"Bound" residues from biomass and CO₂ in soils – formation, fate and stability during biotic incubation

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This thesis is dedicated to the memory of my father, Prof. Dr. habil. Grzegorz Nowak (1946 – 2002), who inspired my fascination in environmental science

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LIST OF ABBREVIATIONS

¹³C Labelled (stable isotope)

¹⁴C Labelled (radioactive isotope)

2,4-Dichlorophenoxyacetic acid

2,4-DCP 2,4-Dichlorophenol

2-OH-ibu Hydroxyibuprofen

a (for FA) Anteiso

AA Amino Acids
AC Acetylchloride

ACN Acetone
Ala Alanine

ASE Accelerated Solvent Extraction

Asp/Asg Aspartate/Asparagine

BAME Bacterial Acid Methyl Esters

bioAA Biomass Amino Acids

br (for FA) Branched

C_{org} Organic Carbon

cy (for FA) cyclopropyl

DCM Dichloromethane

Dept Department

DT₅₀ Half-life in soil

DW Dry Weight

EA-C-irMS Elemental Analyser-Combustion-isotope ratio Mass Spectrometry

EDTA Ethylenediaminetetraacetic acid

FA Fatty Acids

FAME Fatty Acid Methyl Esther

FAME 21:0 Heneicosanoate methyl ester

GC-C-irMS Gas Chromatography Combustion-isotope ratio Mass Spectrometry

GC-MS Gas Chromatography Mass Spectrometry

Glu/Gln Glutamate/Glutamine

Gly Glycine

i (for FA) Iso

Ibu Ibuprofen

Ile Isoleucine
IP Isopropanol

IUPAC The International Union of Pure and Applied Chemistry

Koc, logK_{oc} Soil-sorption coefficient

Kow, logK_{ow} Octanol/water partition coefficient

Leu Leucine Lys Lysine

M Molar concentration
Me (for FA) Mid-chain branched

MeOH Methanol

MM Minimum Medium Na₂CO₃ Sodium bicarbonate

n.d. not detectable

NER Non-Extractable Residues
NH₄OH Ammonium Hydroxide

 $\begin{array}{ll} nm & Nanometre \\ \\ N_{Total} & Total \ Nitrogen \\ OD & Optical \ Density \end{array}$

OECD Organization for Economic Co-operation and Development

OM Organic Matter

PAH Polycyclic Aromatic Hydrocarbons

PB Phosphate Buffer

PEG 600 Polyethylenglycol 600

Phe Phenylalanine

PLFA Phospholipid Fatty Acids

Pro Proline Ser Serine

SOM Soil Organic Matter

SPE Solid Phase Extraction

tAA Total Amino Acids

TCC Tricarboxylic acid Cycle

tFA Total Fatty Acids

TFAA Trifluroacetic Acid Anhydride

Threo Threonine

TM Tryptone Medium

TMCS Trimethylchlorosilane

TNT Trinitrotoluene

TOC Total Organic Carbon

UFZ Centre for Environmental Research

USA United States of America

UV/VIS Ultraviolet-Visible Spectrophotometry

Val Valine

WHC Water Holding Capacity

WWTP Wastewater treatment plant

SUMMARY X

SUMMARY

Biodegradation of organic pollutants in soil generally results in the formation of metabolites, microbial biomass, mineralisation products and bound or non-extractable residues (NER). It is speculated that NER can pose a risk for humans due to the remobilisation and further distribution of active parent compounds or their metabolites over the food web (BARRACLOUGH ET AL., 2005). Although there are many studies on NER formation available, their chemical structures are still unknown. However, without this knowledge, a proper risk assessment for the pollutant and the related NER during its transformation in soil is impossible. Part of the NER may be biogenic, since the pollutant-derived C and CO₂ released from its mineralisation are assimilated by microorganisms into their cellular components [e.g. fatty acids (FA) and amino acids (AA)], which are subsequently incorporated into soil organic matter (SOM) after cell death

In order to study the microbial biomass contribution to NER formation, soil was incubated with either ¹³C-labelled 2,4-Dichlorophenoxy acetic acid (2,4-D) or ibuprofen (Ibu) for 64 and 90 days, respectively. At different sampling dates, the soil was analysed for the presence of the ¹³C label in FA and AA. The ¹³C label assimilated in biomolecules was determined in both the living and the non-living SOM fractions. Moreover, for investigating the relevance of heterotrophic CO₂ fixation for the incorporation of ¹³C into biogenic NER in soil and to distinguish between formations of NER directly from 2,4-D and via CO₂, an experiment with unlabelled 2,4-D and labelled CO₂ was conducted. In addition, to understand both the processes of the biogenic NER formation and the incorporation of ¹³C label into FA and AA from either ¹³C₆-2,4-D or ¹³CO₂ in the complex soil environment, a simple biological model system with the known 2,4-D degrader (*Cupriavidus necator* JMP 134), was studied.

After 64 days of ¹³C₆-2,4-D incubation in soil, the total contents of ¹³C detected in AA in SOM indicated that 44% of the initially applied ¹³C₆-2,4-D equivalents had been converted to microbial biomass and finally to biogenic residues. The intermediate maximum of ¹³C-FA in SOM indicated a 20% conversion of ¹³C₆-2,4-D to biomass. However, ¹³C-FA in the non-living SOM fraction decreased to 50% indicating their metabolisation and their further distribution within the food web. Contrary to the ¹³C-FA, ¹³C-AA in non-living SOM pool were surprisingly stable.

The soil experiment with ¹³CO₂ showed that after 16 days of incubation, the heterotrophic CO₂ fixation was relevant in the assimilation of ¹³C label into biomass components, in particular FA. In addition, the liquid culture experiment with *C. necator* JMP 134 demonstrated the high importance of CO₂ fixation in the incorporation of ¹³C label into

SUMMARY XI

the biomass components in the later phase of incubation. In this study about 4% of 13 C derived from 13 C₆-2,4-D was assimilated in the biomass of this bacterial strain via CO₂ fixation.

A total content of biogenic NER of 54% of ¹³C₆-ibu equivalents at the end of the ¹³C₆-ibu soil experiment was indicated by the amount of ¹³C found in AA in SOM. From the maximum content of ¹³C-FA detected in this experiment, at least 24% (of ¹³C₆-ibu equivalents) biomass must have been formed. Contrary to the ¹³C₆-2,4-D soil incubation experiment, the ¹³C-FA remained stable until the end. However, the formation of NER in the soil incubated with ¹³C₆-ibu started much later than that in the ¹³C₆-2,4-D study, due to the longer lag phase and the interruption of the incubation before the destabilisation of biogenic NER. The ¹³C-AA in SOM remained also stable as found in ¹³C₆-2,4-D experiment.

The results from these soil biodegradation experiments provide the first evidence that nearly all NER both from 2,4-D and Ibu are biogenic, containing only natural microbial biomass components stabilised in SOM. However, biogenic residues are formed if the respective organic contaminant is readily degraded by microorganisms under significant formation of CO_2 . Depending on the yield coefficients of C conversion into biomass, we expect ratios of biomass plus biogenic residues to CO_2 of about 0.2 to 1. In the $^{13}C_6$ -2,4-D study, the ratio was ~ 0.8 and in the $^{13}C_6$ -ibu was 1.2.

Biogenic residues are clearly excluded from the definition of NER according to IUPAC. Due to biogenic residue formation, the potential risk of NER from the readily metabolised organic contaminants in soils is thus highly overestimated in many cases. Therefore, the formation of biogenic residues must be taken into consideration, when determining the mass balances of contaminants during their microbial degradation in soil.

ZUSAMMENFASSUNG

Biologischer Abbau von organischen Schadstoffen im Boden führt generell zur Mineralisierung und zur Bildung von Metaboliten und gebundenen oder nicht-extrahierbaren Rückständen (NER). Es wird spekuliert, dass NER ein Risiko für den Menschen darstellen können, da sie möglicherweise remobilisiert werden können und die aktiven Ausgangsverbindungen oder deren Metabolite weiter in der Nahrungskette verbreitet werden können (BARRACLOUGH AT AL., 2005). Trotz vieler Studien über die NER-Bildung sind ihre genauen chemischen Strukturen noch nicht bekannt. Allerdings ist ohne dieses Wissen eine korrekte Risikobewertung für den Schadstoff und die damit verbundenen NER während der Transformation im Boden unmöglich. Ein Teil der NER könnten biogen sein, weil der schadstoffbürtige Kohlenstoff und das durch Mineralisierung freigesetzte CO₂ von Mikroorganismen in ihre zellulären Bestandteile eingebaut werden, z.B. Fettsäuren (FS) und Aminosäuren (AS), die nach dem Zelltod in der organischen Bodensubstanz (OBS) festgelegt werden können.

Um den Beitrag der mikrobiellen Biomasse zur NER-Bildung zu untersuchen, wurde ein Boden mit ¹³C-markierter 2,4-Dichlorphenoxyessigsäure (2,4-D) bzw. Ibuprofen (Ibu) für 64 bzw. Tagen inkubiert. Die Bodenproben wurden zu verschiedenen Probenahmezeitpunkten auf die ¹³C-Anreicherung in FS und AS analysiert. Die in den Biomolekülen gebundene ¹³C-Markierung wurde sowohl in der lebenden als auch in der toten OBS Fraktionen bestimmt. Außerdem wurden Experimente mit unmarkiertem 2,4-D und markiertem CO2 durchgeführt, um den Beitrag der heterotrophen CO2-Fixierung zum Einbau von ¹³C in biogene NER in Böden zu untersuchen. Damit kann zwischen direkter NER-Bildung aus 2,4-D und indirekter über CO₂ unterschieden werden. Zudem wurde ein einfaches biologisches Modellsystem mit dem bekannten 2,4-D Abbauer Cupriavidus necator JMP 134 untersucht, um sowohl die an der Bildung von biogenen NER beteiligten Prozesse als auch den Einbau von ¹³C aus ¹³C₆-2,4-D bzw. ¹³CO₂ in FS und AS in der komplexen Bodenumwelt besser zu verstehen.

Die Menge der gesamten ¹³C-AS in SOM nach 64-tägiger Inkubation mit ¹³C₆-2,4-D im Boden deutet darauf hin, dass 44% der anfänglichen ¹³C₆-2,4-D-Äquivalente in mikrobielle Biomasse und schließlich in biogene Rückstände umgesetzt wurden. Das zwischenzeitliche Maximum der ¹³C-FS in der organischen Bodensubstanz (SOM) zeigte, dass 20% des ¹³C₆-2,4-D in mikrobielle Biomasse umgesetzt wurden. Allerdings sanken die ¹³C-FS in der toten SOM Fraktion wegen ihrer Metabolisierung und ihrer weiteren Verbreitung im Nahrungsnetz

auf 50% ab. Im Gegensatz zu den ¹³C-FS waren ¹³C-AS im toten SOM Pool überraschend stabil.

Das Bodenexperiment mit ¹³CO₂ zeigte nach 16-tägiger Inkubation die Relevanz der heterotrophen CO₂-Fixierung bei der Assimilation von ¹³C in die Bestandteile der mikrobiellen Biomasse, insbesondere in die FA. Zudem wies das Flüssigkultur Experiment mit *C. necator* JMP 134 auf die hohe Bedeutung der CO₂-Fixierung für den Einbau der ¹³C-Markierung in die Biomasse-Bestandteile während der späteren Phase der Inkubation hin. In dieser Studie wurde etwa 4% des 2,4-D-bürtigen Kohlenstoffs über CO₂-Fixierung in die Biomasse dieses Bakterienstamms assimiliert.

Am Ende des Bodenexperiments mit ¹³C₆-Ibu wies die Menge der ¹³C-AS in der SOM auf einen gesamten biogenen NER-Gehalt von 54% der ¹³C₆-Ibu-Äquivalente hin. Auf der Basis des maximalen ¹³C-FA Gehalts, der in diesem Experiment gefunden wurde, müssen mindestens 24% (der ¹³C₆-Ibu-Äquivalente) Biomasse gebildet worden sein. Im Gegensatz zur Inkubation des Bodens mit ¹³C₆-2,4-D blieben die ¹³C-FA im Falle von ¹³C₆-Ibu bis zum Ende des Experiments stabil. Allerdings begann die NER-Bildung in dem Boden, der mit ¹³C₆-Ibu inkubiert wurde, viel später als in der ¹³C₆-2,4-D-Studie, was auf der längeren Lag-Phase und dem Abbruch der Inkubation vor der Destabilisierung der biogenen NER beruht. Die ¹³C-AA in der SOM blieben ebenfalls stabil, wie bereits im ¹³C₆-2,4-D Experiment.

Die Ergebnisse dieser Abbauexperimente im Boden liefern den ersten Beweis dafür, dass sowohl im Falle des 2,4-D als auch des Ibu fast die gesamten NER biogen sind. Sie enthalten deshalb nur natürliche Bestandteile der mikrobiellen Biomasse, die in der SOM stabilisiert sind. Die Bildung biogener Rückstände ist vor allem dann von Bedeutung, wenn der jeweilige organische Schadstoff durch Mikroorganismen leicht abbaubar ist und dabei erhebliche Mengen CO_2 gebildet werden. Abhängig vom Ertragskoeffizienten der C Umwandlung in Biomasse, erwarten wir Verhältnisse von Biomasse plus biogene Rückstände zu CO_2 von etwa 0,2 bis 1. In der $^{13}C_6$ -2,4-D-Studie war dieses Verhältnis \sim 0,8 und in der $^{13}C_6$ -ibu betrug es 1,2.

Biogene Rückstände sind eindeutig von der NER-Definition nach IUPAC ausgeschlossen. Da sie aber analytisch in der Regel nicht von direkt schadstoffbürtigen NER abgetrennt werden, wird das potenzielle Risiko von NER aufgrund der biogenen Rückstandsbildung aus leicht abbaubaren organischen Schadstoffen in Böden oft weit überschätzt Deswegen sollte die biogene Rückstandsbildung bei der Erstellung von Massenbilanzen für Schadstoffe während ihres mikrobiellen Abbaus im Boden berücksichtigt werden.

1 INTRODUCTION

Soil as a very complex medium with a large number of interaction sites is a major sink for xenobiotics (Kästner, 2000), which are released in huge amounts due to human activities. Organic contaminants, which enter the soil environment harbouring an enormous number and high diversity of bacteria, are generally subject to microbial degradation. Microbial degradation of these contaminants is often the main pathway of their disappearance from soils (Waldman and Shevah, 1993; Edgehill and Fin, 1983).

Microbial degradation of organic pollutants in soil generally results in the formation of metabolites, microbial biomass, mineralisation products (CO₂ and H₂O) and bound or Non-Extractable Residues (NER; KÄSTNER, 2000). The process of the NER formation from organic contaminants as well as their stability in soils over time has gained strong interest as a subject of research in the recent years. It is generally believed that NER are formed as a result of the various physicochemical interactions between parent compound or its metabolites and soil organic matter [(SOM); BOLLAG ET AL., 1992; SENESI, 1992; VERSTRAETE AND Devliegher, 1996; Führ, 1998; Alexander, 2000; Gevao et al., 2000; Loiseau and BARRIUSO, 2002; MORDAUNT ET AL., 2005]. The formation of these residues in soil decreases the bioavailability of contaminants and thereby reduces their toxicity (FÜHR, 1998; NORTHCOTT AND JONES 2000). Enhanced organic contaminants transformation into NER is thus actually suggested as highly desirable process for the natural elimination of the contaminants from soils (BOLLAG, 1992; BERRY AND BOYD, 1985; VERSTRAETE AND DEVLIEGHER, 1996). On the other hand, recent studies on the stability of these residues have revealed that they are not always irreversibly bound to SOM and can be remobilised (BURAUEL AND FÜHR, 2000; BOIVIN ET AL., 2005; GEVAO ET AL., 2005; LERCH ET AL., 2009B). Thus toxic acive compounds and their derivatives of unknown structure can be spread to other compartments of the environment, such as surface- and ground waters, which may finally result in their distribution over the food web (BARRACLOUGH ET AL., 2005). In terms of the public health safety these potential risks related to NER formation in soils make this natural detoxification process one of the hot topics of the world-wide scientific debate.

Most of the available studies on the formation and the fate of NER are performed with radioactive tracer compounds. However, this labelling technique only allows the quantitative distribution analyses and does not provide any information of the chemical nature of tested compounds (RICHNOW ET AL., 1999). Therefore it is not clear, how NER formed during biodegradation of contaminants in soils are composed and if they really pose a risk for the environment and humans. Stable isotopes allow tracing the flux of pollutant-derived C into

different chemical structures using more sophisticated analytical techniques such as gaschromatography-mass spectrometry (RICHNOW ET AL., 1999). Therefore, detailed elucidation of the formation and chemical nature of NER during the biodegradation of organic pollutants in soil is possible using these isotopes as a tracer.

Some part of NER formed during the microbial degradation of organic contaminants may be biogenic. It is generally known that a wide variety of microorganisms utilise the pollutant-derived C during its biodegradation in soils for their growth and biomass formation (KÄSTNER AND RICHNOW, 2001). The biomass components are incorporated into the non-living SOM fraction after their death and cell lysis (KINDLER ET AL., 2006, 2009), and thus may contribute to the formation of "biogenic" form of NER.

Contrary to the xenobiotic-derived NER, biogenic residues are composed only of non-toxic microbial components stabilised in the SOM pool, which do not pose any risks for the environment. However, soil is a very complex system, which is not fully understood, thus it is difficult to analyse the NER in SOM and the mechanisms of their formation are not elucidated yet. In addition, all available studies on NER formation during biodegradation of organic contaminants are limited to quantitative analyses. Therefore, it is essential to study the NER structure in detail in order to assess properly the risks related with the NER formation during biodegradation of contaminants in soil.

The overall objective of the present study was to trace the NER formation during biodegradation of organic contaminants in soil, and 2,4-Dichlorophenoxyacetic acid (2,4-D) and Ibuprofen (Ibu) were selected as the model compounds. In order to investigate in detail the fate of these contaminants in the complex soil system, stable isotope tracers (13 C) were used for the proper quantitative analyses and for tracing the flux of the pollutant-derived C. The transformation of 13 C label from 13 C₆-2,4-D or 13 C₆-ibu into CO₂, biomass components, metabolites, biogenic NER and non-biogenic NER was determined. For assessing the risks related to the NER formation from 13 C₆-2,4-D or 13 C₆-ibu, the estimated biogenic NER contents were compared with the non-biogenic NER amounts.

2 STATE OF THE ART

2.1 Complexity of the soil system

Organic contaminants, which enter the soil system, are subject to various interactions within the complex soil matrix (KÄSTNER, 2000). Basically, these contaminants can be degraded by microorganisms, immobilised in the form of non-extractable residues (NER) via binding processes to soil components, volatilised, leached to the groundwater or taken up by living organisms (**Figure 1**). Besides the physico-chemical properties of the contaminant itself, also the soil components, which create the soil environment, affect the fate of contaminants in this system. These respective components present in soil define the overall structure and the physico-chemical properties of this whole complex system.

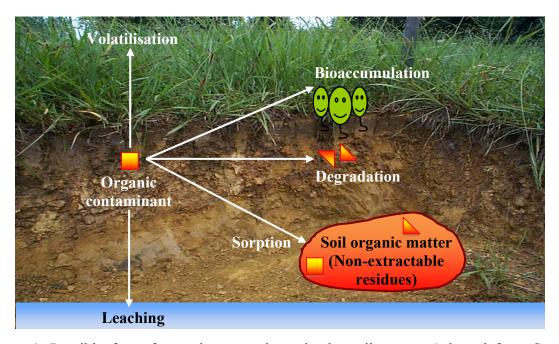


Figure 1. Possible fate of organic contaminant in the soil system (adapted from SEMPLE ET AL., 2003; STOKES ET AL., 2006)

Soil is a heterogeneous system consisting of four phases; the volumetric proportions of which varies over a wide range depending on the soil type (SIMS ET AL., 1990):

- 1. *inorganic solids* (38–45%): which are represented by larger-sized quartz, sand, silt and very fine clay. The larger mineral particles ensure the drainage in soils via formation of macropores, which enable both the transport of water and the distribution of gases within components of soil aggregates;
- 2. *organic solids* (1–12%): so-called soil organic matter (SOM). Although this component makes up the smallest part of the soil system, it is very important, because it affects all physical, chemical and biological soil properties (STEVENSON,

- 1994). Furthermore, it plays a crucial role in the protection of soil from degradation and erosion (PICCOLO, 1996);
- 3. *soil water* (15–35%): is present in pore and capillary spaces. Water in pore spaces is a solvent for nutrients and salts, which support microbial activity in soil (HAIDER AND SCHÄFFER, 2009). Therefore, microbial biomass tends to concentrate along water flow paths in the soil (VINTHER ET AL., 1999; BUNDT ET AL., 2003). Moreover, it affects the soil aeration status, the soil water osmotic pressure and the pH of soil solution (PAUL AND CLARK; 1989);
- 4. *soil gases* (15–35%): are distributed within pore spaces and are necessary for microbial and plant root respiration.

The distribution and type of solid particles in the soil system defines its overall structure and thus the active surface area affecting the fate of organic compounds (GAVRILESCU, 2005). The interaction of SOM with clay in presence of polyvalent cations (e.g. Fe³⁺, Ca²⁺, Al³⁺) results in the formation of stable clay-organic complexes and thus microaggregates (BRONICK AND LAL, 2005; HAIDER AND SCHÄFFER, 2009). Microaggregates are joined with larger soil particles (e.g. quartz and sand) forming macroaggregates (TISDALL, 1996). This aggregation process plays an important role in shaping the physical, chemical and biological properties of the soil system. Adhesive exopolymers excreted by living microorganisms on soil particles and fungal hyphae on roots stabilise macroaggregates in the form of "sticky string bags" (KÄSTNER, 2000, see **Figure 2**).

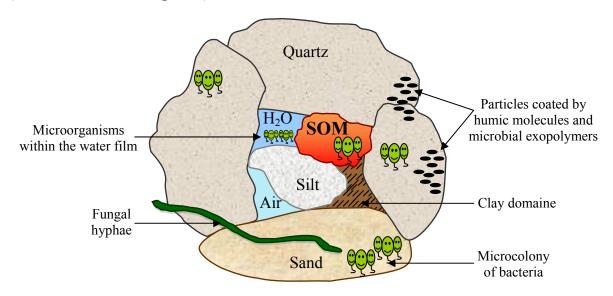


Figure 2. Aggregation of soil constituents (KÄSTNER 2000, ATLAS AND BARTHA 1997)

The water content, the presence of soil biota and the composition of solid components (in particular clays and SOM) affect the aggregate size (BRONICK AND LAL, 2005). Pore sizes and

relative proportions of water and air in the soil system have an impact on the mobility of contaminants (SIMS ET AL., 1990).

The main actors in the interactions with organic contaminants in soils are small-sized soil particles clays and soil organic matter (CALDERBANK, 1989; SIMS ET AL., 1990; STEVENSON, 1994). Clay particles have a large reactive charged surface, thus they are believed to be involved in physico-chemical interactions with organic contaminants (SIMS ET AL., 1990). Interactions with pesticides and other organic chemicals have also been reported (WHITE, 1976; THENG, 1982; MORTLAND, 1986; ZIELKE ET AL., 1989).

SOM possessing many highly reactive functional groups (SIMS ET AL., 1990) was proposed as the dominant factor affecting the organic contaminants interactions within soil systems (BRUSSEAU ET AL., 1991; CORNELISSEN ET AL., 1998). This soil fraction, called "non-living" SOM is a complex mixture of organic materials derived from fragments and decomposed products of plant residues (e.g. proteins, starch, polysaccharides, lignins, lipids), microbial cell components (e.g. lipids, carbohydrates, proteins) and humic substances (KÖGEL-KNABNER, 2002; STEVENSON, 1994; SOLLINS ET AL., 1996; ZECH ET AL., 1997). Because of their small size, microorganisms cannot be separated from SOM and thus form a living fraction of it. The living fraction (1–5% of total SOM) containing a wide variety of microorganisms, higher animals and plant roots (KLEBER ET AL., 1998; OADES, 1995), constitutes a primary source for SOM formation (KELLEHER ET AL., 2006).

Humic substances are considered to be dominant components of SOM (50–60%, PAUL AND CLARK, 1996, SCHNITZER, 1978; STEVENSON, 1994) and show molecular weights ranging from a few hundred to several hundred thousand Daltons (Huang et al., 2003). They are believed to form from decaying plant biomass in a process called "humification" (Haider, 1998). The main components of the dead plant biomass, high-molecular-weight polysaccharides and lignin are supposed to affect the size of humic substances (Haider, 1998). Humic substances are commonly classified depending on their solubility in alkaline or acidic solutions into humins (insoluble in both solvents), fulvic acids (soluble in acid and alkali) and humic acids [(soluble in alkali and insoluble in acid); Northcott and Jones, 2000; Stevenson, 1994]. This classification into these three fractions is only operational and does not indicate any chemical behaviour or structure of humic acids (Hayes et al., 1989; Stevenson, 1994). Due to the fact that humic substances are very complex heterogenous organic compounds, their overall chemical structure has not been clarified yet (Ziechmann, 1994). Pauli (1967) suggested a model, in which humic substances are present as complex soil colloids with micellar structure. In this model, hydrophobic aromatic and aliphatic

building blocks, which are linked by covalent bonding, carry reactive functional groups with a hydrophilic character. Both hydrophilic and hydrophobic sites of humic substances are commonly considered to be involved in various interactions with organic contaminants (Kästner, 2000). Other authors have recently described humic substances as supramolecular associations of low-molecular-mass organic biomolecules (Sutton and Sposito, 2005), which have been excluded from traditional definitions of humic substances (Stevenson, 1994; Haider, 1998). These low-molecular-mass organic molecules include for example branched and linear alkanes, alkenes, fatty acids, dicarboxylic acids, and long chain alcohols/ethers, ketones/aldehydes, amino acids and esters (Fabbri et al., 1996; Schulten, 1999; Schnitzer, 2000; Kramer et al., 2001; Chefetz et al., 2002; Grasset et al., 2002; Stenson et al., 2002; Mugo and Bottaro, 2004). In addition, the hydrophobic properties of humic substances, which have been suggested in the proposed models of humic substances, have not yet been identified (Sutton and Sposito, 2005). Therefore, Sutton and to redefine the concept of "humification".

KÄSTNER (2000) proposed an other definition of "humification", which is in reality a prolonged stabilisation process of organic components from non-living components of SOM leading to the formation of refractory SOM. This refractory SOM is formed by complex processes, in which metabolites and fragments of microbial and plant biomass metabolised by microorganisms rearrange to form macromolecular aggregates. The continuous process of biomolecules incorporation results in increasing molecular weight and thus in the formation of macromolecules, which are characterised commonly as humins, humic acids and fulvic acids. These macromolecules rearrange with each other, fragments of biomass and clay minerals and finally form larger aggregates. In cultivated soils, 50–75% of SOM is associated with clay-sized organo-mineral particles (Christensen, 2001). Due to the fact that SOM is tightly bound to clays in the form of clay-organic complexes (HAIDER AND SCHÄFFER, 2009), it is difficult to distinguish between the contributions of clay and SOM to the interactions with organic contaminants (KÄSTNER, 2000).

2.2 Microbial activity in the soil

Microorganisms are known to play a crucial role in biogeochemical cycles and in sustainable development of the biosphere (VAN HAMME, 2004; ADRIANO AND BOLLAG, 1999). Soil microbes need to cover their energy expenditure by uptake of the products of enzymatic breakdown of organic substrate in order to survive in an active stage (EKSCHMITT ET AL., 2005). The microbial transformation of organic compounds into various metabolites, which

become integral parts of the soil after stabilisation processes, can ultimately define the overall structure and thus the physico-chemical properties of this system (YOUNG AND CRAWFORD, 2004).

Microorganisms, which might be detected in the soil system, are eubacteria, actinomycetes, archaea, fungi, algae, protozoa and viruses (KÄSTNER, 2000). Scientists have estimated that in one gram of fertile soil between 5000–7000 different bacterial species can exist and that their populations can often exceed one hundred million individuals (GAVRILESCU, 2005). The soil systems are heterogeneous habitats, therefore it is difficult to characterise their microbial community in detail and only 1–10% of the microscopically visible soil bacteria were estimated to be actually culturable by artificial media (PICKUP, 1991; LEADBETTER, 1997).

Due to the high abundance and diversity of microorganisms present in soils, most of the natural and anthropogenic organic compounds are subject to microbial attack (Kästner, 2000). Owing to the presence of small-sized microorganisms (few µm) at hardly accessible places within the soil system, organic contaminant, even if occluded within pore spaces, can be degraded. Therefore, soil microbes are the major agents in the disappearance of organic contaminants from soil in the so-called biodegradation process. This decontamination mechanism is considered to be environmentally friendly (Golovleva et al., 1990; Singh and Walker, 2006; Müller et al., 2007), because they are capable of degrading of a vast diversity of compounds without hazardous by products. Among the microbial communities, bacteria, fungi, archaea and actinomycetes are the main organic compound degraders (Häggblom, 1992; Müller et al., 2007).

The microbial degradation (biodegradation) of organic compound in soil is dependent on many factors (BOLLAG AND LIU, 1990; SIMS AT AL., 1990; GAVRILESCU, 2005):

- *Soil conditions:* temperature, aeration, pH, moisture and SOM content, that influence the activity of soil microorganisms, their size and diversity. In addition, the organic compound availability for microorganisms is closely linked to the soil conditions (e.g. decrease by sorption to SOM or lower pH);
- Organic compound characteristics, which include solubility in water, tendency to adsorb to the soil matrix and persistence in the soil (half-life);
- Frequency of organic compound application: repeated application of a compound may result in the development of a microbial community capable of the organic compound degradation;

• Concentration of organic compound: very low concentrations may not provide maintenance energy levels sufficient for microorganisms. Very high concentrations may either be toxic and thus inhibit the microbial activity or, as an easily available substrate, stimulate their growth.

The organic compound characteristics are important in determining its transformation, fate and persistence in the soil system. Among others these characteristics include (AGA AND THURMAN, 2001; ANDREU AND PICÓ, 2004; GAVRILESCU, 2005):

- water solubility, which strongly depends on temperature and pH of the solution. It controls the mobility of organic compound in soil;
- soil-adsorption coefficient (K_{oc} , log K_{oc}), which describes the tendency of a compound to be adsorbed to soil particles. The higher the K_{oc} value, the more strongly the contaminant is sorbed, which leads to reduced availability and mobility;
- octanol/water partition coefficient (Kow, log K_{ow}), which is related to water solubility and soil/sediment sorption coefficients (K_{oc}), for instance: log $K_{ow} < 2.5$ indicates low sorption potential, values log K_{ow} between 2.5 and 4.0 indicate intermediate sorption potential, whereas $log K_{ow} > 4.0$ indicates high sorption potential (JONES-LEPP AND STEVENS, 2007). In addition, K_{ow} , log K_{ow} is widely accepted to describe bioconcentration in aquatic organisms, which is indicative of the tendency for bioaccumulation of a contaminant in organisms;
- half-life in soil (DT_{50}), is the amount of time necessary for disappearance of 50 % of parent compound from soil.

Depending on the chemical nature of a compound, the degradation rate can be different. Easily degradable compounds are usually degraded immediately accompanied by the formation of biomass and mineralisation products (H₂O and CO₂), whereas readily and hardly degradable xenobiotics are degraded at slow rates from the beginning or at higher rates, but after only prolonged phases of adaptation (Kästner, 2000). During degradation, the C derived from the pollutant is used as energy and C source by microorganisms to form their biomass components (Müller et al., 2007). However, many organic contaminants are also degraded cometabolically (Müller et al., 2007). In this cometabolic process, the organic contaminants do not serve as C or energy source and are metabolised together with another substrate used for growth (SIMS AT Al., 1990; Häggblom, 1992). This type of transformation is frequently based on metabolic reactions catalysed by extracellular enzymes present in soils (Müller et al., 2007; SIMS et al., 1990).

