d ⁻¹)
				(g kg ⁻¹)					(mg per 100 g)					
				Clay	Sand				P	K	Mg			
Quellendorf	8.7	526	Chernozem	130	450	loam	14	7.4	6.8	22.1	10.4	11	1.29	158
Buttelstedt	8.4	541	Chernozem	280	40	silty clay loam	21	6.9	3.5	51.3	24.1	15	1.14	157
Rothenberga	8.5	500	Haplic Luvisol	130	60	silt loam	13	6.7	5.9	19.0	6.0	17	1.10	137
Kranichborn	8.5	500	Mollic Fluvisol	70	530	sandy loam	51	7.4	12.4	13.9	12.1	30	0.82	157

T = average annual temperature; N = average annual precipitation; TOC = total organic carbon; P = phosphorus; K = potassium; Mg = magnesium; N_{min} = mineralizable nitrogen; BD = dry bulk density; K_s = saturated hydraulic conductivity

All parameters except BD and K_s were determined by Eurofins Agraranalytik Deutschland GmbH, Jena, Germany.

^a FAO (1998)

^b USDA classification scheme (Gee & Bauder, 1986)

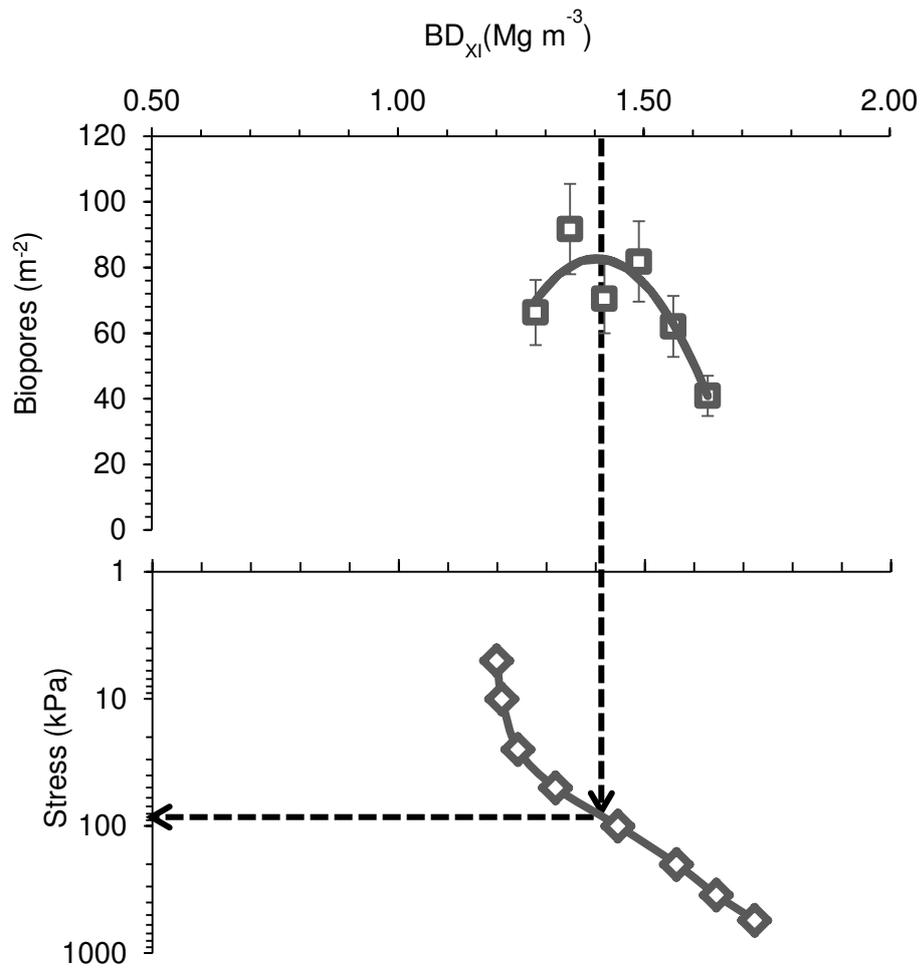


Figure 1: Scheme to derive critical stress value for earthworm activity. BD_{opt} ($1.41\ Mg\ m^{-3}$) is targeted on the abscissa of the $BD - Biopores$ diagram. Then the critical stress value (80 kPa) at which BD_{opt} value is reached can be read of the ordinate of the stress - BD_{xi} diagram.