2.2.1 Natural organic compounds biodegradation in soil

The high-molecular-weight organic compounds derived from plant and microbial biomass can be mineralised to CO₂ and H₂O with the formation of microbial biomass (see **Figure 3**). After the death of microorganisms, low-molecular-weight microbial compounds are incorporated into SOM and form so-called biogenic residues. Thereafter, these residues are stabilised in SOM, which leads to the formation of refractory SOM (humic substances in "humification"). Biogenic residues, their metabolites and biomass components are also probably parent materials for renewed degradation and conversion reactions (KÄSTNER AND RICHNOW, 2001).

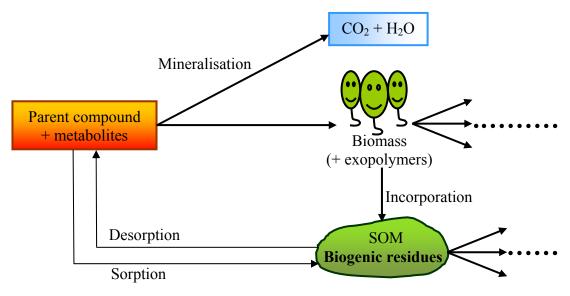


Figure 3. Scheme of C flow during microbial degradation of natural organic compounds in soil (adapted from KÄSTNER AND RICHNOW, 2001)

Formation of biogenic residues (up to 35%) in soil was observed during the biodegradation of plant biomass residues (STOTT ET AL., 1983). In addition, even after a one-year incubation of soil with ¹⁴C-glucose, residues derived from this easily degradable compound were still detected in SOM (10%; BALDOCK ET AL. 1989).

Plant residues are suggested to be the primary source for refractory SOM formation (STEVENSON, 1994; SCHOLES ET AL., 1997; HAIDER, 1998; KÖGEL-KNABNER, 2002). This is caused by the relative high inputs to soil and the fact that plant components, in particular high-molecular-weight compounds (e.g. lignin) are believed to be highly resistant to microbial degradation (STEVENSON, 1994; HAIDER, 1998; KÖGEL-KNABNER, 2002). However, recent findings show that lignin is less persistent than the average of bulk SOM (VON LÜTZOW ET AL., 2008; MARSCHNER ET AL., 2008). The importance of microbial biomass C in the formation of SOM is considered to be minor (SCHOLES AND SCHOLES, 1995), because of both its small pool size and fast turnover (JENKINSON AND LADD, 1981; COLEMAN ET AL., 1983).

However, studies on the molecular composition of SOM proved that plant-derived organic matter (OM) was not stored in arable soils, but was transformed to microbial residues, i.e. carbohydrates and proteins, which were kept in soil by organo-mineral interactions (BOL ET AL., 2009). Several recent studies on SOM showed that the contribution of microbial biomass to SOM formation in the generally accepted range of 1-5% is highly underestimated (SIMPSON ET AL., 2007). For instance, soil incubation with ¹³C-labelled *Escherichia coli* indicated that microbial biomass-derived C significantly contributed to the formation of refractory SOM (LÜDERS ET AL., 2006; KINDLER ET AL., 2006, 2009; MILTNER ET AL., 2009). In this experiment about 56% of the bulk C of *E. coli* was mineralised, the residual 44% was stabilised in the soil after 224 days (KINDLER ET AL., 2006). Further research reported that SOM was predominantly of microbial origin (> 50% of extractable humic acids; SIMPSON ET AL., 2007). The high content of *E. coli* biomass-derived C stabilised in SOM after 224 days thus indicates that microorganisms may form significant amounts of biogenic residues in soil and finally the refractory SOM during biodegradation of organic compounds.

2.2.2 Anthropogenic organic compounds biodegradation in soil

The microbial degradation of organic contaminants in soil is generally understood as their transformation into mineralisation products, metabolites, microbial biomass and non-extractable residues (NER; see mass balance in **Figure 4**).

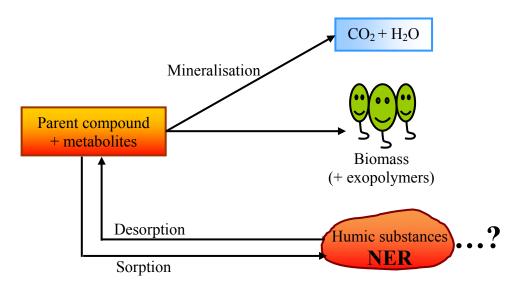


Figure 4. Conventional model of C flow during microbial degradation of organic contaminants in soil (mass balance)

The organic contaminant or products of its partial biodegradation are believed to sorb to SOM via various physico-chemical mechanisms leading to the NER formation (ALEXANDER, 2000; GEVAO ET AL., 2000; GAVRILESCU, 2005; LOISEAU AND BARRIUSO, 2002; MORDAUNT ET AL.,

2005). The sorption of contaminants to SOM is considered as the major factor preventing their complete biodegradation in soil (BÜYÜKSÖNMEZ ET AL., 1999). The mass balance of organic contaminant within soil fractions during its biodegradation is usually determined using radioactive tracers (14C-labelled compounds; GERST AND KLIGER, 1990; KUBIAK ET AL., 1990; RICHNOW, 1999). The rate of microbial transformation of an organic compound is estimated only by the determination of the residual concentrations in soil (FOGARTY AND TUOVINEN, 1991). The analyses of parent compounds and their metabolites remaining in soil solution are usually accomplished using different extraction methods such as batch solvent shaking, Soxhlet extraction, supercritical fluid extraction or accelerated solvent extraction (HAWTHORNE ET AL., 2000; NORTHCOTT AND JONES, 2000) with aqueous or organic solvents. However, these conventional and enhanced solvent extraction techniques do not extract contaminants that are strongly bound or sequestered into the components of soil matrix like SOM (NORTHCOTT AND JONES, 2000).

Therefore, the total amount of unextracted contaminant residues is mostly quantified as ¹⁴CO₂ released from combustion of soil samples (WAIS, 1998; BARRIUSO ET AL., 2008). Using this destructive approach, it is not possible to check if NER are intact contaminants, their metabolites or ¹⁴C accumulated in microbial biomass or in humic substances (BARRIUSO ET AL., 2008). Alternatively, the distribution of the ¹⁴C in humic acids, fulvic acids and the insoluble humins of soil containing non-extractable contaminant residues is often analysed after the extraction of both fulvic and humic acids (NORTHCOTT AND JONES, 2000). A major obstacle for deeper analysis of these residues is that no proper method for the identification of the chemical nature of the transformation products bound to complex soil matrix using radiotracers is available. Therefore, due to the problems with NER analyses, the detailed biodegradation pathways of many organic contaminants in these complex soil systems have not yet been elucidated (FOGARTY AND TUOVINEN, 1991).

2.3 Bioavailability of organic contaminants in soil

Bioavailability of contaminants is the key factor, which controls their overall fate in soil, in particular their biodegradability and toxicity for biota (SEMPLE ET AL., 2007). Bioavailability is affected by many factors such as: the properties of contaminant and soil, aging time in the soil, climate and organisms of concern (KATAYAMA ET AL., 2010). On the whole, the assessment of the bioavailability of contaminants in soil is necessary both for understanding the risks, which may be posed by these contaminants and for the proper choice of the method for soil remediation (SEMPLE ET AL., 2003).

The term bioavailability has been used in many different scientific fields, thus there are many definitions (SEMPLE ET AL., 2007), which generally consider the interactions between an organism and a chemical (SEMPLE ET AL., 2004). Toxicologists term bioavailability as the fraction of chemical absorbed and able to reach systemic circulation in an organism (SEMPLE ET AL., 2004). For instance, the National Research Council reports:

"Bioavailability may represent the fraction of a chemical accessible to an organism for absorption, the rate at which a substance is absorbed into a living system or a measure of the potential to cause a toxic effect" (NRC, 2002).

EHLERS AND LUTHY (2003) have stated that:

"Bioavailability refers to the extent to which humans and ecological receptors are exposed to contaminants in soil or sediment".

Environmental scientists consider bioavailability as the accessibility of soil-bound chemicals for assimilation and possible toxicity (ALEXANDER, 2000). SEMPLE ET AL. (2004) identified the lack of clarity of the term "bioavailability" among environmental scientists, and thus proposed the distinction of two terms bioavailability and bioaccessibility:

"Bioavailability represents, at a given time, the fraction of the chemical that is freely available to cross an organism's membrane from the medium which the organism inhabits." Bioaccessibility was defined as: "that which is available to cross an organism's (cellular) membrane from the environment it inhabits, if the organism had access to it; however, it may be either physically removed from the organism or only bioavailable after a period of time. Bioaccessibility encompasses what is actually bioavailable now and what is potentially bioavailable" (SEMPLE ET AL., 2004).

Not only bioavailable compounds present in the water-soluble fraction of soil are available, but also these actually desorbed from soil during the time when a target organism is in direct contact with the soil (HARMSEN, 2007). Bioaccessibility includes both the readily available contaminants present in the water-soluble fraction of soil and the contaminants which can become available after desorption from the soil matrix (**Figure 5**). This includes also the contaminants, which may be released after longer timescale (slowly reversible). The accessibility depends on the desorption conditions (e.g. shaking, temperature) and desorption time (REICHENBERG AND MAYER, 2006). Both bioavailability and bioaccessibility of contaminants decrease with increasing contact time of contaminants with soil matrix [(so-called "aging"); REID ET AL., 2000].

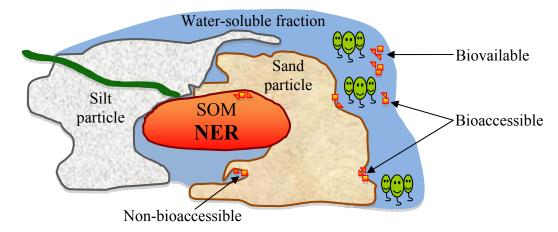


Figure 5. Bioavailability and bioaccessibility of organic contaminants in soil system (adapted from SEMPLE ET AL., 2004); (a) parent compound; (b) metabolites

A chemical immobilised in SOM or present in the soil solution can become bioaccessible. This involves a number of consecutive steps until eventually the compound is absorbed into an organism (see bioavailability processes A-D in **Figure 6**; EHLERS AND LUTHY, 2003; SEMPLE ET AL., 2004).

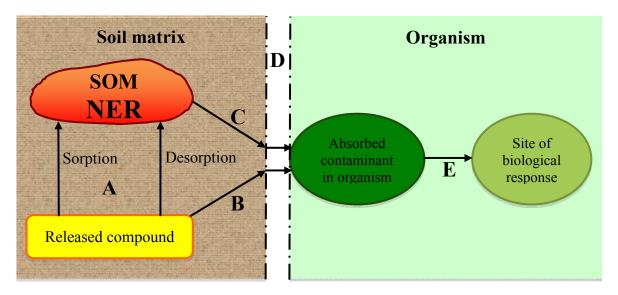


Figure 6. Bioavailability processes in soil system (EHLERS AND LUTHY, 2003; SEMPLE ET AL., 2004)

Process "A" encompasses the possible behaviour of the contaminant within the soil matrix (sorption or desorption). When environmental conditions are changed (e.g. changes in water saturation, pH or temperature) immobilised contaminant in SOM (NER), may be released or transformed into more stable associations over time (aging) and become non-accessible. The release of immobilised contaminant from SOM is the key step toward the assessment of its bioaccessibility and thus the toxicity for living organisms (HARMSEN, 2007). Therefore, the

formation of NER during microbial degradation of organic contaminants in soils is actually discussed in terms of the probability of the risks which they can pose to the environment (BARRACLOUGH ET AL., 2005). Processes B and C include the transport of the contaminant, which is present in the soil solution (B) or bound to soil components (C). D shows the barrier between external environment (soil matrix) and an organism and represents the uptake of the contaminant through the cell membrane into the target organism. E refers to paths taken by contaminant after uptake into the cells (e.g. metabolic transformation or exerting toxic effects within cells) and addresses toxicological bioavailability (EHLERS AND LUTHY, 2003).

Various chemical or biological measurements are employed to assess the bioavailability of an organic contaminant for living organisms. Chemical measurements involve various extraction methods of the contaminant from soil, whereas biological ones are based on monitoring the toxic effects of a contaminant taken up by the target living organism (HARMSEN, 2007).

2.4 Definition of the non-extractable residues (NER)

The term NER has been defined in different ways over the years (MORDAUNT ET AL., 2005). In the past, NER were considered to be pesticide residues remaining in the soil after application until the next growing season or the planting of the following crop (CRAVEN, 2000). Pesticides were and are still of particular interest in this regard, since they have selective activity and are deliberately manufactured for application in terrestrial systems (GEVAO ET AL., 2000).

The first official definition of the term NER was provided in 1975 by the American Institute of Biological Sciences – Environmental Task Group (NORTHCOTT AND JONES, 2000), which stated:

"bound pesticide residues in the soil are unextractable and chemically unidentifiable pesticide residues remaining in the fulvic, humic acids and humin fractions after exhaustive sequential extraction with nonpolar and polar solvents".

The most accepted and widely used NER definition, proposed by the Applied Chemistry Division, Commission on Pesticide Chemistry of the International Union of Pure and Applied Chemistry (IUPAC), is the following:

"bound residues (also referred as "non-extractable" residues or "non-extracted" residues) in plants and soils as chemical species originating from pesticides, used according to good agricultural practice, that are unextracted by methods which do not significantly change the chemical nature of these residues. **These residues are**

considered to exclude fragments recycled through metabolic pathways leading to natural products" (ROBERTS, 1984).

Biogenic residues formed during biodegradation of organic compounds are thus excluded from the IUPAC NER definition, because they are composed only of natural compounds derived from microbial biomass stabilised in SOM. However, with the use of radiotracers for estimating of mass balance of the tested contaminant in soil, these natural compounds cannot be excluded from the quantitative analysis of NER.

The IUPAC NER definition was later slightly modified but not substantially changed (NORTHCOTT AND JONES, 2000) and was extended to include that "the NER formation reduces the bioaccessibility and bioavailability significantly" (CALDERBANK, 1989; FÜHR ET AL., 1998).

2.5 Determination of NER in soil

The determination of NER is based on quantitative analyses accomplished after the extraction of residual parent compound and its metabolites from soil as mentioned in section 2.2.2. The distinction between extractable and non-extractable residues depends on extraction methods and conditions employed (KHAN 1991; NORTHCOTT AND JONES, 2000; MORDAUNT ET AL., 2005). Numerous "exhaustive" (harsh) and "non-exhaustive" (mild) extraction methods employed for the assessment of bioavailability of organic contaminants for living organisms show different quantitative recovery of a target compound (NORTHCOTT AND JONES, 2000; SEMPLE ET AL., 2007). The main aim of harsh extraction methods measuring the total concentration of contaminant in soil is to recover all or as much as possible of the contaminant from environmental samples (HATZINGER AND ALEXANDER, 1995; NOORDKAMP ET AL., 1997; ALEXANDER, 2000; REID ET AL., 2000; STOKES ET AL., 2006). These harsh methods involving mostly heated organic solvents include soxhlet extraction, microwave extraction, supercritical fluid extraction, ultrasonication and accelerated solvent extraction [(ASE); NORTHCOTT AND JONES, 2000]. It is generally assumed that the amount of compound recovered by exhaustive soil extraction is 100% available when assessing its potential risk for environment (ALEXANDER, 2000). However, the measurement of total concentration besides the accessible fraction of contaminant may also include parts or the entire recalcitrant fraction of contaminants (e.g. aged or sorbed strongly to SOM). In addition, the soil matrix by these extraction techniques is altered thus no longer represents a real soil (MORDAUNT ET AL., 2005). This occurs, because used organic solvents remove H₂O and SOM (mostly humic acids) from the soil matrix (MORDAUNT ET AL., 2003). Therefore, these harsh methods were considered to be improper for the assessment of the toxicity and thus the risk of soil

associated organic contaminants for the environment (Kelsey et al., 1997; Gevao et al., 2003). Mild extractions methods that remove only the "labile" pool of the contaminant (water-soluble and loosely adsorbed contaminants on surface of soil particles) are more useful for the measurement of bioaccessibility of contaminants in soils and their potential risks (Kelsey et al., 1997; Reid et al., 2000; Semple et al., 2007). These mild extraction methods involving natural extractants (e.g. ionic solution of CaCl₂) mimic solutions likely to be present in the soil (Mordaunt et al., 2003). These methods are aqueous-based extractions including solid-phase extraction (e.g. Tenax TA) and the cyclodextrin technique (Reid et al., 2000; Liste and Alexander, 2002; Semple et al., 2007).

Besides the physico-chemical properties of compound and soil, the amount of NER also strongly depends on the extraction method (MORDAUNT ET AL., 2005). Therefore, the methods used for the extraction of contaminants from soil affects also the quantitative analyses of NER. For instance, harsh extraction methods result in lower estimates of NER, whereas the mild ones give higher estimates for NER in soils. Hence, the numerous presented data on the total amounts of NER formed during the microbial degradation of the same model organic contaminant in soils show high variations. Summing up the above, the NER term is defined operationally by the extraction method employed; hence it is important to clarify how the method is developed from this definition and what information this method will provide (MORDAUNT ET AL., 2005).

2.6 Formation of NER in soil

It is believed that formation of NER during biodegradation of a contaminant is based on the various physico-chemical interactions between parent compounds or its metabolites and SOM (BOLLAG ET AL., 1992; SENESI, 1992; VERSTRAETE AND DEVLIEGHER, 1996; FÜHR, 1998; ALEXANDER, 2000; GEVAO ET AL., 2000; LOISEAU AND BARRIUSO, 2002; MORDAUNT ET AL., 2005). However, biodegradation of natural organic compounds in soil results in the formation of biogenic residues (as shown in **Figure 3** in section 2.2.1). Organic contaminants similar to the natural organic compounds are also subject to microbial degradation, which can lead to the formation of organic contaminant-derived biogenic residues. These contaminant-derived biogenic residues are only quantified as NER and no detailed information about their chemical structure is available (RICHNOW ET AL., 2000; KÄSTNER AND RICHNOW, 2001; BARRIUSO, 2008). Therefore, it is necessary to consider the formation of biogenic residues during biodegradation of organic contaminants and to distinguish between formations of NER via components of microbial biomass and via parent compound or its metabolites in soils.

The amounts of NER formed during the biodegradation of organic contaminants depend on the previously mentioned physico-chemical properties of both the contaminant and the soil system. In addition, the position of the label in the molecule of the tested contaminant has an impact on the results on the fate of this contaminant in general including the formation of NER (BARRIUSO ET AL., 2008). For instance, if the label is positioned in a labile molecular fragment of contaminant (which is easily evolved as ¹⁴CO₂), the mineralisation will be overestimated, whereas the detected amounts of NER tend to be low (BARRIUSO ET AL., 2008).

2.6.1 Parent compounds and metabolites

Humic acids as major components of SOM were previously regarded as the main agents leading to the formation of NER in soil (SENESI, 1992; WAIS, 1998). This is caused by the fact that it is believed that humic acids are large aromatic polymers with many binding sites for organic contaminants (hydrophilic and hydrophobic) in their molecular structure (SENESI, 1992).

Investigations of the interactions between organic contaminants and SOM involved various extraction techniques basing mostly on the isolation and fractionation of soil humic substances. To characterise NER, several attempts to cleave bonds of the contaminant with SOM were used. For instance, alkaline hydrolysis was used to cleave organic compounds bound to SOM via ester bonds (RICHNOW ET AL., 1998), and acid hydrolysis to release compounds linked to SOM by ether bonds (RICHNOW ET AL., 1997). However, in most cases, the physico-chemical interactions of contaminants with humic substances were studied in simple systems with solutions containing only humic acids (mainly synthetic) and the contaminant (BOLLAG, 1991; BOLLAG 1992; HATCHER ET AL., 1993; PICCOLO ET AL., 2001). These approaches follow only a specific single reaction or binding mechanism of the respective contaminant to humic acids (NORTHCOTT AND JONES, 2000). The physico-chemical interactions were thus described mainly theoretically. Therefore, these data on the mechanisms of NER formation cannot necessarily be extrapolated to the realistic processes, which take place in complex soil systems, especially when considering the fact that the molecular structure of humic acids is not yet clear.

Organic contaminants are believed to bind to humic acids via various mechanisms of adsorption, covalent bonds and sequestration (BOLLAG ET AL., 1992; WAIS, 1998; GEVAO ET AL., 2000).

Adsorption

Adsorption is the binding of the organic compound to the surface of solid particles through weak and reversible bonds (BÜYÜKSÖNMEZ ET AL., 1999). Minerals, in particular clays and SOM can take part in the sorption of the organic contaminants. The clay minerals are responsible for adsorption of polar and hydrophilic compounds, whereas SOM has both hydrophilic and hydrophobic sites and therefore can interact with polar, charged as well as apolar and lipophilic contaminants (VERSTRAETE AND DEVLIEGHER, 1996). This type of association is reversible and leads to the formation of unstable NER subject to microbial degradation (VERSTRAETE AND DEVLIEGHER, 1996). This can be caused by the fact that microorganisms, which colonise surfaces of soil particles, might have direct access to adsorbed chemicals and degrade them.

The extent of adsorption depends on the properties of the soil (e.g. SOM content) and the nature of contaminant e.g. if it is acidic or alkaline (GEVAO ET AL., 2000). Adsorption occurs via several mechanisms like ionic and hydrogen bonding, charge-transfer, ligand exchange, van der Waals forces and hydrophobic bonding (KHAN, 1978; PIGNATELLO, 1989; GEVAO ET AL., 2000).

Ionic binding: this type of binding occurs with those contaminants and their metabolites that exist in the cationic form or can become cationic upon protonation (SENESI, 1992; WAIS, 1998; GEVAO ET AL., 2000). These contaminants react with ionised or easily ionisable carboxylic and phenolic hydroxyl groups of humic acids or hydroxyl groups of minerals (SENESI, 1992).

Hydrogen binding: nonionic contaminants can interact with oxygen- and nitrogen-containing functional groups at the surfaces of humic substances or oxygen- and nitrogen-containing functional groups of the contaminant can react with hydrogen containing groups of SOM (WAIS, 1998).

Charge-transfer complexes: also called donor-acceptor complexes are formed when molecules with a high electron density (e.g. π -electrons in aromatic systems) react with electron-deficient molecules (e.g. quinones; WAIS, 1998).

Ligand exchange: in this type metal ions complexed by humic acids are usually associated with water molecules, which can be replaced by functional groups (e.g. carboxylic or amino groups) of organic contaminants (WAIS, 1998). However, this bonding is rare and weak and thus plays minor role.

Van der Waals forces: are relatively weak short-range dipolar or induced-dipolar attractions that exist in presence of stronger binding forces, in all adsorbent-adsorbate interactions (GEVAO ET AL., 2000).

Hydrophobic bonding: is the interaction mechanism between hydrophobic groups of humic substances and non-polar contaminants (WAIS, 1998). Active partners are aliphatic side chains, fat constituents or lignin components of humic substances (SENESI, 1993).

Covalent bonding

Covalent bonding is a chemical interaction between a contaminant or its metabolites and humic substances driven by strong bonding forces (> 300kJ/mol; BLASCHETTE, 1974). This type of bonding is generally accepted to be an irreversible and very stable association (SENESI, 1992; WAIS, 1998), which results in the formation of persistent contaminant NER in soil. These interactions can be mediated by enzymatic, chemical or photochemical catalysts (SENESI, 1992; GEVAO ET AL, 2000) and result in ester, ether or carbon-carbon linkages (KÄSTNER AND RICHNOW, 2001). The organic contaminants, which are bound covalently to soil humic acids, are integral components of humic substances (BOLLAG ET AL., 1992; SENESI, 1992). Thereafter, they are subject to all further transformation processes involved in humification (WAIS, 1998). The residues bound chemically to soil humic acids are considered to be toxicologically inactive (CALDERBANK, 1989), because they are not bioavailable. However, the exact chemical nature and structure of this type of NER in SOM have not yet been elucidated for all xenobiotics (KATAYAMA ET AL., 2010). The studies on NER formation by covalent bonding are limited to simple humic acids-contaminant systems (BOLLAG, 1991, 1992; HATCHER ET AL., 1993).

Sequestration (Aging)

Sequestration (aging) contrary to the adsorption mechanisms does not include reactions that alter the structure of an organic contaminant (HATZINGER AND ALEXANDER, 1995). Sequestration is also referred to a slow sorption/diffusion (PIGNATELLO AND XING, 1996; DEC AND BOLLAG, 1997) of non-polar and hydrophobic compounds, which in comparison to adsorption is a very long-term process (GEVAO ET AL., 2000). During aging, the molecule becomes progressively more tightly bound or entrapped in SOM and correspondingly less bioavailable (BARRACLOUGH ET AL., 2005). The size distribution of micropores is a factor in aging: the smaller the micropore, the slower aging proceeds, but the effect is stronger (ALEXANDER, 2000; KATAYAMA ET AL., 2010). Aging may also result from covalent bonding of contaminant with soil humic acids after sorption inside or within micropores in soil aggregates (ALEXANDER, 2000; KATAYAMA ET AL., 2010).

2.6.2 Components of microbial biomass

Several studies showed that the formation of NER from degradation of most pesticides is often related to soil biological activity and to the amount of SOM present in the soil (KAUFMANN AND BLAKE, 1973; ABDELHAFID ET AL., 2000A, B). For instance, soil microbial activity, which is generally higher in topsoil layers, induces the formation of NER in higher amounts than in deeper soil horizons (SCHIAVON, 1988; BALUCH ET AL., 1993; STOLPE AND SHEA, 1995; RICE ET AL., 2002). In addition, soil amendments with organic materials such as manure and straw enhanced the formation of NER and the dissipation of pesticides (DOYLE ET AL., 1978; PRINTZ ET AL., 1995). For example, during the incubation of composted non-sterile straws with 2,4-dichlorophenoxyacetic acid (2,4-D), 2,4-dichlorophenol (2,4-DCP) or 4-chlorophenol, the contents of NER were high, whereas in sterile ones these values were negligible (BENOIT AND BARRIUSO, 1997). The addition of glucose, which usually induces the microbial activity, increased the rate of NER formation in soil amended with atrazine in comparison to a soil without glucose application (ABDELHAFID ET AL., 2000A, B).

Also for Polycyclic Aromatic Hydrocarbons (PAH) adsorption and covalent bonding were proposed as possible mechanisms of NER formation in soils. However, the molecular analyses of SOM by pyrolysis gas chromatography-mass spectrometry after alkaline hydrolysis of ester bonds revealed that the contents of both parent compounds and known metabolites of PAH were very low (0.5% of the initial amount of PAH; RICHNOW, ET AL., 1994). Other study on the ¹⁵N-labelled simazine-derived NER in ¹⁵N-depleted plant compost indicated that these NER contained no parent compound and were composed of degradation products resulting from N-dealkylation and triazine ring destruction (BERNS ET AL., 2005). KÄSTNER ET AL. (1999) suggested for the first time that NER formed during biodegradation of PAH (9-[14C]-anthracene) could be of "biogenic" origin. In their experiments, addition of compost led to higher mineralisation rate of anthracene and lower amounts of NER in comparison with the native soil (see **Table 1**). The amount of NER in the experiment with 9-[14C]-anthracene was relatively high, in spite of the fact that the C at the labelled position is subject to release as CO₂. In addition, an experiment with ¹⁴CO₂ (in the amount corresponding to the final mineralisation of 9-[14C]-anthracene) showed that the label was found mostly in NER. This indicates that they were not directly formed from anthracene, but via CO₂ fixation and microbial biomass. In addition, the partitioning of ¹⁴C in humines, humic and fulvic acids in the ¹⁴CO₂ experiment was similar to the one in the 9-[¹⁴C]-anthracene experiment. The amounts of ¹⁴C originating from ¹⁴CO₂ in NER of soil inhibited with CHCl₃ were significantly lower than in soil without CHCl₃ treatment.

Table 1. Distribution of ¹⁴C after microbial degradation of 9-[¹⁴C]-anthracene and in native soil and soil-compost-mixture (after 176 days) and after incubation under ¹⁴CO₂-atmosphere [(after 90 days); from KÄSTNER ET AL. 1999]

Mass balance	Applied radioactivity [%] 9-[14C]-anthracene			
	Native soil	Soil-compost ^a	Soil-compost a,b	+CHCl ₃ ^{a,b,c}
Total recovery	98.7	91.9	99.8	98.7
CO_2	43.8	67.2	4.6 e	11.3 ^e
Total soil	54.9	24.6	95.2 ^f	87.4
Extraction+alk. hydrolysis	9.5	3.9	9.4	44.2
"bound" residues	45.4	20.7 ^d	85.8	43.2

soil-compost-mixture (80% : 20% dry weight); ^b 90d; ^c microflora inhibited by fumigation with CHCl₃ d amount represents 25.5 Bq/g of soil; ^e recovered ¹⁴CO₂ (initially applied amount: 24.8 Bq/g of soil) f 31% bound in acid labile carbonates, 67% bound to SOM, < 2% bound to clays and silicates

The above-mentioned studies clearly indicate that in many cases microorganisms mediate NER formation in soils. As already mentioned, the C derived from organic contaminants can be used for the formation of microbial biomass during their biodegradation in soil (MÜLLER ET AL., 2007). The CO₂ evolved during mineralisation of organic contaminants can be utilised as an additional C source by soil microorganisms, since even heterotrophic microorganisms need CO₂ for their growth (KREBS, 1941). PEREZ AND MATIN (1982) showed that during heterotrophic growth about 10% of the cell C originated from CO₂. Another experiment with labelled CO₂ in the atmosphere proved that soil microorganisms assimilated C and incorporated the label into SOM (MILTNER ET AL. 2004, 2005). Therefore, NER formed during biodegradation of organic contaminants in soil, can in reality be biogenic residues (KÄSTNER AND RICHNOW, 2001; **Figure 7**), the formation of which is also observed during biodegradation of natural organic compounds (**Figure 3** in section 2.2.1; KÄSTNER, 2000).

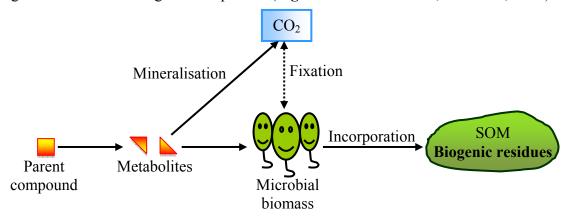


Figure 7. Biogenic residues formation during microbial degradation of organic contaminants in soil

In addition, when labelled contaminants are used for soil biodegradation studies, a certain amount of this label incorporated into microbial biomass components, which thereafter are stabilised in SOM, will be detected as NER. Biogenic residues can be formed via direct incorporation of C from pollutant or indirectly via fixation of CO₂ released during mineralisation of the contaminant (see **Figure 7**).

Fatty acids and amino acids as representatives for biogenic residues

Fatty acids (FA) and amino acids (AA) that are known microbial biomarkers (Boschker and MIDDELBURG, 2002) can be incorporated into SOM and thus form biogenic residues. Free and bound lipids, amino acids and carbohydrates form the major fraction of analytically recognisable compounds in SOM of various origin (ALLARD, 2006). In addition, the extraction of known microbial biomarkers (e.g. FA and AA) from soil after addition of C isotope tracer allows an estimation of microbial activity in the biogeochemical cycling of C (PELZ ET AL., 1998). Biomarker analysis is thus also useful in studying the microbial transformation of labelled organic contaminants in soil (Boschker and Middelburg, 2002). Therefore, the analysis of C isotope label distributions in FA and AA within the living and non-living SOM fraction enables understanding the formation and the fate of biogenic residues during biodegradation of a labelled organic contaminant in soil.

FA are major constituents of the soil lipids (PAUL AND CLARK, 1996) and may originate from both plant residues and soil organisms (STEVENSON, 1994). This biomarker is used for identification of microbial populations involved in specific geochemical transformations, e.g. organic contaminant degradation in soils after deliberately added tracers (BOSCHKER AND MIDDELBURG, 2002). Short-chain FA (C:4-C:20) are typical for microorganisms (SCHNITZER ET AL., 1986). On average, FA represent about of 5% of the dry weight (DW) of the microbial biomass (BAS ET AL., 2003). Phospholipid fatty acids (PLFA) are characteristic for the lipids of living cells (ZELLES, 1999). PLFA are known to be stable only in intact cells and are hydrolysed within weeks after their death (VESTAL AND WHITE ET AL., 1989). They therefore can be used as biomarkers for representative groups of the living microorganisms in the environment (GREEN AND SCOW, 2000; KAUR ET AL., 2005; ZELLES, 1999). The characteristic chain lengths of ester-linked saturated FA of PLFA range from C_{14:0} to C_{18:0} for bacteria and from C_{14:0}-C_{24:0} for fungi (ZELLES, 1997). Several authors (GREEN AND SCOW, 2000; KAUR ET AL., 2005; ZELLES, 1999) proposed that the following five classes of PLFA are typical for specific groups of microorganisms: (1) saturated straight chain PLFA for all microorganisms, (2) saturated methyl branched PLFA for Gram-positive bacteria, (3) monounsaturated PLFA for Gram-negative bacteria, (4) polyunsaturated PLFA for fungi and (5) saturated cyclopropyl PLFA for starving Gram-negative bacteria.

AA, which account for 10–20% of total C and 30–50% of total N in SOM, are mainly present as polymers e.g. proteins, protein-humic complexes and peptides (STEVENSON, 1982). AA are also the most abundant components (55% of dry weight) of bacterial cells (MADIGAN AND MARTINKO, 2006). Total free AA concentrations in soil are very low representing < 0.05% of total soil N (JONES AT AL., 2005) and have very short turnover times (JONES, 1999), because they are easily available and degraded immediately by soil microorganisms.

Two biodegradation studies with labelled organic contaminants showed that the label was incorporated into AA. In the first biodegradation experiment with tar oil contaminated soil spiked with 1-[\(^{13}\text{C}\)]-phenanthrene significant amounts of the label were found in hydrolysable AA, which represented 11% of total NER (RICHNOW ET AL., 2000). The second experiment studying the fungal degradation of [\(^{15}\text{N}\)]-Trinitrotoluene (\(^{15}\text{N}\)-TNT) indicated that 1.7% of the \(^{15}\text{N}\) label was detected as "biogenic" AA in the wheat straw containing the fungus (Weiß et Al., 2004). In contrast to the wheat straw with fungus, no incorporation of the label from \(^{15}\text{N}\)-TNT into AA was observed in the soil layer without fungus. A study on the fate of \(^{13}\text{C}\)-labelled *E. coli* demonstrated high stability of \(^{13}\text{C}\)-amino acids, whose contents remained constant even after 224 days of incubation (MILTNER et Al., 2009). This is caused by the fact that proteins are amphiphilic molecules with a strong tendency to bind to surfaces (Hlady AND Buils, 1996), especially to mineral surfaces, over long periods of time (Kleber et Al., 2007).

2.7 Stability of NER in soil

The stability of NER formed during microbial degradation of organic contaminants in soils has been investigated since the late 1960s (HATZINGER AND ALEXANDER, 1995). Since then, it has been observed that aged organic compounds in NER are not always immobilised irreversibly in the soil. Therefore, many scientists focused on the possible release of NER from soils under certain conditions (BURAUEL AND FÜHR, 2000; BOIVIN ET AL., 2005; GEVAO ET AL., 2005; LERCH ET AL., 2009) and the long-term environmental impact of these residues (GEVAO ET AL., 2001).

The release of organic contaminant or its various metabolites immobilised in SOM is related closely to SOM decomposition as a result of either biochemical processes or physicochemical mechanisms (GEVAO ET AL., 2000; BARRACLOUGH ET AL., 2005). However, the activity of microorganisms is considered as the primary factor responsible for the release of NER (GEVAO ET AL., 2000). To explore the possible scenarios of long-term behaviour of contaminant NER in soils during SOM turnover, several biodegradation studies were conducted. Their overall aim was to simulate enhanced SOM decay by stimulation of

microorganisms. Microbial activity was induced by modification of the physicochemical properties of soil like pH or by an increase of SOM content. The addition of fresh soil to soil containing [U-14C]3,4-dichloroaniline-derived NER induced a slight mineralisation of these NER to ¹⁴CO₂ (HSU AND BARTHA, 1974). A similar effect was also observed in a soil experiment with [14C]prometryn, in which the addition of the fresh soil inoculum remobilised 27% of the initial amount of [14C]prometryn-derived NER, which could be extracted with methanol (KHAN AND IVARSON, 1981). No release of [14C]prometryn-derived NER in the sterile control experiment highlighted the importance of microorganisms in the destabilisation of NER derived from this contaminant (KHAN AND IVARSON, 1981). Another experiments with prometryn, in which the effect of pH was studied, revealed that an increase of pH from 4 to 8 caused a release of up to 25% of the initial amount of prometryn NER from soil (YEE ET AL., 1985). Addition of either glucose or cow manure, both known to induce the microbial activity, resulted in higher release of [14C]parathion NER from soil compared to controls (RACKE AND LICHTENSTEIN, 1985). In an inverse experiment, the microbial activity in soil was supressed by either the bactericide chloramphenicol or the fungicide captafol. Both treatments led to the significant decrease of the mineralisation rate of [14C]parathion NER in soil (RACKE AND LICHTENSTEIN, 1985).

ESCHENBACH ET AL. (1998) and WEISS ET AL. (2004) investigated the stability of NER formed during microbial degradation of 9-[¹⁴C]-anthracene or ¹⁴C-TNT by simulation of extreme physical, chemical or biological situations. They used the following treatments:

- 1. Physical treatment = simulation of climatic effects by changing the soil texture via grinding, freezing and thawing of soil;
- 2. Chemical treatment = extraction of soil with the metal complexing agent EDTA, in order to estimate the effects of bivalent cations on the aggregation of macromolecular organic compounds and their potential release from soil. The extraction of soil with acidified H₂O was used to simulate the acid rain impact on the release of ¹⁴C-TNT NER;
- 3. Biological treatment = simulation of increased turnover of SOM by addition of the compost and incubation of the soil containing NER with ligninolytic fungi. The ligninolytic fungi are capable of depolymerising lignin and humic substances (KÄSTNER AND HOFRICHTER, 2001). Additionally, the impact of the plants on the uptake or mobilisation of NER was studied.

Neither physical nor biological treatments caused a mobilisation of "bound" $9-[^{14}C]$ -anthracene and ^{14}C -TNT residues in soils. In contrast, a low (< 15% of initial amounts

in soil) release of NER from soil containing residues from microbial degradation of either 9-[¹⁴C]-anthracene or ¹⁴C-TNT was observed after addition of the complexing agent EDTA. Simulation of acid rain also mobilised ¹⁴C-TNT residues in soil. Neither 9-[¹⁴C]-anthracene nor ¹⁴C-TNT was released from SOM after EDTA or the acid rain treatment.

However, the NER released by these mobilisation treatments were analysed only for the presence of the parent compounds and their known main transformation products, which are supposed to be responsible for their formation. A detailed characterisation of the chemical nature of other compounds, which may form NER, is still lacking. However, the absence of the relevant toxic components clearly indicates that they were further transformed into different kind of compounds during residue formation. On the whole, the NER formed during microbial degradation of ¹⁴C-TNT or 9-[¹⁴C]-anthracene were very stable (ESCHENBACH ET AL., 1998; WEISS ET AL., 2004). The stability of NER is strongly dependant on their "age", as it was shown in soil experiments with 15 years-old NER and 90 years-old NER derived from 2,4-D (LERCH ET AL., 2009). The addition of fresh soil to younger NER induced their mineralisation, whereas the older ones were stable and turned over at similar rates as SOM (LERCH ET AL., 2009).

2.8 Risk assessment of NER

Independent of the binding mechanism, the formation of NER leads to the limitation of bioavailability of contaminants (NORTHCOTT AND JONES 2000). In addition, the binding of contaminants to humic substances was reported to reduce the amounts which may interact with the biota and thus their complexed products should be less toxic than the free parent compounds (BOLLAG ET AL., 1992). Therefore, the enhanced transformation of organic contaminants into NER has been proposed as an alternative remediation method for polluted soils (BERRY AND BOYD, 1985; BOLLAG, 1992; VERSTRAETE AND DEVLIEGHER, 1996). On the other hand, under certain conditions the release of NER was observed; therefore, it is speculated that the compounds of unknown structure released from SOM can become bioavailable and thus toxic for the living organisms (BARRACLOUGH ET AL., 2005). Due to the difficulties in the identification of NER, the complications of multiple applications of a range of compounds and the uncertainties around their bioavailability it has been suggested to include them in the "conventional" environmental risk assessments (BARRACLOUGH ET AL., 2008).

The "conventional" risk assessment comprises the followings steps (ISO/DIS, 2006):

1. The hazard identification: recognition of the potential of the compound, which may cause harm to human health or the environment.

- 2. The exposure assessment: estimation whether and how much exposure will occur between a receptor and a contaminated source.
- 3. The dose-response assessment: characterisation of the relationship between the dose of a chemical and the anticipated incidence of health and environmental effect in an exposed population.
- 4. The risk characterisation: description of nature and magnitude of health or environmental risk.

In step 1, to identify the hazard, respective samples containing the contaminant are taken from the investigation site and their total concentrations are measured. For that purpose, different soil extraction methods are employed. These concentrations are compared with threshold values from dose-response relationships (step 3). Crossing the threshold is considered to be a risk. This procedure leads to overestimation of the risks (HARMSEN, 2007). In step 2, both the bioavailability of the contaminant and the exposure of a target receptor to the contaminant are analysed. In the assessment of the environmental risk, which may be posed by the contaminant, the receptor or protection goal is a leading principle (HARMSEN, 2007). These receptors can be human and higher animals or soil-living organisms. Furthermore, soils and sediments act as buffers, since they may absorb contaminants and prevent their transport to groundwater, surface water and the terrestrial food chain (HARMSEN, 2007).

Using this model for the assessment of environmental risk associated with NER is more complex than for human health risk assessment (EHLERS AND LUTHY, 2003). Therefore, EHLERS AND LUTHY (2003) proposed to redefine "bioavailability" into "bioavailability processes" in order to allow the accurate assessment of the environmental risk posed by a chemical in soil (as shown in **Figure 6**). Up to date, however, there are still many factors, which may impede the proper risk assessment of a target compound in soils, for example:

- 1. The complex and heterogeneous soil matrix, makes the direct analysis of the chemical nature of NER without application of any specific techniques and separation methods impossible;
- 2. Measurement of bioavailability of organic contaminants in soils is often inappropriate. The prediction of the potential risk of chemicals in soils is conducted by using various solvent extraction methods (GEVAO ET AL., 2003) for the assessment of "bioavailability". However, these chemical methods used for extraction of the target compound from soil show still large uncertainties in the estimated NER content (EHLERS AND LUTHY, 2003; total concentration or bioaccessibility are measured). Furthermore, apart from parent compounds and their metabolites

extracted from soils, other unidentified small molecular fragments may be present in extracts, in forms that do not pose ecological risks (KELSEY ET AL., 1997; CHUNG AND ALEXANDER, 1998; ROBERTSON AND ALEXANDER, 1998);

3. Proper knowledge about possible mechanisms of NER formation during microbial degradation of organic contaminants in the soil system is missing.

Available studies on the NER formation during biodegradation of organic contaminants in soil are mostly limited to quantitative analyses using radioactive tracers (GERST AND KLIGER, 1990; KUBIAK ET AL., 1990; RICHNOW, 1999). It is suggested that NER are formed by sorption or sequestration of parent compound or metabolites within SOM components and their possible release is considered as a potential risk for both environment and humans (ALEXANDER, 2000; BARRACLOUGH ET AL, 2005), since their chemical nature is still not clear (see conventional model in **Figure 8**).

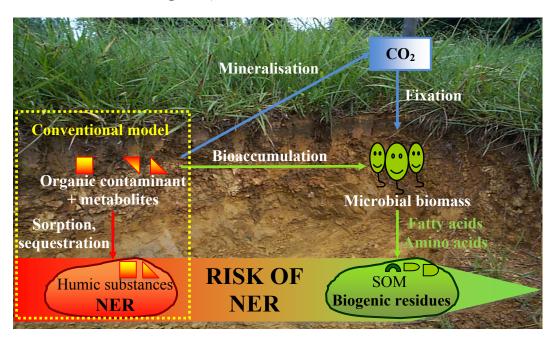


Figure 8. Proposed concept and the conventional model of the risk assessment related to NER formation during the microbial degradation of organic contaminants in soil

In case of readily biodegradable contaminants, mainly the misunderstanding of the processes of NER formation in soil is thus the key problem affecting the proper assessment of the risk which they can pose. It should be kept in mind that NER formed during microbial degradation of organic contaminants can be also biogenic residues (**Figure 8**). Biogenic residues are composed only of natural compounds derived from microbial biomass, such as FA and AA, which are incorporated into the SOM pool. Over time these residues are more and more irreversibly bound to SOM by the stabilisation process. Residues from biomass, however, do not present any toxicological hazard and thus would lead to an overestimation of the risk

related to NER in soil. Although biogenic residues are clearly excluded by the IUPAC definition (ROBERTS ET AL. 1984), they are generally not analysed and considered in studies on the fate and in the risk assessment concepts.

Therefore, the detailed study on the formation of biogenic residues during the microbial degradation of organic contaminants in soil is needed to establish both their realistic degradation rates and a proper assessment of their potential risk.

3 AIMS OF THE STUDY

Studies on the formation of NER during microbial degradation of various organic contaminants in soils presented in the literature are mostly limited to quantitative analyses. Therefore, the chemical nature of NER is not clear yet. However, the soil environment is a very complex system, thus the detailed analysis of NER structure needs a thorough understanding of the mechanisms of their formation and specific extraction methods. In the present study, it has been assumed that some part of NER formed during biodegradation of organic contaminants in soils is composed of microbial components. This is based on the fact that microorganisms utilise C derived from organic compounds or from CO₂ for biomass synthesis and after death their cell constituents are incorporated into SOM and form biogenic residues. This mechanism occurs for both natural organic compounds and readily degradable contaminants in soil.

The overall aim of this study was to elucidate the pathways of biogenic residue formation originating from microbial degradation of organic contaminants in soil and to study their stability in time course. The formation and fate of two representatives for biogenic residues, FA and AA during biodegradation of either ¹³C₆-2,4-Dichlorophenoxyacetic acid (2,4-D) or ¹³C₆-Ibuprofen (Ibu) in soil were investigated. In addition, an incubation experiment with ¹³CO₂ and unlabelled 2,4-D was carried out in order to study the incorporation of the labelled C into FA and AA via heterotrophic CO₂ fixation. A simple model experiment with a well-known 2,4-D degrader in the pure culture was performed to obtain a better view on the pathways of biogenic residue formation in complex soil systems. FA and AA were extracted from soil and analysed for their concentration and isotopic composition.

The specific objectives of this study were:

- 1. To quantify the extent and the kinetics of biogenic residue formation during microbial degradation of two model organic contaminants. Up to date, there is no report on the contribution of microorganisms to biogenic residue formation. This is caused by the fact that it is generally accepted that the population of microorganisms colonising soil particles is negligible in comparison with the large surfaces of both clays and SOM with many binding sites. The main questions adressing this point are:
 - Are biogenic residues formed during microbial degradation of organic contaminants in soil?

- If yes, how much biogenic residues are formed and what is their contribution to the total NER content in soil? Are these values really important for the environment?
- 2. To evaluate if specific biomass components are preferentially incorporated into NER. Are FA or AA more relevant for biogenic residues formation in soil?
- 3. To assess if heterotrophic CO₂ fixation plays any role for the incorporation of C into biogenic residues in soil.
- 4. To check the stability of biogenic residues in soil and their possible remobilisation in time course. Several studies have revealed that some part of the NER was released as CO₂ even after the target compound was completely depleted. Therefore in this study, besides the quantification of the total content of biogenic residues at respective sampling dates, we also focused on the monitoring of the labelled C distribution within the living and the non-living SOM fraction. The non-living fraction provides information about the stabilisation process of biogenic residues in the SOM pool.

4 MODEL COMPOUNDS

2,4-Dichlorophenoxyacetic acid (2,4-D) and Ibuprofen (Ibu) were selected as model compounds for the biodegradation studies. These compounds are aromatic and acidic molecules, and are used extensively world-wide. In addition, both 2,4-D and Ibu are reported to degrade readily in soil and to form NER, which potentially may pose a risk for the environment. Detailed description and current state of the art of these compounds are presented in two next sections.

4.1 2,4-Dichlorophenoxyacetic acid (2,4-D)

2,4-Dichlorophenoxyacetic acid (2,4-D) as a member of the phenoxy group compounds is among the most widely used herbicids worldwide to control broad leaf weeds in cereal and grass seed crops (McGhee and Burns, 1995; Voos and Groffman, 1997; Boivin et al., 2005) with an annual production of more than 100,000 tons (Merini et al., 2008). This herbicide is a weak by acidic molecule (**Figure 9**) with a relatively low molecular mass of 221 g/mol and a high solubility of 0.6 mg/L in H₂O (Villaverde at al., 2008). The octanol/water partition coefficient (log K_{ow}) 2.83 indicates a moderate sorption potential of 2,4-D (Boivin at al., 2005). The persistence of 2,4-D is considered low to moderate (DT₅₀ 5-59 days; Villaverde at al., 2008).

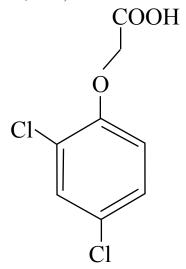


Figure 9. The molecular structure of 2,4-Dichlorophenoxyacetic acid

The microbial degradation of 2,4-D is well described (FULTHORPE ET AL., 1996; KAMAGATA ET AL., 1997; VOOS AND GROFFMAN, 1997B; CRESPÍN ET AL., 2001) and this compound is classified as readily biodegradable. 2,4-D is used as C and energy source (SANDMANN ET AL., 1988; SMITH AND LAFOND, 1990 LERCH ET AL., 2009A) by a wide range of microorganisms even in pristine soils (FULTHORPE ET AL., 1996; KAMAGATA ET AL., 1997). Bacteria genera, which have been reported to degrade 2,4-D include: *Pseudomonas* (KILPI ET AL., 1980),

Alcaligenes (PIEPER ET AL., 1988), Arthrobacter (BEADLE AND SMITH, 1982) and Flavobacterium (CHAUDHRY AND HUANG, 1988). Microbial breakdown of 2,4-D in soils is initiated by the removal of the carboxyl side chain or the cleavage of the ether bond (FORSTER AND MCKERCHER, 1973; ROBERTS ET AL., 1998) leading to the formation of 2,4-dichlorophenol (2,4-DCP) and other phenolic metabolites that are further degraded by cleavage of the phenyl ring (SMITH AND AUBIN, 1991; ROBERTS ET AL., 1998). The final products of 2,4-D degradation are succinyl-CoA and Acetyl-CoA, which enter the tricarboxylic acid cycle [(TCC); MANZANO ET AL., 2007; see Figure 10). Besides 2,4-DCP, the removal of the two-C side chain from 2,4-D also results in the formation of glyoxylate (AMY ET AL., 1985), which joins the TCC.

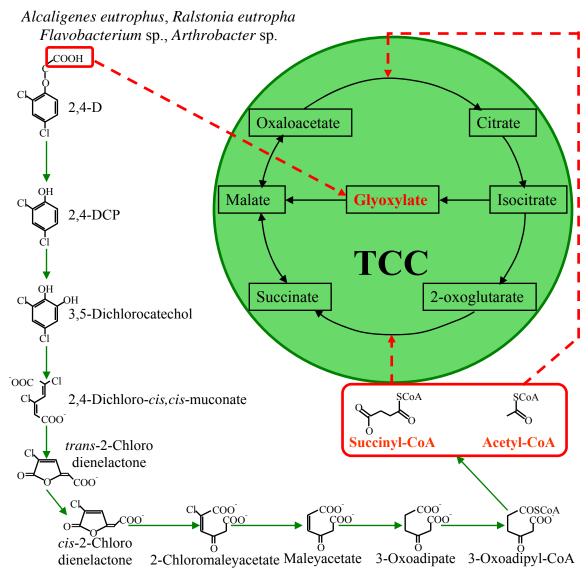


Figure 10. Scheme of 2,4-D degradation pathway (adapted from E. Young; source: map http://umbbd.msi.umn.edu/2,4-d/2,4-d_map.html)

Most laboratory studies on the fate of 2,4-D in soils are limited to radiotracer experiment with [ring-U-¹⁴C] 2,4-D (BENOIT AND BARRIUSO, 1997; BOIVIN ET AL., 2005; GAULTIER ET AL., 2008). Under non-sterile conditions the mineralisation of 2,4-D in soil reached about 50–60% of the initially applied compound (BENOIT AND BARRIUSO, 1997; BOIVIN ET AL., 2005). The degradation rate of 2,4-D in soils with higher total organic C (TOC) content was significantly higher than in soils with lower amount of TOC (GETENGA ET AL., 2004). It is generally known that the TOC content of a soil affects the microbial activity, which is higher in soils containing more TOC (GETENGA ET AL., 2004). However, it is speculated that 2,4-D is only partially degraded in soil, since a portion of 2,4-D and its metabolites are "bound" to SOM forming NER, which become stable over time (BARRIUSO ET AL., 1997; BENOIT AND BARRIUSO, 1997; BOIVIN ET AL., 2005). The amounts of NER formed during microbial degradation of 2,4-D in soil were reported to be relatively high, in a range of 20–56% of the initially added pollutant (BARRIUSO ET AL, 1997; BENOIT AND BARRIUSO, 1997; XIE ET AL., 1997).

A very reactive metabolite of 2.4-D degradation 2.4-DCP, is considered to be a major precursor for NER formation by binding to SOM (BARRIUSO ET AL., 1997; BENOIT AND BARRIUSO, 1997). Binding of 2,4-DCP to SOM occurs mainly by covalent bonding, which is catalysed by oxidoreductive enzymes, such as laccases and peroxidases (BOLLAG, 1991; BOLLAG 1992; HATCHER ET AL., 1993; DEC AND BOLLAG, 2000; XU AND BHANDARI, 2003A, B). The complexes resulting from covalent binding of 2,4-DCP to humic material were reported to be highly persistent (SAXENA AND BARTHA, 1983; BOLLAG, 1991, 1992). KLIBANOV ET AL. (1983) thus proposed using horseradish peroxidase for an in-situ stabilisation of phenolic compounds at contaminated sites. Horseradish peroxidase affected the sorption potential of 2,4-DCP in soil slurries; more 2,4-DCP was bound to SOM in presence of horseradish peroxidase than in soil slurries without this enzyme (PALOMO AND BHANDARI, 2005, 2006). However, horseradish peroxidase is a very reactive enzyme, which may react differently than peroxidases of microbial origin. In addition, soils used in these experiments were autoclaved several times; both the soil structure and physicochemical properties are changed during autoclaving. SHAW ET AL. (1999) reported that autoclaving soil increased dissolved organic C concentration 20-fold and reduced soil pH. This resulted in higher 2,4-DCP adsorption to autoclaved soil than to non-autoclaved one in their experiments. They also suggested using both re-inoculated sterile and sterile soils for a proper investigation of sorption-biodegradation interactions. Other study also demonstrated that autoclaving changes also the aggregation state of the soil, which may result in a higher surface area of soils and thus may lead to an increased sorption potential (BERNS ET AL., 2008).

The humification of straw was reported to increase the sorption potential of pollutants by increasing the proportion of constituents with sorption capacities relative to cellulosic constituents with low sorption capacities (BENOIT ET AL., 1996). However, even the presence of highly humified sterile straw did not induce a sorption of [U-ring-¹⁴C] 2,4-D and [U-ring-¹⁴C] 2,4-DCP to SOM (BENOIT AND BARRIUSO, 1997). XIE ET AL. (1997) fractionated [U-ring-¹⁴C] 2,4-D residues bound to humin after previous removal of soluble fulvic and humic acids from soil. They showed that most of the label was associated with the aqueous bound-humic acid solution and a very low amount was bound to the mineral fractions. Most recent studies have revealed that aged NER formed during 2,4-D biodegradation in soils are still available to microorganisms (BOIVIN, 2005; LERCH ET AL., 2009B) and that the stability was dependent on their age (LERCH ET AL., 2009B).

All studies on NER formation by sorption of either 2,4-D or its main metabolite 2,4-DCP are limited to simple humic acids-contaminant models (Bollag, 1991; Bollag 1992; Hatcher et al., 1993; Piccolo et al., 2001) or to autoclaved soil slurries (Palomo and Bhandari, 2005; 2006). In addition, NER formed during 2,4-D biodegradation in the complex soil systems were only quantified, thus their chemical structure is still unknown. It is difficult to compare the processes of NER formation occurring in the simple systems with the ones which may take place in the complex soil environment. Moreover, the comparison of biotic experiment with abiotic one is perhaps not appropriate, since these soils have different soil physico-chemical properties and a contaminant may thus behave differently. To conclude, in spite of many studies on NER formation during microbial degradation of 2,4-D in soil, a satisfying explanation on the mechanism of their formation is still missing.

4.2 Ibuprofen (Ibu)

Ibuprofen (Ibu) represents a novel environmentally relevant compound; is an important anti-inflammatory and analgesic non-prescription drug (BUSER ET AL., 1999), with a global annual production of several kilotons and a high daily dose of 600–1200 mg (ZWIENER ET AL., 2002). Ibu is derived from propionic acid (**Figure 11**), has a relatively low molecular mass of 206.28 g/mol and a moderate solubility of 21 mg/L in H₂O (YALKOWSKY AND DANNENFELSER, 1992). The log K_{ow} of Ibu is 3.5 (STUER-LAURIDSEN ET AL., 2000), indicating its moderate sorption potential. However, a study on the environmental fate of ibuprofen [carboxyl-¹⁴C] in water/sediment systems pointed to a low affinity for sorption onto sediment

(LÖFFLER ET AL., 2005). Ibuprofen displayed a low persistence in soils with DT₅₀ between 0.91 and 6.09 days (XU ET AL., 2009).

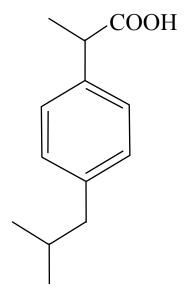


Figure 11. The molecular structure of ibuprofen

After normal therapeutic use, large amounts of Ibu are excreted as the parent compound or its various metabolites, e.g. hydroxyibuprofen, carboxyibuprofen and carboxyhydratropic acid to the wastewater treatment plant (WWTP; ZWIENER ET AL., 2002). The main mechanisms involved in the removal of pharmaceuticals in WWTP are biodegradation, sorption on sludge or particulate matter, filtration and chemical oxidation (MIÈGE ET AL., 2009). Ibu is eliminated in WWTP efficiently (between 75% to more than 98% [CLARA ET AL., 2005; JOSS ET AL., 2005; Castiglioni et al., 2006; Gómez et al., 2007; Jones et al., 2007; Kosjek et al., 2007; MATAMOROS ET AL., 2008]). However, presence of this pharmaceutical was analysed only in the water phase of effluents, thus its exact removal mechanism cannot be fully elucidated. After chemical step of WWTP using FeCl₃ only a small decrease (25%) of Ibu concentration was reported (ZORITA ET AL., 2009). The last step of WWTP sand filtration eliminated up to 20% of the initial concentration of Ibu (ZORITA ET AL., 2009). Despite the fact that Ibu is mostly removed during biological step of WWTP (JONES ET AL., 2007; ZORITA ET AL., 2009), this pharmaceutical is still detected at high concentration in biosolids (246–750 ng/g of dry weight of dewatered municipal biosolids [EDWARDS ET AL., 2009; MCCLELLAN AND HALDEN, 2010]). Application of biosolids in agriculture as a fertiliser may introduce Ibu into soils (EDWARDS ET AL., 2009).

In comparison with 2,4-D, up to date, there are only few studies on the fate of Ibu in soils after sewage sludge application. Ibu was easily degraded in the water/sediment system (77% was evolved as CO₂ after 100 days) and its metabolite 2-hydroxyibuprofen (2-OH-ibu)

dissipated already after 28 days of incubation (LÖFFLER ET AL., 2005). The degradation of Ibu was also quick in four agricultural soils with different OM contents; its dissipation kinetics depended on OM content and was much slower in soils containing more OM (XU ET AL., 2009). The degradation of this pharmaceutical in non-sterile soils was 34.5-fold faster than in sterile ones (XU ET AL. 2009) clearly indicating the contribution of microorganisms to its transformation. Another study focussing on the leaching potential of Ibu from soil demonstrated the absence of this compound in the leachate already after 20 hours (LAPEN ET AL., 2008), suggesting its microbial degradation or sorption to soil components. The first study on the label distribution within the soil system during 14 C₃-ibu biodegradation was presented by RICHTER ET AL. (2007). In their studies, after 100 days, \sim 35 % of 14 C₃-ibu equivalents was mineralised, \sim 50% of label was found in 14 C₃-ibu-derived NER and 11% was found as unidentified extractable 14 C₃-ibu residues. The formation of 14 C₃-ibu residues were detectable in soil. Thereafter, their contents decreased at a very slow rate (\sim 5% of 14 C₃-ibu equivalents after 100 days).

Microbial degradation of Ibu and NER formation in soil, up to date, are not studied in detail. In addition, neither data on the chemical structure of NER formed by Ibu nor information about the mechanisms of their formation in soil is available.

5 MATERIAL AND METHODS

5.1 Chemicals and materials

All chemicals and materials, except where otherwise specified (see **Table 2** and **Table 3**) were purchased from VWR/Merck (Darmstadt, Germany) at the highest quality commercially available.

Table 2. Chemicals

Chemical	Company				
¹³ C ₆ -2,4-D	Alsachim, Illkirch, France; (isotopic purity: 99 at%				
	chemical purity: ~ 98%; ¹³ C label in 6 positions)				
$Na_2^{13}CO_3$	Chemotrade, Leipzig, Germany; (isotopic purity:				
	99 at%; chemical purity: ~ 98%)				
¹³ C ₆ -ibu	Alsachim, Illkirch, France; (isotopic purity: 99 at%;				
	chemical purity: ~ 98%; ¹³ C label in 6 positions)				
2-OH-ibu	LGC GmbH (Luckenwalde, Germany); chemical				
	purity 99.8%)				
Bacterial Acid Methyl Esters (BAME)	Supelco, München, Germany				
Heneicosanoate methyl ester (FAME	Sigma-Aldrich, Munich, Germany				
21:0)					
Sodium deoxycholate	Sigma-Steinheim, Germany				
Polyethylenglycol (PEG 6000)	Carl Roth, Karlsruhe, Germany				

Table 3. Materials

Material	Company			
CHROMABOND® EASY, 3ml	Macherey-Nagel, Düren, Germany			
200 mg (for 2,4-D + metabolites)				
Silicagel (for PLFA-analysis)	Unisil, Clarkson Chromatography Products, South			
	Wiliamsport, USA			
Silicagel (for tFA-analysis)	Mallinckrodt Baker, Griesheim, Germany			
glass fiber membrane filters	0.7 μm, Puradisc 25 GF/F, Whatman Inc.,			
(for AA-analysis)	Cliffton, USA			
Cupriavidus necator JMP 134	Environmental Microbiology Department (UFZ,			
	Leipzig); DMSZ 4058, Braunschweig, Germany			

5.2 Liquid culture experiments

In order to understand the processes of biogenic residue formation during microbial degradation of 2,4-D in the complex soil system, a simple model experiment with the bacterial strain *Cupriavidus necator* JMP 134 was performed. *C. necator* JMP 134, which is a well-described 2,4-D degrader in soil, was selected for biodegradation studies in pure culture. As mentioned in section 2.6.2, biogenic residues can be formed directly via the incorporation of C derived from an organic contaminant or indirectly from CO₂ evolved during contaminant mineralisation (CO₂ fixation). To distinquish between these two processes, different ¹³C-labelled compounds were used. To study ¹³C incorporation into biomass of *C. necator* directly from a labelled contaminant, ¹³C₆-labelled 2,4-D was used as a model compound. An incubation experiment with *C. necator* JMP 134 grown on unlabelled 2,4-D and in the presence of ¹³CO₂ was prepared to investigate the incorporation of ¹³C into biomass components via CO₂ fixation.