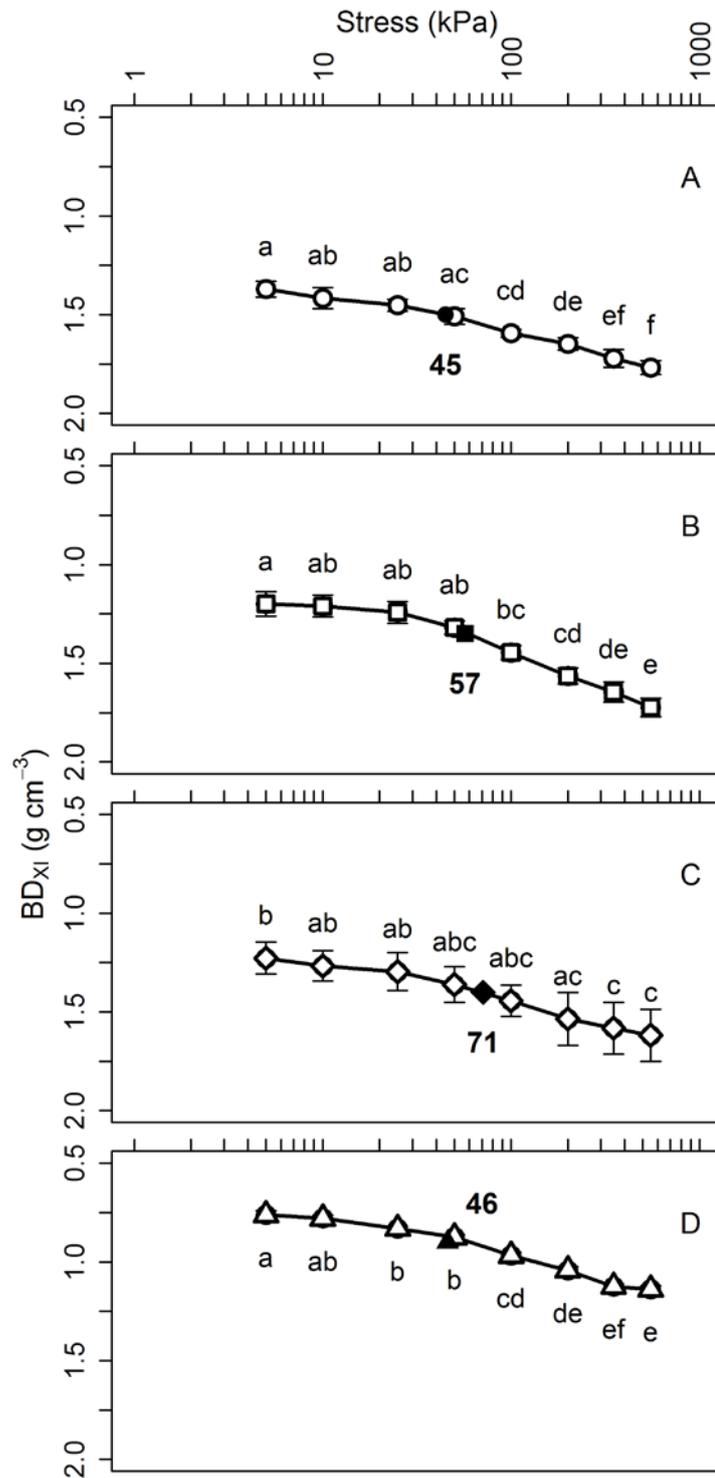


Figure 2: Dry bulk density (BD_{xi}) from sequential load application (stress) to soil cores from (A) loam, (B) silty clay loam, (C) silt loam and (D) sandy loam. Error bars show the standard deviations. Statistically significant differences ($p \leq 0.05$) are indicated by lower case letters. Black symbols and numbers indicate the values (kPa) of the mechanical pre-compression stress.