5.2.1 Strain

The strain used for medium inoculation was *Cupriavidus necator* JMP 134. *C. necator* JMP 134 (formerly *Alcaligenes eutrophus, Ralstonia eutropha, Wautersia eutropha*) was isolated from an Australian soil by its ability to grow on 2,4-D (Don And Pemberton, 1981) and up to date is the best-studied 2,4-D-degrading soil \(\beta\)-proteobacterium (Manzano et al., 2007). *C. necator* JMP 134 is gram-negative bacterium with short rod-shaped cells [see microscopic cells visualised at 100x magnification using a laboratory microscope (Carl Zeiss, Jena, Germany) in **Figure 12**].

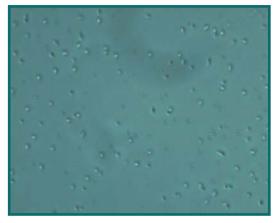


Figure 12. Microscopic view of single cells of Cupriavidus necator JMP 134

C. necator JMP 134 was tested in bioaugmentation procedures performed to remediate 2,4-D polluted environments (DEJONGHE ET AL., 2000; DIGIOVANNI ET AL., 1996). However, limited success of bioremediation with *C. necator* JMP 134 in soil was observed (DIGIOVANNI ET AL.,

1996). This is caused by the fact that various biotic (e.g. presence of eukaryotic microbiota) and abiotic factors in the complex soil environment may affect both the survival and catabolic performance of this strain in bioremediation processes (MANZANO ET AL., 2007).

5.2.2 Medium

The cultures were grown in a Minimal Medium (MM) prepared according to LERCH ET AL. (2007). 2,4-D was added as the sole C source at a total concentration of 73 mg/L (1978.3 µmol/L). MM components are presented in **Table 4**. MM was autoclaved (Systec Autoclave, Wettenberg, Germany) one time at 121°C for 20 min before use for the incubation experiment.

Table 4. MM components for cultivation of *C. necator* JMP 134 (LERCH ET AL., 2007)

Component	Concentration
K ₂ HPO ₄	1.5 g/L
KH_2PO_4	0.5 g/L
$(NH_4)_2SO_4$	1 g/L
$MgSO_4(H_2O)_7$	207 mg/L
$ZnSO_4 (H_2O)_7$	$200~\mu g/L$
$MgCl_2(H_2O)_4$	10 μg/L
H_3BO_3	5 μg/L
$CoCl_2(H_2O)_6$	25 μg/L
CuSO ₄	$100 \mu g/L$
$NiCl_2(H_2O)_6$	5 μg/L
$FeSO_4(H_2O)_7$	$250~\mu g/L$
EDTA	125 μg/L

5.2.3 Preculture incubation

Prior to the incubation experiment, to increase the biomass content and to adapt *C. necator* JMP 134 for 2,4-D degradation, this strain was pre-cultivated on tryptone medium (TM) as described previously by LERCH ET AL. (2007), which consisted of 5 g/L tryptone, 3 g/L yeast extract and 50 mg/L 2,4-D. Unlabelled 2,4-D dissolved in acetone (ACN) was placed in 500 mL Erlenmeyer flasks. In order to prevent killing the bacteria, the solvent was evaporated under the gentle air stream in a sterile bench. Both the tryptone and the yeast extract were dissolved in de-ionised water, autoclaved and transferred into the Erlenmeyer flasks after removal of ACN. After inoculation of the TM with cells of *C. necator* JMP 134, the batch preculture was incubated on a rotary shaker (~ 70 min⁻¹; Infors HT, Bottmingen, Switzerland) for 48 hours, in the dark and at 30°C. After 48 hours of incubation, an aliquot of 10 mL of the culture was centrifuged (11,000 min⁻¹ for 15 min at 4°C; Jouan KR25i, DBJ Labcare Ltd, Buckinghamshire, UK). After centrifugation, the biomass pellets were rinsed once with

Phosphate Buffer (PB) prior to the inoculation of the MM to avoid carryover of C from TM to MM.

5.2.4 Incubation experiment with ¹³C₆-2,4-D

Three different incubation experiments were prepared, in which MM (for details see section 5.2.2) was spiked with:

- 1. ¹³C₆-2,4-D (271 μmol/L) and ¹²C₆-2,4-D (1707.3 μmol/L). The ¹³C-labelled 2,4-D was mixed with the unlabelled one in order to increase the 2,4-D concentration for formation of the minimum biomass necessary for analysis (¹³C-labelled compounds are expensive, therefore ¹³C₆-2,4-D was diluted with unlabelled 2,4-D);
- 2. ¹²C₆-2,4-D (1978.3 μmol/L; control);
- 3. abiotic (sterile) containing ${}^{13}C_{6}$ -2,4-D (271 µmol/L) and ${}^{12}C_{6}$ -2,4-D (1707.3 µmol/L).

Control samples provided the information on the natural abundance of 13 C in the medium after addition of unlabelled 2,4-D. The abiotic system was performed to prove that the incorporation of 13 C-label into the biomass during 13 C₆-2,4-D degradation is a result of biotic processes. 13 C labelled or unlabelled 2,4-D were dissolved separately in ACN and then placed in 500 mL of Duran glass bottles. The ACN was evaporated and 250 mL (final volume of culture) of MM was placed in the glass bottles and the content of the bottle was mixed for 10 min. The medium in each bottle was inoculated with cells of *C. necator* JMP 134 (~ 10 mg/bottle) from the preculture on TM (see previous section 5.2.3). The abiotic systems were then autoclaved once. Duran glass bottles were sealed with Teflon-lined caps and incubated on the rotary shaker for 14 days under controlled laboratory conditions (in the dark and constant temperature $30 \pm 2^{\circ}$ C). To ensure a sufficient amount of O₂ for bacterial respiration, the systems were flushed periodically with humidified (bottle with de-ionised H₂O) and CO₂-free air. During aeration of systems, CO₂ evolved during 2,4-D mineralisation was trapped in two vials containing 2 M NaOH (see Figure 13).

The 13 C label incorporation into biomass components such as fatty acids (FA) and amino acids (AA) and into total biomass of *C. necator* during 13 C₆-2,4-D biodegradation were determined after 1, 2, 3, 7 and 14 days of incubation. In addition, to set up a detailed mass balance of 13 C in the system, the 13 C in the CO₂ produced during mineralisation of 2,4-D in the gas space of the bottle and in the medium was also monitored. At each sampling date, the optical density (OD) indicating bacterial biomass growth was determined using a Lambda 2S spectrophotometer UV/VIS ($\lambda = 560$ nm; Perkin Elmer, Massachutes, USA).

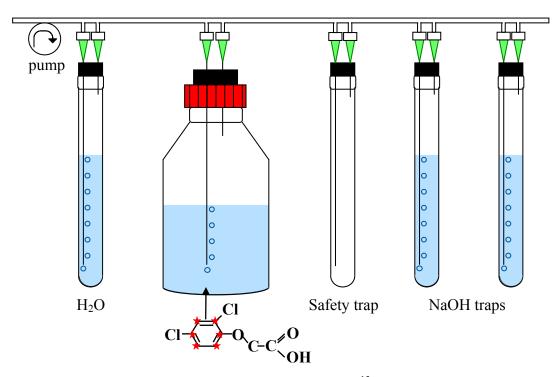


Figure 13. The liquid culture incubation experiment with ${}^{13}C_{6}$ -2,4-D in intermittent aeration systems

5.2.5 Incubation experiment under ¹³CO₂ atmosphere

To study the incorporation of C into biogenic residues via CO_2 released during 2,4-D mineralisation (CO_2 fixation), MM in Duran glass bottles spiked previously with unlabelled 2,4-D at a concentration of 1978.3 μ mol/L was incubated as three different incubation experiments under:

- 1. ¹³CO₂ (33.93 μmol/bottle);
- 2. ¹²CO₂ (33.93 µmol/bottle; control);
- 3. ¹³CO₂ (33.93 µmol/bottle) in an abiotic system (sterile).

Small tubes with either labelled Na₂¹³CO₃ or unlabelled Na₂CO₃ dissolved in de-ionised H₂O, were placed in each Duran glass bottle (**Figure 14**). To ensure conditions of ¹³CO₂ fixation experiment as close as possible to those in the ¹³C₆-2,4-D incubation experiment, the same concentration of unlabelled 2,4-D was used. The initial amount of CO₂ in each bottle was adjusted to 33.93 μmol and corresponded to an assumed 50% mineralisation of the initial amount of 2,4-D in the ¹³C₆-2,4-D experiment (67.75 μmol/bottle). After closing the Duran bottle, CO₂ was released from Na₂CO₃ by adding 6M HCl drop wise into this tube using a syringe. Thereafter, the bottle was incubated in the same way as for the ¹³C₆-2,4-D incubation experiment. The bottles were destructively sampled after 2, 3, 7 and 14 days and the ¹³C label was analysed in FA, AA and in the total biomass of *C. necator*. However, no samples were taken after 1 day of incubation as it was done for ¹³C₆-2,4-D incubation experiment. This is

caused by the fact that CO₂ fixation plays a secondary role at the initial phase of incubation. One triplicate set of liquid culture systems with ¹³CO₂ (33.93 µmol/bottle) and without 2,4-D for 7 days was incubated to check if CO₂ fixation takes place in the absence of the main C source. At each sampling date, the OD of the bacterial biomass suspension was measured.

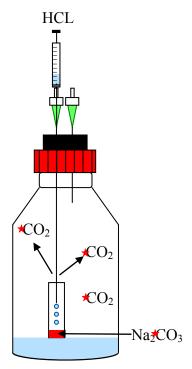


Figure 14. The preparation of ¹³CO₂ fixation liquid culture incubation experiment treated with unlabelled 2,4-D

5.3 Soil experiments

The formation of biogenic residues in soil systems during microbial degradation of two most widely used organic contaminants 2,4-D and Ibu was investigated. The herbicide 2,4-D was selected as a first model compound, due to numerous studies on its microbial degradation and reported high NER content in the soil. In spite of many studies on NER formation during 2,4-D biodegradation, their chemical structure is still not known and there is no report on the contribution of microorganisms to biogenic residues formation. Therefore, to trace biogenic residues formation directly from 2,4-D, ¹³C₆-2,4-D incubation experiments were prepared. In addition, also the indirect incorporation of C into biogenic residues via CO₂ fixation process during 2,4-D biodegradation was studied in the soils amended with unlabelled 2,4-D and incubated in presence of ¹³CO₂. The second model compound Ibu is a emerging contaminant, which both microbial degradation and the processes of NER formation are not studied in detail compared to 2,4-D. In order to investigate biogenic residues formation via the direct incorporation of C from Ibu, soil systems were incubated with ¹³C₆-ibu.

Summing up the above, the mechanisms of NER formation and thus the risks for an environment related to these two extensively used organic contaminants are not elucidated yet. Therefore, the knowledge on biogenic residues formation is necessary for a proper risk assessment and a realistic degradation rate of these contaminants in the soil.

5.3.1 Soil

The soil was collected from A horizon of the agricultural long-term experiment "Statischer Düngungversuch" located in Bad Lauchstädt, Germany (BLAIR ET AL., 2006). It is a Haplic Chernozem, which has been cultivated continuously with a crop rotation (sugar beet, summer barley, potatoes and winter wheat) and fertilised every second year with a farmyard manure (30 t/ha) since 1902. Detailed soil characteristics are presented in **Table 5**.

Table 5. Soil characteristics used for an incubation experiment (KÖRSCHENS ET AL., 2000)

Parameter	Content
Clay	21%
Silt	68%
Sand	11%
N_{total}	0.17%
C_{org}	2.1%
pН	6.6
WHC	37.5%

5.3.2 Incubation experiment

Three different biotic soil experiments with ¹³C₆-2,4-D, ¹³C₆-2,4-ibu or ¹³CO₂ were prepared. In order to distinguish between the abiotic and the biotic NER formations, one or two triplicate sets of abiotic systems for each experiment were also incubated. All soil incubation experiments were prepared according to the OECD guideline 307 addressing the aerobic and anaerobic transformation of chemicals in the soil (OECD, 2002). The experimental setup of the soil systems (**Figure 15**) was the same as for the liquid culture experiments (**Figure 13**; section 5.2.4).

To prevent killing all microorganisms, first only a small amount (10% of the total amount used for an experiment) of soil (13 C₆-2,4-D experiment) or soil-sludge mixture (13 C₆-ibu experiment) was spiked. Either 13 C-labelled or unlabelled model compound was dissolved in MeOH. This initially spiked portion was then mixed with the rest of the soil to yield a final compound concentration of 20 mg/kg. The water content of the soil was adjusted to 60% of its maximum water holding capacity (WHC). Finally 40 gram of this soil was weighed using laboratory balance (Sartorius AG, Göttingen, Germany) and placed in 1000 mL Duran glass bottles sealed with Teflon-lined caps.



Figure 15. Incubation experiment of soil treated with organic contaminant (according to OECD 307)

The soil systems were incubated in modified intermittent aeration systems under controlled laboratory conditions in the darkness and a constant temperature of 20°C (\pm 2°C). Similar to the liquid culture systems, soil systems were aerated periodically and the CO₂ was trapped in 2 M NaOH solutions.

5.3.3 Soil incubation experiment with ¹³C₆-2,4-D

Three different soil incubations were prepared:

- 1. soil spiked with 13 C₆-2,4-D,
- 2. soil spiked with ¹²C₆-2,4-D (control),
- 3. non-amended (blank, without 2,4-D).

Blank and control samples allow the correction for the natural abundance of 13 C in soil. The final concentration of spiked 2,4-D in the soil was 20 mg/kg (542.5 μ mol 13 C/kg) of soil. The soil was incubated for 64 days. After 2, 4, 8, 16, 32 and 64 days of incubation, the bottles were sampled destructively and the soil was analysed for the amount and the isotopic composition (13 C) of FA and AA as representatives for biogenic residues. In addition, two triplicate sets of abiotic soil systems with 13 C₆-2,4-D were also incubated and sampled destructively after 32 and 64 days.

5.3.4 Soil incubation experiment under ¹³CO₂ atmosphere

In order to investigate CO₂ fixation (from mineralisation of 2,4-D residues) by soil microorganisms and its incorporation into biogenic residues, three soil incubations were prepared:

- 1. with unlabelled 2,4-D (20 mg/kg) and $^{13}CO_2$ (12.6 μ mol/bottle);
- 2. with unlabelled 2,4-D (20 mg/kg) and unlabelled CO_2 (12.6 μ mol/bottle);

3. with neither 2,4-D nor additional CO₂.

The preparation of an incubation experiment under CO₂ atmosphere was already presented in detail in section 5.2.5 (**Figure 14**). The initial amount of CO₂ in each bottle was adjusted to the final concentration of ¹³CO₂ evolved during mineralisation of ¹³C₆-2,4-D. To estimate the incorporation of ¹³C-label into FA, the soil was extracted and analysed after 8, 16, 32 and 64 days of incubation. No samples were taken on days 2 and 4 due to a minor role of CO₂ fixation in the initial phase of incubation. In addition, because of a low level of the isotopic enrichment (in a range of the natural isotopic abundance) in the extracted AA and tFA from soils, they were not further quantified. One triplicate set of soil systems with ¹³CO₂ (12.6 μmol/bottle) without 2,4-D and abiotic systems with ¹³CO₂ and unlabelled 2,4-D were also incubated for 64 and 32 days, respectively. The soil systems without 2,4-D were prepared to check if the presence of this readily available organic substrate stimulates CO₂ fixation.

5.3.5 Soil incubation experiment with ¹³C₆-ibu

Similar to the soil incubation experiment with ¹³C₆-2,4-D, also three types of soil systems were prepared:

- 1. soil spiked with ¹³C₆-ibu;
- 2. soil spiked with ¹²C₆-ibu (control);
- 3. non-amended (blank, without Ibu).

Contrary to 13 C₆-2,4-D soil incubation experiment, the soil was amended additionally with stabilised sewage sludge (Klärwerk Rosental, Leipzig, Germany), because Ibu can only be introduced into the soil by sewage sludge application (EDWARDS ET AL., 2009). Soil and sludge were mixed at a 9:1 (DW/DW) ratio. The initial Ibu concentration was 20 mg/kg (581.5 μ mol/kg) of soil. The soil systems were incubated for 90 days and sampled destructively after 2, 7, 14, 28, 59 and 90 days of incubation and analysed for the amount and the isotopic composition of FA and AA. Additionally, two triplicate sets of abiotic soil systems with 13 C₆-2,4-ibu were also incubated and sampled destructively after 28 and 90 days of incubation.

5.4 Chemical analyses

The main focus of all experiments was to study the ¹³C incorporation into two representatives for biogenic residues, fatty acids (FA) and amino acids (AA). FA and AA are known microbial biomarkers (BOSCHKER AND MIDDELBURG, 2002), which represent the major fraction of the analytically recognisable compounds in SOM (ALLARD, 2006). Owing to FA

and AA extraction and their amount and isotopic composition analyses, we can outline the formation and the fate of biogenic residues in the complex soil system.

However, for a better understanding of biogenic residue formation, additional information on the ¹³C label distribution within a system (mass balance) is necessary. In order to set up a detailed mass balance of ¹³C in a system, the amount and the isotopic composition of the following components should be considered: CO₂, solvent-extractable parent compound and its metabolites and finally either ¹³C incorporation into total biomass (liquid culture experiment) or total NER (soil experiments).

5.4.1 CO₂ measurement

The $^{13}\text{CO}_2$ trapped in 2 M NaOH solution was quantified by measurement of the total inorganic C using a Total Organic Carbon Analyser (Schimadzu TOC-5050, Duisburg, Germany). The isotopic composition (atom %) of CO₂ was determined using a GC-Combustion-isotope ratio Mass Spectrometry (GC-C-irMS; Finnigan MAT 252, Thermo Electron, Bremen, Germany, coupled to Hewlett Packard 6890 GC, Agilent Technologies, Germany), equipped with Porabondt Q-HT Plot FS column (50 m × 0.32 m × 5 μ m; Chrompack, Middburg, Netherlands). In order to measure the isotopic composition, 2 ml of the NaOH solution was transferred to 15 ml crimp cap vials and acidified with 400 μ l of phosphoric acid (85%). Headspace samples of 100-250 μ l were then injected into the GC-C-irMS. CO₂ was separated isothermally at a temperature of 40°C. The temperature of the oxidation oven was 940°C and the one of the reduction oven was 640°C. Both the C content and isotopic composition of CO₂ trapped in MM after removal of *C. necator* biomass by centrifugation was determined in the same way as CO₂ trapped in the NaOH solutions.

5.4.2 Parent compound and metabolites measurement

¹³C₆-2,4-D or ¹³C₆-ibu residues were extracted from soil using an Accelerated Solvent Extraction (ASE) 200 system (Dionex, Sunnyvale, CA) equipped with 11 mL stainless steel extraction cells according to RADJENOVIĆ ET AL. (2009) with few modifications. 5 gram of soil was placed in the extraction cell and mixed with diatomaceous earth (Hydromatrix[™], Varian Associates, Inc., Hansen Way, Palo Alto, USA) and 20 μg of the internal standard 2-methyl-4-chlorophenoxyacetic acid (MCPA) used for quantitative analyses. The soil samples were extracted three times with methanol-water (1:1, v/v) at the following operating conditions: extraction temperature, 100°C; pressure 100 bar; preheating period, 5 min; static extraction period, 10 min; solvent flush, 150% of the cell volume; and nitrogen purge, 150 s.

The biomass of *C. necator* was separated from MM by cetrifugation (11,000 min⁻¹ for 15 min at 4°C, Jouan KR52i). The supernatant containing ¹³C₆-2,4-D and ¹³C₆-2,4-D-derived metabolites was then purified.

MM after centrifugation and soil extracts were acidified to pH 2 prior to purification by Solid Phase Extraction (SPE, Macherey Nagel, Düren, Germany). The SPE cartridges (CHROMABOND® EASY) were conditioned with 3 mL of MeOH and 3 mL of de-ionised H₂O. After the conditioning step, the acidified sample was passed slowly through the column applying a slight vacuum. The column then was washed with 10 mL of de-ionised H₂O and dried under vacuum using a vacuum pump (KNF Neuberger, Frankfurt, Germany) for 15 min. The elution of 2,4-D and its metabolites from column was performed 2 times with 5 mL of methanol (MeOH)-ACN mixture (1:1, v/v). A mixture of MeOH-tetrahydrofuran (1:1, v/v; 5 mL; 2 times) was used for the elution of ibu and ibu-derived metabolites. The eluates were then evaporated under a gentle stream of N₂. Dried sample was resuspended in the mixture of acetonitrile/trimethylsilyltrifluoroacetamide (1:2, v/v) for silylation. This mixture was then incubated for 10 min at 60 °C, next cooled down at room temperature and finally transferred into a GC vial.

Parent compounds and their metabolites were quantified and identified using Gas Chromatography-Mass Spectrometry (GC-MS, Hewlett Packard 6890 GC, Agilent Technologies, Germany) equipped with a BPX-5 column (30 m × 0.32 m × 0.25 μm, SGE, Darmstadt, Germany) with the following temperature program: initial temperature 60°C (1.5 min), heat to 120°C (0 min) at 20°C/min, to 160°C (0 min) at 4°C/min and finally to 260°C (5 min) at 16°C/min (ZWIENER AT EL., 2002). The injector was set at 260°C and the He flow was 1.5 mL/min. The identification of 2,4-D, 2,4-DCP, Ibu and 2-OH-ibu in sample was done by comparison of their retention times with those of a standard.

5.4.3 Total biomass and NER measurement

At each sampling date, bacterial cells of *C. necator* were harvested from MM by centrifugation (11,000 min⁻¹ for 15 min at 4°C; Jouan KR25i) and then freeze-dried (2 h; Christ RVC 2-25, Osterode am Harz, Germany). About of 2-5 mg of either air-dried soil sample (after removal of solvent-extractable parent compound and its metabolites) or freeze-dried biomass was weighed. Both the amount of ¹³C and isotopic composition in either total biomass or total NER were determined using Elemental Analyser-Combustion-isotope ratio Mass Spectrometry (EA-C-irMS; Finnigan MAT 253, Thermo Electron, Bremen, Germany) coupled to Euro EA 3000, Eurovector, Milano, Italy). The temperature of the oxidation oven

was 1020°C and the one of the reduction oven was 650°C. In addition, bulk extracts from soil before purification by SPE were also analysed for ¹³C content and the isotopic composition.

5.4.4 Fatty acids (FA) analyses

We analysed the ¹³C label incorporation into FA of living biomass of *C. necator* and into FA of living soil microbial biomass. To distinguish between ¹³C distributions in FA within the living (PLFA) and the non-living SOM fraction, total fatty acids (tFA) in the soil were also determined. The tFA contains FA both from the living and the dead biomass stabilised in SOM (non-living SOM; DRENOVSKY ET AL., 2004).

However, the PLFA extraction procedure from the soil microbial biomass is different than from the biomass of *C. necator* and includes two purification steps. Therefore, PLFA extractions from *C. necator* and the soil are presented below separately.

Biomass of *C. necator* was harvested from MM by centrifugation (3300 min⁻¹ for 15 min at 4°C; Jouan KR25i) prior to PLFA extraction. The PLFA were extracted with a mixture of PB/MeOH/chloroform from the *C. necator* (0.5 mL/1 mL/2 mL) according to BLIGH AND DYER (1959). The PB/MeOH/chloroform mixture containing cells of *C. necator* was votexed for 3 min (Vibromix®, Sartorius Stedim Biotech, Aubagne Cedex, France). After addition of H₂O (0.5 mL), the mixture was shaken again for 1 min and finally separated into two phases by centrifugation for 10 min at 1000 min⁻¹ (Jouan KR25i). The bottom phase (chloroform) containing the PLFA was evaporated under a gentle stream of N₂ prior to derivatisation

The PLFA were extracted from a soil using also the PB/MeOH/chloroform mixture (2 mL/5 mL/2.5 mL; BLIGH AND DYER, 1959). The soil suspended in the PB/MeOH/chloroform mixture was shaken for 2 h on a rotary shaker (Infors HT, Bottmingen, Switzerland). After shaking, 2.5 mL of each H₂O and chloroform was added to the mixture in order to separate into two phases (overnight). The bottom chloroform phase was dried using sodium sulfate and evaporated under N₂. This phase was then fractionated into neutral lipids, glycolipids and PLFA according to MILTNER ET AL. (2005) by chromatography over silica columns (Unisil) impregnated previously with ammonium acetate (0.02 M). To elute neutral lipids, glycolipids and PLFA chloroform (5 mL), ACN (5 mL) and MeOH (15 mL) were used, respectively. The MeOH was evaporated under N₂ and the PLFA were then derivatised.

The dried samples containing PLFA extracted from either *C. necator* or the soil were resuspended in the mixture of MeOH/Trimethylchlorosilane [(TMCS); 9:1; v:v] and methylated at 60°C in a oven (Memmert, Schwabach, Germany) for 2 h as described by THIEL ET AL., (2001). After cooling to room temperature, the excess of derivatisation reagent

was removed in the N_2 -gasstream. Finally, the dried sample was redissolved in 50 μ L of hexane prior to further gas chromatographic analysis.

The tFA in the soil were analysed by methylation with the same derivatisation reagent as for PLFA, (MeOH/TMCS; 1.8 mL/0.2 mL) added directly to the dry soil (THIEL ET AL., 2001). After derivatisation, methylated tFA were extracted from soil according to MILTNER, ET AL. (2005) with diethyl ether and purified over silica gel columns (Baker). The eluate was then evaporated under N_2 and dried sample was finally dissolved in 100 μ L of hexane prior to GC/MS and GC-C-irMS analyses.

The Fatty Acid Methyl Esters (FAME) in PLFA and tFA fractions were quantified and identified by GC/MS equipped with BPX-5 column as described in section 5.4.2. FA are designated as A:BωC where A is the number of C atoms, B is the number of double bonds and C is the position of the double bond relative to the aliphatic end of the molecule. The prefix "cy" assigns a cyclopropyl fatty acid, and branched fatty acids are indicated by i-, a- or 10Me.

For quantification, heneicosanoate methyl ester (FAME 21:0) was added to each sample as an internal standard prior to injection into the GC/MS. The FAME were separated using the following temperature program: initial temperature 50°C (hold for 1 min), heat to 250°C (0 min) at 4°C/min and finally to 300°C (10 min) at 2°C/min. The injector was set at 280°C, the He flow was 2.5 mL/min. FAME were identified by comparison of the retention times and the mass spectra with those of a standard mixture BAME (Bacterial Acid Methyl Esters). The isotopic composition of the FAME was determined GC-C-irMS (Finnigan MAT 253 coupled to Trace GC, Thermo Electron, Bremen, Germany) equipped with BPX-5 column (50 m \times 0.32 m \times 0.5 µm, SGE International, Darmstadt, Germany). The FAME were separated with the following temperature program for GC-C-irMS: initial temperature 70°C (1 min), heat to 130°C (0 min) at 20°C/min, to 150°C (5 min) at 2°C/min, to 165°C (5 min) at 2°C/min, to 230°C (0 min) at 2°C/min and finally to 300°C (5 min) at 20°C/min. The injector was set at 250°C and the He flow was 2 mL/min.

With the methylation, one additional C is introduced resulting in a shift of the isotopic composition. This shift was corrected for the original isotopic composition of the respective FA and calculated according to BOSCHKER (2004) as follows:

$$ICcorr = \frac{(n_{\textit{FAME}} + n_{\textit{MeOH}}) \times IC_{\textit{meas}} - IC_{\textit{MeOH}}}{n_{\textit{FAME}}}$$

where:

n_{FAME} – number of C in respective FAME without the methyl group

n_{MeOH} – number of C of MeOH

IC_{corr} – corrected isotopic composition of FAME without the methyl group

IC_{meas} – measured isotopic composition of FAME with the methyl group

IC_{MeOH} – C isotopic composition of MeOH

5.4.5 Amino acids (AA) analyses

¹³C-amino acids (AA) hydrolysed from proteins were analysed in soil in two fractions: in the total AA fraction (tAA) and in the living microbial biomass AA fraction (bioAA). As considered for the fatty acids, the difference between the total and the biomass AA fractions represents the refractory AA in the non-living SOM. AA in *C. necator* were hydrolysed directly from the biomass similarly as the tAA from the total soil, whereas bioAA in the soil were determined in the extracted living microbial biomass.

Prior to hydrolysis, bacterial cells were harvested from MM by centrifugation (3300 min⁻¹ for 15 min at 4°C; Jouan KR25i). AA in the biomass of *C. necator* and in the total soil (tAA) were hydrolysed with 6 M HCl for 22 hours at 110°C as described by MACKO AND UHLE (1997) and SILFER ET AL. (1991). After hydrolysis, biomass pellets and soil particles were discarded by filtering through glass-fibre membrane filters. The HCl was then evaporated in an N₂-gasstream and AA from biomass of *C. necator* were further derivatised, whereas the soil hydrolysate was purified.

After evaporation of the HCl excess, soil sample was resuspended in 4 mL of 0.1 M HCl and purified over cation exchange resin (DOWEX 50W-X8) according to AMELUNG AND ZHANG (2001). Cation exchange resin was prepared according to BOAS ET AL. (1953), by prerinsing with 2 M NaOH, conditioning with 2 M HCl, and washing with de-ionized H₂O until the pH of the eluate was neutral. The sample containing the AA was pipetted on top of the column. The column then was washed with oxalic acid (25 mL, pH 1.6-1.8) followed by 0.01 M HCl (5 mL) and de-ionised H₂O (5 mL). Finally the AA were eluted with 2.5 M NH₄OH (30 mL) and NH₄OH was evaporated. The dried samples were re-dissolved in 4 mL 0.1 M HCl and centrifuged for 15 min at 4200 min⁻¹ (Jouan KR25i). After centrifugation, the 0.1 M HCl excess was evaporated under N₂ prior to derivatisation.

The carboxyl groups of AA in dried sample were esterified with a mixture of Isopropanol/Acetylchloride (IP/AC; 1 mL/250 μ L) at 110°C for 1 h and the amino groups trifluoroacetylated with a mixture of Dichloromethane/Trifluroacetic Acid Anhydride (DCM/TFAA; 500 μ L/500 μ L) at 60°C for 1 h as described previously by MILTNER ET AL. (2005). After derivatisation, the samples were dried under N₂ and purified using a mixture of PB (1 mL) and chloroform (0.5 mL) as described previously by UEDA ET AL. (1989). The

mixture was shaken, centrifuged (Centrifuge 5415R eppendorf; Eppendorf, Hamburg, Germany) at 600 min⁻¹ for 10 min and the chloroform phase was finally evaporated. After drying, the samples containing the derivatised AA were re-dissolved in 100 μ L DCM for injection into the GC-MS.

For the determination of the AA in the living biomass (bioAA), the biomass was first extracted from soil as described by JACOBSEN AND RASMUSSEN (1992). The efficiency of the bacterial biomass extraction has been reported to be in the range of 40% by independent studies (JACOBSEN AND RASMUSSEN, 1992; MILTNER ET AL., 2009). Hence, bioAA are underestimated but the relative trends of bio-AA are correct because of the relatively constant recoveries even in different soils. The cells were extracted from soil by shaking using an end-over-end shaker (2h; GFL 3040, Fachhandel für Labor- und Medizintechnik GmbH, Kothen, Germany) according to MILTNER ET AL. (2009) with 1 g Chelex 100 (chelating cation exchange resin) and 10 mL of sodium deoxycholate/Polyethylenglycol 600 solution (0.1%/2.5%). After shaking, soil particles were removed using low speed centrifugation (960 min⁻¹, 15 min; Jouan KR25i) and finally the biomass in supernatant was harvested using high-speed centrifugation (11,000 min⁻¹, 1h, 4°C; Jouan KR25i). Biomass pellets containing AA were further hydrolysed, purified and derivatised in the same way as tAA.

The identity, quantity and isotopic composition of the AA were determined by GC/MS and GC-C-irMS. The temperature program of GC/MS was as follows: 70°C (5 min) to 100°C (30 min) at 30°C/min, to 175°C (5 min) at 10°C/min, to 250°C (5 min) at 10°C/min, to 340°C (15 min) at 30°C/min. The injector was set at 280°C; the He flow was 1.7 mL/min. For quantification and identification of respective AA in samples, an external standard containing all detectable AA in the samples [alanine (Ala), glycine (Gly), threonine (Thr), serine (Ser), valine (Val), leucine (Leu), isoleucine (Ile), proline (Pro), aspartate/asparagine (Asp), glutamate/glutamine (Glu), phenylalanine (Phe), and lysine (Lys)] was used. Asparagine and glutamine are desaminated during the acid hydrolysis, resulting in aspartate and glutamate, respectively. Therefore they are shown as Asp and Glu. The internal standard L-norleucine was added to each sample before hydrolysis to estimate the losses of AA in soil during the two purification steps. The isotopic composition of each AA was determined using GC-C-irMS with the same column as for the FAME and the following temperature program: 50°C (0 min) to 100°C (10 min) at 15°C/min, to 130°C (0 min) at 2°C/min, to 220°C (0 min) at 10°C/min, to 250°C (10 min) at 30°C/min. The injector was set at 250°C; the He flow was 2 mL/min.

The original isotopic enrichment of respective AA in samples, after shifting their isotopic composition during two derivatisation steps, was corrected using the following equation (SILFER ET AL., 1991):

$$IC_{corr} = \frac{IC_{meas} \times (n_{AA} + n_{TFAA} + n_{IP}) - IC_{TFAA} \times n_{TFAA} - IC_{IP} \times n_{IP}}{n_{AA}}$$

where:

n_{AA}– number of C in respective AA

n_{TFAA} - number of C of TFAA

n_{IP} - number of C of IP

IC_{corr} - corrected isotopic composition of AA

IC_{meas} – measured isotopic composition of AA

IC_{TFAA} – C isotopic composition of TFAA

IC_{IP} – C isotopic composition of IP

5.5 Data analyses

A stable isotope tracer (¹³C) was used to investigate the formation and the fate of biogenic residues during biodegradation of a target contaminant in a system. The application of ¹³C-labelled compounds in biodegradation studies allows the analysis of their structural assignments using more sophisticated analytical techniques such as GC-MS (RICHNOW ET AL., 1999). By extraction of known ¹³C-labelled biogenic residues representatives (FA and AA) formed during ¹³C-labelled compound biodegradation and inspection of their chemical structures using GC-MS, we can thus prove the microbial origin of NER in the soil.

However, the soil is also naturally abundant in 13 C ($\sim 1\%$), therefore blank (without compound application) and control (12 C-labelled compound) samples were used for the correction of 13 C abundance in a soil amended with the tested 13 C-labelled compound. Similarly to the soil, 13 C abundance in the liquid culture experiments was also corrected accordingly due to addition of 12 C-2,4-D to 13 C-2,4-D incubation experiments .

The contents of ¹³C-FA, ¹³C-AA, evolved ¹³CO₂ during mineralisation, ¹³C-parent compound and its ¹³C-metabolites and ¹³C in either total biomass or NER were presented as a percentage of the initially applied ¹³C-labelled compound to an experiment. In addition, a mass balance of ¹³C in the system was taken into consideration when interpreting the data on the ¹³C label incorporation into biogenic residues.

All incubation experiments were prepared in triplicates and results are shown as averages. The error bars represent the standard deviation of these triplicates.

6 RESULTS

The incorporation of ¹³C label into biogenic residue representatives (FA and AA) during microbial degradation of ¹³C-labelled compounds was studied in two different systems: simple liquid pure culture experiments and complex soil systems.

The main objective of liquid culture experiments was to trace the fate of 13 C label derived from 13 C₆-2,4-D in the simple biological system with a particular focus on 13 C-incorporation into the biomass of the 2,4-D degrader. Due to simplicity, these experiments were also designed to estimate the ratio of 13 C-FA and 13 C-AA to the total biomass 13 C content, which are necessary for a proper quantification of the total amount of biogenic residues in the soil systems.

The formation and the fate of biogenic residues during biodegradation of two different model compounds ($^{13}C_6$ -2,4-D and $^{13}C_6$ -ibu) were investigated in complex soil experiments. The general objective was to estimate the extent of biogenic residue formation and their amount in relation to total NER and finally to assess the risk for the environment related to NER formation from readily degradable organic contaminants in soil.

6.1 Liquid culture experiments

Cupriavidus necator JMP 134 was selected for the biodegradation studies, because it is a well-described 2,4-D soil degrader. C. necator grows on 2,4-D as C source. The ¹³C label incorporation into FA and AA of C. necator during microbial degradation of 2,4-D was investigated directly from ¹³C₆-2,4-D and indirectly via CO₂ fixation from ¹³CO₂

The detailed description of PLFA classes typical for specific groups of microorganisms, their designation and AA abbreviations used for data interpretation have already been presented in the sections 2.6.2.1, 5.4.4 and 5.4.5, respectively.

6.1.1 Incubation experiment with ¹³C₆-2,4-D

The ¹³C incorporation into the biomass of *C. necator* during biodegradation of ¹³C₆-2,4-D was studied over 14 days. The ¹³C content in FA, AA, in total biomass and the ¹³C mass balance in the system were analysed after 1, 2, 3, 7 and 14 days of incubation. The abiotic experiment with *C. necator* and ¹³C₆-2,4-D showed neither ¹³C₆-2,4-D biodegradation nor ¹³C incorporation into FA and AA of *C. necator* after 14 days, indicating the importance of bacterial activity in the degradation of this compound and in the formation of ¹³C-labelled FA and AA.

6.1.1.1 Mass balance of ¹³C₆-2,4-D in the system

The Duran glass bottle (500 mL) was aerated with humidified air prior to CO_2 measurement in order to trap the evolved CO_2 in a headspace into NaOH solution. However, due to reported high overpressure in the small headspace volume of about 250 mL and the need to connect the bottle to a pump, some CO_2 was lost before aeration. The experiment with *C. necator* JMP134 grown on unlabelled 2,4-D and under $^{13}CO_2$ atmosphere, in which the initial $^{13}CO_2$ concentration was known, revealed that at each sampling date $\sim 70\%$ of the initial $^{13}CO_2$ was lost (data not shown). Therefore, the data on the mineralisation kinetics presented in **Figure 16** are highly underestimated.

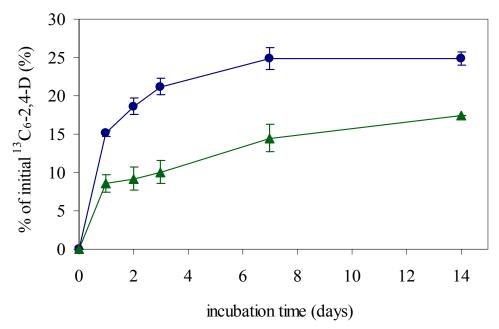


Figure 16. Distribution of 13 C label in an incubation experiment with A *C. necator* JMP 134 grown on 13 C₆-2,4-D medium. Mineralisation () and biomass (). Extractable 13 C₆-2,4-D and metabolites not detectable since day 1

After the first day of incubation, 15% of 13 C derived from 13 C₆-2,4-D was already mineralised (50% of 13 C₆-2,4-D equivalents after correction for 70% loss). At that time, neither 13 C₆-2,4-D nor its known 13 C-aromatic ring metabolites were detected, indicating their complete utilisation. After day 1, the content of 13 CO₂ evolved during 2,4-D mineralisation increased slowly and finally on the day 7 reached a plateau ($\sim 83\%$ of 13 C₆-2,4-D equivalents after correction for 70% loss). The incorporation of 13 C from 13 C₆-2,4-D into the biomass started immediately (after 1 day) and their contents remained constant until day 3 ($\sim 10\%$ of 13 C₆-2,4-D equivalents; see **Table 6**). Thereafter, the amount of 13 C found in the total biomass increased slowly reaching finally 17.4% of the added 13 C₆-2,4-D equivalents. However, based on the OD measurement of a culture suspension, the biomass growth of *C. necator* JMP134

reached a plateau already after 1 day of incubation (see **Figure 17**), demonstrating clearly that ¹³C-incorporation into the biomass was not growth-depending.

Table 6. Incorporation of	¹³ C into the biomass o	f C. necator JMP	134 grown on	$^{13}C_6$ -2,4-D
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	Label incorporation [% of ¹³ C ₆ -2,4-D]				
Incubation time (days)	1	2	3	7	14
biomass	9.5 (± 1.1)	9.2 (±1.5)	$10.0 (\pm 1.5)$	14.5 (± 1.8)	$17.4 (\pm 0.06)$
PLFA	$0.3 (\pm 0.05)$	$0.5 (\pm 0.02)$	$0.6 (\pm 0.01)$	$0.8 (\pm 0.05)$	$0.6 (\pm 0.04)$
AA	$6.0 (\pm 0.07)$	$4.7(\pm 0.2)$	$6.1 (\pm 0.3)$	$7.3 (\pm 0.2)$	$8.1 (\pm 0.4)$
PLFA/Biomass AA/Biomass	3.1 63.2	5.4 51.1	6 61	5.5 54.3	3.4 46.6

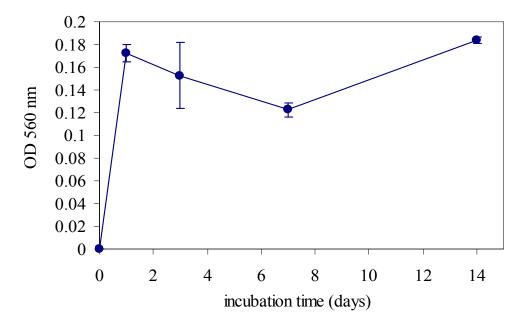


Figure 17. Time course of the biomass growth of *C. necator* JMP 134 on $^{13}C_6$ -2,4-D medium based on the OD measurement (560 nm)

 13 C incorporation into FA and AA of *C. necator* JMP 134 was observed in the 13 C₆-2,4-D incubation experiment. The contents of 13 C-PLFA in *C. necator* represented about 3–6% of the total 13 C in the biomass (see **Table 6**), which is in a good accordance to ~ 5% in FA in the microbial biomass reported by BAS ET AL. (2003). The amounts of 13 C-AA of *C. necator* were in the range of 46–63% of the total amount of 13 C in the biomass, which is consistent with the protein content in microbial biomass (50–55%) reported by MADIGAN AND MARTINKO (2006).

6.1.1.2 Incorporation of ¹³C into PLFA of C. necator JMP 134 grown on ¹³C₆-2,4-D

The 13 C label was rapidly incorporated into PLFA (on day 1), all classes of PLFA received 13 C label and remained enriched until the end of the experiment (see **Figure 18**). The amounts of 13 C in the PLFA increased strongly reaching a maximum on day 7 (0.8% of 13 C₆-2,4-D equivalents). Thereafter, their contents decreased until day 14 (0.6% of the initially 13 C₆-2,4-D added). At the beginning (day 1), the monounsaturated PLFA contained slightly more 13 C label than the saturated straight chain PLFA and the saturated cyclopropyl PLFA. On the day 2, the contents of 13 C in the saturated cyclopropyl PLFA increased strongly reaching a maximum on the day 7 (~ 0.2% of 13 C₆-2,4-D equivalents) and this class of PLFA remained dominant until the end of the incubation period.

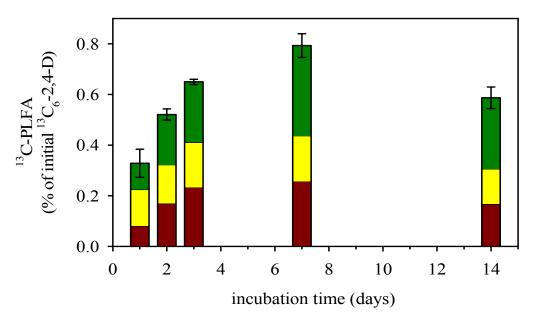


Figure 18. Distribution of 13 C label in lipid classes of PLFA of *C. necator* JMP 134 grown on 13 C₆-2,4-D medium. [saturated straight-chain (\blacksquare), monounsaturated (\square), and saturated cyclopropyl PLFA (\blacksquare)]

6.1.1.3 Incorporation of ¹³C into AA of C. necator JMP 134 grown on ¹³C₆-2,4-D

The incorporation of ¹³C label into AA started already on the first day (**Table 7**) and the contents of ¹³C-AA were constant until day 3 (excluding the ¹³C-AA on day 2). Thereafter, the amount of ¹³C in AA increased slightly, finally reaching 8.13% of ¹³C₆-2,4-D equivalents. The ¹³C label was distributed within all AA and all ¹³C-AA contained ¹³C label at the end of the incubation. Both ¹³C-Asp and ¹³C-Glu were dominant AA throughout the incubation period. After 3 days, the amount of ¹³C in Glu and Asp decreased slightly, whereas ¹³C-Thr, ¹³C-Val, ¹³C-Leu, ¹³C-Ile and ¹³C-Pro increased strongly. After 7 days, the contents of ¹³C in

Threo and Val increased again, while ¹³C-Phe decreased. At the end, ¹³C-Thr and ¹³C-Gly decreased (~ 2 times and 4 times, respectively), while the amount of ¹³C in Leu, Ile, Pro, Asp, Glu and Phe increased. At the same time, a very low content of ¹³C label was also detected in Ser.

Table 7. Distribution of the 13 C label in various 13 C-AA of *C. necator* JMP 134 grown on 13 C₆-2.4-D medium

	¹³ C-AA [% of ¹³ C ₆ -2,4-D]				
Incubation time (days)	1	2	3	7	14
Alanine	n.d.	n.d.	$0.43 (\pm 0.6)$	$0.46 (\pm 0.07)$	$0.28 (\pm 0.08)$
Glycine	$0.07 (\pm 0.01)$	$0.16 (\pm 0.02)$	n.d.	$0.53 (\pm 0.04)$	0.13 ↓ (± 0.02)
Threonine	$0.25 (\pm 0.004)$	$0.24 (\pm 0.11)$	0.65 ↑ (± 0.08)	0.86 ↑ (± 0.002)	0.46 ↓ (± 0.05)
Valine	$0.09 (\pm 0.09)$	$0.12 (\pm 0.06)$	0.27 ↑ (± 0.03)	0.6 ↑ (± 0.004)	$0.56 (\pm 0.09)$
Serine	n.d.	n.d.	n.d.	n.d.	$0.003 (\pm 0.11)$
Leucine	$0.61 (\pm 0.003)$	$0.62 (\pm 0.04)$	0.94 ↑ (± 0.02)	$0.82 (\pm 0.1)$	1.2 ↑ (± 0.04)
Isoleucine	$0.14 (\pm 0.02)$	$0.13 (\pm 0.01)$	0.30 ↑ (± 0.009)	$0.32 (\pm 0.01)$	0.38 ↑ (± 0.01)
Proline	$0.55 (\pm 0.02)$	$0.39 (\pm 0.03)$	0.51 ↑ (± 0.07)	$0.59 (\pm 0.02)$	0.73 ↑ (± 0.03)
Aspartate ^a	$1.19 (\pm 0.03)$	1.17 (\pm 0.24)	1.01 \downarrow (± 0.07)	$1.07 (\pm 0.02)$	1.56 ↑ (± 0.32)
Glutamate ^b	$1.52 (\pm 0.08)$	0.98 ↓ (± 0.06)	$0.89 \downarrow (\pm 0.18)$	$1.07 (\pm 0.004)$	1.53 ↑ (± 0.33)
Phenylanine	$0.77 (\pm 0.02)$	$0.59 (\pm 0.04)$	$0.76 (\pm 0.14)$	0.58 ↓ (± 0.03)	0.77 ↑ (± 0.04)
Lysine	$0.85 (\pm 0.005)$	$0.25 (\pm 0.02)$	$0.35 (\pm 0.17)$	$0.44 (\pm 0.12)$	$0.49 (\pm 0.01)$
TOTAL	6.06 (± 0.07)	4.67 (± 0.17)	6.13 (± 0.34)	7.34 (± 0.19)	8.13 (± 0.41)

^a incl. asparagine; ^b incl. glutamine; n.d. - not detectable; values are shown as averages ± standard deviation values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective AA compared to the preceeding sampling time

6.1.2 Incubation experiment under ¹³CO₂ atmosphere

The incorporation of ¹³C label from ¹³CO₂ into the biomass of *C. necator* JMP 134 grown on unlabelled 2,4-D, but in the presence of ¹³CO₂ was studied over 14 days. The content and the isotopic composition of ¹³C in PLFA, AA and in total biomass were analysed after 2, 3, 7 and 14 days of incubation. The main objective of this experiment was to study the relevance of heterotrophic CO₂ fixation in ¹³C label incorporation into the biomass of *C. necator* during 2,4-D biodegradation. However, the first sampling date (day 1) considered in the incubation experiment with ¹³C₆-2,4-D was excluded and not analysed, because CO₂ fixation plays usually a minor role in the initial phase of biodegradation experiment. One triplicate set of the biotic incubation experiment with ¹³CO₂ and without 2,4-D showed no incorporation of ¹³C into the biomass of *C. necator*, demonstrating strong dependence of ¹³CO₂ fixation on the presence 2,4-D in the medium. The abiotic incubation experiment also revealed no ¹³C incorporation into the PLFA and AA of *C. necator* after 7 days, indicating the importance of biological activity for CO₂ fixation.

6.1.2.1 Incorporation of ¹³C into the biomass of C. necator JMP 134 grown under ¹³CO₂

A very low ¹³C incorporation into the biomass of *C. necator* JMP 134 from ¹³CO₂ was observed already after 2 days of incubation (0.07% of the initially ¹³CO₂ added; see **Table 8**). Thereafter, the amount of ¹³C in the total biomass was increasing progressively throughout the incubation time, finally reaching 4.2% of the added ¹³CO₂ equivalents. The biomass grew rapidly from day 0 until day 3 (see **Figure 19**), indicating the high importance of CO₂ fixation in the later phase of the incubation experiment. Thereafter, the biomass growth reached a plateau.

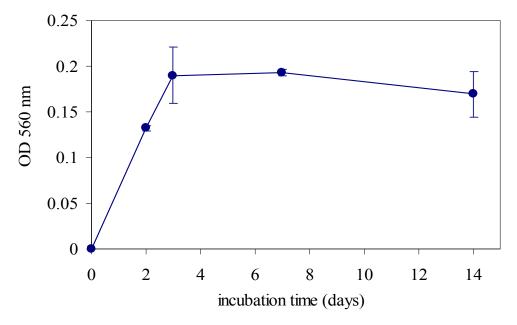


Figure 19. Time course of the biomass growth of *C. necator* JMP 134 on unlabelled 2,4-D medium and under ¹³CO₂ based on the OD measurement (560 nm)

The ¹³C label was incorporated into PLFA and AA of *C. necator* JMP 134 in the ¹³CO₂ fixation experiment. The amount of ¹³C-PLFA in *C. necator* JMP 134 was in the range of 3–10% of the total ¹³C in the biomass (excluding the day 2 where the amount of PLFA was very low; see **Table 8**), what is in good agreement with the data of BAS ET AL. (2003) and with those of the ¹³C₆-2,4-D experiment with *C. necator*. ¹³C-AA in *C. necator* JMP 134 represented 42–62% of the total ¹³C in the biomass, which is close to the percentage of proteins in biomass reported by MADIGAN AND MARTINKO (2006) and to the ¹³C₆-2,4-D experiment with *C. necator*.

Table 8. Incorporation of 13 C into biomass of *C. necator* JMP 134 grown on unlabelled 2,4-D medium and under 13 CO₂

	¹³ C-AA [% of ¹³ CO ₂]				
Incubation time (days)	2	3	7	14	
biomass	$0.07 (\pm 0.5)$	1.6 (±0.4)	$3.5 (\pm 0.2)$	4.2 (± 0.2)	
PLFA	$0.01 (\pm 0.004)$	$0.17 (\pm 0.02)$	$0.15 (\pm 0.008)$	$0.12 (\pm 0.003)$	
AA	$0.02 (\pm 0.07)$	$1.0 (\pm 0.4)$	$1.5 (\pm 0.3)$	$1.8 (\pm 0.08)$	
PLFA/Biomass AA/Biomass	14.3 28.6	10.5 62.5	4.3 42.9	2.9 42.9	

6.1.2.2 Incorporation of ¹³C into PLFA of C. necator JMP 134 grown under ¹³CO₂

The ¹³C derived from ¹³CO₂ was incorporated into the PLFA immediately on day 2, only the monounsaturated PLFA and the saturated straight chain PLFA carried ¹³C label (see **Figure 20**).

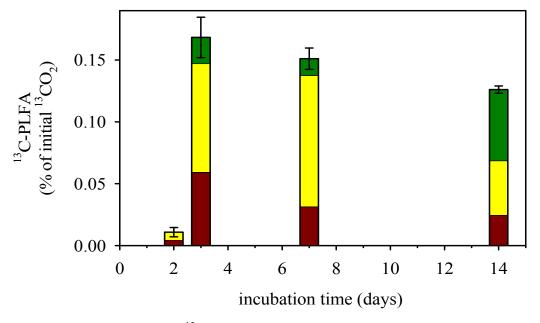


Figure 20. Distribution of ¹³C-label in lipid classes of PLFA of *C. necator* JMP 134 grown on unlabelled 2,4-D medium and under ¹³CO₂.

[saturated straight-chain (), monounsaturated (), and saturated cyclopropyl PLFA ()

The contents of ¹³C-PLFA increased rapidly, reaching their maximum on the day 3 (0.17% of the initially ¹³CO₂ added) and remained constant until day 7. At the end, the ¹³C-PLFA decreased to 0.12% of the initially ¹³CO₂ added. From day 3 onwards, the ¹³C label was distributed within all classes of PLFA and on both day 3 and 7, the monounsaturated PLFA dominated over other classes of PLFA. Similarly as observed in the ¹³C₆-2,4-D experiment,

the amount of ¹³C in the cyclopropyl PLFA increased strongly at the end of experiment, indicating the onset of starvation (KAUR ET AL., 2005). At the same time, the content of ¹³C-monounsaturated PLFA decreased to about half of the value at day 3.

6.1.2.3 Incorporation of ¹³C into AA of C. necator JMP 134 grown under ¹³CO₂

The incorporation of ¹³C started on the second day of incubation and this label was detected only in Asp (see **Table 9**). After 2 days, a continuous increase of ¹³C in AA until the end of incubation was observed and finally on day 14, ¹³C-AA reached 1.85% of the initially added ¹³CO₂. After 3 days, Thr contained most ¹³C and dominated over all other AA until the end, whereas ¹³C-Glu and ¹³C-Asp were also significantly enriched. After 7 days, the amounts of ¹³C in Thr, Asp, Ile increased, but the ¹³C in Glu and Phe decreased. At the end, an increase of ¹³C in Val and Leu was observed, while the amount of the label in Thr decreased.

Table 9. Distribution of the ¹³C label in various ¹³C-AA of *C. necator* JMP 134 grown on unlabelled 2,4-D medium and under ¹³CO₂

	¹³ C-AA [% of ¹³ C ₆ -2,4-D]						
Incubation	2	3	7	14			
time (days)							
Alanine	n.d.	n.d.	n.d.	$0.07 (\pm 0.04)$			
Glycine	n.d.	$0.02 (\pm 0.01)$	$0.04 \ (\pm \ 0.04)$	n.d.			
Threonine	n.d.	0.32 ↑ (± 0.13)	0.66 ↑ (± 0.03)	0.46 ↓ (± 0.03)			
Valine	n.d.	n.d.	$0.03 \ (\pm \ 0.01)$	0.26 ↑ (± 0.01)			
Serine	n.d.	n.d.	n.d.	n.d.			
Leucine	n.d.	$0.07 (\pm 0.02)$	$0.08 (\pm 0.06)$	0.16 ↑ (± 0.004)			
Isoleucine	n.d.	$0.01 (\pm 0.002)$	0.12 ↑ (± 0.01)	$0.12 (\pm 0.01)$			
Proline	n.d.	$0.05 (\pm 0.02)$	$0.05 (\pm 0.009)$	n.d.			
Aspartate ^a	$0.02 (\pm 0.07)$	0.19 ↑ (± 0.05)	0.29 ↑ (± 0.005)	$0.26 (\pm 0.02)$			
Glutamate ^b	n.d.	$0.19 (\pm 0.02)$	0.14 \downarrow (± 0.02)	$0.13 (\pm 0.0008)$			
Phenylanine	n.d.	$0.13 (\pm 0.09)$	0.07 ↓ (± 0.002)	$0.11 (\pm 0.007)$			
Lysine	n.d.	n.d.	n.d.	$0.25 (\pm 0.04)$			
TOTAL	$0.02~(\pm~0.07)$	$0.97 (\pm 0.42)$	1.49 (± 0.29)	1.85 (± 0.08)			

a incl. asparagine; b incl. glutamine; n.d. - not detectable; values are shown as averages ± standard deviation values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective AA compared to the preceeding sampling time

6.2 Soil experiments

The formation and the fate of biogenic residues in soil systems were investigated with two various organic contaminants, ${}^{13}C_6$ -2,4-D and ${}^{13}C_6$ -ibu. Soil samples were analysed for the amounts and isotopic compositions of FA and AA. The supplemental data on the ${}^{13}C$ mass balance in the soil, which includes mineralisation, solvent-extractable ${}^{13}C_6$ -2,4-D or ${}^{13}C_6$ -ibu and their metabolites, and total NER were kindly provided by Cristobal Girardi (Dept. Environmental Biotechnology, UFZ, Leipzig; for details see also GIRARDI ET AL., submitted).

The labelled biomarkers were extracted from the living part of SOM as ¹³C-PLFA or as ¹³C-bioAA, whereas the total fractions were directly analysed from soil (¹³C-tFA and ¹³C-tAA). The data presented in this section as FA or AA in the non-living SOM fraction were calculated by subtracting the amounts of the living fractions (PLFA or bioAA) from the amounts of total fraction (tFA or tAA). The detailed description of FA classes typical for specific groups of microorganisms, their designations and AA abbreviations used for the data interpretation have already been presented in the sections 2.6.2.1, 5.4.4 and 5.4.5, respectively.

6.2.1 Soil incubation experiment with ¹³C₆-2,4-D

The formation and the fate of FA and AA during microbial degradation of 2,4-D in the soil system were studied in two different soil experiments. In the first experiment, which focused on tracing 13 C label incorporation into FA and AA directly from the labelled herbicide, the soil was amended with 13 C₆-2,4-D. In the second experiment, the soil was treated with unlabelled 2,4-D and incubated under 13 CO₂ in order to study the indirect 13 C incorporation into PLFA via heterotrophic CO₂ fixation.

The formation and fate of FA and AA in the soil experiment with $^{13}C_6$ -2,4-D was studied over 64 days. These microbial biomarkers were analysed for their ^{13}C content and isotopic composition after 2, 4, 8, 16, 32 and 64 days of incubation. No incorporation of ^{13}C into FA and AA after 32 and 64 days was observed in sterile soil experiments incubated with $^{13}C_6$ -2,4-D, indicating their biotic formation. The results on the formation of FA and AA and their fate in the soil incubated with $^{13}C_6$ -2,4-D were published in Nowak ET AL. (in press).

6.2.1.1 Mass balance of ${}^{13}C_{6}$ -2,4-D in the soil

The kinetic of 13 C distribution within the soil fractions during microbial degradation of 13 C₆-2,4-D is presented in **Figure 21** (GIRARDI ET AL., submitted). The degradation of 13 C₆-2,4-D in soil started immediately without an apparent lag phase, and the readily available 2,4-D was present at high concentrations until day 4 (30.4% of the 13 C₆-2,4-D added). From day 8 onwards, only residual amounts of 13 C₆-2,4-D were detected in soil (0.5–2.3% of the initially 13 C₆-2,4-D added). At the end of incubation, $\sim 58\%$ of the initial 13 C label in the soil was evolved as 13 CO₂, $\sim 36\%$ was measured as NER and only 0.5% constituted the solvent extractable 13 C₆-2,4-D (after purification by SPE). The total amounts of unidentified solvent extractable (before purification) were relatively high until day 8 (40–90% of 13 C₆-2,4-D equivalents) and decreased to $\sim 8\%$ of 13 C₆-2,4-D equivalents on the day 64. The solvent extractable 2,4-D residues were extracted from soil samples with a MeOH/de-ionised H₂O

mixture (1:1, v/v) using ASE, which is a relatively harsh extraction method (NORTHCOTT AND JONES, 2000). Therefore, the high solvent extractable 2,4-D residues could be assigned to easily extractable biogenic residues before their stabilisation in SOM. The total recovery of 13 C from 13 C₆-2,4-D on day 64 was ~ 102.5%.

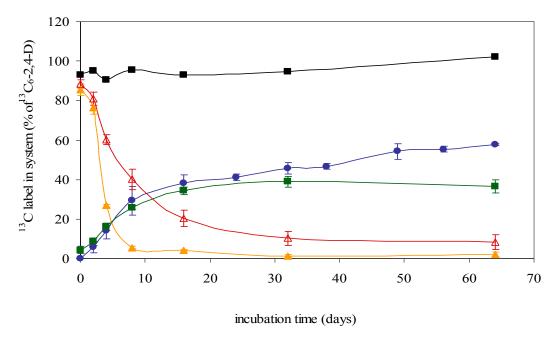


Figure 21. Distribution of the initial 13 C label from 13 C₆-2,4-D (%) in the system within 64 days of incubation experiment (GIRARDI ET AL., submitted). Mineralisation (\bigcirc), extractable amount before purification (\triangle), extractable amount after purification (\triangle), NER (\blacksquare) and recovery (\blacksquare)

No mineralisation was found in the abiotic soil experiments after 64 days, the solvent extractable ${}^{13}\text{C}_6\text{-}2,4\text{-D}$ and ${}^{13}\text{C}_6\text{-}2,4\text{-D}$ CP accounted for about of 60% of ${}^{13}\text{C}_6\text{-}2,4\text{-D}$ equivalents (GIRARDI ET AL., submitted). At the same time, NER amounted to about 19% of the ${}^{13}\text{C}_6\text{-}2,4\text{-D}$ equivalents, indicating for their formation even in absence of living microrganisms.

6.2.1.2 Formation of FA and their fate in soil incubated with ¹³C₆-2,4-D

The 13 C label was found only in short chain FA (C_{14} - C_{20}), which are typical for bacteria (CRANWELL, 1974). These organisms were thus involved in 2,4-D degradation in soil. The incorporation of 13 C label into the PLFA during the degradation of 13 C₆-2,4-D started already after 2 days of incubation and the PLFA remained enriched in 13 C until the end of the incubation experiment (**Figure 22A**).