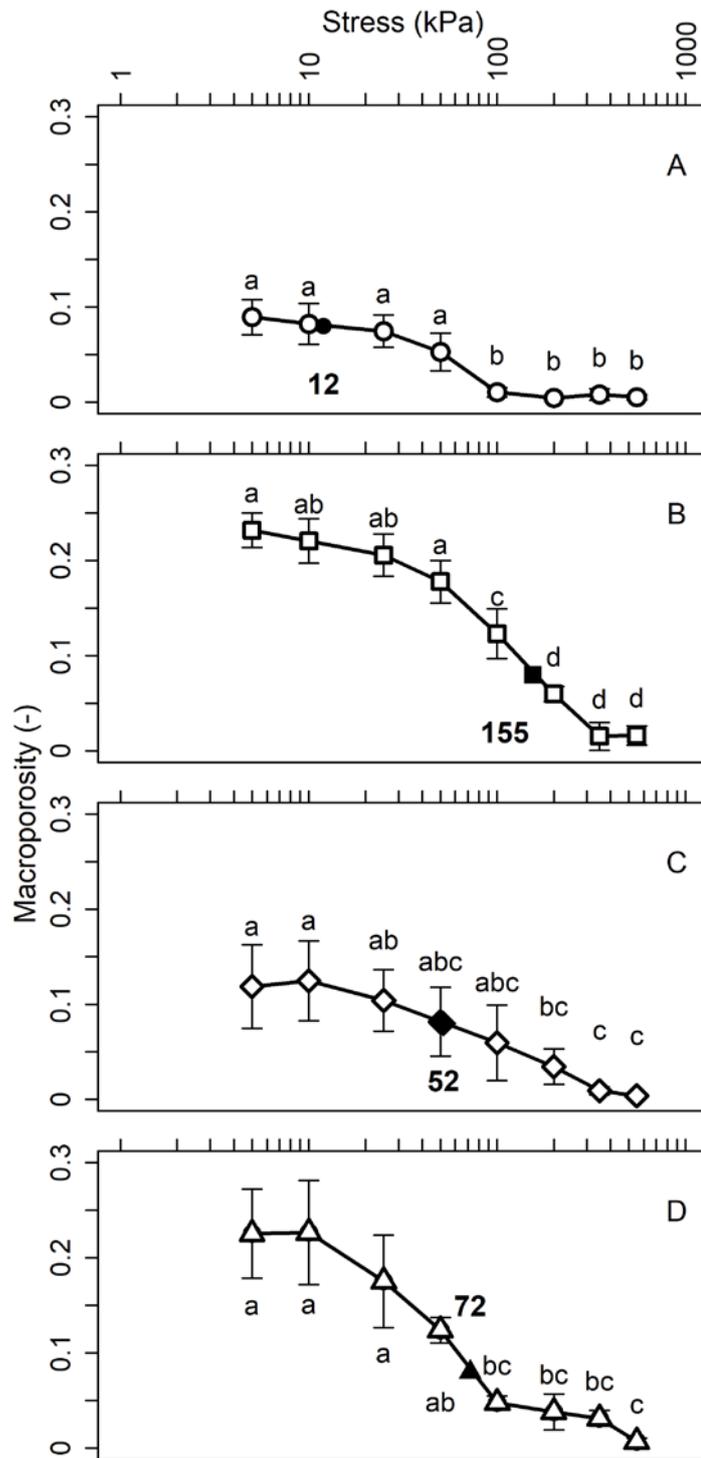


Figure 3: Macroporosity from sequential load application to soil cores for (A) loam, (B) silty clay loam, (C) silt loam and (D) sandy loam. Error bars show the standard deviations. Statistically significant differences ($p \leq 0.05$) are indicated by lower case letters. Black symbols and numbers indicate the critical stress values (kPa) which correspond to a macroporosity of 0.08.

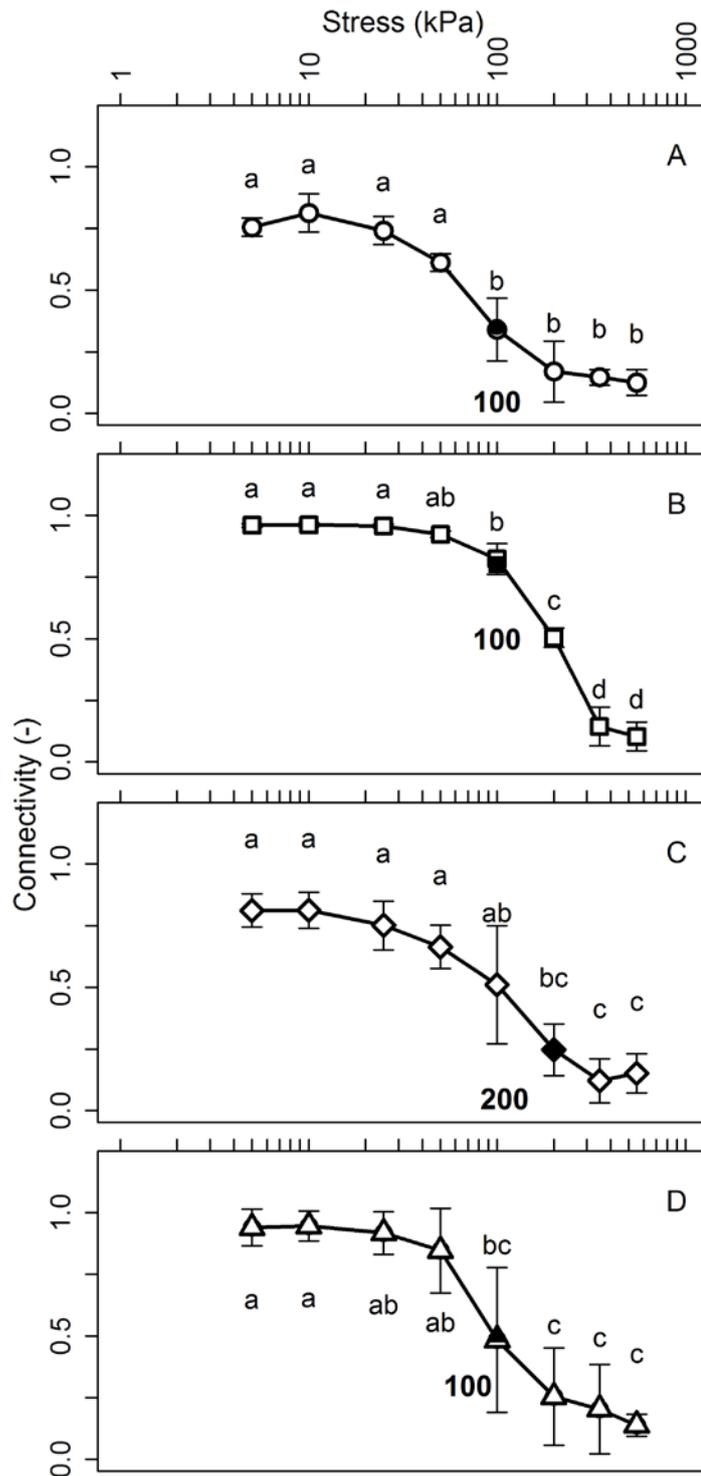


Figure 4: Pore connectivity from sequential load application to soil cores for (A) loam, (B) silty clay loam, (C) silt loam and (D) sandy loam. Error bars show the standard deviations. Statistically significant differences ($p \leq 0.05$) are indicated by lower case letters. Black symbols and numbers indicate the critical stress values (kPa) which correspond to the first significant change in pore connectivity.

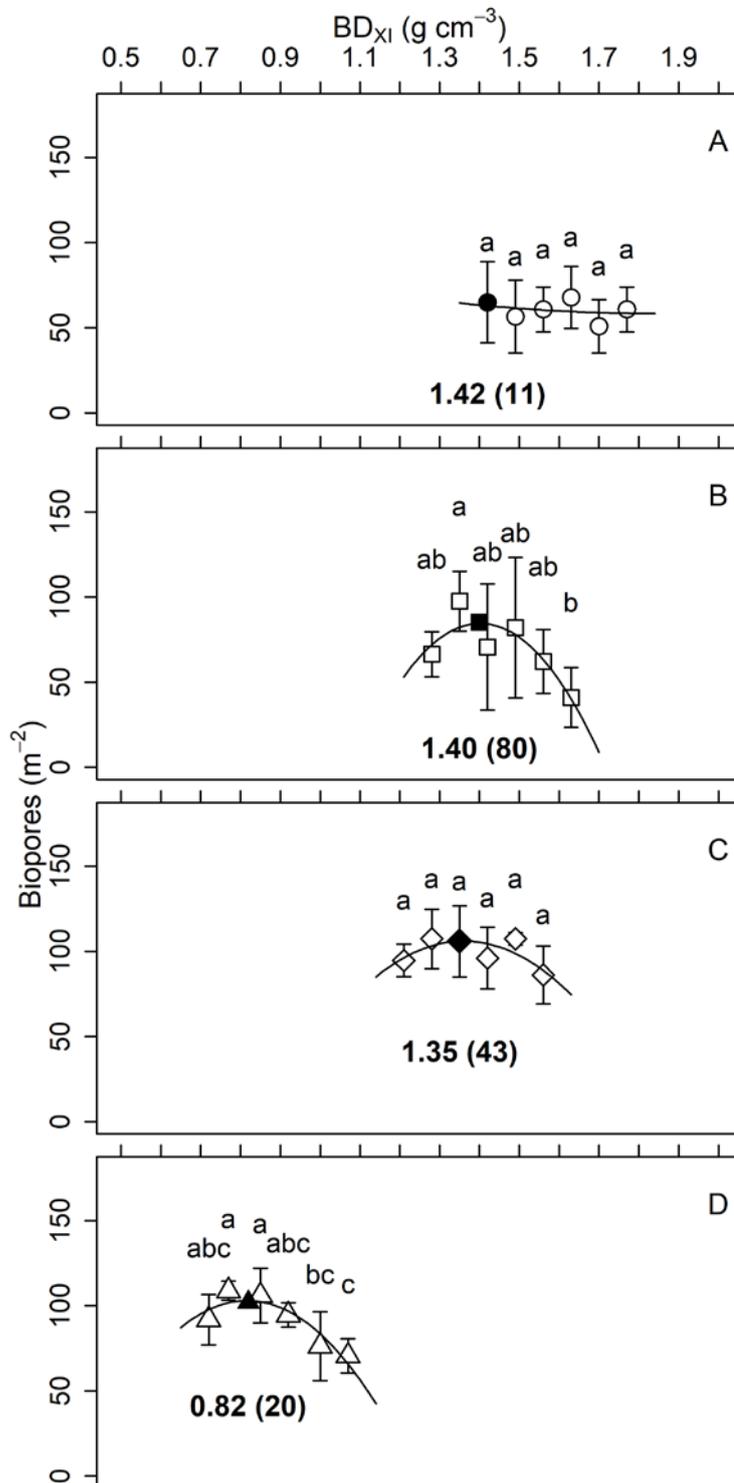


Figure 5: Biopores from *Lumbricus terrestris* as a function of dry bulk density (BD_{xi}) for (A) loam, (B) silty clay loam, (C) (silt loam and (D) sandy loam. Error bars show the standard deviations. Statistically significant differences ($p \leq 0.05$) are indicated by lower case let-

ters. Black symbols and numbers indicate the optimum BD (Mg m^{-3}) at which the biopores are at maximum and their critical stress value (kPa) in brackets.