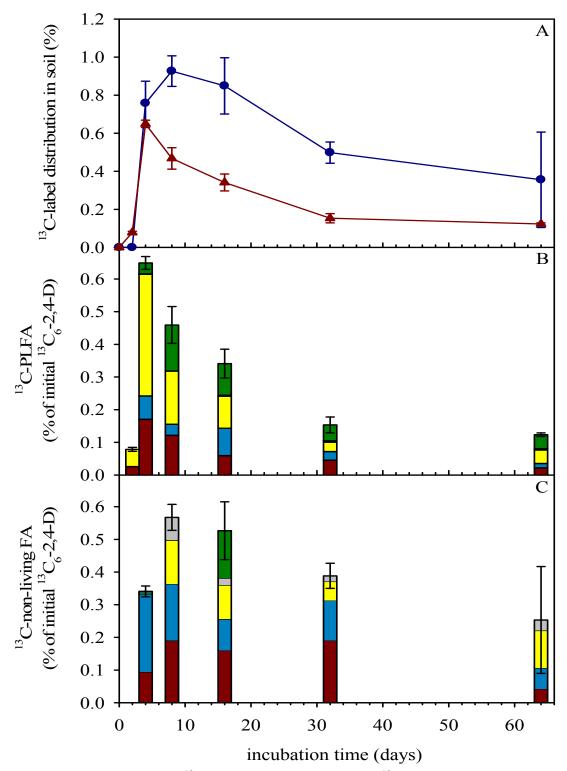


Figure 22. Time course of 13 C label incorporation during 13 C₆-2,4-D microbial degradation in soil (**A**) within tFA (**O**) and PLFA (**A**). Distribution of 13 C-label in lipid classes of PLFA (**B**) and t-FA in non-living SOM (**C**).

[saturated straight-chain (), saturated methyl-branched (), monounsaturated (), polyunsaturated () and saturated cyclopropyl FA ()

The maximum content of 13 C in PLFA ($\sim 0.7\%$ of the initially added 13 C₆-2,4-D equivalents) was detected already on day 4. Thereafter, the amounts of 13 C-labelled PLFA from membranes of the living microbial biomass decreased continuously until day 32 ($\sim 0.2\%$ of 13 C₆-2,4-D equivalents). After 4 days, a simultaneous flux of 13 C from the living biomass to the non-living fraction of SOM was observed. On days 8 and 16, the contents of 13 C-FA in the non-living fraction of SOM reached a maximum ($\sim 0.5\%$ of the initial 13 C₆-2,4-D amount in soil). The 13 C-FA in the non-living SOM fraction started to decline after 16 days. After day 32, already about of 70% of the label found in the 13 C-tFA could be assigned to the non-living SOM. At the same time, the contents of both 13 C-PLFA and 13 C-FA in the non-living SOM fraction remained nearly constant until the end of incubation. At the end of the experiment, the 13 C-FA in the non-living fraction declined to 50% and the 13 C-PLFA to 20% of their maximum, indicating clearly the degradation of biomass-derived residues in SOM.

During the initial phase of ¹³C₆-2,4-D degradation, the ¹³C label was found only in the saturated straight chain and in the monounsaturated PLFA, indicating that Gram-negative bacteria were the initial degraders of this herbicide (**Figure 22B**; see also **Table 10**). On day 4, when the content of ¹³C-PLFA was highest, the monounsaturated PLFA contained most of the label in comparison to other classes of PLFA, and 16:1ω7c PLFA was the dominant PLFA. On day 4, when a continuous decrease of ¹³C-PLFA in soil was observed, the incorporation of ¹³C label into the saturated cyclopropyl PLFA started. After 8 days, when the amounts of ¹³C-monounsaturated PLFA (e.g. 16:1ω7c, 16:1ω5c, 18:1ω7c and 18:1ω5c) declined strongly, the ¹³C-cy 17:0 carried most of the label (except for 16:0) and dominated until the end of incubation period. At the same time, the labelled saturated straight-chain PLFA, saturated methyl-branched PLFA and the monounsaturated PLFA from the living biomass were already partly incorporated into the non-living SOM (**Figure 22C**). However, the ¹³C cyclopropyl PLFA were not incorporated into the non-living SOM (apart from day 16), what might be explained by the fact that this class of FA became labelled later than other FA.

Table 10. The classification and composition of ¹³C label in respective PLFA during microbial degradation of ¹³C₆-2,4-D in soil

Incubation	¹³ C-PLFA [10 ⁻⁴ ; % of ¹³ C ₆ -2,4-D]						
time (days)	2	4	8	16	32	64	
Saturated straigl	nt chain						
14:0	$23.6 (\pm 2.4)$	$25.0 (\pm 25.3)$	n.d.	n.d.	n.d.	n.d.	
15:0	n.d. ^a	n.d.	$3.6 (\pm 2.7)$	n.d.	n.d.	n.d.	
16:0	$232 (\pm 1.7)$	$1615 (\pm 19.6)$	$1185.3 (\pm 161.5)$	$552.6 (\pm 41.4)$	$403.6 (\pm 23.9)$	$220.4 (\pm 11.7)$	
18:0	n.d.	$57 (\pm 31.9)$	$25.9 (\pm 42.5)$	$32.2 (\pm 6.5)$	$31.6 (\pm 29.7)$	n.d.	
20:0	n.d.	$6.8 (\pm 10.5)$	n.d.	$6.6 (\pm 9.3)$	$19.7 (\pm 9.3)$	n.d.	
Saturated methy	l branched					_	
i ^b 14:0	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.	
i 15:0	n.d.	$227.8 (\pm 58.6)$	$42.7 (\pm 42.6)$	$25.6 (\pm 10.5)$	$23.7 (\pm 37.5)$	$51.7 (\pm 5.1)$	
a ^c 15:0	n.d.	$142.1 (\pm 46.7)$	$12.2 (\pm 38.6)$	$43.4 (\pm 3.3)$	$7.8 (\pm 22.9)$	$28.7 (\pm 5.7)$	
i 16:0	n.d.	$64.8 (\pm 12.7)$	$4.9 (\pm 49)$	n.d.	n.d.	n.d.	
i 17:0	n.d.	$99.6 (\pm 5.2)$	$20.2 (\pm 4.2)$	$22.3 (\pm 0.13)$	$7.2 (\pm 15.6)$	$26.2 (\pm 9.6)$	
a 17:0	n.d.	$33.9 (\pm 11.1)$	$13.6 (\pm 0.8)$	n.d.	n.d.	$26.3 (\pm 29.4)$	
10 Me ^d 16:0	n.d.	n.d.	$158.8 (\pm 2)$	n.d.	$19.4 (\pm 30.6)$	n.d.	
10 Me 17:0	n.d.	n.d.	$28.5 (\pm 6)$	$32.9 (\pm 10.4)$	$2.5 (\pm 0.2)$	n.d.	
10 Me 18:0	n.d.	$77.4 (\pm 18)$	$32.4 (\pm 28.9)$	$39.7 (\pm 3.5)$	$2.5 (\pm 6.8)$	n.d.	
18:0 br ^f	n.d	$71.6 (\pm 0.07)$	$25.6 (\pm 3.2)$	$674.2 (\pm 531.4)$	$199.9 (\pm 24.9)$	n.d.	
Monounsaturate	d						
$16:1\omega7c^{c}$	$139.3 (\pm 38.4)$	2100.6 (± 72)	671.9 \downarrow (± 70.4)	$215.9 (\pm 9.0)$	$115 (\pm 22)$	$136.3 (\pm 7.7)$	
16:1ω5c	$56.9 (\pm 43)$	$292.2 (\pm 5.1)$	73.1 \downarrow (± 15.4)	n.d.	n.d.	$30.5 (\pm 2.2)$	
18:1ω9c	$59.7 (\pm 10.2)$	$201.9 (\pm 39.6)$	$42.4 (\pm 4.3)$	$143 (\pm 29.7)$	n.d.	n.d.	
18:1ω7c	$252.6 (\pm 2)$	$988.7 (\pm 18.9)$	779.1 \downarrow (± 89.8)	$589.5 (\pm 30.2)$	$173.2 (\pm 70.2)$	$232.8 (\pm 23.6)$	
18:1ω5c	$20.1 (\pm 5)$	$63.2 (\pm 9.3)$	9.1 ↓ (± 3.5)	n.d.	n.d.	n.d.	
Monounsaturate	d branched						
15:1 br	n.d.	$27.7 (\pm 11.5)$	$16.2 (\pm 6.8)$	$10.4 (\pm 8.2)$	$0.6 (\pm 5.4)$	$4.4 (\pm 2.1)$	
17:1 br	n.d.	$51.0 (\pm 5.5)$	$36.0 (\pm 2.7)$	$28.1 (\pm 10)$	n.d.	10.7 (± 18)	
Polyunsaturated							
18:2	n.d.	$1.5 (\pm 12)$	$1.6 (\pm 0.02)$	$7.1 (\pm 0.8)$	n.d.	n.d.	
18:2ω6,9	n.d.	$10.3 (\pm 12.5)$	n.d.	$15.7 (\pm 23)$	39.9 (± 12.9)	35.2 (± 73.7)	
Saturated cyclop	ropyl						
cy 17:0	n.d.	239 (± 59.2)	1292 ↑ (± 80)	$893.8 (\pm 49.2)$	$457.8 (\pm 11.4)$	$335.7 (\pm 6.4)$	
cy 19:0	n.d.	$94.2 (\pm 24.8)$	118.8 (± 27.5)	$78.7 (\pm 4.0)$	$30.2 (\pm 26.9)$	94.7 (± 33.2)	
TOTAL	784 (± 102.8)	6486 (± 511)	4677 (± 683.7)	3412 (± 775.2)	1535 (± 350.3)	1234 (± 228.6)	

^aa- anteiso; ^bbr – branched; ^cc - cis isomer; ^di – iso; ^en.d. - not detectable; ^fMe - methyl; values in brackets (±) represent the standard deviation of the average of triplicate; values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective FA compared to the preceding sampling time

6.2.1.3 Formation of AA and their fate in soil incubated with ¹³C₆-2,4-D

The incorporation of the ¹³C label into the living biomass AA fraction (bioAA) during microbial degradation of 2,4-D in soil started 2 days later than into the PLFA (**Figure 23**).

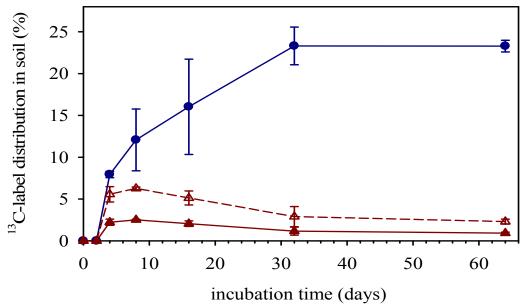


Figure 23. Time course of 13 C label incorporation within tAA (\bigcirc), bioAA (\triangle) and bioAA (\triangle) recalculated based on 40% extraction efficiency (for details see section 5.4.5) during microbial degradation of 13 C₆-2,4-D in soil

The amounts of 13 C-bioAA were highest on days 4 and 8 (2.2% and 2.5%, respectively, of the initially 13 C₆-2,4-D equivalents added). From day 16, a decay of 13 C-bioAA was observed and at that time the majority ($\sim 70\%$) of the 13 C-tAA was already detected in the non-living part of SOM, although it clearly derived from the living biomass. A 13 C flux of AA from the living to the non-living fraction of SOM started after 4 days as it was also observed for PLFA. After 32 days, the contents of both 13 C-bioAA and 13 C-AA in the non-living SOM did not change significantly until the end of the experiment. At the end of the incubation (day 64), the contents of 13 C-AA in the living biomass fraction of SOM were 64% lower (0.9% of 13 C₆-2,4-D equivalents) than their maximum detected in this experiment. Finally, 90% of the 13 C-tAA in soil was stabilised in the non-living SOM pool, reaching 22% of the initial content of 13 C₆-2,4-D equivalents added to soil. On day 4, when the incorporation of 13 C label into bioAA fraction was initiated, Glu carried most of this label (1.9% of the initial amount 13 C₆-2,4-D in soil; see **Table 11A**) among the 13 C-AA. At the same time, much lower amounts of 13 C were found in Asp and Val (each $\sim 0.2\%$ of the initially 13 C₆-2,4-D added).

Table 11. Distribution of the ¹³C label in various ¹³C-bioAA (**A**) and ¹³C-AA in the non-living SOM (**B**) during microbial degradation of ¹³C₆-2,4-D in soil

<u>A</u>	¹³ C-AA [% of ¹³ C ₆ -2,4-D]								
Incubation time (days)	2	4	8	16	32	64			
Alanine	n.d.	n.d.	n.d.	$0.10 (\pm 0.05)$	$0.04 (\pm 0.10)$	$0.05 (\pm 0.03)$			
Glycine	n.d.	n.d.	n.d.	$0.19 (\pm 0.09)$	$0.02 (\pm 0.14)$	n.d.			
Threonine	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.			
Valine	n.d.	$0.16 (\pm 0.04)$	$0.13 (\pm 0.04)$	$0.34 (\pm 0)$	$0.10 (\pm 0.11)$	$0.04 (\pm 0.01)$			
ß-alanine	n.d.	n.d.	n.d.	n.d.	$0.002 (\pm 0.01)$	$0.004 (\pm 0.007)$			
Leucine	n.d.	n.d.	$0.38 (\pm 0.03)$	$0.15 (\pm 0.03)$	$0.10 (\pm 0.93)$	$0.15 (\pm 0.01)$			
Isoleucine	n.d.	$0.02 (\pm 0.05)$	$0.08 (\pm 0.04)$	$0.04 (\pm 0)$	$0.12 (\pm 0.44)$	$0.07 (\pm 0.03)$			
Proline	n.d.	n.d.	$0.35 (\pm 0.03)$	$0.26 (\pm 0.1)$	$0.21 (\pm 0.66)$	$0.18 (\pm 0.05)$			
Aspartate ^a	n.d.	$0.14 (\pm 0.14)$	0.60 ↑ (± 0.17)	$0.27 (\pm 0.47)$	n.d.	n.d.			
Glutamate ^b	n.d.	$1.90 (\pm 0.50)$	0.23 \downarrow (± 0.06)	$0.25 (\pm 0.58)$	$0.11 (\pm 0.17)$	n.d.			
Phenylanine	n.d.	n.d.	$0.31 (\pm 0.04)$	$0.16 (\pm 0.17)$	$0.15 (\pm 0.50)$	$0.34 (\pm 0.09)$			
Lysine	n.d.	n.d.	$0.42 (\pm 0.01)$	$0.30 (\pm 0.14)$	$0.31 (\pm 0.29)$	$0.08 (\pm 0.03)$			

 $2.50 (\pm 0.07)$

 $2.06 (\pm 0.33)$

 $1.16 (\pm 0.48)$

 $0.91 (\pm 0.12)$

 $2.22 (\pm 0.37)$

n.d.

TOTAL

	¹³ C-AA [% of ¹³ C ₆ -2,4-D]								
Incubation time (days)	2	4	8	16	32	64			
Alanine	n.d.	$0.5 (\pm 4.4)$	$3.9 (\pm 2.7)$	$1.0 (\pm 7.4)$	1.0 (± 1.3)	$1.6 (\pm 4.0)$			
Glycine	n.d.	n.d.	n.d.	n.d.	n.d.	n.d.			
Threonine	n.d.	n.d.	n.d.	n.d.	$0.01 (\pm 0.01)$	$0.3 (\pm 0.1)$			
Valine	n.d.	$2.6 (\pm 1.8)$	n.d.	$0.7 (\pm 2.1)$	n.d.	$0.9 (\pm 0.7)$			
ß-alanine	n.d.	n.d.	n.d.	$0.3 (\pm 2.0)$	$0.2 (\pm 0.9)$	$0.3 (\pm 0.2)$			
Leucine	n.d.	n.d.	n.d.	$3.4 (\pm 0.4)$	$1.8 (\pm 5.5)$	$3.7 (\pm 0.5)$			
Isoleucine	n.d.	$2.3 (\pm 1.5)$	n.d.	$0.5 (\pm 1.7)$	$0.6 (\pm 1.7)$	$1.8 (\pm 0.3)$			
Proline	n.d.	n.d.	n.d.	n.d.	$3.8 (\pm 5.1)$	$2.5 (\pm 0.02)$			
Aspartate ^a	n.d.	$1.6 (\pm 3.7)$	$0.5 (\pm 0.8)$	$1.3 (\pm 0.1)$	$1.6 (\pm 0.7)$	$1.3 (\pm 0.5)$			
Glutamate ^b	n.d.	n.d.	n.d.	$3.6 (\pm 0.03)$	$2.7 (\pm 1.1)$	$1.4 (\pm 0.2)$			
Phenylanine	n.d.	n.d.	$0.2 (\pm 0.6)$	$0.5 (\pm 0.2)$	$5.4 (\pm 1.6)$	$4.8 (\pm 0.2)$			
Lysine	n.d.	n.d.	$1.5 (\pm 2.3)$	$0.02 (\pm 0.01)$	$1.5 (\pm 2.9)$	$1.2 (\pm 0.3)$			
TOTAL	n.d.	7.0 (± 0.3)	6.1 (± 1.9)	11.3 (± 4.1)	18.6 (± 1.8)	19.8 (± 0.6)			

a incl. asparagine; b incl. glutamine; n.d. - not detectable; values are shown as averages ± standard deviation values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective AA compared to the preceeding sampling time

After 8 days, when the content of ¹³C in Glu decreased strongly, the ¹³C-Asp increased rapidly and this bioAA was dominant over the other ¹³C-AA, indicating heterotrophic CO₂ fixation (MILTNER ET AL., 2005). In addition, on day 8, the ¹³C label was also distributed to lesser extent within Val, Leu, Ile, Pro, Phe and Lys. On day 16, when the ¹³C-bioAA decreased, the incorporation of ¹³C into two additional AA Ala and Gly, was observed. The ¹³C-AA from the living biomass fraction were also transferred into the non-living SOM pool (see **Table 11B**). After 32 days of incubation, when the contents of both ¹³C-bioAA and ¹³C-AA in the non-living SOM did not change significantly, the distribution of the ¹³C label in the individual AA in the non-living and living SOM fraction remained nearly constant until the end of the experiment.

6.2.1.4 Biogenic residues in soil incubated with ¹³C₆-2,4-D

The FA represent about 5% of the microbial biomass (BAS ET AL., 2003); this was also proved in the liquid experiments with C. necator JMP 134 grown on 2,4-D, which clearly showed that 13 C-labelled FA represented $\sim 5\%$ of the total 13 C in the biomass (see sections 6.1.1.2 and 6.1.2.1). Therefore, considering this content of FA in bacterial cells, we can conclude from the highest amounts of the 13 C-tFA (about 1%) determined on day 8 that at least 20% biomass in 13 C₆-2,4-D experiment had been formed. The proteins containing AA are the most abundant components (50–55% of the microbial biomass) of bacterial cells (MADIGAN AND MARTINKO, 2006), which is also consistent with the results obtained in the experiment with C. necator (sections 6.1.1.1 and 6.1.2.1). Therefore, taking into consideration this conversion factor of 2, we can conclude from the amounts of 22% found in the 13 C-tAA at the end of soil incubation, that there were 44% biogenic residues in this experiment.

6.2.2 Soil incubation experiment under ¹³CO₂ atmosphere

The ¹³CO₂ evolved during mineralisation of ¹³C₆-2,4-D in the soil may be used as an additional C source by microorganisms for the synthesis of their cellular components (KREBS, 1941), which are also finally incorporated into SOM. Therefore, the objective of this soil experiment was to study the contribution of heterotrophic CO₂ fixation to the formation of biogenic residues during 2,4-D biodegradation and to distinguish from ¹³C incorporation into these residues directly from the labelled 2,4-D.

The formation and the fate of ¹³C-PLFA in soil spiked with non-labelled 2,4-D and incubated under ¹³CO₂ was studied over 64 days. The labelled CO₂ was added in an amount equivalent to the final release of CO₂ evolved during biodegradation of labelled 2,4-D in the soil. One triplicate set of abiotic incubation experiments with ¹³CO₂ and with unlabelled 2,4-D showed no incorporation of the label into the PLFA, indicating the relevance of microorganisms in the CO₂ fixation process in the soil.

6.2.2.1 Formation of PLFA and their fate in soil incubated under ¹³CO₂ atmosphere

GIRARDI ET AL. (submitted; see also section 6.2.1.1) reported that the readily available $^{13}C_6$ -2,4-D was at high concentrations in soils until day 4; thereafter, only traces of this herbicide were detected. The CO_2 fixation usually plays a minor role during first days of incubation, when the readily available C source is still present at high concentrations. Therefore, two sampling dates (days 2 and 4), considered in the investigation of the formation and fate of PLFA in soils incubated with $^{13}C_6$ -2,4-D were excluded and not analysed in the $^{13}CO_2$ fixation experiments. The incorporation of ^{13}C label into PLFA during heterotrophic

CO₂ fixation was compared with the amounts of ¹³C-PLFA found in the ¹³C₆-2,4-D experiment. The distribution of this label derived from ¹³C₆-2,4-D or ¹³CO₂ during 2,4-D degradation in these two experiments is presented in **Figure 24A**.

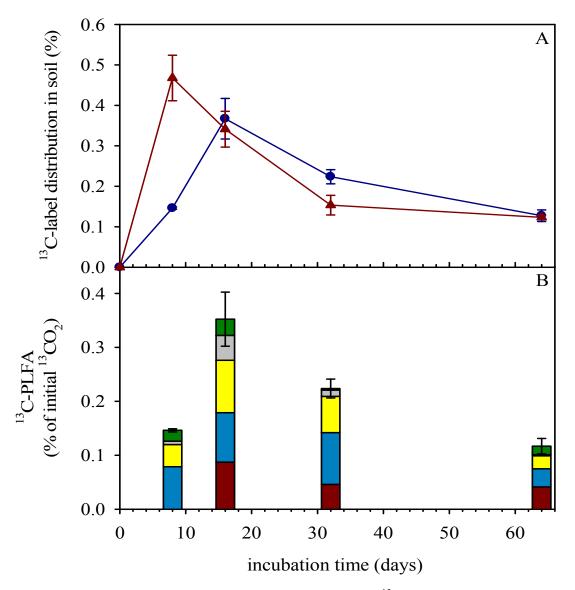


Figure 24. Comparison of the distribution of the 13 C label within PLFA (**A**) from biodegradation of 13 C₆-2,4-D (\blacktriangle) and from 13 CO₂ fixation (\bullet) and within classes of PLFA after 13 CO₂ fixation (**B**).

[saturated straight-chain (), saturated methyl-branched (), monounsaturated (), polyunsaturated () and saturated cyclopropyl FA ()

After 8 days of incubation, the incorporation of 13 C into the PLFA from 13 C₆-2,4-D was still three times higher (~ 0.5% of the initial 13 C₆-2,4-D equivalents added) than that from 13 CO₂ fixation (0.15% of the initial 13 CO₂ in system). The highest incorporation of 13 C into the

PLFA via heterotrophic CO_2 fixation was observed on day 16 (~ 0.4% of the initial $^{13}CO_2$) and was in the same range as the ^{13}C incorporation into PLFA from $^{13}C_6$ -2,4-D at that time. On day 16, a decrease of both ^{13}C -PLFA and ^{13}C -FA in the non-living SOM fraction was observed in the experiment with $^{13}C_6$ -2,4-D. This indicates that in the absence of the readily degradable $^{13}C_6$ -2,4-D in soil, the ^{13}C derived from these FA was recycled via heterotrophic CO_2 fixation, which reached a maximum on that day. From day 32, the CO_2 fixation decreased slowly until the end of the incubation period.

The ¹³C label derived from ¹³CO₂ was incorporated to all classes of PLFA (**Figure 24B**). After 8 days, the saturated methyl-branched PLFA (e.g. i:15:0, a:15:0, i:17:0 and 10Me16:0; see **Table 12**) contained more ¹³C label than the monounsaturated PLFA (except for 18:1ω7c), in particular in comparison to 16:1ω7c, which was the dominant PLFA detected in the soil incubated with ¹³C₆-2,4-D (**Table 10**). After 16 days of incubation, the amounts of ¹³C in the monounsaturated PLFA and saturated methyl-branched PLFA were nearly similar, indicating the contribution of both Gram-negative and Gram-positive bacteria to CO₂ fixation. At that time, however, more ¹³C from ¹³CO₂ was incorporated into the monounsaturated 18:1ω7c and 18:1ω9c and into the polyunsaturated 18:2ω6,9 in comparison to the other ¹³C-PLFA. From day 32, the methyl-branched PLFA contained slightly more ¹³C from ¹³CO₂ than the monounsaturated PLFA until the end of experiment.

The presence of a readily available organic substrate, e.g. acetic acid, stimulates CO_2 fixation (MILTNER ET AL., 2005). In order to estimate this effect on the fixation process, we also incubated one triplicate set of soil systems with $^{13}CO_2$ without the herbicide for 64 days. After 64 days of incubation, the incorporation of ^{13}C into the PLFA from CO_2 in the absence of 2,4-D was 10 times lower (0.013% of the initial $^{13}CO_2$) than the amounts fixed by soil microorganisms grown in the presence of this readily degradable compound (0.13% of the initial $^{13}C_6$ -2,4-D).

Table 12. The classification and composition of ¹³C label in respective PLFA during ¹³CO₂ fixation in soil.

Incubation	¹³ C-PLFA [10 ⁻⁴ ; % of ¹³ CO ₂]							
time (days)	8	16	32	64				
Saturated stra	ight chain							
14:0	n.d.e	n.d.	$37.0 (\pm 35.5)$	$25.3 (\pm 7)$				
15:0	n.d.	$62.8 (\pm 45.9)$	n.d.	n.d.				
16:0	n.d.	$459.6 (\pm 209.1)$	$348.5 (\pm 14.6)$	209.1 (±19.3)				
18:0	n.d.	$200.9 (\pm 99.3)$	$10.1 (\pm 15.2)$	122.1 (±22.3)				
20:0	$8.5 (\pm 1.4)$	$149.8 (\pm 84.2)$	$64 (\pm 12.4)$	55.9 (±18.6)				
Saturated met	hyl branched			_				
^d i 14:0	n.d.	n.d.	n.d.	n.d.				
i 15:0	231.9 (± 31.7)	$132.8 (\pm 146.2)$	$167 (\pm 33.3)$	$33.3 (\pm 5.5)$				
^a a 15:0	136.2 (± 19.4)	$78.8 (\pm 200.1)$	$90.3 (\pm 20.5)$	$41.1 (\pm 2.2)$				
i 16:0	n.d.	$110.2 (\pm 64.8)$	$190.7 (\pm 34.1)$	$8.5 (\pm 7)$				
i 17:0	107.6 (\pm 23.9)	n.d.	$97 (\pm 4.8)$	$9.4 (\pm 1.4)$				
a 17:0	$68.3 (\pm 15.6)$	$189.8 (\pm 121.9)$	n.d.	$1.5 (\pm 23.9)$				
10 Me ^f 16:0	105.3 (\pm 33.8)	$217.5 (\pm 89.6)$	$247.2 (\pm 61.7)$	$39.7 (\pm 6)$				
10 Me 17:0	$28.6 (\pm 26.3)$	$34.0 (\pm 4.9)$	$48.9 (\pm 0.9)$	$56.1 (\pm 0.8)$				
10 Me 18:0	$65.3 (\pm 3.9)$	$70.1 (\pm 99.8)$	$54 (\pm 3.8)$	$101.9 (\pm 29.2)$				
18:0 br ^b	$36.1 (\pm 25.9)$	$84.2 (\pm 69.5)$	$64.2 (\pm 0.9)$	45.1 (± 13.3)				
Monounsatura								
$16:1\omega7c^{c}$	$56.6 (\pm 25.1)$	$52.4 (\pm 104.7)$	$75 (\pm 25.7)$	$46.7 (\pm 39.8)$				
16:1ω5c	$41.4 (\pm 19.7)$	$63.2 (\pm 72.7)$	$69.2 (\pm 0.6)$	$27.9 (\pm 12.3)$				
18:1ω9c	n.d.	326.3 (± 164.3)	$113.3 (\pm 65.4)$	n.d.				
18:1ω7c	261.2 (± 35.2)	347.5 (\pm 70.2)	$237.1 (\pm 118.2)$	$87.7 (\pm 138.9)$				
18:1ω5c	n.d.	22.31 (± 18.8)	$109.6 (\pm 0.5)$	62.9 (± 13.2)				
Monounsatura								
15:1br	$12.8 (\pm 8.1)$	$80.4 (\pm 16.6)$	$1.6 (\pm 4.8)$	$6.0 (\pm 9.7)$				
17:1br	38.9 (± 14.1)	$77.8 (\pm 55.7)$	66.4 (± 13.4)	$10.1 (\pm 3.9)$				
Polyunsaturat								
18:2	$18.3 (\pm 19.9)$	n.d.	$118.3 (\pm 13.3)$	$19.3 (\pm 17.0)$				
18:2ω6,9	$45 (\pm 48.8)$	462.4 (± 92.9)	n.d.	n.d.				
Saturated cycl								
cy 17:0	$91.5 (\pm 13.4)$	$121.6 (\pm 18.8)$	$1.6 (\pm 14.6)$	$159.7 (\pm 18.5)$				
cy 19:0	$107.8 (\pm 50.9)$	$179 (\pm 202.6)$	$25.8 (\pm 7.5)$	n.d.				
TOTAL	1462 (± 29.7)	$3524 (\pm 501.4)$	2237 (± 175.3)	$1169 (\pm 143.1)$				

a- anteiso; br – branched; c - cis isomer; i – iso; n.d. - not detectable; Me - methyl; ; values in brackets (±) represent the standard deviation of the average of triplicates; values printed in bold represent characteristic values

6.2.3 Soil incubation experiment ¹³C₆-ibu

The formation and fate of biogenic residues during microbial degradation of $^{13}C_6$ -ibu in the soil were studied over 90 days. The content and the isotopic composition of FA and AA were determined after 2, 7, 14, 28, 59 and 90 days of incubation. The supporting data on the mass balance of ^{13}C derived from $^{13}C_6$ -ibu in the soil system were provided by Cristobal Girardi (see also GIRARDI ET AL., submitted). Similar to the $^{13}C_6$ -2,4-D experiment, no ^{13}C incorporation into FA and AA was observed in the abiotic soil experiments with $^{13}C_6$ -ibu.

6.2.3.1 Mass balance of ${}^{13}C_6$ -ibu in the soil

The distribution of the 13 C in the soil system during microbial degradation of 13 C₆-ibu is shown in **Figure 25** (GIRARDI ET AL., submitted).

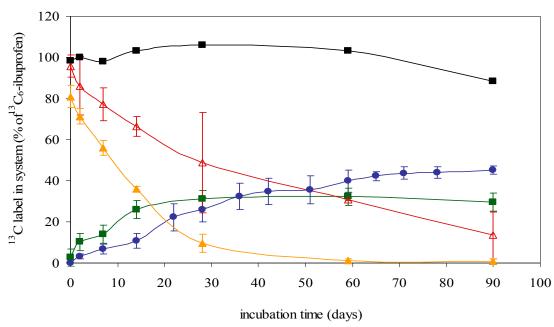


Figure 25. Distribution of the initial 13 C label from 13 C₆-ibu (%) in the system over 90 days of incubation experiment (GIRARDI ET AL., submitted). Mineralisation (\bigcirc), extractable amount before purification (\triangle), extractable amount after purification (\triangle), NER (\blacksquare) and recovery (\blacksquare)

The degradation of 13 C₆-ibu started later than 13 C₆-2,4-D degradation. The solvent extractable 13 C₆-ibu and its metabolites were detected until day 59 ($\sim 1\%$ of 13 C₆-ibu equivalents). At the end, 45% of the initial 13 C label in soil was mineralised, 30% was found as NER and only 0.5% was identified as 13 C₆-ibu (after purification by SPE). The content of solvent bulk extractable 13 C (before purification) was high until day 59 (48% of 13 C₆-ibu equivalents) and its decrease was observed on day 90 (13% of 13 C₆-ibu equivalents). The solvent extractable Ibu residues were extracted from soil using the same method as used for 13 C₆-2,4-D. The total recovery of 13 C at the end was much lower (88% of the initial 13 C in the system) than in the 13 C₆-2,4-D soil experiment.