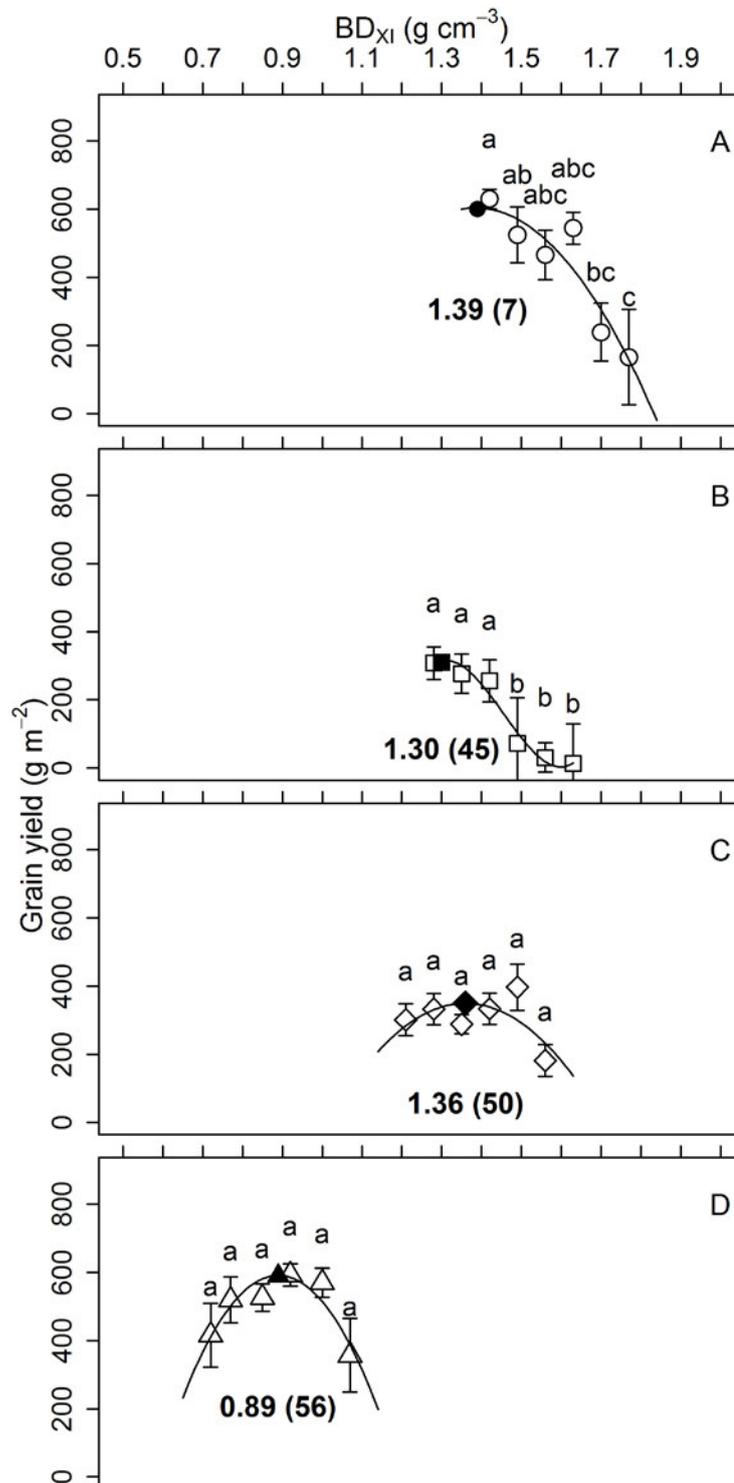


Figure 6: Grain yield of *Hordeum vulgare* as a function of dry bulk density (BD_{xi}) for (A) loam ($1.42\text{-}1.77 \text{ Mg m}^{-3}$), (B) silty clay loam ($1.28\text{-}1.63 \text{ Mg m}^{-3}$), (C) silt loam (1.21-

1.56 Mg m⁻³) and (D) sandy loam (0.72-1.07 Mg m⁻³). Statistically significant differences ($p \leq 0.05$) are indicated by lower case letters. Black symbols and numbers indicate the optimum BD (Mg m⁻³) at which the grain yield of *Hordeum vulgare* is at maximum and their critical stress value (kPa) in brackets.

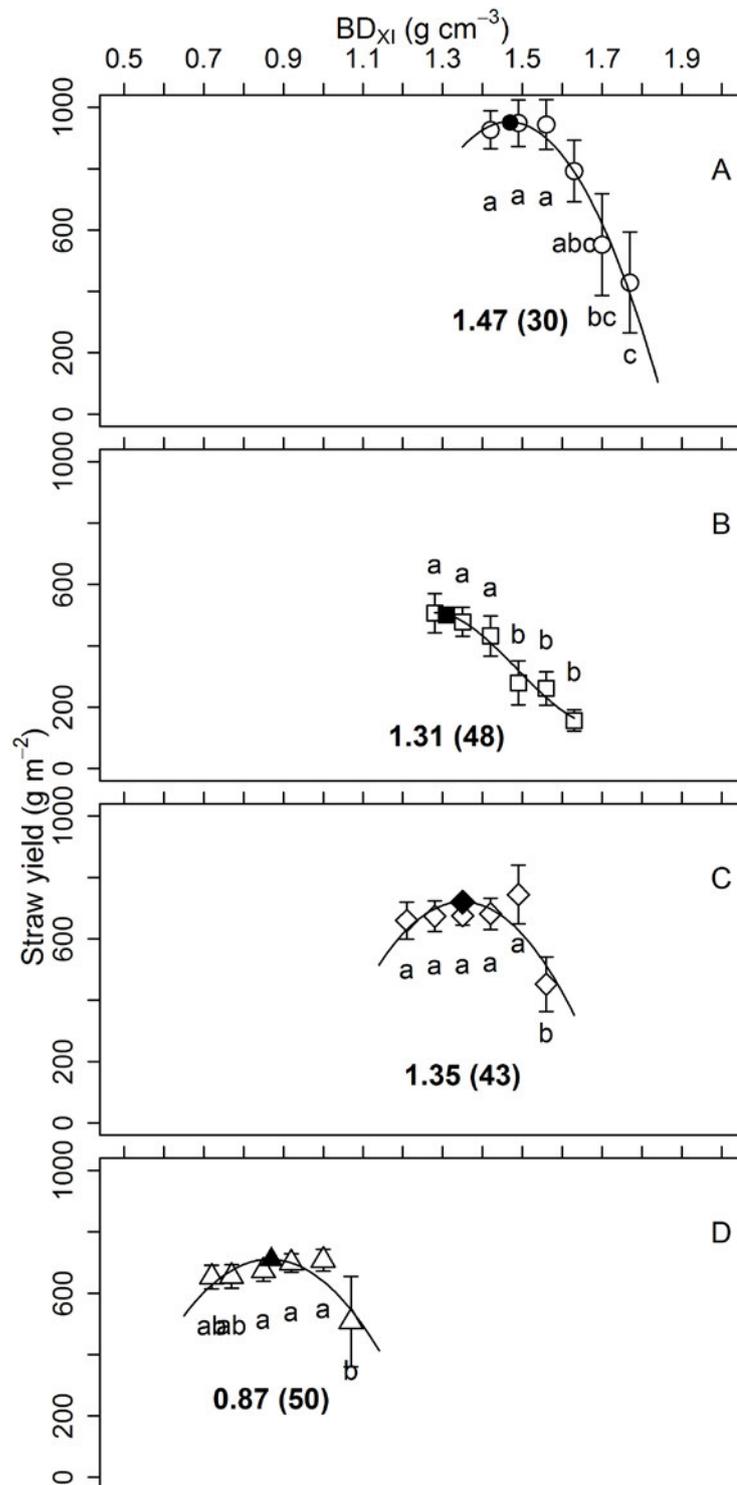


Figure 7: Straw yield of *Hordeum vulgare* as a function of dry bulk density (BD_{xi}) for (A) loam (1.42-1.77 Mg m⁻³), (B) silty clay loam (1.28-1.63 Mg m⁻³), (C) silt loam (1.21-1.56 Mg m⁻³) and (D) sandy loam (0.72-1.07 Mg m⁻³). Error bars show the standard deviations. Statistically significant differences (p ≤ 0.05) are indicated by lower case letters. Black symbols and numbers indicate the optimum BD (Mg m⁻³) at which the straw

yield of *Hordeum vulgare* is at maximum and their critical stress value (kPa) in brackets.