No mineralisation of $^{13}C_6$ -ibu after 90 days was seen in the abiotic soil experiments, the solvent extractable fraction accounted for 65% of the initially added, $^{13}C_6$ -ibu and $\sim 16\%$ of $^{13}C_6$ -ibu equivalents was found as NER (GIRARDI ET AL., submitted).

6.2.3.2 Formation of FA and their fate in soil incubated with ¹³C₆-ibu

The incorporation of ¹³C label into PLFA started immediately and they were enriched until the end of the incubation time. During the first 28 days of incubation, the ¹³C label was detected only in the membranes of living biomass (PLFA, see **Figure 26A**).

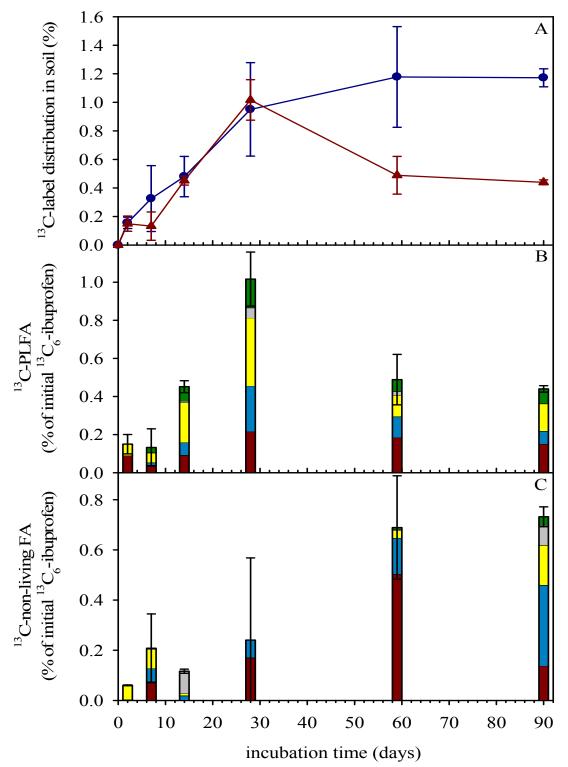


Figure 26. Time course of 13 C label incorporation during microbial degradation of 13 C₆-ibu in soil (**A**) within tFA () and PLFA (). Distribution of 13 C label in lipid classes of PLFA (**B**) and tFA in non-living SOM (**C**).

[saturated straight-chain (■), saturated methyl-branched (□), monounsaturated (□), polyunsaturated (□) and saturated cyclopropyl FA (■)]

On day 28, the ¹³C-PLFA reached their maximum (1% of ¹³C₆-ibu equivalents). After 28 days, the ¹³C-PLFA decreased significantly and a ¹³C flux from the living part to the non-living fraction of SOM was observed. After 59 days, the content of ¹³C-PLFA decreased to 50% of their maximum (~ 0.5% of ¹³C₆-ibu equivalents), whereas the ¹³C-FA in the non-living fraction were highest and reached 0.7% of the initial ¹³C₆-ibu amounts in soil. After day 59, about of 60% of the label in the t-FA could be already assigned to the non-living SOM. From this time, the amounts of both ¹³C-PLFA and ¹³C-FA in the non-living SOM were nearly constant until the end of the experiment.

At the beginning of the incubation, the ¹³C label was found in the saturated straight chain and in the monounsaturated PLFA (16:1ω7c, 16:1ω5c, 18:1ω9c and 18:1ω7c, **Figure 26B**, see also **Table 13**). From day 7 until day 28, the monounsaturated PLFA carried more ¹³C label than the other PLFA classes. During that time the amounts of ¹³C-monounsaturated (in particular 16:1ω7c 18:1ω9c and 18:1ω7c) and also the ¹³C-saturated cyclopropyl (cy 17:0) increased rapidly. However, incorporation of ¹³C into the saturated methyl branched PLFA started already on day 7 and their contents increased strongly until day 28. On day 28, when the amount of ¹³C-PLFA reached its maximum, the highest amount (except for 16:0) of the ¹³C label was found in the monounsaturated 16:1ω7c and 18:1ω7c and the cyclopropyl cy 17:0. In addition, a high incorporation of ¹³C into the polyunsaturated 18:2ω6,9 and the saturated methyl branched i:16:0 was also observed. After 59 days, the label was distributed nearly equally within the monounsaturated PLFA and the saturated methyl branched PLFA, whereas at the end of experiment, the monounsaturated PLFA contained more ¹³C than saturated branched PLFA.

On day 59, since when ¹³C-FA were detected in the non-living SOM, saturated strain chain FA carried most of ¹³C label (**Figure 26C**). From that time onwards, the saturated methyl branched contained more ¹³C than monounsaturated FA. Finally, ¹³C-cyclopropyl FA and ¹³C-polyunsaturated FA were also incorporated into the non-living SOM.

Table 13. The classification and composition of ¹³C label in respective PLFA during microbial degradation of ¹³C₆-ibu in soil

Incubation		¹³ C-PLFA [10 ⁻⁴ ; % of ¹³ C ₆ -ibu]							
time (days)	2	7	14	28	59	90			
Saturated straig	ht chain								
14:0	n.d.	$29.6 (\pm 98.6)$	$32.3 (\pm 8.4)$	$301.4 (\pm 12.5)$	$81.7 (\pm 94.7)$	$4.7 (\pm 0.8)$			
15:0	$24.7 (\pm 12.4)$	$3,2 (\pm 24.2)$	$38.8 (\pm 4.3)$	$51.2 (\pm 30.6)$	$104.2 (\pm 85.3)$	n.d.			
16:0	$74.1 (\pm 138.1)$	$377.6 (\pm 655.4)$	$728.9 (\pm 82.5)$	$1361.3 (\pm 130)$	$1499 (\pm 719.4)$	$934.8 (\pm 171.2)$			
18:0	$141.8 (\pm 345.1)$	n.d.	$135.2 (\pm 47.7)$	$470.6 (\pm 14.9)$	$116.9 (\pm 158.7)$	$560.6 (\pm 126.6)$			
20:0	n.d.	n.d.	n.d.	n.d.	$70.2 (\pm 33.5)$	$22.6 (\pm 6.5)$			
Saturated methy									
i ^b 14:0	n.d.e	n.d.	$7.4 (\pm 0.8)$	$8,3 (\pm 0)$	$21.7 (\pm 13.0)$	$5.4 (\pm 2)$			
i 15:0	n.d.	n.d.	$61.1 (\pm 5.4)$	$387 (\pm 64)$	$287.5 (\pm 188.3)$	$41 (\pm 136.4)$			
a ^c 15:0	n.d.	$2.8 (\pm 109.9)$	$160.8 (\pm 8.3)$	$190 (\pm 51.2)$	$15.1 (\pm 100.2)$	$12 (\pm 84.7)$			
i 16:0	n.d.	n.d.	$74.5 (\pm 9.2)$	523.9 (± 48)	$57.9 (\pm 98.6)$	$48.8 (\pm 55.6)$			
i 17:0	n.d.	$20.4 (\pm 125.7)$	$41.7 (\pm 19.4)$	$148.2 (\pm 27.7)$	$99.4 (\pm 96.8)$	$61.5 (\pm 52)$			
a 17:0	$3.2 (\pm 3.4)$	$31.1 (\pm 116)$	$19.2 (\pm 30.9)$	$110.5 (\pm 30.4)$	$253.4 (\pm 249.6)$	$141 (\pm 16.7)$			
10 Me ^d 16:0	n.d.	n.d.	$120.5 (\pm 11.8)$	$382.3 (\pm 48.6)$	$191.1 (\pm 164.5)$	$278.5 (\pm 54.5)$			
10 Me 17:0	n.d.	n.d.	n.d.	n.d.	n.d.	$16.4 (\pm 4)$			
10 Me 18:0	n.d.	$54.7 (\pm 98.9)$	$76.8 (\pm 0.9)$	$231.6 (\pm 26)$	$95.5 (\pm 57.6)$	n.d.			
18:0 br ^f	n.d.	$18.5 (\pm 140.7)$	$72.5 (\pm 33.9)$	$193.8 (\pm 8)$	$62 (\pm 46.4)$	$55.6 (\pm 24.4)$			
Monounsaturate	ed								
$16:1\omega7c^{c}$	$359 (\pm 28.4)$	$196.7 (\pm 10.1)$	855.7 ↑ (± 41.9)	1041 ↑ (± 98.7)	$232.5 (\pm 162.2)$	$49 (\pm 83.6)$			
16:1ω5c	$66.2 (\pm 9.7)$	n.d.	$25.1 (\pm 13.8)$	$242.3 (\pm 63.3)$	n.d.	$39.3 (\pm 57.4)$			
18:1ω9c	$89.5 (\pm 181.5)$	$136 (\pm 371.8)$	189.3 \uparrow (± 92.5)	539.6 ↑ (± 76.9)	$28.1 (\pm 219)$	$473.8 (\pm 104.1)$			
18:1ω7c	$67.6 (\pm 243.3)$	$179.4 (\pm 358.9)$	971.8 ↑ (± 47)	1679 ↑ (± 377.3)	$872.2 (\pm 378)$	$861 (\pm 64.8)$			
18:1ω5c	n.d.	n.d.	$57.7 (\pm 0)$	$68.5 (\pm 45)$	n.d.	$18.7 (\pm 3.9)$			
Monounsaturate	ed branched								
17:1 br	n.d.	n.d.	$29.7 (\pm 13.1)$	$217.7 (\pm 13.1)$	n.d.	$10.9 (\pm 72.7)$			
Polyunsaturated									
18:2	n.d.	$4.4 (\pm 209.4)$	n.d.	n.d.	n.d.	n.d.			
$18:2\omega 6,9$	n.d.	$18 (\pm 481.9)$	$78.3 (\pm 79.1)$	525.1 (\pm 3.7)	$188.2 (\pm 204.3)$	n.d.			
Saturated cyclor	oropyl								
cy 17:0	n.d.	$153.5 (\pm 169.4)$	649 ↑ (± 36.4)	944.5 ↑ (± 238.8)	$317.3 (\pm 153)$	$616.7 (\pm 107.5)$			
cy 19:0	n.d.	$96.4 (\pm 249.9)$	$89 (\pm 45.8)$	$550.7 (\pm 75)$	294.1 (± 245.7)	$145.2 (\pm 75.8)$			
TOTAL	1498.2 (± 509)	1325 (± 980.1)	4515.7 (± 312.4)	10168 (± 1416.6)	4887.5 (± 1322.2)	4397.6 (± 168.7)			

^aa- anteiso; ^bbr – branched; ^cc - cis isomer; ^di – iso; ^en.d. - not detectable; ^fMe - methyl; values in brackets (±) represent the standard deviation of the average of triplicates; values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective FA compared to the preceeding sampling time

6.2.3.3 Formation of AA and their fate in soil incubated with ¹³C₆-ibu

The incorporation of 13 C into the living biomass AA fraction (bioAA) started later (day 7, see **Figure 27**) than into PLFA (day 2). A flux of 13 C-bioAA to the non-living SOM fraction was already detected on day 7 and continued throughout the incubation. The amounts of 13 C-bioAA reached a maximum on day 28 ($\sim 3.2\%$ of 13 C₆-ibu equivalents) and then declined gradually until the end of the experiment, whereas the contents of 13 C-AA in the non-living fraction remained stable already after 58 days. The content of 13 C-bioAA decreased to 23% of its maximum, reaching on day 90 about 0.7% of the initially added 13 C₆-ibu. At the end (day 90), 93% of the total 13 C-AA in soil was stabilised in the non-living SOM pool, reaching finally 27% of 13 C₆-ibu equivalents.

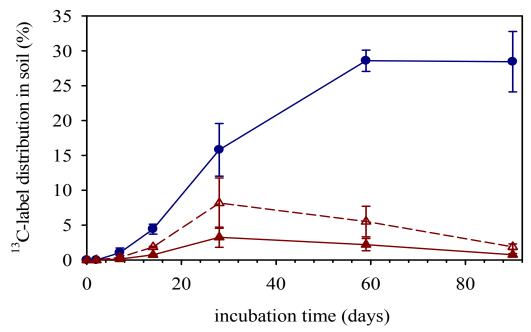


Figure 27. Time course of 13 C label incorporation within tAA (\bigcirc), bioAA (\triangle) and bioAA (\triangle) recalculated based on 40% extraction efficiency during microbial degradation of 13 C₆-ibu in soil

On day 7, only Asp carried 13 C (see **Table 14A**). After 14 days, the 13 C-Asp dominated (0.4% of 13 C₆-ibu equivalents) over other bioAA. However, Gly and Lys (each $\sim 0.2\%$ of 13 C₆-ibu equivalents) also contained high amount of 13 C. Only traces of 13 C were detected in Thr and Ile. When the amounts of 13 C-bioAA were highest (day 28), both 13 C-Asp (0.9% of 13 C₆-ibu equivalents) and 13 C-Glu (0.8% of 13 C₆-ibu equivalents) dominated, but the label was also incorporated to a lesser extent into other AA (e.g. Gly, Thr, Leu, Ile, Pro, Glu, Phe and Lys). Interestingly, at that time, a very low amount of 13 C in the nonprotein amino acid 13 C₆-ibu equivalents) was found. On day 59, when the contents of two

dominant¹³C-bioAA (¹³C-Glu and ¹³C-Asp) decreased strongly, the label was incorporated into all microbial AA. In addition, at that time, ¹³C in β -Ala increased significantly (0.05% of ¹³C₆-ibu equivalents). At the end, ¹³C-Asp carried most of the label (0.3% of the initial ¹³C₆-ibu added). In addition to Asp, the label was also detected in five other AA (Leu, Ile, Pro, Phe and Lys).

Table 14. Distribution of the ¹³C label in various ¹³C-bioAA (**A**) and ¹³C-AA in the non-living SOM (**B**) during microbial degradation of ¹³C₆-ibu in soil

Name of	Incubation time (days) [% of ¹³ C ₆ -ibu]							
AA	2	7	14	28	59	90		
Alanine	n.d.	n.d.	n.d.	n.d.	$0.27 (\pm 0.08)$	n.d.		
Glycine	n.d.	n.d.	$0.16 (\pm 0.04)$	$0.16 (\pm 0.02)$	$0.27 (\pm 0.04)$	n.d.		
Threonine	n.d.	n.d.	$0.02 (\pm 0.006)$	$0.12 (\pm 0.04)$	$0.11 (\pm 0.01)$	n.d.		
Valine	n.d.	n.d.	n.d.	n.d.	$0.17 (\pm 0.006)$	n.d.		
ß-alanine	n.d.	n.d.	n.d.	$0.008 (\pm 0.03)$	0.05 ↑ (± 0.01)	n.d.		
Leucine	n.d.	n.d.	n.d.	$0.15 (\pm 0.03)$	$0.21 (\pm 0.04)$	$0.06 (\pm 0.03)$		
Isoleucine	n.d.	n.d.	$0.0009 (\pm 0.01)$	$0.04 (\pm 0.003)$	$0.11 (\pm 0.007)$	$0.04 (\pm 0.01)$		
Proline	n.d.	n.d.	n.d.	$0.21 (\pm 0.02)$	$0.11 (\pm 0.04)$	$0.10 (\pm 0.007)$		
Aspartate ^a	n.d.	$0.15 (\pm 0.11)$	0.42 \uparrow (± 0.002)	0.92 \uparrow (± 0.14)	$0.26 \downarrow (\pm 0.34)$	$0.30 (\pm 0.12)$		
Glutamate ^b	n.d.	n.d.	n.d.	0.81 ↑ (± 0.43)	$0.50 \downarrow (\pm 0.33)$	n.d.		
Phenylanine	n.d.	n.d.	n.d.	$0.33 (\pm 0.29)$	$0.09 (\pm 0.06)$	$0.18 (\pm 0.06)$		
Lysine	n.d.	n.d.	$0.16 (\pm 0.02)$	$0.5 (\pm 0.5)$	$0.06 (\pm 0.16)$	$0.08 (\pm 0.09)$		
TOTAL	n.d.	$0.15 (\pm 0.11)$	$0.76 (\pm 0.04)$	3.25 (± 1.5)	2.20 (± 0.88)	$0.75 (\pm 0.16)$		

D								
Name of	Incubation time (days) [% of ¹³ C ₆ -ibu]							
AA	2	7	14	28	59	90		
Alanine	n.d.	n.d.	n.d.	n.d.	n.d.	$2.5 (\pm 1.9)$		
Glycine	n.d.	n.d.	n.d.	$2.5 (\pm 0.5)$	$0.6 (\pm 0.5)$	$0.6 (\pm 4.8)$		
Threonine	n.d.	n.d.	n.d.	$0.3 (\pm 1.2)$	n.d.	n.d.		
Valine	n.d.	n.d.	n.d.	n.d.	n.d.	$3.5 (\pm 1.7)$		
ß-alanine	n.d.	n.d.	n.d.	n.d.	n.d.	$0.2 (\pm 0.2)$		
Leucine	n.d.	n.d.	n.d.	$0.8 (\pm 0.6)$	$0.9 (\pm 0.1)$	$3.1 (\pm 1.5)$		
Isoleucine	n.d.	n.d.	$0.3 (\pm 0.035)$	$0.6 (\pm 0.6)$	$0.4 (\pm 0.8)$	$1.7 (\pm 0.3)$		
Proline	n.d.	n.d.	n.d.	$0.3 (\pm 0.3)$	$0.1 (\pm 0.2)$	$1.8 (\pm 3.4)$		
Aspartate ^a	n.d.	0.3 (± 0.3)	3.0 (\pm 0.5)	3.2 (± 0.3)	9.9 (± 1.0)	$4.0 (\pm 10.3)$		
Glutamate ^b	n.d.	n.d.	n.d.	$0.6 (\pm 0.6)$	$5.9 (\pm 1.0)$	$3.4 (\pm 16.0)$		
Phenylanine	n.d.	n.d.	n.d.	$0.6 (\pm 0.1)$	$1.9 (\pm 1.0)$	$1.5 (\pm 1.6)$		
Lysine	n.d.	n.d.	n.d.	n.d.	$4.3 (\pm 0.5)$	$4.1 (\pm 0.7)$		
TOTAL	n.d.	$0.3 (\pm 0.3)$	$3.3 (\pm 0.5)$	8.9 (± 3.8)	24.0 (± 1.3)	26.4 (± 4.0)		

a incl. asparagine; b incl. glutamine; n.d. - not detectable; values are shown as averages ± standard deviation values printed in bold represent characteristic values, arrows visualise increase or decrease of the respective AA compared to the preceeding sampling time

The microbial AA were also incorporated into the non-living SOM fraction (see **Table 14B**). On day 14, when ¹³C flux to the non-living SOM was observed, the label was found only in the dominant ¹³C-Asp (3.0% of ¹³C₆-ibu equivalents) and in ¹³C-Ile (0.3% of ¹³C₆-ibu equivalents). After 59 days, ¹³C-Asp still dominated (9.9% of ¹³C₆-ibu equivalents), but Glu was also highly enriched in ¹³C (5.9% of ¹³C₆-ibu equivalents). At that time, the ¹³C label was incorporated to lesser extent into other AA (Gly, Leu, Ile, Pro, Phe and

Lys). At the end, all ¹³C-microbial AA except for Thr were partly stabilised in the non-living part of SOM and the label was distributed within these AA nearly equally.

6.2.3.4 Biogenic residues in soil incubated with ¹³C₆-ibu

The content of 13 C-tFA reached 1.2% of 13 C₆-ibu equivalents at the end of the experiment; considering a conversion factor of 20 we can conclude that at least 24% of 13 C label from 13 C₆-ibu went through the microbial biomass. From the amount of 13 C-tAA of 27% (of 13 C₆-ibu equivalents) at the last day, we can estimate that there were 54% biogenic residues in this experiment (using the conversion factor of 2).

7 DISCUSSION

The soil biodegradation studies with ¹³C₆-2,4-D and ¹³C₆-ibu proved that microorganisms used C directly from these contaminants or indirectly via CO₂ fixation to form their biomass components. After the death of microorganisms, their biomass components were incorporated into SOM and finally formed biogenic residues, which lead to the overestimation of the risk related to NER formed during biodegradation of these contaminants in soils.

The simple liquid culture experiment with *C. necator* JMP 134 provided a general view on the processes of ¹³C-FA and ¹³C-AA formation during 2,4-D biodegradation. In addition, this experiment showed the relative abundance of FA and AA in the total biomass of this strain, necessary for an estimation of the total biogenic residues content in the soil experiments.

7.1 Liquid culture experiments

Two liquid culture experiments with ¹³C₆-2,4-D and with ¹³CO₂ showed that the bacterial strain *C. necator* JMP 134 used the ¹³C derived either from this herbicide or from ¹³CO₂ for its PLFA and AA synthesis. In addition, the abiotic systems clearly showed that ¹³C label incorporation from labelled 2.4-D or from CO₂ were relevant only for biotic incubation experiments. ¹³C₆-2,4-D was degraded rapidly and dissipated already after the first day of incubation. Therefore, ¹³C label incorporation into PLFA and AA directly from the labelled 2,4-D was observed only until day 1. In the later phase of experiment, when neither 2,4-D nor its known aromatic metabolites were detected in the medium, ¹³C was incorporated into PLFA and AA only via CO₂ fixation.

Incorporation of ¹³C into biomass of C. necator JMP 134 grown on 2,4-D medium

The ¹³C label incorporation into the biomass of *C. necator* JMP 134 in the ¹³C₆-2,4-D experiment started immediately and remained constant until day 3. Thereafter, the ¹³C content in the biomass still increased, although already after the first day of incubation ¹³C₆-2,4-D was completely depleted and no biomass growth was observed (see results in section 5.1.1.1 and 5.1.2.1). This is contrary to the results in the ¹³C₆-2,4-D experiment presented by LERCH ET AL. (2007), which indicated a low content of ¹³C in the biomass of *C. necator* JMP 134 at the beginning and its strong increase in the later phase of incubation. This biomass growth was also consistent with the OD measurement of the culture (LERCH ET AL., 2007). However, in their experiment the initial concentration of 2,4-D was much higher (250 mg/L) than in the present study (77 mg/L) and this herbicide was still detectable after the first day of incubation.

Although the amounts of ¹³C label in PLFA increased continuously until day 7, this however, was not related to growth. At the beginning of incubation, the ¹³C-monounsaturated PLFA were dominant over other PLFA, their amounts then decreased strongly, whereas the content of ¹³C in saturated cyclopropyl PLFA increased rapidly throughout the incubation time. The cyclopropyl PLFA are known starvation biomarkers, formed by the methylation of the double bond in monounsaturated PLFA (KAUR ET AL., 2005). This modification makes these FA more stable and minimizes the membrane lipid losses under stress conditions (KAUR ET AL., 2005). Therefore, this mechanism was employed by C. necator JMP 134 in the absence of 2,4-D in the medium. The continuous increase of the ¹³CO₂ evolution observed until day 7 in the ¹³C₆-2,4-D study and of the ¹³C incorporation into PLFA in the ¹³CO₂ fixation experiment until day 3 clearly indicates the continuous metabolisation of ¹³C incorporated into biomass residues via CO₂ fixation. The presence of CO₂ is reported to be essential for the growth of many microorganisms (KREBS, 1941), since CO₂ plays a key role in the synthesis of the tricarboxylic acid cycle (TCC) intermediates malate and oxalacetate from pyruvate or phosphoenolpyruvate (MILTNER ET AL., 2004; FEISTHAUER ET AL., 2008; see Figure 28). At the end of the incubation, the contents of ¹³C in PLFA decreased strongly in both experiments and the ¹³C-saturated cyclopropyl PLFA were dominant over other PLFA (in particular in the ¹³CO₂ experiment), indicating the decreasing contribution of CO₂ fixation.

At the beginning of the ¹³C₆-2,4-D experiment ¹³C was highest in Glu. This AA is synthesised from 2-oxoglutarate formed in the TCC (see Figure 28). At that time, Asp also contained significant amounts of ¹³C in comparison to other ¹³C-AA. Asp is derived from oxaloacetate, which is the acceptor molecule for acetate in the TCC (MICHAL, 1999). When oxaloacetate is removed from the TCC for Asp synthesis, it has to be replenished by various anaplerotic reactions, which include the carboxylation of phosphoenolpyruvate; thus Asp is a direct product of heterotrophic CO₂ fixation (FEISTHAUER ET AL., 2008; MILTNER ET AL., 2004). This is in good accordance to the ¹³CO₂ experiment, in which the ¹³C label was found only in Asp at the beginning (day 2), when the CO₂ fixation started. Both Glu and Asp formed within the TCC are important precursors for other AA, Pro and Thr, Ile, Lys, respectively (arginine and methionine were not detected in experiments). Therefore, these ¹³C-labelled AA were found in the later phase of incubation period. Interestingly, the contents of ¹³C in Thr increased strongly from day 3 until day 7 in both experiments and then decreased rapidly at the end. The simultaneous decline of ¹³C in both Thr and PLFA at the end thus indicates the decrease of CO₂ fixation contribution to ¹³C incorporation into the biomass of C. necator. Other AA, such as Ala, Val and Leu formed from pyruvate and Phe from phosphoenolpyruvate were also enriched in the later phase of incubation period. However, the trends of both decline and increase of their ¹³C contents could be assigned to the different food demands of this bacterial strain throughout the incubation time. The low incorporation of ¹³C into Ser found only at the end of ¹³C₆-2,4-D experiment demonstrates the formation of this AA from ¹³C-Gly. The ¹³C-Gly was presumably converted directly from glyoxylate, which is a cleavage product of isocitrate (MICHAL, 1999) or of 2,4-D metabolite (see **Figure 10** in section 4.1).

At the end of the 13 C₆-2,4-D incubation experiment, ~ 17% of initially added 13 C₆-2,4-D was detected in the total biomass of *C. necator* JMP 134, 8% was found in AA and 0.6% was incorporated into the PLFA. However, also a relatively high amount of 13 C was incorporated into the biomass of *C. necator* via CO₂ fixation (~ 4.2% of the initial 13 CO₂ amount). Therefore, it should be kept in mind that CO₂ fixation can contribute significantly to the incorporation of C into biomass components and thus to the NER formation in soils. In addition, the results from both 13 C₆-2,4-D and 13 CO₂ incubation experiments showed that 13 C-labelled AA represented ~ 50% of the total 13 C in the biomass, demonstrating their relatively high abundance in microbial cells. Therefore, this microbial biomarker can be helpful in the quantification of biogenic residues in complex soil systems.

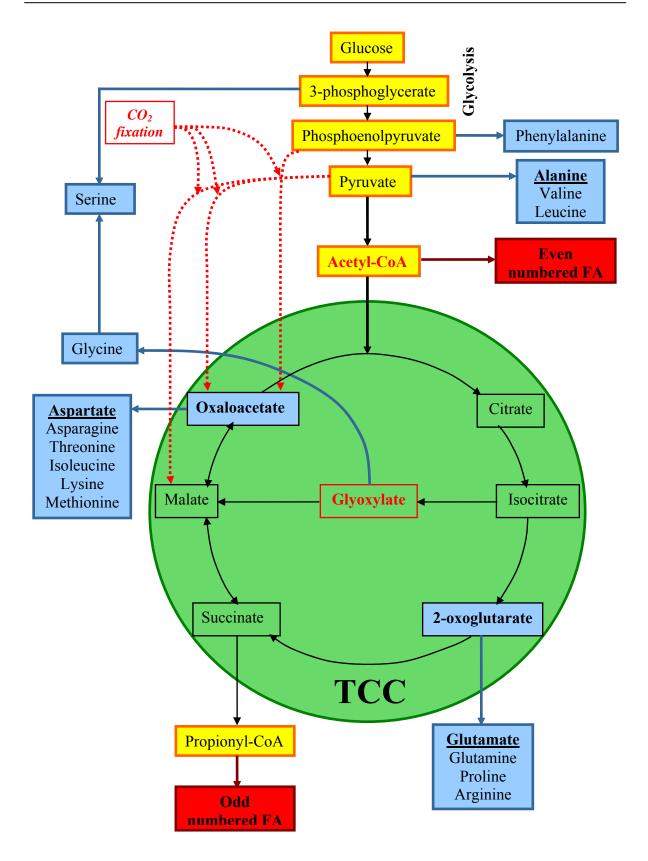


Figure 28. Scheme of the relevant pathways for AA and FA synthesis in TCC and glycolysis; and the anaplerotic reactions (...) replenishing the TCC (adapted from MICHAL, 1999; FEISTHAUER ET AL., 2008)

7.2 Soil experiments

Two soil degradation experiments with ¹³C₆-2,4-D and ¹³C₆-2,4-ibu prove for the first time that biogenic residues formation can be a relevant process during microbial degradation of many organic contaminants in soil. In case of the readily biodegradable 2,4-D and Ibu, the majority of NER are formed from biomass components. However, the kinetics of biogenic residue formation in the biodegradation experiments with these various organic contaminants was different and depended strictly on the degradation kinetics in soil.

7.2.1 Soil incubation experiment with ¹³C₆-2,4-D

The results from the soil incubation experiment with ¹³C₆-2,4-D provide detailed insight into the C flux from the ¹³C-labelled pollutant 2,4-D or from labelled CO₂ (evolved during 2,4-D mineralisation) via microbial biomass to non-living SOM. It has been clearly shown that microorganisms use the readily available C derived from 2,4-D for their biomass synthesis, which was proven by identifying the ¹³C label within the biomarkers FA and AA. After cell lysis and the decay of microbial biomass, both ¹³C-labelled PLFA and bioAA were distributed within the microbial food web and incorporated into the non-living fraction of SOM, which was finally stabilised. These results are also in good agreement with recent studies on the fate of ¹³C labelled *E. coli* in soil (KINDLER ET AL., 2006, 2009; MILTNER ET AL., 2009), which showed that biomass-derived C was distributed within the microbial food web and finally contributed to the formation of refractory non-living SOM (KINDLER ET AL., 2006, 2009; LÜDERS ET AL., 2006). The majority of NER formed during microbial degradation of 2,4-D in soil was composed of biomass-derived components.