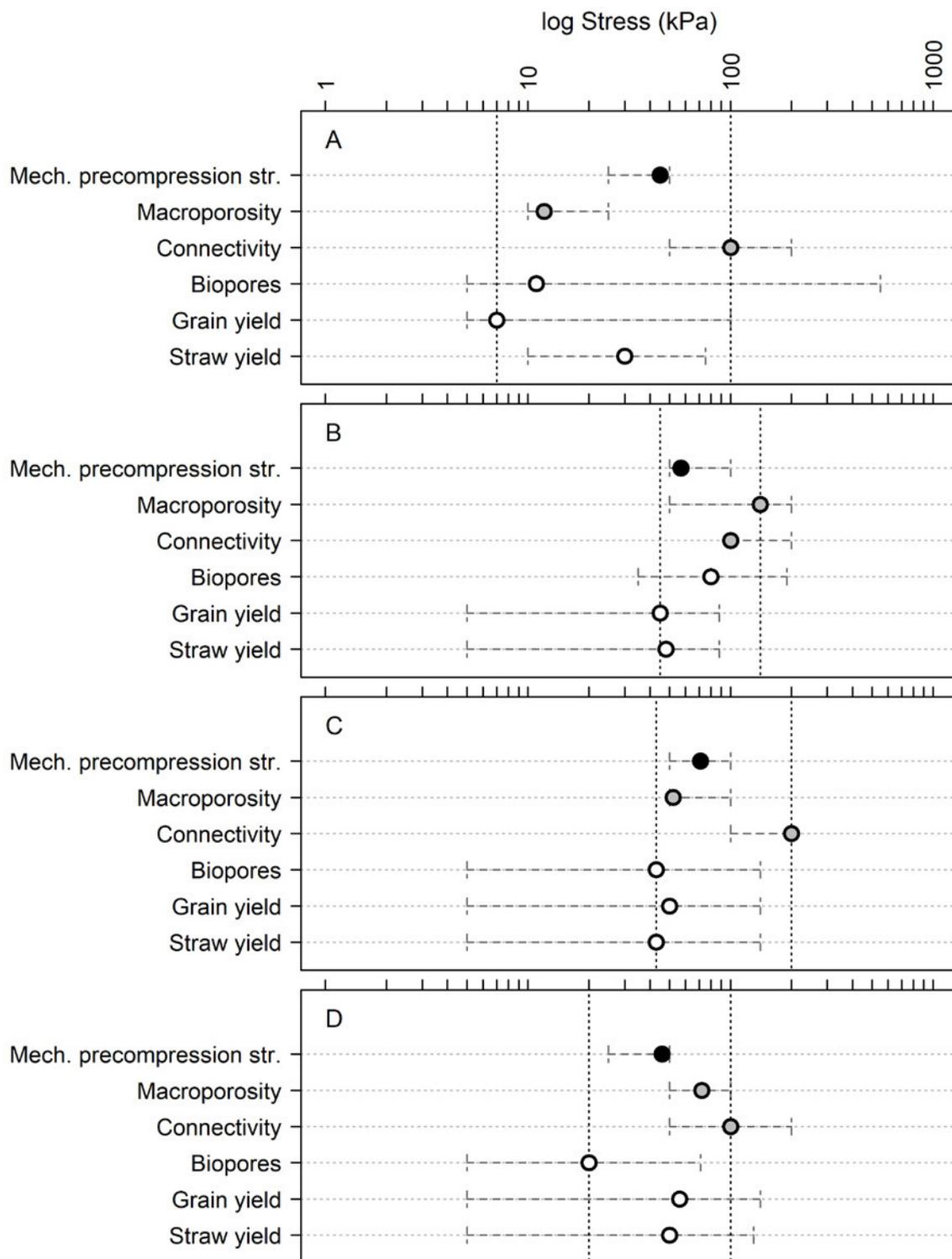
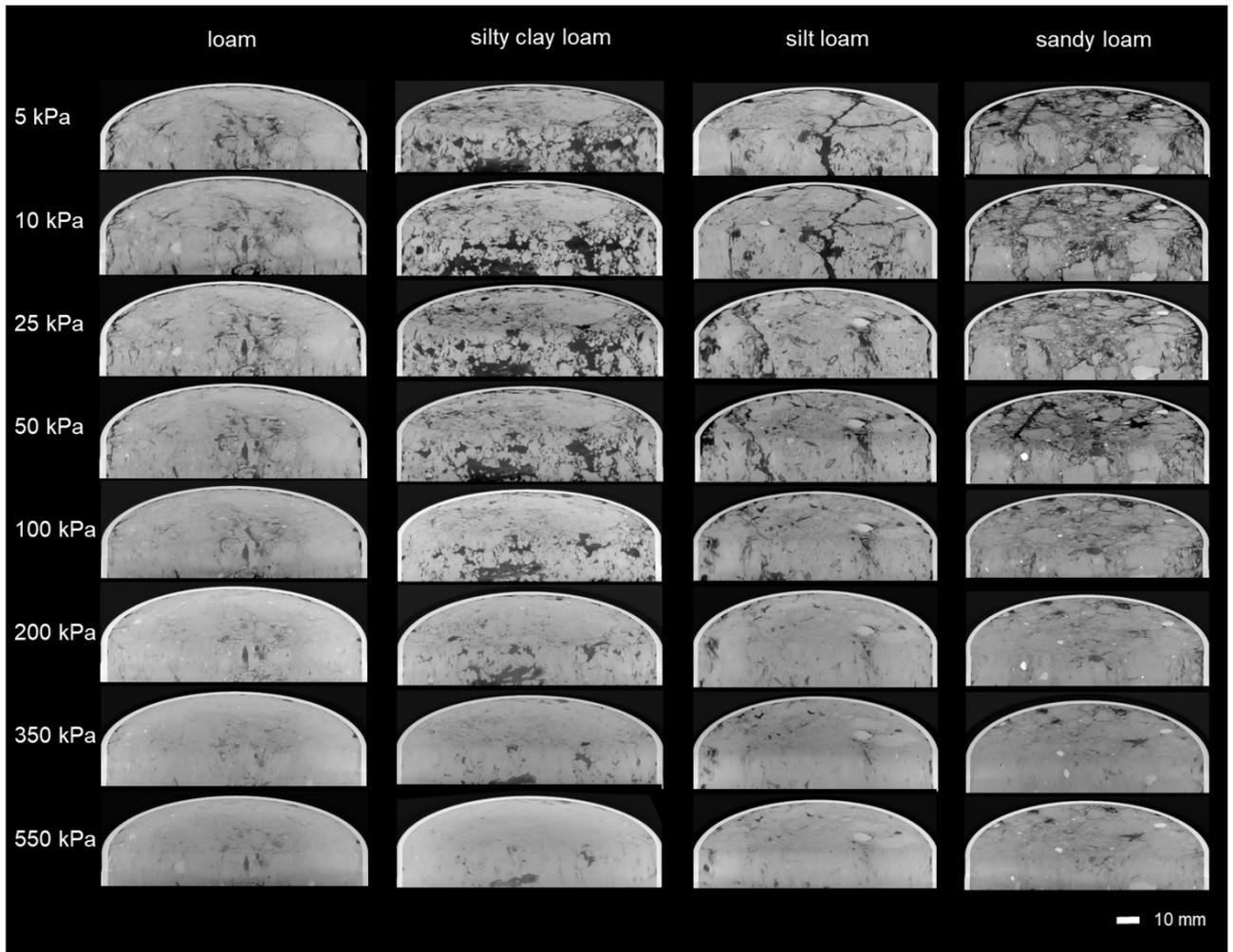


Figure 8: Critical stress ranges for (A) loam, (B) silty clay loam, (C) silt loam and (D) sandy loam based on precompression stress (black circle), critical stress values of macroporosity and pore connectivity (grey circle), and optimum values for biopores, grain yield and straw yield (white circles). The dotted vertical lines indicate the lower and upper limit of

the critical stress range for a site. The dashed horizontal lines (bars) indicate the spread (min-max) of the critical stress values among replicates for a given parameter.

Appendix 1: Examples of CT cross sectional images from load application for a load of 5, 10, 25, 50, 100, 200, 350 and 550 kPa loam (L), silty clay loam (SICL), silt loam (SIL) and sandy loam (SL).



Appendix 2: Dry bulk density (BD_{xi}), macroporosity, pore connectivity, and logarithmic precompression stress ($\log \sigma_P$) for loam (L), silty clay loam (SICL), silt loam (SIL) and sandy loam (SL). Statistically significant differences ($p \leq 0.05$) are indicated by lower case (load step within each site), and upper case letters (sites within each load step).

Parameter	Texture	Load step (kPa)									$\log \sigma_P$
		5	10	25	50	100	200	350	550		
BD_{xi} ($Mg\ m^{-3}$)	L	1.37 aC	1.42 abC	1.45 abC	1.51 acC	1.59 cdC	1.65 deB	1.72 efB	1.77 fB	1.65 a	
	SICL	1.20 aB	1.21 abB	1.24 abB	1.32 abB	1.45 bcB	1.56 cdB	1.65 deB	1.72 eB	1.76 a	
	SIL	1.23 bB	1.27 abB	1.30 abB	1.36 abcB	1.44 abcB	1.53 acB	1.58 cB	1.62 cB	1.85 a	
	SL	0.76 aA	0.78 abA	0.83 bA	0.87 bA	0.97 cdA	1.04 deA	1.12 efA	1.14 eA	1.66 a	
Macroporosity (-)	L	0.09 aA	0.08 aAB	0.07 aA	0.05 aA	0.01 bA	0.00 bA	0.01 bB	0.01 bA		
	SICL	0.23 aB	0.22 abCB	0.21 abC	0.18 aB	0.12 cB	0.06 dB	0.02 dAB	0.02 dA		
	SIL	0.12 aA	0.12 aAB	0.10 abAB	0.08 abcA	0.06 abcAB	0.03 bcAB	0.01 cAB	0.00 cA		
	SL	0.23 aB	0.23 aC	0.18 aBC	0.12 abAB	0.05 bcA	0.04 bcAB	0.03 bcAB	0.01 cA		
Pore Connectivity	L	0.76 aA	0.81 aA	0.74 aC	0.61 aA	0.34 bA	0.17 bA	0.15 bA	0.13 bA		
	SICL	0.96 aB	0.96 aB	0.96 aB	0.92 abB	0.82 bA	0.50 cB	0.14 dA	0.10 dA		

(-)	SIL	0.81	aAB	0.81	aA	0.75	aAC	0.66	aA	0.51	abA	0.25	bcAB	0.12	cA	0.15	cA
	SL	0.94	aB	0.95	aAB	0.92	abAB	0.85	abAB	0.48	bcA	0.25	cAB	0.20	cA	0.14	cA

Appendix 3: Biopores from *Lumbricus terrestris*, grain yield, straw yield, and above ground biomass of *Hordeum vulgare* for different dry bulk densities (BD) for loam (L), silty clay loam (SICL), silt loam (SIL) and sandy loam (SL). Statistically significant differences ($p \leq 0.05$) are indicated by lower case letters.

Texture	BD (Mg m ⁻³)	Biopores (m ²)	Grain yield (g m ²)	Straw yield (g m ²)
L	1.42	65 a	630 a	927 a
L	1.49	56 a	525 ab	949 a
L	1.56	61 a	466 abc	944 a
L	1.63	68 a	544 abc	792 abc
L	1.70	51 a	239 bc	552 bc
L	1.77	61 a	166 c	430 c
SICL	1.28	66 ab	308 a	507 a
SICL	1.35	97 a	277 a	478 a
SICL	1.42	71 ab	256 a	433 a
SICL	1.49	82 ab	72 b	279 b
SICL	1.56	62 ab	31 b	261 b
SICL	1.63	41 b	13 b	157 b
SIL	1.21	95 a	301 a	660 a
SIL	1.28	107 a	333 a	674 a
SIL	1.35	106 a	288 a	676 a
SIL	1.42	96 a	333 a	682 a
SIL	1.49	107 a	397 a	745 a
SIL	1.56	86 a	182 a	453 b
SL	0.72	92 abc	416 a	653 ab
SL	0.78	109 a	520 a	655 ab

SL	0.85	106 a	526 a	673 a
SL	0.92	95 abc	593 a	699 a
SL	1.00	76 bc	570 a	708 a
SL	1.07	71 c	357 a	508 b