7.2.1.1 Biogenic residues as NER in soil incubated with ¹³C₆-2,4-D

At the end of the incubation experiment, the amount of 13 C found in tAA (non-living SOM + bioAA fraction) was high and reached 22% of the initially added 13 C₆-2,4-D equivalents corresponding to 61% of the NER based on isotope mass balance. It should be kept in mind that these AA do not represent free AA, which are reported to be degraded very rapidly in soil (JONES, 1999; VINOLAS ET AL., 2001), but they were hydrolysed from peptides and proteins. However, it is impossible to extract all biomolecules from soil microbial biomass, therefore only part of them could be analysed. By the extraction of some known microbial biomolecules e.g. FA and AA, we can estimate the actual amounts of biogenic residues in soil. The particular focus however, has been laid on AA from proteins and peptides due to their high abundance in microbial biomass. Two experiments with *C. necator* JMP 134 grown on 2,4-D medium showed that 13 C-labelled AA represented $\sim 50\%$ of the total 13 C in the

biomass (see results in sections 6.1.1.1 and 6.1.2.1). In addition, the ratio of ¹³C-AA to total ¹³C in biomass was relatively stable at the different sampling dates; thus AA are relevant for estimating the NER content in soil. Therefore, both the results from this experiment and a conservative conversion factor of 2 (MADIGAN AND MARTINKO, 2006) was considered for quantification of the actual incorporation of C derived from 2,4-D into the biomass, which resulted in a total amount of 44% (of ¹³C₆-2,4-D equivalents) for biogenic residues. The amount of these residues is close to the total NER amount of $\sim 36\%$ (of $^{13}C_6$ -2,4-D equivalents) in soil measured by GIRARDI ET AL. (submitted; see also results in section 6.2.1.1). This small discrepancy in the NER contents could be related to the different methods used for NER determination. NER are generally quantified as CO₂ released from combusted soil after the pollutant residues extraction (BARRIUSO ET AL., 2008; WAIS, 1998). This analytical approach was also implemented by GIRARDI ET AL. (submitted). 2,4-D residues were extracted from soil by ASE, which is a harsh extraction method (NORTHCOTT AND JONES, 2000). At the end of experiment, about 8% of the initial ¹³C₆-2,4-D equivalents was extracted from soil by ASE, but from that amount only 0.5% was identified as 2,4-D. Therefore, a small part of biogenic residues stabilised in SOM could be still extracted from soil and thus affected the lower total NER content measured by GIRARDI ET AL. (submitted). It should be noted that the hydrolysis of AA from proteins is a very efficient extraction method, which enables also the extraction of AA stabilised in SOM; this could thus affect the higher biogenic NER content than the total NER content. However, the estimated amounts of biogenic residues indicate that the biomass compounds were finally converted into the biogenic residues already after 32 days and represented the majority of 2,4-D-derived NER in the soil (see Figure 29).

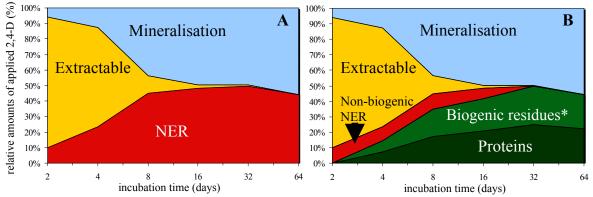


Figure 29. Scheme of the ¹³C conversion over microbial degradation of ¹³C₆-2,4-D in soil. **A: conventional mass balance and B: new mass balance considering biogenic residues formation.** *The biogenic residues were estimated from AA using a conversion factor of 2 (for details see text).

Studies with anthracene (KÄSTNER ET AL., 1999), phenanthrene (RICHNOW ET AL., 2000) and TNT (WEISS ET AL., 2004) proved that C or N from these contaminants was incorporated into soil microbial biomass, which contributed to the NER formation in soils. The NER amounts formed during degradation of [ring-14C-U]-labelled 2,4-D or 2,4-DCP in non-sterile composted straws were much higher than in sterile ones, where the parent compounds remained unchanged (BENOIT AND BARRIUSO, 1997). The high NER content found in the present study, where the ¹³C₆-2,4-D was nearly mineralised already after 16 days (GIRARDI ET AL., submitted) in comparison with the negligible mineralisation in sterile soils is in a good accordance to those in non-sterile straws (BENOIT AND BARRIUSO, 1997). Therefore, high contents of NER in non-sterile composted straw or soil underlines the key role of soil microorganisms in the formation of residues from 2,4-D in native soil. Another soil biodegradation experiment with [ring-14C-U] 2,4-D showed that the amount of 14C label from this herbicide detected in the microbial biomass was similar to that observed for [1-14C] glucose, suggesting the utilisation of 2,4-D as a C source (ROBERTSON AND ALEXANDER, 1994). In addition, the well-known 2,4-D soil degrader C. necator JMP 134 also used C from 2,4-D for their biomass components synthesis in the liquid culture experiments (see discussion in section 7.1.1).

There is also evidence for abiotic residue formation, because 18.6% of ¹³C₆-2,4-D equivalents were recovered as NER in sterile soils at the end of incubation. However, in the sterile soils, the biogenic residue representatives ¹³C-FA and ¹³C-AA were not formed, indicating that these NER were exclusively formed by the reported abiotic sequestration mechanisms (PIGNATELLO AND XING, 1996; ALEXANDER, 2000; GEVAO ET AL., 2000; KATAYAMA ET AL., 2010). It was reported that autoclaving changes the soil physico-chemical properties (SHAW ET AL., 1999; BERNS ET AL., 2008), therefore the contaminant behaviour in sterile soils can be different than in native soils. In addition, the rapid mineralisation of ¹³C₆-2,4-D in the biotic soil experiment, will reduce the extent of abiotic NER formation because rapid 2,4-D degradation competes with abiotic residue formation. The fact that the non-living AA and FA can explain most of the NER formation strongly suggests a high contribution of biotic NER formation. Several studies demonstrated that microorganisms can also mediate the enzymatic covalent oxidative coupling of 2,4-DCP to humic substances (BOLLAG, 1991; PALOMO AND BHANDARI, 2005, 2006). However, it should be noted that these studies were limited to simple (BOLLAG, 1991; HATCHER ET AL., 1993), or abiotic systems (PALOMO AND BHANDARI, 2005; 2006) and covalently bound 2,4-DCP in SOM has never been detected.

In addition, all available studies on NER formation during microbial degradation of 2,4-D in soil are limited to quantitative analyses with radioactive tracers (BENOIT AND BARRIUSO, 1997; BOIVIN ET AL., 2005). In contrast to these available results, where no information on the chemical nature of NER in SOM was presented, the present study shows that nearly all NER formed during biodegradation of $^{13}C_6$ -2,4-D in soil were composed from microbial components stabilised in the SOM pool during the later phase of experiment (aging process).

Finally, the high content of ¹³C-AA detected in the present study contradicts the generally accepted view that the parent compounds or their primary metabolites) are the main precursors of phenoxyacetic acid herbicides derived NER, in particular via sorption processes (Bollag et al., 1991; Bollag et al., 1992; Palomo and Bhandari, 2005, 2006).

7.2.1.2 Fate and stability of FA in soil incubated with ¹³C₆-2,4-D

The formation of ¹³C-PLFA during microbial degradation of ¹³C₆-2,4-D in soil was rapid during the first four days of incubation, since the amount of the readily available ¹³C₆-2,4-D in soil was still high (GIRARDI ET AL., submitted). Taking into consideration a relatively rapid turnover of FA in bacterial cells and a total FA content of ~ 5% (see results in section 6.2.1.4), for the maximum of about 1% of ¹³C₆-2,4-D equivalents detected in the tFA fraction on the day 8, it can be concluded that more than 20% of the ¹³C initially added must have passed through the microbial biomass. The high amount of ¹³C found in the monounsaturated PLFA indicates that Gram-negative bacteria were initially involved in 2,4-D transformation. This is in good accordance with the previous study reported by LERCH ET AL. (2009A), in which the content of ¹³C in monounsaturated tFA fraction (including the PLFA) was highest. The high incorporation of ¹³C label into even-numbered PLFA (such as: 16:1ω7c; 16:0; $18:1\omega$ 7c; $18:1\omega$ 9c; $16:1\omega$ 5c) in the initial phase of incubation experiment clearly indicates for their direct synthesis from Acetyl-CoA (see Figure 28 in section 7.6.1), which is the end product of 2,4-D degradation (see Figure 10 in section 4.1). In addition, a lower incorporation of ¹³C within the saturated methyl-branched PLFA, which are typical for Grampositive bacteria, was also observed in this experiment. This indicates that this group of bacteria took indirectly part in the transformation of 2,4-D via fixation of the CO₂ evolved during mineralization of this pollutant or by uptake of metabolites formed by Gram-negative bacteria. When the amounts of ¹³C in monounsaturated PLFA started to decrease after 8 days, the amount of ¹³C in saturated cyclopropyl PLFA was highest. This high of ¹³C content in the cyclopropyl PLFA indicates that bacteria employed starvation mechanism (KAUR ET AL., 2005), which agrees with the low residual concetrations of 2,4-D at that time (GIRARDI ET AL., submitted). In the later phase of the experiment, when the microbial biomass started to decay, a flux of 13 C from living cells to non-living SOM was observed. After 16 days, besides 13 C-PLFA of the decaying biomass, the 13 C-FA of the non-living SOM were also continuously metabolised at a low level and distributed within the food web. Therefore, when 13 C₆-2,4-D was nearly depleted, the 13 C released from these 13 C-FA was thus rapidly recycled via CO₂ fixation by Gram-negative and Gram-positive bacteria and by fungi.

Two independent studies showed that the NER contents formed during 2,4-D biodegradation in soil (BOIVIN ET AL., 2005) and 15 years-old NER after fresh soil amendment (LERCH ET AL., 2009B) decreased slightly over time. BOIVIN ET AL. (2005) suggested that this decline could be related to the temporal immobilisation of ¹⁴C from 2,4-D in the protoplasm of some soil microorganisms. The decrease of NER in these abovementioned studies thus could be assigned to the release of FA from both the decaying biomass and biomass residues in the non-living SOM, as it was observed in the present study.

At the end of the ¹³C₆-2,4-D soil incubation experiment, the amount of ¹³C-PLFA decreased to 20% of its maximum value, indicating the metabolisation of cell components of the initial 2,4-D degraders via the microbial food web. However, the ¹³C-FA in the non-living SOM decreased to 50% of their maximum value, demonstrating that the compounds in non-living fraction were partly stabilised in SOM against microbial attack. KINDLER ET AL. (2009) also reported the decline of ¹³C-tFA derived from *E. coli* to 24% (of the initial amount of ¹³C-E. coli) in soil after eight months of incubation. In spite of the long incubation time, still about 50% of microbial biomass C remained in soil (KINDLER ET AL., 2006), demonstrating the high stability of microbial biomass residues in soil.

7.2.1.3 Fate and stability of AA in soil incubated with ¹³C₆-2,4-D

Compared to PLFA, the incorporation of ¹³C into the bioAA fraction started 2 days later, indicating that PLFA as components of microbial cell membranes are turned over faster and thus received the ¹³C label earlier than the AA. However, the flux of ¹³C from the decaying biomass AA fraction into the non-living SOM was higher than that observed for the PLFA. At the beginning of the incubation, when the readily degradable ¹³C₆-2,4-D was still present at high amounts (GIRARDI ET AL., submitted), Glu carried highest label of all ¹³C-AA detected in this experiment. However, on day 8, the content of ¹³C in Asp was highest compared to the other ¹³C-AA. Glu is formed from 2-oxoglutarate derived from the TCC; see **Figure 28** in section 7.1.1. The dominant incorporation of ¹³C into Glu thus clearly indicates for a different route of the label into AA in first days of incubation, which was followed by CO₂ fixation from mineralised CO₂ and from decaying microbial biomass, as it was shown by the high amount of ¹³C-Asp. Asp has been reported to be a direct product of heterotrophic CO₂

fixation in soil (MILTNER ET AL., 2005), which is also in good accordance to the pure culture experiment with *C. necator* JMP 134 grown in presence of ¹³CO₂ atmosphere (see section 6.1.1). In addition, CO₂ fixation as a relevant process for biogenic residue formation in the later phase of 2,4-D degradation experiment was also proven by the detection of ¹³C label in the PLFA of the ¹³CO₂ soil experiment. Overall, both Asp and Glu are important precursors for other AA, e.g. Glu for Pro and Asp for Lys and Ile (see **Figure 28** in section 7.1.1). Therefore, lower amounts of ¹³C-Lys, ¹³C-Ile and ¹³C-Pro were also detected later in this experiment. On day 16, when only traces of readily available 2,4-D residues were left and ¹³C-bioAA started to decline, the ¹³C label also was distributed within Ala, Gly and Phe, which would point to their biosynthesis via different intermediates of glycolysis (MADIGAN AND MARTINKO, 2006; MICHAL, 1999).

At the end of the incubation, the content of the ¹³C-bioAA fraction decreased to 36% of its maximum found in the experiment, indicating the decrease of the contribution of the living biomass fraction to SOM pool over the incubation time. Surprisingly, the amounts of ¹³C-AA in the non-living SOM from day 32 onwards remained stable and finally reached 19.8% of the initial amount of ¹³C₆-2,4-D in soil, demonstrating the stabilisation of proteins from the decaying microbial biomass. MILTNER ET AL. (2009) also demonstrated the high stability of ¹³C-labelled bacterial proteins in the non-living part of SOM; their contents decreased only marginally even over a period of several months. Proteins have been reported to be stabilised in SOM by various sorption processes (Kleber et Al., 2007), which protect them from the microbial degradation (RILLIG et Al., 2007) and affect their stability for a very long period of time (KNICKER et Al., 1993; Kögel-Knabner, 2002). These results thus underline the high importance of AA in NER formation from 2,4-D due to their high abundance in the microbial biomass and the stability in contrast to FA, which were turned over relatively fast.

In a study on the fractionation of [U-ring-¹⁴C] 2,4-D residues XIE ET AL. (1997) showed that most of the radioactivity was associated with humic acids. It should be noted that proteins can bind to humic acids by hydrophobic and ionic interactions, which protect them from chemical and microbial attack (ZANG ET AL., 2000). Therefore, the dominance of ¹⁴C label in the bound-humic acid fraction could be assigned to AA, as observed in the present study. Moreover, in the soil experiments with ¹³C₆-2,4-D, the amount of labelled biogenic residues did not change after 32 days of incubation, which was caused by their stabilisation in SOM. This could also explain the constant mineralisation rate of 90-years-old NER after addition of the fresh soil reported by LERCH ET AL. (2009B), when taking into consideration the fact that these NER are mostly biogenic.

7.2.2 Soil incubation experiment with ¹³C₆-ibu

This part of the study provides insight into the C flux from ¹³C₆-ibu during its biodegradation via microbial biomass components to the non-living part of SOM. The microorganisms used the C derived from this contaminant to synthesise their biomass, as it was shown by ¹³C incorporation into the microbial biomarkers FA and AA. After cell death, biomass constituents such as ¹³C-PLFA and ¹³C-bioAA from the decaying cells were distributed into the food web and finally incorporated into the non-living SOM fraction. These results are in good agreement with the studies on the fate of both labelled *E. coli* in soil (KINDLER ET AL., 2006, 2009; MILTNER ET AL., 2009) and ¹³C-biogenic residue in the soil incubated with ¹³C₆-2,4-D (NOWAK ET AL., in press). However, the kinetics of biogenic residues formation in soil incubated with ¹³C₆-ibu was initially slower than that observed in the ¹³C₆-2,4-D experiment. Nevertheless, biomass-derived compounds such as FA and AA also contributed significantly to NER formation from ¹³C₆-ibu in soil. Biomass residues represented the majority of ¹³C₆-ibu-derived NER, which is also consistent with the ¹³C₆-2,4-D experiment.

7.2.2.1 Biogenic residues as NER in soil incubated with ¹³C₆-ibu

The content of ¹³C-tAA detected at the end of the ¹³C₆-ibu experiment was high, reaching 27% of the initial ${}^{13}C_6$ -ibu equivalents and corresponding to 90% of the NER based on isotope mass balance. Considering a conservative conversion factor of 2 (for details on the conversion factors see section 6.1.1.1), these 27% in ¹³C-tAA corresponds to a total amount of 54% (of ¹³C₆-ibu equivalents) for biogenic residues. However, the total ¹³C-NER content measured by GIRARDI ET AL (submitted; see also section 6.2.3.1) was much lower (30% of ¹³C₆-ibu equivalents) than the estimated amount of biogenic residues. It should be noted that the total NER content was determined by EA-C-irMS after the removal of Ibu residues using a harsh extraction method (ASE). The estimated amount of NER strongly depends on the extraction method used for the bioavailability measurement of a target compound (KHAN 1991; NORTHCOTT AND JONES, 2000; MORDAUNT ET AL., 2005). GIRARDI ET AL. (submitted; see also results in section 6.2.3.1) reported that the solvent bulk extractable amount prior to purification by SPE accounted for $\sim 14\%$ of $^{13}C_6$ -ibu equivalents; and from that only 0.5% was identified as Ibu. In addition, the contents of the bulk solvent extractable varied over time. At the beginning of biogenic residue formation, a high content of the solvent bulk extractable (~ 50% of ¹³C₆-ibu equivalents on the day 28) was observed, which decreased strongly at the end; this could be assigned to the aging processes of NER. Therefore, the ~ 14% extracted at the end from soil prior to quantitative NER analyses could still contain a

certain part of living biomass residues (e.g. bioAA and PLFA), which are much easier to extract before their stabilisation in SOM. In addition, the hydrolysis of AA from proteins is definitely a more efficient extraction method than ASE and enables the extraction of all AA, also those stabilised in SOM, which results in the higher NER estimation in the present study. Nevertheless, the high content of biogenic residues derived from ¹³C₆-ibu detected at the end of the incubation period clearly indicates that nearly all ¹³C₆-ibu-derived NER were based on the stabilised compounds of microbial origin (see **Figure 30**), which is also in good accordance to the data obtained in the ¹³C₆-2,4-D experiment (NOWAK ET AL., in press).

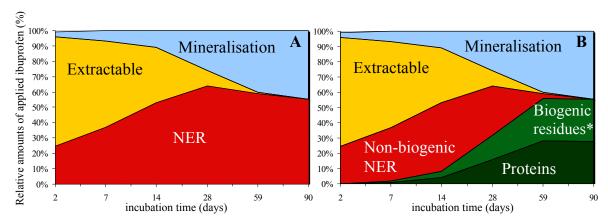


Figure 30. Scheme of the ¹³C conversion over microbial degradation of ¹³C₆-ibu in soil. **A: conventional mass balance** and **B: new mass balance considering biogenic residues formation.** *The biogenic residues were estimated from AA using a conversion factor of 2 (for details see text).

GIRARDI ET AL (submitted) clearly presented that - in contrast to the biologically active soils - NER formation and mineralisation in the sterile soils incubated with ¹³C₆-ibu were negligible during the first 28 days of incubation. However, after 90 days, the NER contents in the inactive soils reached about 16% of ¹³C₆-ibu equivalents (GIRARDI ET AL., submitted), indicating abiotic NER formation by ibu entrapment into SOM after prolonged incubation times (PIGNATELLO AND XING, 1996; ALEXANDER, 2000; GEVAO ET AL., 2000; KATAYAMA ET AL., 2010). However, no ¹³C incorporation into FA, AA and thus into microbial biomass was observed, clearly indicating the crucial role of microorganisms in biogenic NER formation in non-sterile soils.

The 14 C₃-ibu-derived NER content after 100 days of incubation ($\sim 50\%$ of 14 C₃-ibu equivalents; RICHTER ET AL., 2007) was in the same range as the biogenic NER amount detected in the present study (54% of 13 C₆-ibu equivalents). However, 14 C₃-ibu-derived NER were formed rapidly and reached a maximum already after 4 days; their contents then decreased only slightly (RICHTER ET AL., 2007). Contrary to these results, in the present

¹³C₆-ibu experiment, the NER reached a maximum after 59 days and their contents remained stable until the end of the incubation period. This is related to the different labelling positions of the compound used in these two studies. In the present study, the compound was labelled in the aromatic ring, which is cleaved much later during biodegradation than the methyl group labelled in the ¹⁴C₃-ibu experiment is removed (RICHTER ET AL., 2007), which is reflected by different kinetics of NER formation. NER in the ¹³C₆-2,4-D experiment were also formed very quickly (NOWAK ET AL., in press; see also section 7.2.1.1), but this contaminant was degraded much faster than ¹³C₆-ibu (GIRARDI ET AL., submitted). This was caused by the ability of the present soil microorganisms to degrade 2,4-D after previous 2-methyl-4-chloro phenoxyacetic acid (MCPA) application to the soil (INES MERBACH, personal communication), which also belongs to the group of phenoxy compounds as 2,4-D.

7.2.2.2 Fate and stability of FA in soil incubated with ¹³C₆-ibu

Incorporation of the ¹³C-label into PLFA started immediately and increased rapidly until day 28, when the readily available ¹³C₆-ibu was still present at relatively high concentrations in the soil (GIRARDI ET AL., submitted; see also section 6.2.3.1). The monounsaturated PLFA carried most of the ¹³C label and dominated over other ¹³C-labelled PLFA until day 28, indicating that Gram-negative bacteria were important initial Ibu degraders in soil. This is in good accordance to the study on NER formation from ¹³C₆-2.4-D in soil (NOWAK ET AL., in press; see section 7.2.1.2). As the same soil was used in these two experiments, the same groups of microorganisms might have been involved in the degradation of both acidic aromatic molecules ¹³C₆-ibu and ¹³C₆-2,4-D. In the later phase of incubation period, also the saturated methyl branched PLFA were highly enriched in ¹³C and their importance increased over time, in particular when only traces of ¹³C₆-ibu were left. This class of PLFA is typical for Gram-positive bacteria, which indirectly take part in the degradation of ¹³C₆-ibu, as it was also demonstrated previously in the ¹³C₆-2,4-D soil experiment. The observed continuous increase of ¹³C incorporation into the cyclopropyl PLFA throughout the incubation time indicates starvation of the microorganisms (KAUR ET AL., 2005). In addition, the continuous decline of ¹³C-PLFA in the living biomass, the distribution of ¹³C within all classes of PLFA and the contaminant depletion in soil over the incubation period indicated that the Ibu-derived C was distributed within the food web, as found in the ${}^{13}C_{6}$ -2,4-D study. In addition to bacteria, fungi also took part in the turnover of ¹³C-labelled PLFA from dead cells, as proven by the ¹³C found in the fungal biomarker PLFA 18:2\omega6,9. From the content of ¹³C-tFA of 1.2% (of ¹³C₆-ibu equivalents) detected at the end of experiment and considering a conversion factor of 20 (for details see 6.1.1.1) we can conclude that at least 24% of biomass must have been formed. This is close to the amount found in the 2,4-D soil incubation experiment (22% of 13 C₆-2,4-D equivalents).

At the end of the experiment, ¹³C-PLFA decreased to 50% of their maximum, whereas ¹³C-FA in the non-living SOM fraction remained unchanged. This is contradictive to the study with ¹³C₆-2,4-D, where apart from the observed decline of ¹³C-labelled PLFA to 20% (of ¹³C₆-2,4-D equivalents), the contents of ¹³C-FA in the non-living fraction also decreased (50% of ¹³C₆-2,4-D equivalents; see section 6.2.1.2). However, ¹³C₆-ibu dissipated slower than ¹³C₆-2,4-D. The PLFA thus received the ¹³C label much later. Therefore, both their decline and ¹³C flux to the non-living fraction were also observed later than in the ¹³C₆-2,4-D experiment. Additionally, both the lower decrease of ¹³C-PLFA and the stability of ¹³C-FA in the non-living SOM until the end of ¹³C₆-ibu experiment may be related to the duration of the incubation, which might have been stopped before destabilisation of part of Ibu residues started, as it was seen in the 2,4-D study. In addition, RICHTER ET AL. (2007) also reported a slight release of rapidly formed ¹⁴C-NER derived from ¹⁴C₃-ibu over the incubation time (~5% of ¹⁴C₃-ibu equivalents after 100 days).

7.2.2.3 Fate and stability of AA in soil incubated with ¹³C₆-ibu

AA in biomass-living fraction received ¹³C label later than PLFA, which is in good agreement with the data obtained in the experiment with ¹³C₆-2,4-D (NOWAK ET AL., in press; see section 7.2.1.3). However, ¹³C-bioAA derived from ¹³C₆-ibu started to decline at the same time as ¹³C-PLFA, what is different from the soil experiment with ¹³C₆-2,4-D, where the decrease of ¹³C-bioAA content was observed later than that of ¹³C-PLFA. This discrepancy in the decline of bioAA is difficult to explain and could be related to the different dissipation kinetics of the used contaminants and also to timing of the sampling events relative to processe in soil in these two soil incubation experiments. After 58 days, the contents of ¹³C-AA in the non-living fraction derived from ¹³C₆-ibu did not change and remained stable until the end, whereas the ¹³C-bioAA decreased to 23% of their maximum after 90 days (of ¹³C₆-ibu equivalents). This is in good accordance to the ¹³C₆-2,4-D soil experiment, where ¹³C-AA in the non-living SOM were also constant, indicating their high importance in the NER formation.

In the initial phase of ¹³C₆-ibu degradation, the label was incorporated only into Asp. Thereafter, the content of ¹³C-Asp increased significantly and this AA was dominant until day 28, when ¹³C₆-ibu was present at a low concentration in soil (GIRARDI ET AL., submitted; see also 6.2.3.1). The fast appearance of the label in Asp and its later dominance over the other ¹³C-bioAA suggest that ibu-derived ¹³C is incorporated into the bioAA via heterotrophic CO₂ fixation (MILTNER ET AL., 2004; FEISTHAUER ET AL., 2008). After 14 days, a lower amount of

¹³C was found also in Lys and Ile, for which Asp is an important precursor. This however is, in contrast to the study with ¹³C₆-2,4-D, where the label was first incorporated into Glu, at the time when the amount of readily available ¹³C₆-2,4-D in soil was still high. Although the molecular structures of ¹³C₆-ibu and ¹³C₆-2,4-D are similar, these compounds are degraded by soil microorganisms via different pathways. In the soil incubated with ¹³C₆-ibu, ¹³C-Glu was also formed, but later, when the ¹³C-bioAA started to decline and ¹³C₆-ibu was nearly depleted. Interestingly, on days 28 and 58, the non-protein AA ¹³C-β-Ala was found. The ratio of Asp to its decomposition product β-Ala informs about the intensity of SOM decomposition (Dauwe And Middelburg, 1998). The ratio of Asp/β-Ala on day 28 was 121, whereas on day 58 it decreased to 5.4, which clearly shows that the ¹³C-AA in the microbial biomass were decaying. The high contents of ¹³C-Glu found in the later phase of the experiment would thus point to a different degradation pathway of ibu-derived ¹³C. Furthermore, similar to the ¹³C₆-2,4-D study, in the later phase of ¹³C₆-ibu-amended soil incubation, a lower label incorporation into other AA (Pro, Ala, Gly, Thr, Leu, Phe, and Val) was also observed.

8 CONCLUSIONS

In the present study, the incorporation of ¹³C into biomass constituents (FA and AA) and their contribution to NER formation in soil during microbial degradation of ¹³C₆-2,4-D and ¹³C₆-ibu were studied and discussed. Both experiments were carried out with stable isotope labelled compounds, which enabled a detailed analysis of the chemical structure of the extracts containing either AA or FA. The detection of the ¹³C label in the biomass components (PLFA and bioAA) and the comparison of their amounts to the contents of ¹³C in the non-living SOM part (tFA and tAA) in soil clearly showed that these biomass components, which were stabilised in SOM pool, contributed to NER formation in soil. Special attention was given to the AA due to their high abundance (~ 50%) in microbial cells, as it was proven in the simple pure culture system with *C. necator* JMP 34.

Two soil experiments with ¹³C₆-2,4-D and ¹³C₆-ibu proved for the first time that biogenic residue formation contributes to a major extent to the formation of NER from these contaminants. For both 2,4-D and Ibu, nearly <u>all</u> of the NER derived from microbial biomass. The incorporation of the label into the biomass components was very fast in the 2,4-D experiment and slower at the beginning in the case of Ibu study due to a prolonged phase of microorganisms adaptation. However, in both soil experiments in the later phase of incubation, the biomolecules derived from the decaying biomass, in particular AA, were stabilised in SOM for longer periods.

In general, major contributions of biogenic residues in NER formation are to be expected if the respective organic contaminant is readily degraded by microorganisms under significant formation of CO₂. The exact pathways of biogenic residue formation during microbial degradation of organic contaminants in soil are summarised in **Figure 31**. Depending on the yield coefficients of C conversion into biomass, we expect ratios of biomass plus biogenic residues to CO₂ of about 0.2 to 1. In the ¹³C₆-2,4-D study, the ratio was ~ 0.8 and in the ¹³C₆-ibu was 1.2. However, in case of highly reduced C-containing organic compounds, this ratio may be higher as it was seen in the ¹³C₆-ibu biodegradation experiment. However, the position of the label in the parent molecule requires consideration in terms of incorporation into biomass. For instance, due to the oxidation state of the C atoms in the triazine ring of atrazine they will not be incorporated into microbial biomass (STRUTHERS ET AL., 1998) but will be released as CO₂, which then might be assimilated by microorganisms. Similar effects are observed for the C at the 9 position of anthracene (KÄSTNER ET AL., 1999) or phenanthrene (RICHNOW ET AL., 2000). CO₂ fixation has also been reported to be a relevant process in soils (MILTNER ET AL., 2004), which also contributes to

NER formation, as it was shown in two experiments, in the $^{13}\text{CO}_2$ experiment with unlabelled 2,4-D and *C. necator* JMP 134 experiment, where $\sim 4\%$ of ^{13}C assimilated in the biomass was derived from $^{13}\text{CO}_2$.

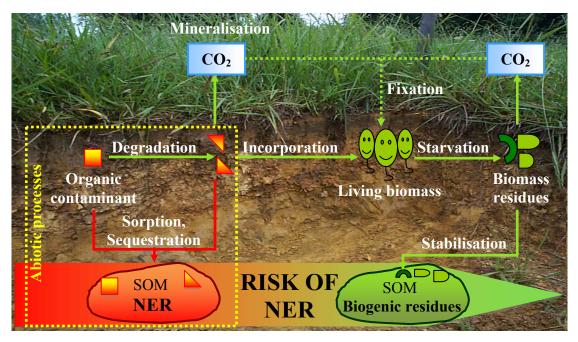


Figure 31. Scheme of the abiotic and the biotic NER formation during microbial degradation of organic contaminants in soil

The results from the two soil experiments indicate that biogenic residues formation can be relevant for many biodegradable pesticides and chemicals of environmental concern. These residues in soil contain only microbial components, which are natural products and explicitly excluded from the IUPAC definition of NER (ROBERTS, 1984). This is contrary to the generally accepted view that NER originating from the microbial degradation of organic contaminants consist mainly of the parent compounds or their metabolites, which are sorbed or sequestered within SOM components (BOLLAG ET AL., 1992; ALEXANDER, 2000; BARRACLOUGH, ET AL., 2005; BARRIUSO ET AL., 2008; see also abiotic processes in Figure 31). However, abiotic NER formation and biogenic residues formation are the competitive processes and do not occur together in a similar extent, because in the biotic treatment, the rapid mineralisation of an organic compound, reduces the extent of abiotic NER formation. The physico-chemical interactions of a contaminant with the SOM were reported mainly in sterile soils (PIGNATELLO AND XING, 1996; ALEXANDER, 2000; PALOMO AND BHANDARI, 2005; 2006), and were also observed in the abiotic treatments with either ${}^{13}C_{6}$ -2,4-D or ¹³C₆-ibu (GIRARDI ET AL., submitted). In addition, this abiotic NER formation can also occur, when a target compound is toxic for soil microorganisms, and thus inhibits the biotic NER formation.

The difficulties in proper identification of NER in soils are caused by the limitations of radiotracers, which only enable proper quantitative analyses. In addition, other factors may have led to the fact that the formation of biogenic residues has not been observed for decades. Firstly, AA in soil are generally bound in peptides or proteins, which are difficult to extract using organic solvents or water (even when harsh methods like ASE are used), in particular if they are stabilised in SOM. Secondly, native proteins in soil are surprisingly resistant to microbial degradation (RILLIG ET AL., 2007) and remain there for long period of time (KNICKER ET AL., 1993; KÖGEL-KNABNER, 2002), because they are stabilised in SOM by sorption to both organic and mineral surfaces (KLEBER ET AL., 2007).

Therefore, it is essential to distinguish the formation of xenobiotic-derived bound residues via various abiotic processes according to IUPAC from its non-toxic biogenic counterpart when assessing the risks of organic contaminants in soil. This differentiation is of utmost importance for assessing the risk of easily biodegradable active compounds, as shown by the results on the degradation of 2,4-D and Ibu in soil.

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